

University of Nevada, Reno

**Detrital zircon uranium-lead geochronology and hafnium-isotope analyses of  
passive margin and Roberts Mountains allochthon strata:  
Interpreting the Early Paleozoic tectonic evolution of western Laurentia**

A dissertation submitted in partial fulfillment of the requirements for the degree of  
Doctor of Philosophy in Geology

By  
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## ABSTRACT

This dissertation investigated Neoproterozoic–Devonian units of the western Laurentian passive margin and Roberts Mountains allochthon (RMA) and determined U-Pb detrital ages and Hf isotope zircon analyses that provide new insights into the early Paleozoic tectonics of western Laurentia. The three chapters investigate several difficult questions and contradictions in the understanding of early Paleozoic tectonism in western Laurentia through analysis of sedimentary units. The provenance, depositional histories, and tectonic evolution of the lower Paleozoic sedimentary strata of north-central Nevada have long been subjects of speculation and debate. Detrital zircon U-Pb geochronology and Hf-isotope analyses indicate the provenance, sedimentary distribution patterns, and tectonic evolution of Upper Neoproterozoic–Cambrian passive margin strata and Ordovician – Devonian strata of the RMA, with a special emphasis on the enigmatic Harmony Formation.

The study reported in Chapter 1 uses detrital zircon U-Pb geochronology to determine whether or not the Upper Neoproterozoic–Lower Cambrian Osgood Mountain Quartzite and the Upper Cambrian–Lower Ordovician Preble Formation in the Osgood Mountains of northern Nevada were units of the western Laurentian passive margin. Within the Osgood Mountain Quartzite, U-Pb age populations of the detrital zircons shift with stratal age. This shift indicates that the zircons were shed in different proportions from the source terranes, which suggests a change in provenance within the Osgood Mountain Quartzite. These changes are consistent across a Great Basin transect of coeval passive margin strata. The change in provenance is due to a shift in sedimentary transport patterns, which was caused by the Late Neoproterozoic–Early Cambrian uplift of the

Transcontinental Arch. This study provided independent corroboration of the existence of the Transcontinental Arch and better precision for the timing at which the Arch uplifted. The study also recorded the impact of the uplifted Arch on continent-wide sediment dispersal patterns—the change in predominant source terranes—and confirmed the Arch as a sediment source for passive-margin strata. Regional coeval changes in detrital zircon U-Pb age patterns provide a correlative tool in unfossiliferous sediments and could be useful in future studies.

Chapter 2 describes how detrital zircon U-Pb geochronology and Hf-isotope analyses were used to determine the provenance, sedimentary transport, and tectonic evolution of RMA strata. Workers have speculated for decades, with little agreement, on the origin, depositional basin(s), and subsequent tectonic transport of the RMA. Zircon grains from six Ordovician to Devonian arenite samples were analyzed for U-Pb ages; approximately one-quarter of these grains were further analyzed for Hf isotope ratios. Five of the studied units have similar U-Pb age populations and Hf-isotope ratios, while the U-Pb ages and Hf-ratios of the Ordovician lower Vinini Formation are significantly different. Comparison of these data with known analyses of igneous basement rocks and other sedimentary units of Laurentia reveals that the lower Vinini Formation originated in the north-central Laurentian craton. The other five units, as well as Ordovician passive margin sandstones of the western Laurentian margin, had a common source in the Peace River Arch region of western Canada. All of the RMA strata were deposited near the Peace River Arch region and subsequently tectonically transported south along the Laurentian margin, from where they were emplaced onto the craton during the Antler orogeny. This study determined the origin, location of the depositional basin, and

proposed a subsequent tectonic evolution that accounts for origin, deposition, and current location of the RMA strata.

Chapter 3 describes the origin, age, and tectonic development of the Harmony Formation. The Harmony Formation has always been difficult to explain—it is mostly an immature feldspathic arenite, which would argue for minimal transport from origin to deposition. However, its general position as the top thrust plate in the RMA stack argues for deposition oceanward of other more texturally mature RMA strata. The age of the Harmony Formation is equally contentious—published age determinations range from Cambrian to Mississippian. Zircon grains from ten arenite samples were analyzed for U-Pb ages; grains from eight of these samples were further analyzed for Hf-isotope ratios. Seven of the arenite units have similar U-Pb age peaks and Hf isotope ratios, whereas three differ significantly. The data confirmed the subdivision of the Harmony Formation into two petrofacies, quartzose (Harmony A) and feldspathic (Harmony B). Harmony A originated in the central Laurentian craton. Harmony B had a common source in eastern Alberta–western Saskatchewan, north of the source of the Harmony A. All of the Harmony Formation strata were deposited near eastern Alberta in Late Neoproterozoic through Cambrian time and subsequently tectonically interleaved with the Roberts Mountains allochthon strata. The entire package was tectonically transported south along the Laurentian margin. Subsequently, it was emplaced eastward onto the craton during the Late Devonian to Early Mississippian Antler orogeny. This study proposed a reasonable solution to one of the longest enduring and most puzzling conundrums of the western Cordillera—the origin, age, and transport of the Harmony Formation.

These three studies demonstrated the utility of detrital zircon U-Pb geochronology and Hf-isotope analyses in better understanding difficult sedimentary and tectonic problems. The studies also provided new insights into the Early Paleozoic tectonic evolution of western Laurentia.

## DEDICATION

I dedicate this dissertation to the glory of the Holy Trinity, Father, Son, and Holy Spirit, who make all such efforts possible.

And Jesus came and said to them, "All authority in heaven and on earth has been given to me. Go therefore and make disciples of all nations, baptizing them in the name of the Father and of the Son and of the Holy Spirit, teaching them to observe all that I have commanded you; and behold, I am with you always, to the close of the age."

Matthew 28:18-20 (Revised Standard Version)

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Interpreting the Early Paleozoic tectonic evolution of western Laurentia

Gwen M. Linde  
Ph.D. dissertation

**Introduction**

The sedimentary units of the Lower Paleozoic passive margin and Roberts Mountains allochthon in north-central Nevada are well studied, but their provenance, depositional histories, and tectonic evolution have long been subjects of speculation and debate (e.g., Schuchert, 1923; Kay, 1951; Roberts et al., 1958; Hotz and Willden, 1964; Burchfiel and Davis, 1972; Speed and Sleep, 1982; Madrid, 1987; Burchfiel et al., 1992; Poole et al., 1992; Gehrels et al., 2000). This dissertation uses detrital zircon uranium-lead geochronology and hafnium-isotope analyses to investigate the origin, sedimentary distribution patterns, and tectonic evolution of the Upper Neoproterozoic – Cambrian passive margin strata and Ordovician – Devonian strata of the Roberts Mountains allochthon, with a special emphasis on the enigmatic Harmony Formation.

Early Paleozoic time along the western Laurentian margin has commonly been interpreted as a quiescent interval (e.g., Poole et al., 1992; Dickinson, 2009, and references cited therein). The final Neoproterozoic rifting that separated the Rodinian supercontinent lasted from ca. 570-520 Ma and was followed by a drift phase and the deposition of passive-margin sediments through mid-Devonian time (Poole et al., 1992; Dickinson, 2009; Yonkee et al., 2014). The quiescent interval came to an end with the Antler orogeny, during which the Roberts Mountains allochthon (RMA) was emplaced

onto the western Laurentian margin (Dickinson, 2006). The RMA is an internally disrupted package of Cambrian – Devonian oceanic sediments, primarily composed of chert, argillite, quartzose turbidites, and some pillow basalts and volcanogenic debris flows (Doebrich, 1994; Dickinson, 2006; Dickinson, 2009).

Chapter 1 tests the hypothesis that the uplift of the Transcontinental Arch can be discerned from changes in provenance of western Laurentian passive margin strata, which in turn reflect changing drainage patterns that resulted from the uplift of the arch. Sedimentary strata of the Lower Paleozoic passive margin were investigated; these strata are mostly quartzite, with some siltstone, argillite, and phyllite and carbonate intervals (e.g., Stewart, 1991; Poole et al., 1992). These units have been correlated across a broad region of western North America (e.g., Poole et al., 1992). Detrital zircon U-Pb geochronology of the Upper Neoproterozoic – Lower Cambrian Osgood Mountain Quartzite and the Upper Cambrian – Lower Ordovician Preble Formation in the Osgood Mountains of northern Nevada records a provenance change within the Osgood Mountain Quartzite. Detrital zircons from the samples of the Osgood Mountain Quartzite collected lower in the stratigraphic section are predominantly Mesoproterozoic. Detrital zircons from the samples collected higher in the stratigraphic section and all of the Preble Formation samples are predominantly Upper Mesoproterozoic – Paleoproterozoic. Comparison of these data with previous work (Lawton et al., 2010; Gehrels and Pecha, 2014; Yonkee et al., 2014) reveals that the same change in U-Pb age populations and provenance occurred in correlative strata throughout an east–west transect of the Great Basin. This regional shift in provenance records the uplift of the Transcontinental Arch by Early Cambrian time. The uplifted arch forced a change in sedimentation and

drainage patterns, restricting the transport of sediment from the eastern Laurentian to western Laurentia craton. The uplift of the arch also exposed basement rocks and provided new sources for western Laurentian passive-margin sediments.

Sedimentary units of the Roberts Mountains allochthon were investigated to determine the provenance, sedimentary transport, and tectonic history of the strata (Chapter 2). The RMA consists of internally deformed Cambrian through Devonian rocks that structurally overlie coeval passive margin strata in northeastern and north-central Nevada (Schuchert, 1923; Kay, 1951; Roberts et al., 1958; Madrid, 1987; Burchfiel et al., 1992). RMA rocks include chert, argillite, arenite, quartzite, limestone, and mafic volcanic rocks. The RMA is often thought to have been deposited in an ocean basin outboard of coeval passive-margin strata in western Laurentia and to have been tectonically emplaced onto the passive margin during the Late Devonian to Early Mississippian Antler orogeny (e.g., Roberts et al., 1958; Burchfiel and Davis, 1972; Madrid, 1987).

Zircon grains from six Ordovician to Devonian arenite samples were analyzed for U-Pb ages; these grains were further analyzed for Hf-isotope ratios. Five of the units sampled have similar U-Pb age peaks and Hf-isotope ratios, whereas the data from the Ordovician lower Vinini Formation are significantly different. Comparison of these data with previous U-Pb and Hf analyses of igneous basement rocks and other coeval sedimentary units reveals that the lower Vinini Formation originated in the north-central Laurentian craton (Bickford et al., 1986; Hoffman, 1989; Ross, 1991; Anderson and Morrison, 1992; Bickford and Anderson, 1993; Van Schmus et al., 1993; Lund et al., 2010; Todt and Link, 2013). The other five units sampled, as well as Ordovician passive margin sandstones of the western Laurentian margin, have detrital-zircon ages similar to

the timing of magmatic activity and terrane accretion in the Peace River Arch region of western Canada (Hoffman, 1989; Ross, 1991; Villeneuve et al., 1993). All of the RMA strata were likely deposited near the Peace River Arch region, and subsequently tectonically transported south along the western Laurentian margin to where they were emplaced onto the craton during the Antler orogeny.

Chapter 3 focuses on a similar study but focused on the Harmony Formation (Chapter 3). The Harmony Formation is primarily a texturally immature feldspathic arenite. It has been interpreted as a turbidite deposit (Roberts, 1964; Suczek, 1979) and is often interpreted as a part of the RMA (Hotz and Willden, 1964; Suczek, 1979; Madrid, 1987). The age of the Harmony Formation has long been controversial. When first mapped, the unit was interpreted as no younger than Mississippian (Ferguson et al., 1951). The Harmony Formation was later interpreted as Cambrian, based on trilobite fauna (Hotz and Willden, 1964). Jones (1997a; 1997b) interpreted the Harmony Formation as Devonian, based on a single conodont.

Zircon grains from ten arenite samples were analyzed for U-Pb ages and Hf-isotope ratios. Seven of the arenite units sampled have similar U-Pb age peaks and Hf-isotope ratios, while the remaining three samples differ significantly. The data confirm the previous subdivision of the Harmony Formation into two petrofacies, quartzose (Harmony A) and feldspathic (Harmony B) (Gehrels et al., 2000). Comparison of these data with U-Pb and Hf analyses of igneous basement rocks and other sedimentary units reveals that Harmony A samples originated in the central Laurentian craton (Gehrels and Pecha, 2014; Yonkee et al., 2014; Linde et al., 2014). In comparison, Harmony B had a common source in eastern Alberta – western Saskatchewan, north of the source of the

Harmony A (Collerson et al., 1988; Villeneuve et al., 1993; Gehrels and Ross, 1998; Gehrels and Pecha, 2014; Peterson et al., 2015). Harmony Formation strata were transported from sources in central Laurentia or eastern Alberta – western Saskatchewan, deposited near eastern Alberta in Late Neoproterozoic through Cambrian time, and subsequently tectonically interleaved with the RMA strata. The entire package was tectonically transported south along the Laurentian margin and was subsequently emplaced eastward onto the craton during the Late Devonian to Early Mississippian Antler orogeny.

These three studies interpreted the provenance of passive margin and Roberts Mountains allochthon strata and thus provided new insights into the Early Paleozoic sedimentary dispersal patterns and tectonic evolution of western Laurentia. These studies also demonstrated the utility of detrital zircon U-Pb geochronology and Hf-isotope analyses to better understand difficult sedimentary and tectonic problems concerning the tectonic evolution of western Laurentia.

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## **Chapter 1**

### **Stratigraphic trends in detrital zircon geochronology of Upper Neoproterozoic and Cambrian strata, Osgood Mountains, Nevada and elsewhere in the Cordilleran miogeocline: Evidence for Early Cambrian uplift of the Transcontinental Arch**

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#### **1. Abstract**

U-Pb detrital zircon geochronology provides insight into the provenance of the Upper Neoproterozoic-Lower Cambrian Osgood Mountain Quartzite and the Upper Cambrian – Lower Ordovician Preble Formation in the Osgood Mountains of northern Nevada. A total of 535 detrital zircon grains from six samples of quartz arenite were analyzed by laser-ablation-multicollector-inductively coupled plasma-mass spectrometry (LA-MC-ICP-MS). The detrital zircon age data of these Neoproterozoic-Lower Paleozoic passive margin units record a provenance change within the Osgood Mountain Quartzite. Comparison of these data with the work of others reveals that this change in provenance occurred in correlative strata throughout an East-West transect of the Great Basin. From latest Neoproterozoic through earliest Cambrian time, most grains were shed from the

1.0-1.2 Ga Grenville orogen. After that time, drainage patterns changed and most grains were derived from the 1.6-1.8 Ga Yavapai and Mazatzal provinces; very few grains from the Grenville orogen were found in the younger strata. We suggest that this shift records the uplift, in Early Cambrian time, of the Transcontinental Arch. Our data also support our interpretation that the Osgood Mountain Quartzite and the Preble Formation are correlative to other contemporaneous passive margin strata in western Laurentia.

## **2. Introduction**

The Transcontinental Arch, a region of uplift that extends from the southwestern U.S. to south-central Ontario, Canada (Fig. 1), is a fundamental feature of the Lower Paleozoic Laurentian craton. It was first recognized from broad structures and Phanerozoic sedimentation patterns in the mid-continent (Fig. 1) (Keith, 1928). Sloss (1963; 1988) noted the deposition of the Middle and lowermost Upper Cambrian Sauk II sequence onlapping from the craton margin onto the Transcontinental Arch (Fig. 1). Carlson (1999) proposed, instead of a discrete arch, a platform—a discontinuous zone with highs and lows and flanking basins that give the appearance of an arch (Fig. 1).

In recent U-Pb detrital zircon geochronology studies, researchers have proposed the Transcontinental Arch as a barrier to sediment delivery from the central Laurentian craton to its western margin in early Paleozoic time (Amato and Mack, 2012; Gehrels and Pecha, 2014; Yonkee et al., 2014). Amato and Mack (2012) document evidence from the Bliss Sandstone for the existence of the Transcontinental Arch by at least the Late Cambrian. They explain the differences in detrital zircon populations between the Tapeats Sandstone west of the arch and the Cambrian sandstones east of the arch, with

the uplift of the arch possibly as early as Early Cambrian time (Amato and Mack, 2012).

Gehrels and Pecha (2014) estimated the uplift of the arch by Early Cambrian time.

Others note the possibility of Early Cambrian uplift of the arch as the cause of the differences in detrital zircon age peaks and groups in passive margin strata in Utah (Yonkee et al., 2014).

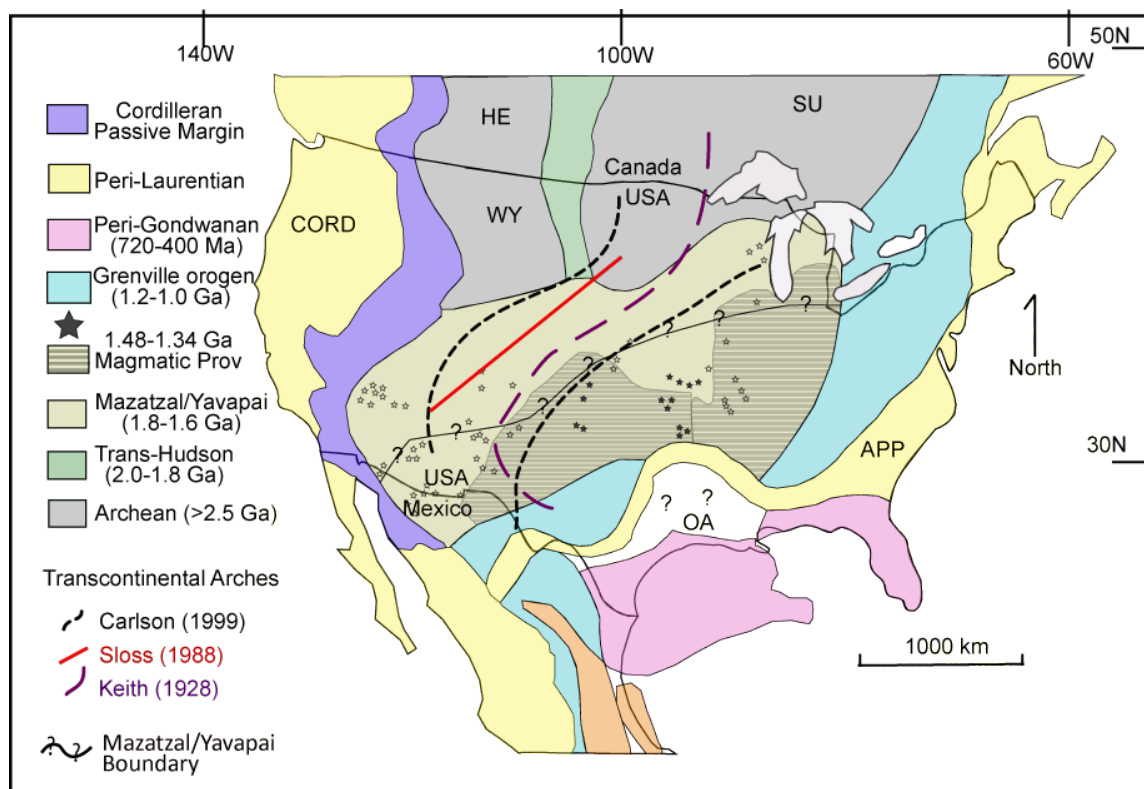


Figure 1: Location of the main age provinces in North America that are potential source terranes for the Late Neoproterozoic-Early Cambrian western Laurentian passive margin. Hypothesized Transcontinental Arches are superimposed (Keith, 1928; Sloss, 1988; Carlson, 1999). WY: Wyoming province; HE: Hearn province; SU: Superior province; CORD: Cordilleran; APP: Appalachian; OA: Ouachita-Marathon. Figure is after Gehrels et al. (2011) and adapted from Anderson and Morrison (1992), Bickford et al. (1986), Hoffman (1989), Burchfiel et al. (1992), Bickford and Anderson (1993), Van Schmus et al. (1993), Dickinson and Lawton (2001), and Dickinson and Gehrels (2009).

Upper Neoproterozoic – Lower Cambrian siliciclastic rocks on the western Laurentian passive margin record sedimentation that initiated after rifting and continental separation (e.g., Stewart, 1972; Poole et al., 1992). These passive margin rocks were deposited on a discontinuously exposed succession of diamictite and volcanic strata that reflect initial rifting (e.g., Poole et al., 1992; Yonkee et al., 2014, and references cited therein). These strata are mostly quartzite, with some siltstone, argillite, and phyllite; carbonate intervals are present in some locations (e.g., Stewart, 1991; Poole et al., 1992). These units have been correlated across a broad region of western North America (e.g., Poole et al., 1992).

Previous detrital zircon studies of Upper Neoproterozoic – Lower Paleozoic passive margin strata record similar changes in detrital zircon age peaks and groups and therefore possibly similar changes in provenance. Zircon ages in Upper Neoproterozoic – Lower Cambrian strata in Utah (Lawton et al., 2010; Gehrels and Pecha, 2014; Yonkee et al., 2014), Idaho (Yonkee et al., 2014), and Nevada (Gehrels and Pecha, 2014; Yonkee et al., 2014) change from predominantly Mesoproterozoic in the older strata to Upper Mesoproterozoic – Paleoproterozoic in the younger strata.

The only previous detrital zircon study of the Osgood Mountain Quartzite was Gehrels and Dickinson (1995); they sampled from the upper part of the formation. The Preble Formation has never been the subject of a published detrital zircon study.

We dated detrital zircons from three localities of the Upper Neoproterozoic – Lower Cambrian Osgood Mountain Quartzite and three localities of the Upper Cambrian – Lower Ordovician Preble Formation in the Osgood Mountains and near Edna Mountain, north-central Nevada (Fig. 2b and Table 1). We show that the detrital zircon ages shift within the Osgood Mountain Quartzite; detrital zircons from the older samples are

predominantly Mesoproterozoic, while detrital zircons from the younger sample, and all of the Preble Formation samples, are predominantly Upper Mesoproterozoic – Paleoproterozoic. Coeval passive margin strata in other studies throughout the Great Basin (e.g., Lawton et al., 2010; Gehrels and Pecha, 2014; Yonkee et al., 2014) show the same shift in ages. This suggests that a change in provenance in these passive margin strata is widely recorded in this region of western Laurentia.

In this paper, we present new U-Pb zircon ages from the Osgood Mountain Quartzite and the Preble Formation. The dates were obtained using laser-ablation-multicollector-inductively coupled plasma-mass spectrometry (LA-MC-ICP-MS). We compare these new data with the detrital zircon ages of coeval passive margin strata throughout the Great Basin to evaluate provenance and sediment transport patterns, and the possibility that these patterns were altered by the uplift of the Transcontinental Arch.

Several unsolved problems are addressed in this study. (1) What is the provenance of the Osgood Mountain Quartzite and the Preble Formation? (2) Are the Osgood Mountain Quartzite and Preble Formation passive margin units, as others have interpreted? (3) Within the Osgood Mountain Quartzite, there is a significant change in detrital zircon grain ages. Are there similar patterns of detrital zircon grain ages varying with time among other coeval passive margin units? (4) If a consistent stratigraphic pattern of detrital zircon ages exists in all Neoproterozoic-Cambrian sections, what caused a widespread change in detrital zircon ages with time in these units?

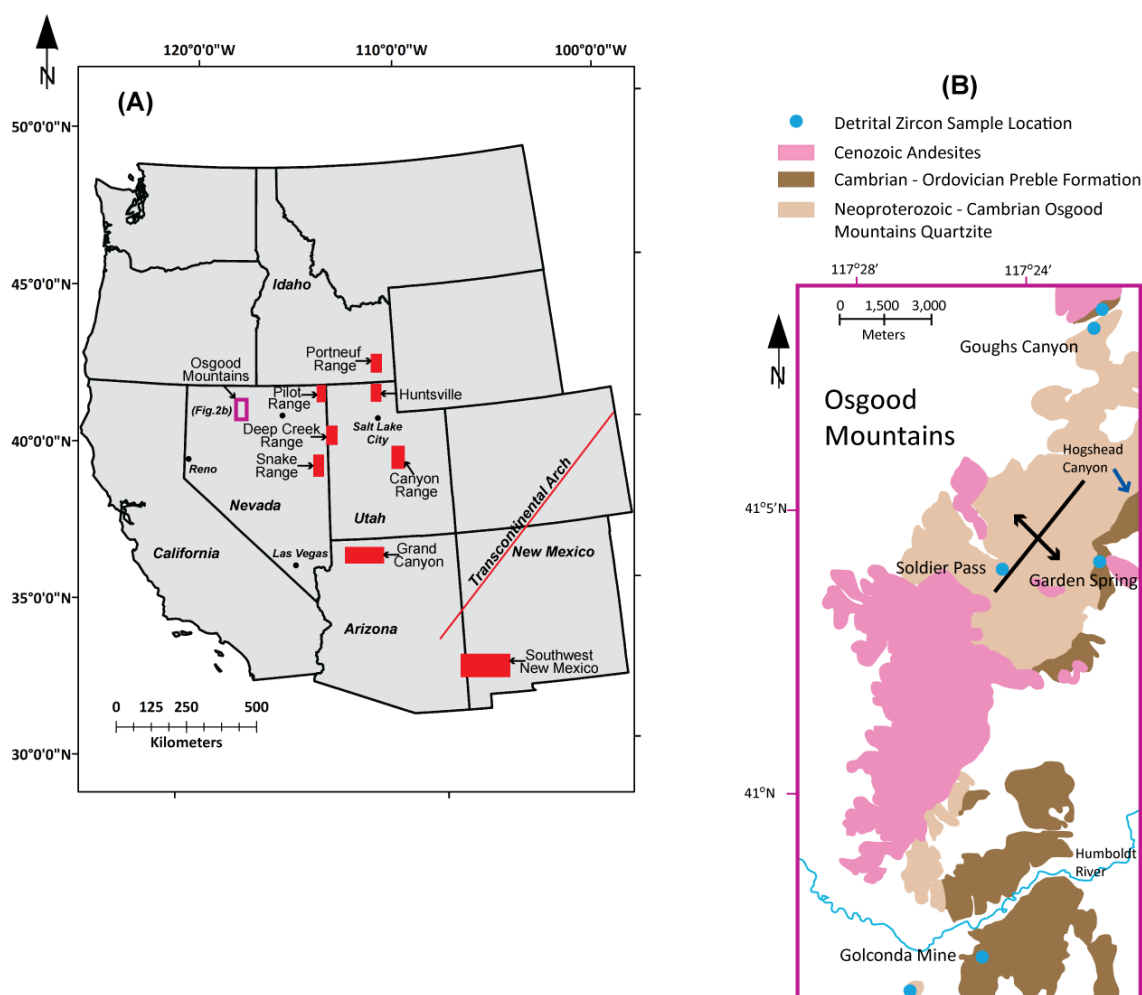


Figure 2. (a) Location of study areas in the vicinity of the Osgood Mountains, and location of the study areas within the Great Basin region. Transcontinental Arch is after Sloss (1988). (b) Geologic map of the Osgood Mountains. The six sample locations are shown. Broad northeast-trending anticline of the range is shown with location of Soldier Pass sample in the core of the anticline and Preble Formation on both flanks of the anticline. Hogshead Canyon, location of unit thickness estimate of the Preble Formation, is shown. Northern part of the map is after Hotz and Willden, 1964. Southern part is after Erickson and Marsh, 1974.

Cambrian - Ordovician Preble Formation

	<u>Sample</u>	<u>Easting</u>	<u>Northing</u>
Golconda Mine	GOL-02-CP	0465271	4533874
Garden Spring	GS-01-CP	0470302	4548660
Gough's Canyon	GC-01-CP	0468513	4554509

Neoproterozoic - Cambrian Osgood Mountains Quartzite

	<u>Sample</u>	<u>Easting</u>	<u>Northing</u>
Golconda Mine	GOL-01-COM	0462354	4531791
Goughs Canyon	GC-03-COM	0468951	4554410
Soldier Pass	SP-01-COM	0463750	4548029

NAD 83 UTM 11N

Table 1. Locations and sample numbers of samples referenced to UTM locations.

### 3. Geologic Setting

The North American craton contains several Proterozoic and Archean age provinces, thus providing geographically distinguishable crustal provinces that are source terranes for the Upper Proterozoic and Lower Paleozoic continental margin sedimentary section (e.g., Gehrels et al., 2011 and references cited therein) (Fig. 1). The Yavapai/Mazatzal Province (1.8-1.6 Ga) forms the core of the craton in the U.S. (Fig.1). It is bounded on the northwest by the Trans-Hudson orogenic terrane (2.0-1.8 Ga) and Archean rocks (> 2.5 Ga) of the Wyoming and Superior Provinces (Fig. 1). It is bounded on the east and southeast by the terranes of the Grenville orogen (1.2-1.0 Ga) (Fig. 1).



The 1.2-1.0 Ga Grenville orogen of southern and eastern North America (Fig. 1) was the dominant sediment source for western Laurentia throughout the Neoproterozoic (Rainbird et al., 1997; 2012), including the Upper Proterozoic passive margin section from the northwest U.S. to Sonora, Mexico (e.g., Lawton et al., 2010; Gehrels and Pecha, 2014; Yonkee et al., 2014). In contrast, the 1.8-1.6 Ga Yavapai-Mazatzal and 1.48-1.34 Ga mid-continent granite rhyolite provinces within the North America craton (Fig. 1) were the dominant sediment sources higher in the passive margin section (e.g., Lawton et al., 2010; Gehrels and Pecha, 2014, Yonkee et al., 2014).

In northern Nevada, the Osgood Mountain Quartzite and Preble Formation have been interpreted as passive margin strata. The Osgood Mountain Quartzite and Preble Formation have an interesting position and tectonic and metamorphic history. They are far to the west of most other passive margin units and are overthrust by both the Roberts Mountains allochthon and the Golconda allochthon (Burchfiel et al., 1992; Poole et al., 1992). The Preble Formation is metamorphosed to greenschist facies, and has refolded folds (Cashman et al., 2011). The Preble Formation has been interpreted as being in conformable stratigraphic succession with the Osgood Mountain Quartzite (Fig. 3), based on map relationships and compositional similarity of an upper member of the Osgood Mountain Quartzite to the Preble Formation (Hotz and Willden, 1964).

Structurally, the Osgood Mountains comprise a large, northeast-trending anticline with a sub-horizontal axis (Fig. 2b). The Preble Formation is exposed only on the flanks of the anticline (Fig. 2b). Late Paleozoic rocks are thrust over the anticline in the northern and western parts of the range, and the southern extent of the Osgood Mountains is overlain by Cenozoic andesite flows (Fig. 2b) (Hotz and Willden, 1964). The Osgood

Mountain Quartzite and Preble Formation are primarily exposed in the central and southern portions of the range (Fig. 2b).

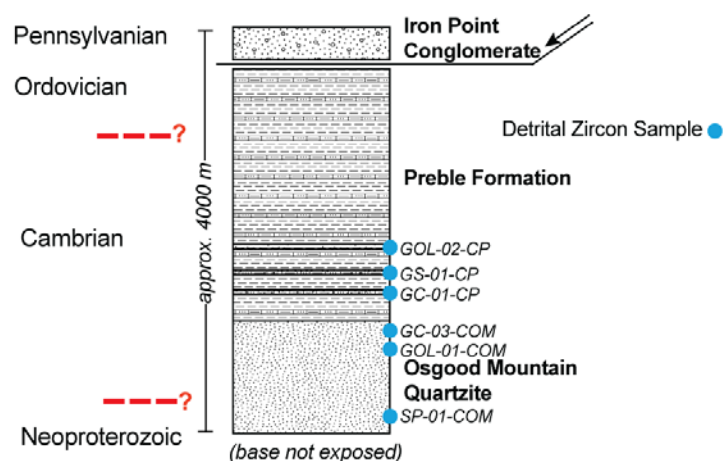


Figure 3: Stratigraphic column of the Osgood Mountain Quartzite and Preble Formation. Subjacent strata are not shown. Red dashed lines are approximate system boundaries. The fault is the Iron Point fault, a mid-Pennsylvanian normal fault. Structural relationships are after Cashman et al. (2011); unit ages are after Hotz and Willden (1964) and Madden-McGuire (1991).

The Osgood Mountain Quartzite consists mostly of fine- to medium-grained quartz arenite, with some silty and shaly beds (Fig. 3). The formation crops out in the Osgood Mountains in an outcrop belt 14 km long, from the northwestern Osgood Mountains near Goughs Canyon, to the Golconda Mine area in the northwestern part of Edna Mountain (Fig. 2b). Goughs Canyon and the Golconda Mine, where two samples were taken, are on opposite flanks of an anticline and are relatively high in the stratigraphic section (Figs. 2b and 3). Soldier Pass, where a third sample was collected, is closer to the core of the anticline and is thus lower in the stratigraphic section (Figs. 2b and 3). The base of the

Osgood Mountain Quartzite is not exposed; the thickness has been estimated at more than 1524 m (5000 ft) (Hotz and Willden, 1964). The upper part of the Osgood Mountain Quartzite is the Twin Canyon Member, which crops out only on the east side of the range, and consists of more silty and shaly material than the rest of the formation. This member has been interpreted as a transition between the Osgood Mountain Quartzite and the overlying Preble Formation (Hotz and Willden, 1964). The Osgood Mountain Quartzite is barren of fossils (Hotz and Willden, 1964). Based on the age of the overlying Preble Formation, the Osgood Mountain Quartzite is Late Neoproterozoic to Early Cambrian in age (Madden-McGuire, 1991).

The Preble Formation consists of phyllite and shale, interbedded limestone, and quartz arenite. The Preble Formation crops out over an area approximately 50 km in length, from the northwestern Osgood Mountain near Goughs Canyon, south to the Sonoma Range. The thickness of the Preble Formation has been estimated at approximately 2350 m (7700 ft) (Ferguson et al., 1952) based upon the estimated thicknesses of the subunits, though the same workers suggested that the structural thickness may exceed 4572m (15,000 ft), due to isoclinal folding. The thickness was estimated at approximately 1524 m (5000 ft) near Hogshead Canyon (Fig. 2b), where both upper and lower contacts are faults (Hotz and Willden, 1964). However, these authors noted that tight folding and lack of distinctive bedding precluded them from making detailed studies of the thickness and stratigraphy of the unit. Based on middle Early Cambrian trilobite fauna collected in the lower part of the Preble Formation, the base of the unit is of Early Cambrian age (Madden-McGuire, 1991). The fossils occur approximately 400 m above the upper contact of the pure quartz arenite Osgood Mountain Quartzite, consistent with a Late

Neoproterozoic age for most of the Osgood Mountain Quartzite (Madden-McGuire, 1991). Graptolites near the youngest subunit of the Preble Formation indicate that the top of the unit is Early Ordovician in age (Madden-McGuire, 1991).

#### **4. Methods**

Quartz arenite samples were collected from six locations and stratigraphic intervals (Figs. 2b and 3 and Table 1). Three samples were analyzed from the Upper Neoproterozoic – Lower Cambrian Osgood Mountain Quartzite and three samples were analyzed from the Lower Cambrian – Lower Ordovician Preble Formation.

Zircon grains were separated and analyzed at the University of Arizona LaserChron facility using standard techniques described by Gehrels (2000, 2011), Gehrels et al. (2006, 2008) and Johnston et al. (2009), to yield a best age distribution reflective of the true distribution of detrital zircon ages in each sample. A split of zircons representative of the final sample yield was mounted in a 1” (2.54 cm) diameter epoxy plug, with the laboratory’s SL (Sri Lanka) zircon standard ( $563.5 \pm 3.2$  Ma, Gehrels et al., 2008). Approximately 100 randomly selected grains were analyzed for each sample. Analyses were conducted by LA-MC-ICP-MS using the New Wave UP193HE laser connected to the Nu Plasma high resolution inductively coupled plasma-mass spectrometry (NU HR ICP-MS).

Analytical results are displayed graphically on normalized probability plots (Figs. 4 and 5). On our plots, we did not include analyses with greater than 10% uncertainty in age. We also discarded analyses with greater than 30% discordance (70% concordance) and greater than 5% reverse discordance (105% concordance). Normalized probability

plots (Figs. 4 and 5) allow visual comparison between zircon populations and display the data from this study and the research of others. The normalized probability plots are generated by summing the ages and uncertainties and normalizing the graphs so that all curves on the same plot have the same area under the curve.

We compared detrital zircon age distributions both visually and statistically. Our initial appraisal was visual comparison of the probability plots. We also compared many age distributions using the Kolmogorov-Smirnov (K-S) statistic (Guynn and Gehrels, 2006). The K-S statistic estimates the probability (P value) that two sample populations could have been derived from the same parent population.  $P > .05$  indicates >95% probability that two U-Pb distributions are not statistically different and could have been derived from the same parent ( $P=1.0$  reflecting effective statistical identity). The K-S statistic is sensitive to proportions of ages present, and a low P value may indicate that the proportions of ages are different, even though the ages are similar (Gehrels, 2012).

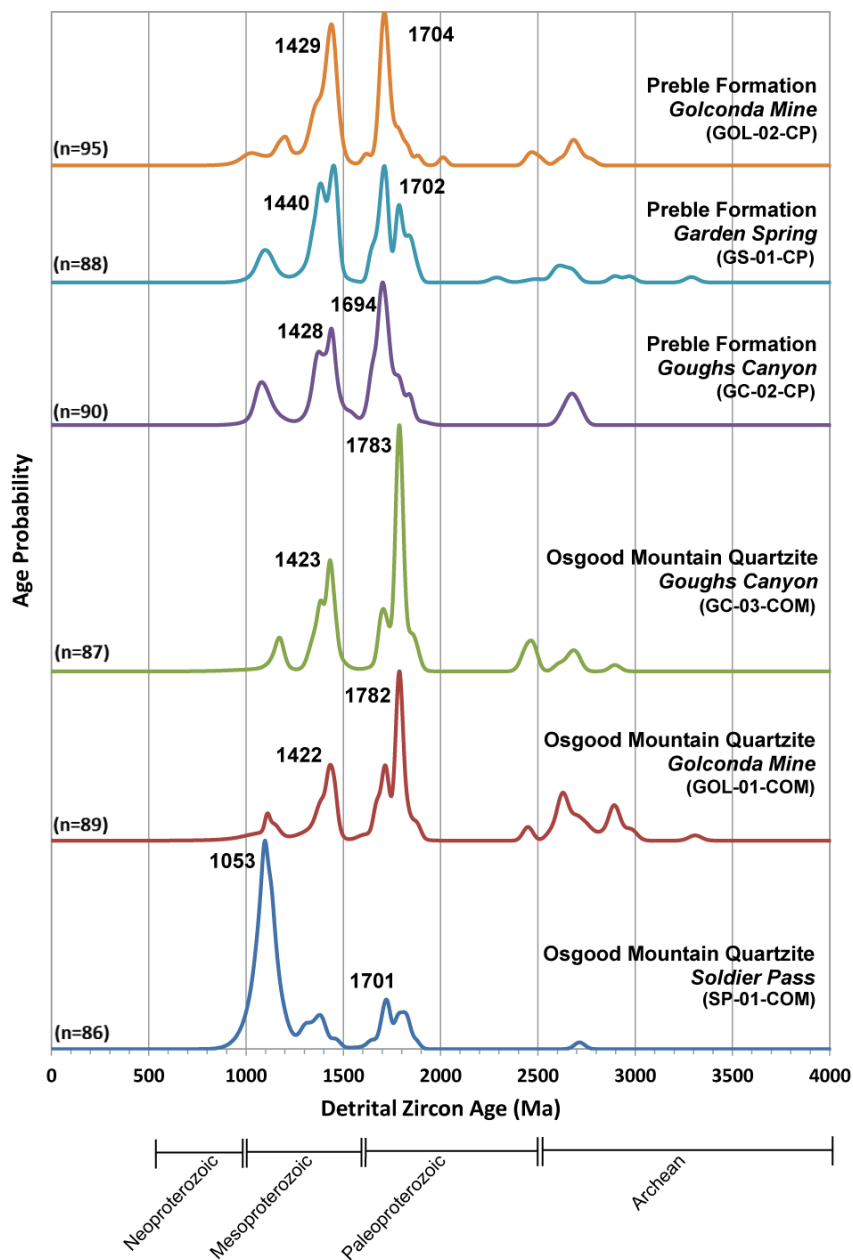


Figure 4: Normalized probability plots of units sampled in and around the Osgood Mountains. The curves contain all analyses for each unit and are normalized such that the area is the same under each curve. Within each location the older units are on the bottom. Horizontal axis is age in millions of years. Number of concordant detrital zircons in each sample is shown in parentheses.

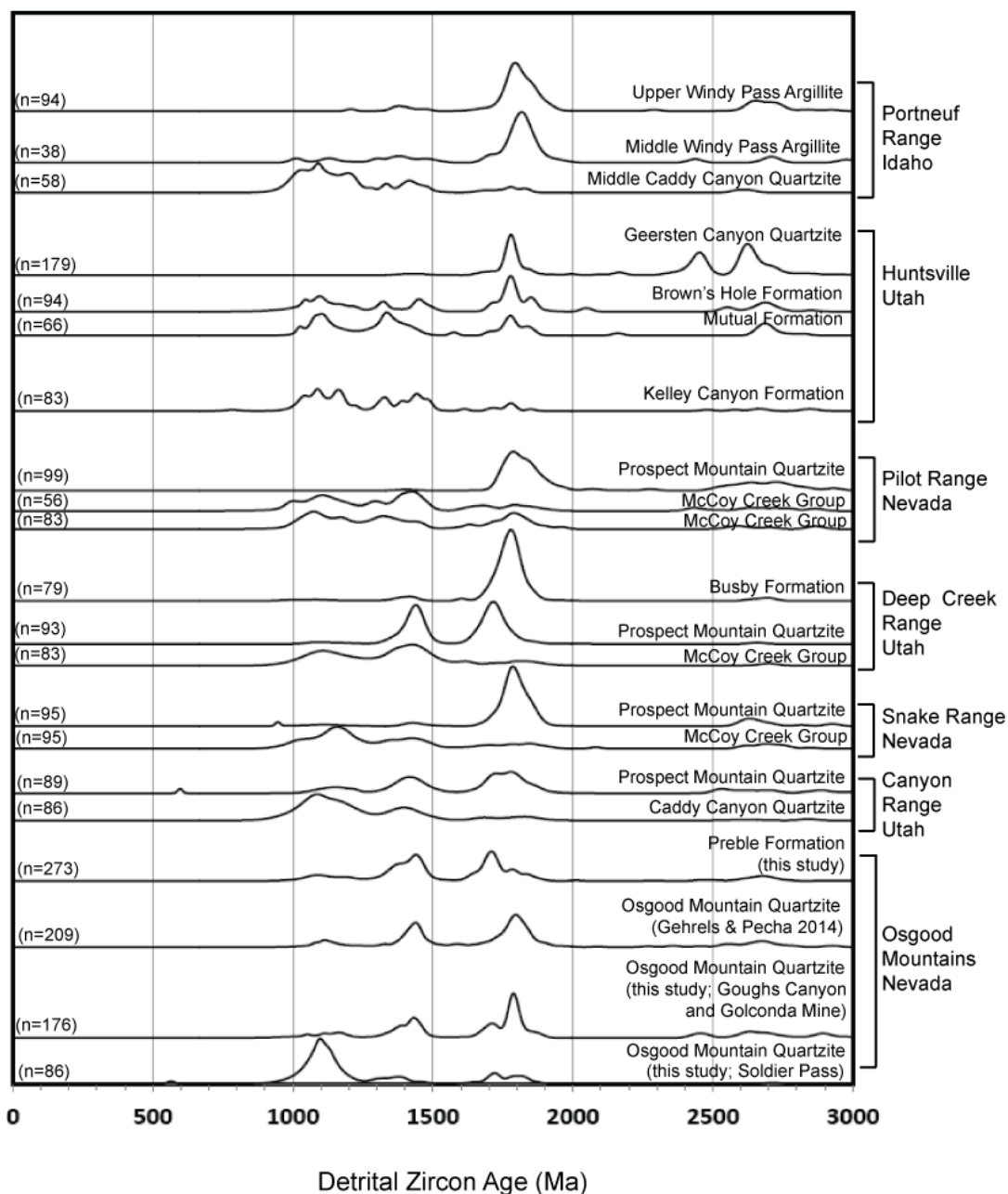


Figure 5. Plots of detrital zircon age of each Neoproterozoic - Cambrian unit organized by locality (see map, Fig. 1) showing the distribution of detrital zircon ages. Curves are normalized probability plots. The number of detrital zircon grains comprising each analysis is shown on the left. Within each location the older units are on the bottom. Canyon Range: Lawton et al., 2010; Snake, Deep Creek, and Pilot Ranges: Yonkee et al., 2014; Huntsville: Yonkee et al., 2014; Portneuf Range: Yonkee et al., 2014; Osgood Mountains: Linde et al., 2012; Gehrels and Pecha, 2014.

## **5. Detrital Zircon Geochronology Results**

### **a. Osgood Mountain Quartzite (Upper Neoproterozoic – Lower Cambrian)**

Two samples from near the top of the Osgood Mountain Quartzite, collected at Goughs Canyon and Golconda Mine (Figs. 2b and 3 and Table 1), and one sample from approximately 400 meters below the top of the unit, collected at Soldier Pass (Figs. 2b and 3 and Table 1), were analyzed. Our visual and statistical analyses indicate that the Mesoproterozoic and Paleoproterozoic detrital zircon age groups of the two younger samples are similar to one another, though the proportions of ages are somewhat different (see Appendix A for statistical analysis discussion). The detrital zircon age groups from these two samples are quite different; the older sample is dominated by Mesoproterozoic grains, while the younger sample is dominated by Paleoproterozoic grains (Fig. 4). K-S test results confirm that these detrital zircon grain populations are dissimilar: P is less than 0.05.

### **b. Preble Formation (Lower Cambrian – Lower Ordovician)**

The three Preble Formation samples were collected low in the formation (Fig. 3) at Goughs Canyon, Garden Spring and Golconda Mine (Fig. 2b and Table 1). The samples all contained Paleoproterozoic and Mesoproterozoic age peaks (Fig. 4). Applying the K-S test, we found that within the Preble Formation, the three different sample pairs have P values of 0.849, 0.881, and 0.937. All sample pairs of the Preble Formation and Osgood Mountain Quartzite have P values less than 0.05.



## **6. Discussion**

### **a. Osgood Mountain Quartzite**

The source of the detrital zircon grains in the Osgood Mountain Quartzite samples is Laurentian, and the samples have two distinct provenances. The change in detrital zircon ages from the older Soldier Pass sample to the younger Goughs Canyon and Golconda Mine samples indicates a significant change in provenance during the deposition of the Osgood Mountain Quartzite. The grains from Soldier Pass are predominantly Mesoproterozoic (Fig. 4) and we interpret that their source was the 1.2-1.0 Ga Grenville orogen (Fig. 1). There are also some Paleoproterozoic grains which we interpret to have been derived from the 1.8–1.7 Ga Yavapai province and the 2.0–1.8 Ga Trans-Hudson orogen (Fig. 1). There are a few Mesoproterozoic grains; these are interpreted to have been shed from the 1.48-1.34 Ga mid-continent granite-rhyolite provinces (Fig. 1). In contrast, the detrital zircon ages in the Goughs Canyon and Golconda Mine samples (Fig. 4) share age peaks with one another, and these peaks are different from those of the Soldier Pass sample (Fig. 4). The Paleoproterozoic grains in these samples are interpreted to have been shed from the 1.8-1.7 Ga Yavapai Province and the 1.7-1.62 Ga Mazatzal Province (Fig. 1). Mesoproterozoic grains are interpreted to have been shed from the 1.48-1.34 Ga mid-continent granite-rhyolite provinces (Fig. 1). Archean grains are interpreted to have been sourced from the Archean craton (Fig. 1). A few Mesoproterozoic grains are interpreted to have been derived from the 1.2-1.0 Ga Grenville orogen (Fig. 1).

### **b. Preble Formation**

The detrital zircon ages in all three Preble Formation samples are similar (Fig. 4). We interpret that these three Preble Formation samples share a common Laurentian source. Mesoproterozoic grains predominate in all three samples (Fig. 4) and are interpreted to have been shed from the 1.48-1.34 Ga mid-continent granite-rhyolite provinces (Fig. 1). There are a large number of Paleoproterozoic grains which are interpreted to have been derived from the 1.7-1.62 Ga Mazatzal Province and the 1.8-1.7 Ga Yavapai Province (Fig 1). A smaller number of Mesoproterozoic grains are interpreted as having their source in the 1.2-1.0 Ga Grenville orogen (Fig. 1). The remaining grains are Archean and their source is interpreted as the Archean craton (Fig. 1).

## **7. Regional Correlation**

Detrital zircon ages in passive margin strata across a transect of the Great Basin vary stratigraphically in a manner similar to that which we documented within the Osgood Mountain Quartzite, recording a major regional change in provenance.

### **a. Nevada-Utah Border: Deep Creek Range, Pilot Range, and Snake Range**

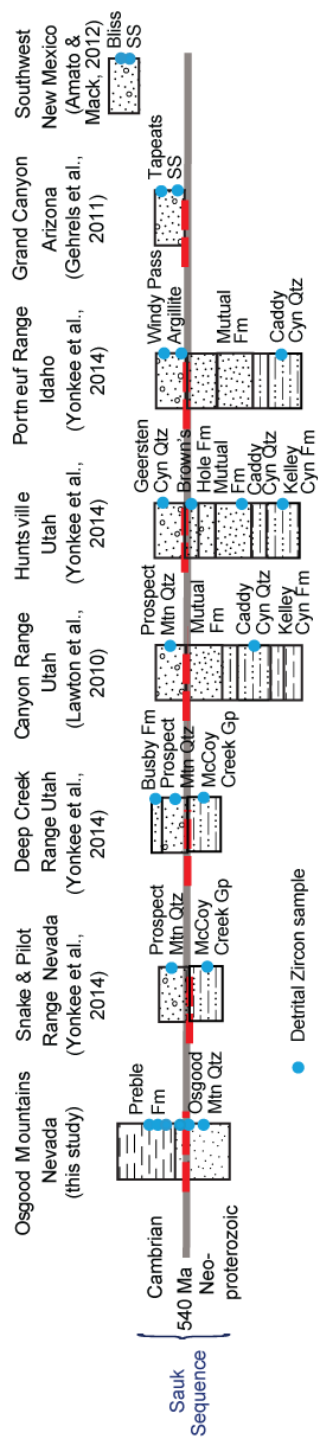
Detrital zircons from rocks in the Deep Creek, Pilot, and Snake Ranges (Fig. 2a) record a similar shift of ages as detrital zircons analyzed in the Osgood Mountains. The older strata are the Neoproterozoic McCoy Creek Group (Fig. 6). In all three ranges, the McCoy Creek group has similar Mesoproterozoic age groups and peaks (Fig. 5) (Yonkee et al., 2014). We interpret these grains as primarily shed from the 1.2-1.0 Ga Grenville orogen and the 1.48-1.34 Ga mid-continent granite-rhyolite province (Fig. 1). This is

quite similar to the older Osgood Mountain Quartzite (Soldier Pass) sample, which we interpreted to be sourced in these same terranes. The younger strata in all three ranges are the Cambrian Prospect Mountain Quartzite, and in the Deep Creek Range, the Cambrian Busby Formation (Fig. 6). In all three ranges, the detrital zircons in these younger strata have similar Paleoproterozoic and Mesoproterozoic age groups and peaks (Fig. 5) (Yonkee et al., 2014). We interpret these grains as shed from the 1.80-1.70 Ga Yavapai Province and the 1.34-1.48 Ga mid-continent granite-rhyolite province (Fig. 1), very similar to the younger Osgood Mountain Quartzite samples.

#### **b. Central Utah: Canyon Range**

Detrital zircons in Canyon Range strata (Fig. 2a) indicate a similar shift of ages as those analyzed in the Osgood Mountains. The older unit analyzed in the Canyon Range is the Neoproterozoic Caddy Canyon Quartzite (Fig. 6). This unit has Mesoproterozoic age groups and peaks (Fig. 5) (Lawton, et al. 2010) which we interpret as primarily shed from the 1.2-1.0 Ga Grenville orogen and the 1.48-1.34 Ga mid-continent granite-rhyolite province (Fig. 1). These age peaks and our source area interpretation are very similar to those of the older Osgood Mountain Quartzite sample. The younger unit in the Canyon Range is the Cambrian Prospect Mountain Quartzite (Fig. 6). This unit has Paleoproterozoic and Mesoproterozoic age groups and peaks (Fig. 5) (Lawton, et al., 2010). We interpret these detrital zircons as shed from the 1.80-1.70 Ga Yavapai Province and the 1.48-1.34 Ga mid-continent granite-rhyolite province (Fig. 1), quite similar to the younger Osgood Mountain Quartzite samples.

Figure 6: Stratigraphic columns of the passive margin strata discussed. Subjacent strata are not shown. Neoproterozoic-Cambrian boundary at 540 Ma is approximated for each location. Strata in the Osgood Mountains are shown in stratigraphic succession and correlative with regional passive margin strata. Only the detrital zircon samples referenced in the text are shown.



### **c. Northeastern Utah: Huntsville**

Detrital zircons from Huntsville samples (Fig. 2a) record a comparable shift of ages to those analyzed in the Osgood Mountains. The older units in the Huntsville section are the Neoproterozoic Kelley Canyon and Mutual Formations and the Neoproterozoic – Cambrian Brown’s Hole Formation (Fig. 6). The Kelley Canyon Formation (Yonkee et al., 2014), the Mutual Formation (Stewart, et al. 2001; Yonkee et al., 2014), and the Brown’s Hole Formation (Yonkee et al., 2014) have similar Mesoproterozoic age groups and peaks (Fig. 5). We interpret these grains as primarily derived from the 1.2-1.0 Ga Grenville orogen and the 1.48-1.34 Ga mid-continent granite-rhyolite province (Fig. 1)—very similar to that of our older Osgood Mountain Quartzite sample. The younger unit in the Huntsville section is the Cambrian Geersten Canyon Quartzite (Fig. 6). The Geersten Canyon Quartzite (Stewart et al., 2001; Yonkee et al., 2014) has Mesoproterozoic, Paleoproterozoic, and Archean age groups and peaks (Fig. 5). We interpret these grains as derived primarily from the 1.80-1.70 Ga Yavapai Province and secondarily from the 1.2-1.0 Ga Grenville orogen and the > 2.5 Ga Archean craton (Fig. 1). These age peaks and our interpreted provenance are similar to that for our younger Osgood Mountain Quartzite samples.

### **d. Southeastern Idaho: Portneuf Range**

Detrital zircons from rocks in the Portneuf Range (Fig. 2a), record a shift of ages similar to those analyzed in the Osgood Mountains. In the Portneuf Range, the older unit is the Neoproterozoic Middle Caddy Canyon Quartzite (Fig. 6). The Middle Caddy Canyon Quartzite has Mesoproterozoic age groups and peaks (Fig. 5) (Yonkee et al.,

2014). We interpret these grains as primarily shed from the 1.2-1.0 Ga Grenville orogen (Fig. 1). These age peaks and our source area interpretation are very similar to that of the older Osgood Mountain Quartzite sample. The younger unit in the Portneuf Range is the Windy Pass Argillite, comprising upper and middle members (Fig. 6). Both members have Paleoproterozoic age peaks and groups (Fig. 5) (Yonkee et al., 2014). We interpret these grains as shed from the 1.80-1.70 Ga Yavapai Province (Fig. 1). This is very similar to the younger Osgood Mountain Quartzite sample.

#### **e. Implications of the regional correlation**

In the five areas of the passive margin examined, the detrital zircon age patterns change in a systematic way (Fig. 7). The detrital zircons of the older strata are predominantly the age of the Grenville orogen and the detrital zircons of the younger strata are predominantly the age of the Yavapai province basement rocks.

The shift in detrital zircon age across the region implies a provenance change across the region at approximately the same time. Continent-spanning river systems carried sands from the Grenville orogenic terrane across the Laurentian continent to the western passive margin in Late Neoproterozoic to Early Cambrian time (e.g., Rainbird et al., 1997; Cawood and Nemchin, 2001; Mueller et al., 2007; Dehler et al., 2010; Kingsbury-Stewart et al., 2013). By Early to Middle Cambrian time, as this study has found, river systems were carrying relatively very few Grenville-aged detrital zircons to the passive margin in the areas that we investigated.

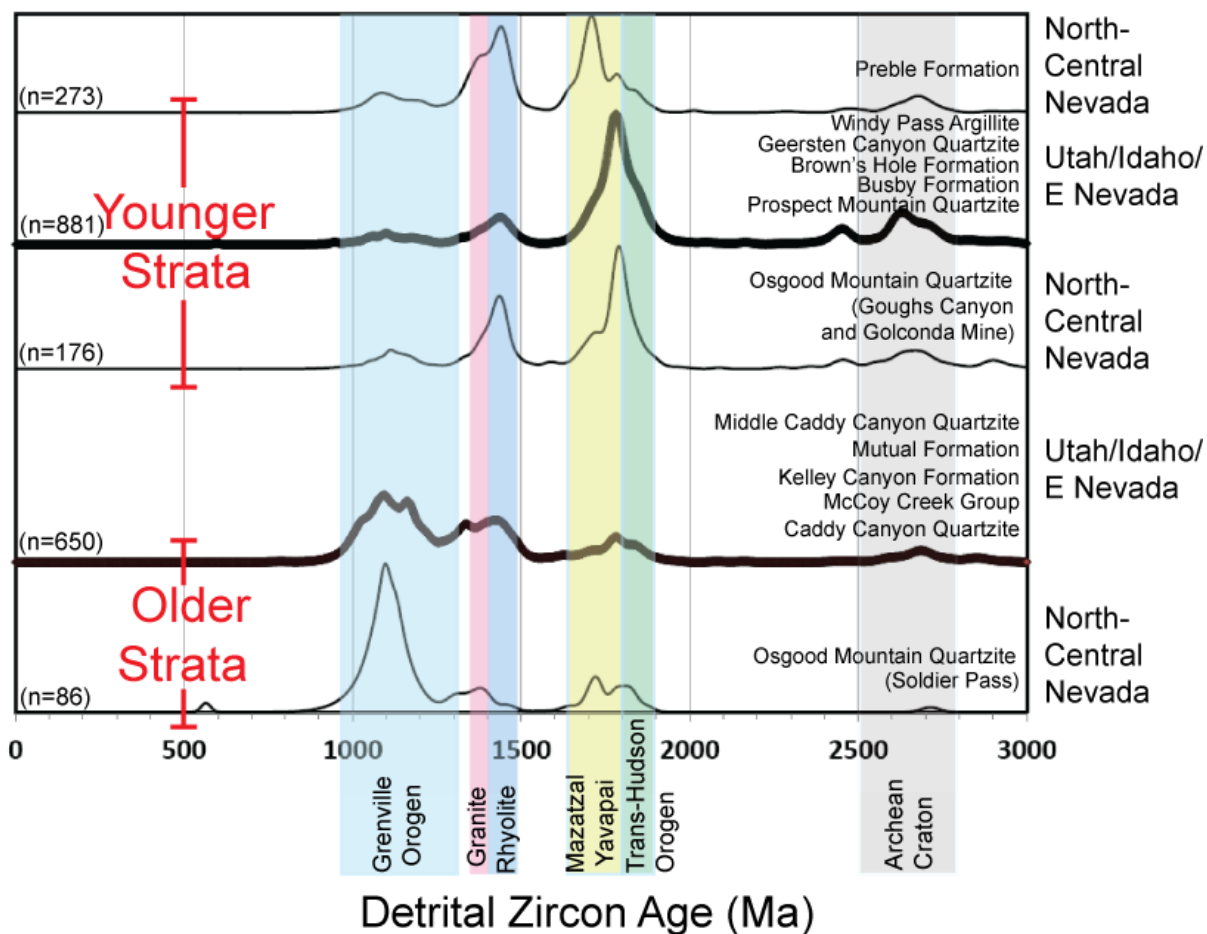


Figure 7. Plots of compilations of units showing the distribution of detrital zircon ages. Curves on the upper part of the figure are compilations of the younger strata throughout the region, compared with the younger strata in our pilot study. Curves on the lower part of the figure are compilations of the older strata throughout the region, compared with the older strata in our pilot study. Curves are normalized probability plots. The number of detrital zircon grains comprising each analysis is shown on the left. The thinner lines are from our pilot study; the thicker lines are the composites of data from elsewhere in the Basin and Range. The color bars indicate the ages of magmatic activity and terrane accretion in Laurentia; the colors are the same as Fig. 1. Canyon Range: Lawton et al., 2010; Snake, Deep Creek, and Pilot Ranges: Yonkee et al., 2014; Huntsville: Yonkee et al., 2014; Portneuf Range: Yonkee et al., 2014; Osgood Mountains: Linde et al., 2012; Gehrels and Pecha, 2012.

Our data support the proposal that the Transcontinental Arch, a continent-scale crustal feature that was not present in the Middle Neoproterozoic, was uplifted in the latest Neoproterozoic or earliest Cambrian. The uplifted arch provided vertical relief, and forced a change in sedimentation and drainage patterns, restricting the transport of sediment from the Grenville orogenic terrane to western Laurentia. This gradual change in drainage patterns caused the observed dearth of Grenville-age zircons. The uplift of the arch also would have exposed basement rocks of the Archean craton, the Trans-Hudson orogen, the Yavapai and Mazatzal terranes, and the mid-continent anorogenic granites (Fig. 1). Sloss (1963) demonstrated the onlap of the Sauk sequence onto basement rocks by Middle Cambrian time; this onlap requires that the basement rocks were exposed. These uplifted and exposed basement rocks became the sources for the sediments draining from the western flank of the Transcontinental Arch to the western Laurentian margin. The uplift of the arch caused the preponderance of zircons of the age of the Trans-Hudson orogen, Yavapai/Mazatal, and mid-continent granite province in the younger passive margin strata investigated in this study.

Across the Great Basin transect investigated for this study, the definitive shift in detrital zircon age spectra records a tectonic event that caused a coeval change in sediment transport patterns. This detrital zircon age shift therefore could be a useful correlation tool. Such a tool would be particularly helpful in quartz arenites, which are often lacking in biostratigraphic markers and thus difficult to date with confidence. We suggest that stratigraphic sections documented by our work and others can be correlated based on this consistent change in detrital zircon age peaks, and that other sections in the region can also be correlated in this way.



## **8. Summary and conclusions**

U-Pb analyses of detrital zircons from the Osgood Mountain Quartzite and the Preble Formation in northern Nevada demonstrate that these units were shed from sources within the North American craton, and that the sources changed with time. The oldest sample of the Osgood Mountain Quartzite, taken near Soldier Pass, was derived primarily from the Grenville orogen. The younger Osgood Mountain Quartzite samples were derived primarily from the Yavapai and Mazatzal provinces. The stratigraphically overlying Preble Formation was also derived primarily from the Yavapai and Mazatzal provinces.

The shift in age peaks and groups of detrital zircons within the Osgood Mountain Quartzite section is also recorded in other passive margin strata in Nevada, Utah, and Idaho. This shift indicates a widespread change in provenance; the older passive margin units were derived primarily from the Grenville orogen and the younger units were derived primarily from the more proximal Yavapai and Mazatzal provinces.

Our data support the proposal that the widespread and time-correlative shift in provenance of passive margin strata across the Great Basin records the uplift of the Transcontinental Arch. In Late Neoproterozoic time, these passive margin units were sourced primarily in the Grenville orogen; by Early Cambrian most passive margin units were sourced primarily from the Yavapai and Mazatzal provinces with very little input from the Grenville orogen. The predominantly east-to-west paleocurrents (Seeland, 1968) carried Grenville sands from the eastern third of the craton to the western passive margin until Early Cambrian time, when the Transcontinental Arch was uplifted and blocked this sediment dispersal pattern, and also provided a new sediment source for

western Laurentia. Our results corroborate previous suggestions that the Transcontinental Arch blocked the transport of sediments in the Early Paleozoic (Mack and Amato, 2012; Gehrels and Pecha, 2014; Yonkee et al., 2014). This study substantiates the previous interpretations by demonstrating coeval shifts in provenance across a broad transect of the Great Basin and confirms that the uplift of the arch was sufficient to change drainage patterns by Early Cambrian time.

We propose that analysis and comparison of detrital zircon age signatures can be used to support other chronostratigraphic correlation tools, especially in fossil-poor and notoriously hard-to-date quartz arenites. When used in conjunction with other correlation methods, shifts in detrital zircon age spectra can provide important information where other stratigraphic techniques are unavailable.

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## **Chapter 2**

### **Detrital zircon U-Pb geochronology and Hf isotope geochemistry of the Roberts Mountains allochthon: New insights into the Early Paleozoic tectonics of western North America**

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#### **1. Abstract**

Detrital zircon U-Pb geochronology and Hf isotope geochemistry provide new insights into the provenance, sedimentary transport, and tectonic evolution of the Roberts Mountains allochthon strata of north-central Nevada. Using laser-ablation inductively coupled plasma mass spectrometry, a total of 1151 zircon grains from six Ordovician to Devonian arenite samples were analyzed for U-Pb ages; of these, 228 grains were further analyzed for Hf isotope ratios. Five of the units sampled have similar U-Pb age peaks and Hf isotope ratios, while the ages and ratios of the Ordovician lower Vinini Formation are significantly different. Comparison of our data with that of igneous basement rocks and other sedimentary units supports our interpretation that the lower Vinini Formation originated in the north-central Laurentian craton. The other five units sampled, as well as Ordovician passive margin sandstones of the western Laurentian margin, had a common source in the Peace River Arch region of western Canada. We propose that the Roberts Mountains allochthon strata were deposited near the Peace River Arch region, and

subsequently tectonically transported south along the Laurentian margin, from where they were emplaced onto the craton during the Antler orogeny.

## **2. Introduction**

The Roberts Mountains allochthon (RMA) consists of internally deformed Cambrian through Devonian rocks, and structurally overlies coeval passive margin strata in northeastern and north-central Nevada (Schuchert, 1923; Kay, 1951; Roberts et al., 1958; Madrid, 1987; Burchfiel et al., 1992) (Figs. 8 and 9). Roberts Mountains allochthon rocks include chert, argillite, arenite, quartzite, limestone, and mafic volcanic rocks. The RMA is often thought to have been deposited in an ocean basin outboard of coeval passive margin strata in western Laurentia and to have been tectonically emplaced onto this margin during the Late Devonian to Early Mississippian Antler orogeny (e.g., Roberts et al., 1958; Burchfiel and Davis, 1972; Madrid, 1987). Various workers have suggested wildly disparate sources for the RMA strata. Some workers (e.g., Roberts et al., 1958; Burchfiel and Davis, 1972; Poole et al., 1992) suggested that the RMA strata originated in western Laurentia (Fig. 8) and deposited in an ocean basin to the west. Speed and Sleep (1982) hypothesized that the RMA strata are the accretionary prism of a far-traveled arc. Gehrels et al. (2000a) proposed that the RMA originated in the Peace River Arch region of western Canada. Wright and Wyld (2006) suggested that the RMA was deposited as far afield as Avalonia or Gondwana and subsequently was tectonically transported to western Laurentia along its southern margin (Fig. 10). Colpron and Nelson (2009) proposed that RMA strata could have originated in the northern Baltica–southern Caledonides region and been tectonically transported along the northwest margin of



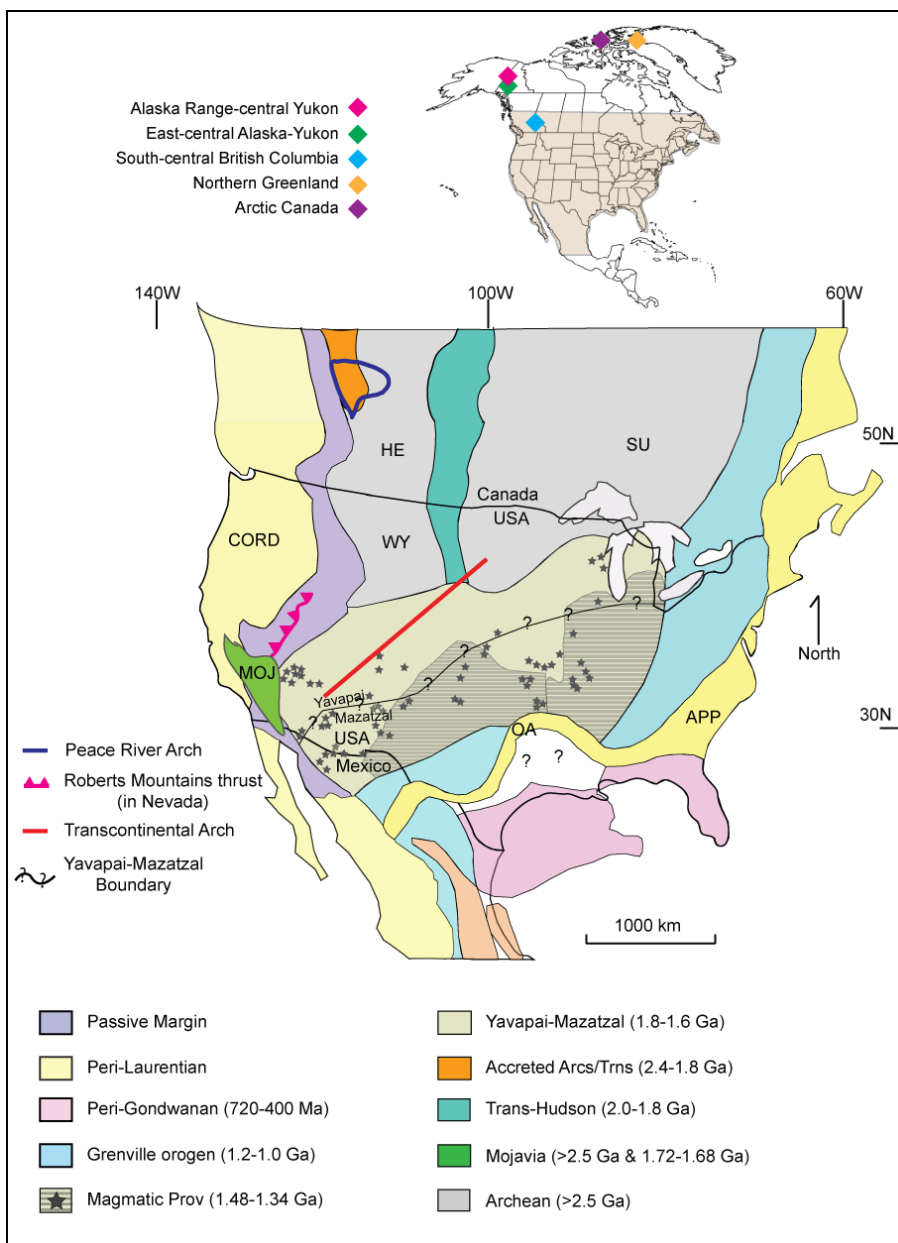


Figure 8. Locations of the main age provinces in North America that are potential source terranes for western Laurentian strata. The location of Transcontinental Arch is shown as a red line (Sloss, 1988); the Peace River Arch is shown as a blue line. The trace of the Roberts Mountains thrust is shown. WY—Wyoming province; HE—Hearn province; SU—Superior province; CORD—Cordilleran; APP—Appalachian; OA—Ouachita-Marathon; MOJ—Mojavia. Figure is after Gehrels et al. (2011) and compiled from Bickford et al. (1986), Hoffman (1989), Ross (1991), Burchfiel et al. (1992), Anderson and Morrison (1992), Bickford and Anderson (1993), Van Schmus et al. (1993), Villeneuve et al. (1993), Dickinson and Lawton (2001), Whitmeyer and Karlstrom (2007), and Dickinson and Gehrels (2009). An inset map of North America shows other locations referred to in the text.

Laurentia (Fig. 10). Determining the provenance of the RMA units will unravel this puzzle and provide new insight into early Paleozoic tectonics in the western Cordillera.

The gaps in understanding about the RMA strata—their provenance, sedimentary transport to depositional basin, and possible subsequent tectonic transport—can be addressed using detrital zircon analyses. We analyzed detrital zircons to obtain both uranium-lead ages and hafnium isotope ratios. U-Pb ages are important for identifying and then characterizing the provenance of sedimentary strata, and for comparison between sedimentary units (Gehrels et al., 2000b; Fedo et al., 2003; Gehrels, 2012, 2014). Hafnium isotope compositions are used to determine the geochemical character of the magma in which the zircons crystallized. When combined with U-Pb ages, Hf isotope composition provides a powerful complement for interpreting sedimentary provenance (Bahlburg et al., 2011; Gehrels and Pecha, 2014).

In this study, we determined the U-Pb ages and Hf isotope compositions of detrital zircons in six samples of RMA strata in north-central Nevada. We use these data to interpret provenance, sedimentary transport to depositional basins, possible subsequent tectonic transport, and relationships between RMA units. Our study builds on an earlier analysis of RMA samples that determined U-Pb ages using isotope-dilution–thermal ionization mass spectrometry (ID-TIMS) (Gehrels et al., 2000a, 2000b). Using detrital zircons from the same samples, the original data set was enlarged and enhanced. We analyzed a significantly larger number of grains per sample, changed and updated grain selection methods, and added Hf isotope composition analyses. We used laser-ablation–inductively coupled plasma mass spectrometry (LA-ICPMS) for all analyses. We report

here 1151 new U-Pb ages and 228 new Hf isotope analyses. Detrital zircon analyses allow us to resolve the original sources of these units. We show that some RMA units in some cases share an origin, while others units do not.

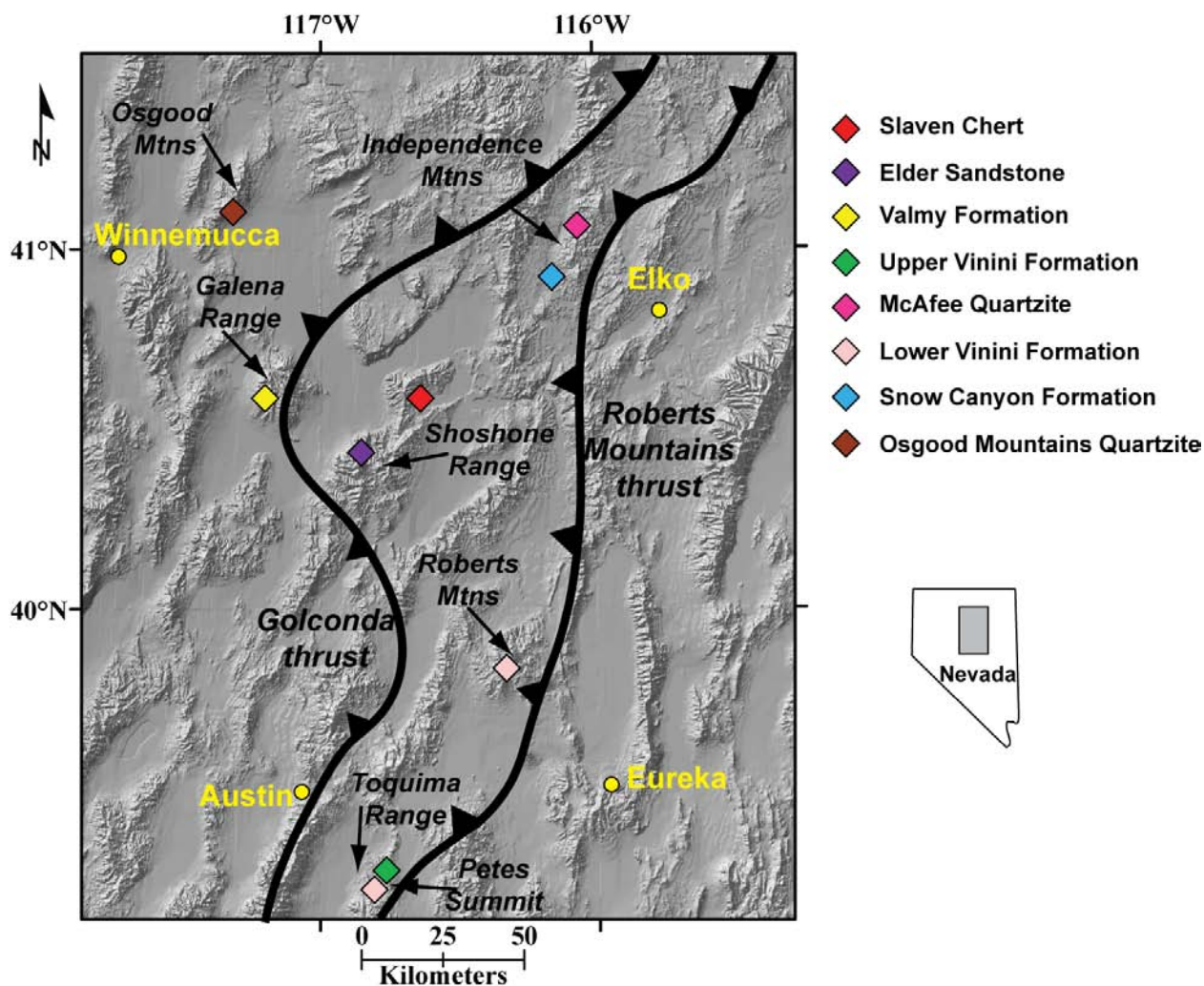


Figure 9. Map of north-central Nevada, showing sample locations (colored symbols) and the traces of the Roberts Mountains and Golconda thrusts. Some Roberts Mountains allochthon (RMA) rocks crop out to the west of the Golconda thrust in tectonic windows through the allochthon. Antler orogenic highlands are the map area to the west of the Roberts Mountains thrust; Antler Foreland Basin is the map area to the east of the Roberts Mountains thrust. Thrust traces are after Dickinson (2006); Antler highlands and basin are after Poole (1974).

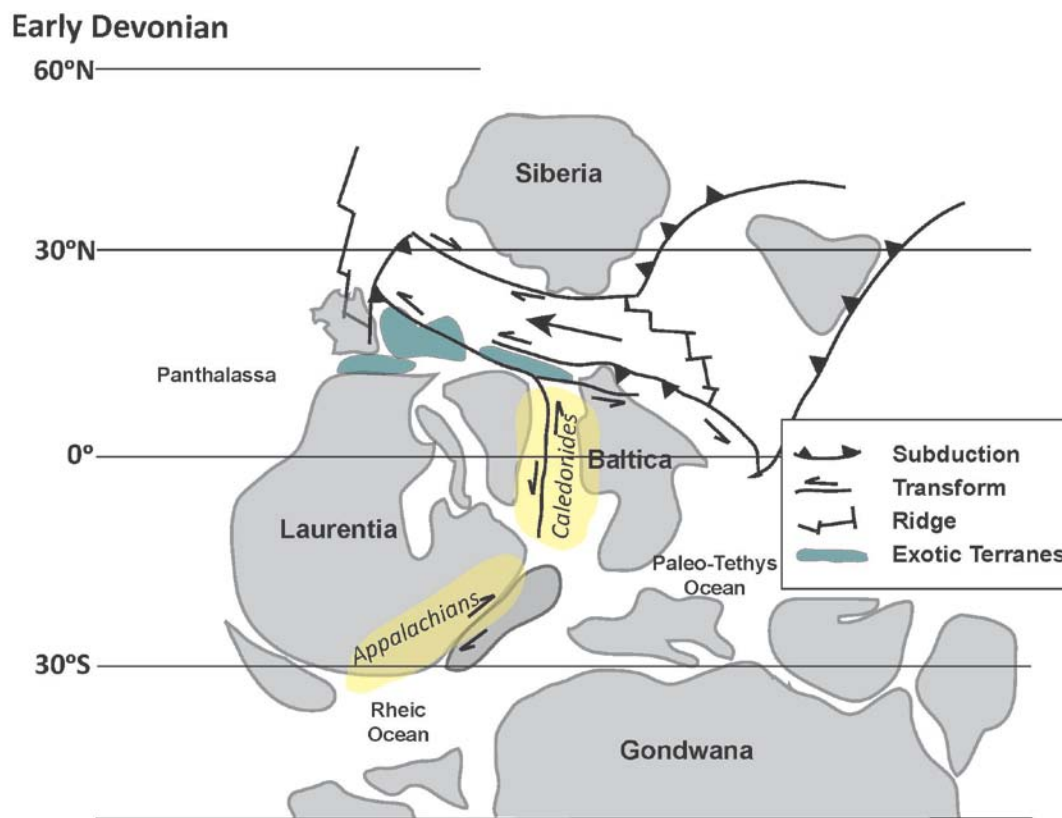


Figure 10. Early Devonian “Northwest Passage” between Laurentia, Baltica, and Siberia proposed by Colpron and Nelson (2009). Exotic terranes include the Alexander, Klamath, and northern Sierran terranes. Map is after Colpron and Nelson (2009).

### 3. Geologic Setting

#### a. Regional Tectonostratigraphic Framework

The North American craton contains several Proterozoic and Archean age provinces, thus providing geologically distinguishable crustal provinces that are source terranes for the upper Proterozoic and lower Paleozoic continental margin sedimentary section (e.g., Gehrels et al., 2011, and references cited therein) (Fig. 8). The Yavapai-Mazatzal Province (1.8–1.6 Ga) extends across central North America (Fig.1). It is bounded on the

north and northwest by the Trans-Hudson orogenic terrane (2.0–1.8 Ga) and Archean rocks (>2.5 Ga) of the Wyoming and Superior Provinces (Fig. 8). It is bounded on the south and east by the terranes of the Grenville orogen (1.2–1.0 Ga) and on the west by the Mojavia terrane (>2.5 Ga with 1.6–1.7 Ga granitoids) (Fig. 8).

Detrital zircon sources for the passive margin section changed in the upper Proterozoic–Lower Cambrian (Linde et al., 2014, and references cited therein). The 1.2–1.0 Ga Grenville orogen of southern and eastern North America (Fig. 8) was a significant sediment source for western Laurentia throughout the Neoproterozoic (Rainbird et al., 1997, 2012), including the upper Proterozoic passive margin section from the northwest United States to Sonora, Mexico (e.g., Lawton et al., 2010; Gehrels and Pecha, 2014; Yonkee et al., 2014; Linde et al., 2014). In contrast, the 1.8–1.6 Ga Yavapai-Mazatzal and 1.48–1.34 Ga mid-continent granite-rhyolite provinces within the North America craton (Fig. 8) were the more predominant sediment sources for strata higher in the passive margin section (e.g., Lawton et al., 2010; Gehrels and Pecha, 2014; Yonkee et al., 2014; Linde et al., 2014).

The RMA is often interpreted as a package of oceanic sediments emplaced structurally eastward onto the western Laurentian craton during the Late Devonian– Early Mississippian Antler orogeny (Roberts et al., 1958; Poole et al., 1992). Roberts Mountains allochthon strata are exposed in north-central Nevada between the Golconda thrust on the west and the Roberts Mountains thrust on the east; some units crop out west of the Golconda thrust in tectonic windows (Fig. 9). Rocks of the allochthon structurally overlie coeval rocks of the western Laurentian passive margin (e.g., Schuchert, 1923; Kay, 1951; Roberts et al., 1958; Madrid, 1987) (Fig. 11). Roberts Mountains allochthon

strata are highly deformed, and include imbricated older-over-younger thrust sheets (Evans and Theodore, 1978; Oldow, 1984; Noble and Finney, 1999). The metamorphic grade of the strata is generally greenschist facies or lower (Gehrels et al., 2000a). The RMA was emplaced along the Roberts Mountains thrust during the Late Devonian to Early Mississippian Antler orogeny (Roberts et al., 1958). The Antler foreland basin, west of the Laurentian craton and east of the Antler orogen, was filled between Devonian and Early Mississippian time by sediments shed from the uplifting Antler highlands (Poole, 1974; Trexler et al., 2003) (Fig. 9).

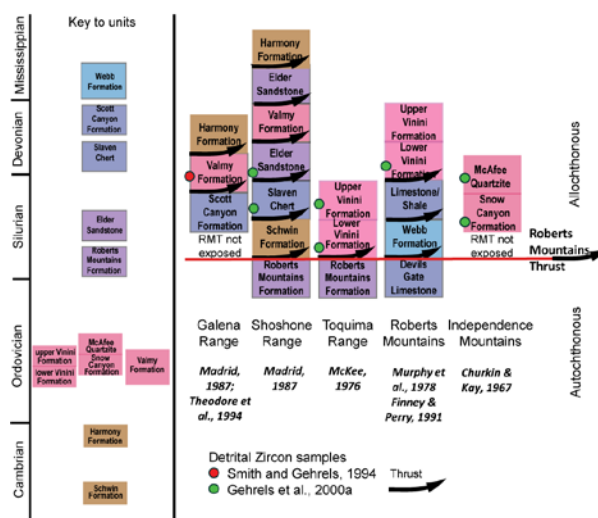


Figure 11. Tectonostratigraphic diagram of units of the Roberts Mountains allochthon (RMA) in selected north-central Nevada mountain ranges, showing locations of detrital zircon samples. Units are shown in their physical, structurally superimposed, order. Most units are internally disrupted with multiple imbricate thrusts not shown on this chart. Units are color coded for geologic period as indicated on left margin of chart.

The plate tectonic setting of the Antler orogeny has been variously interpreted as continent-continent collision, continent-arc collision, backarc thrusting, and polarity reversal of a subduction zone (e.g., Nilsen and Stewart, 1980; Speed and Sleep, 1982;

Dickinson et al., 1983). The RMA is often interpreted as an accretionary prism formed due to plate convergence at the continental margin (Speed and Sleep, 1982; Oldow, 1984; Dickinson, 2000).

Evidence for Antler-age tectonism has been reported along the western Laurentian margin, in Alaska and Canada (e.g., Nilsen and Stewart, 1980; Gehrels and Smith, 1987; Dusel-Bacon et al., 2006; Nelson et al., 2006; Paradis et al., 2006; Piercey et al., 2006; Colpron et al., 2007). Middle to Late Devonian continental arc magmatism occurred in the Alaska Range and central Yukon (Piercey et al., 2006) (Fig. 8). Upper Devonian–Early Mississippian felsic igneous and metaigneous rocks record bimodal volcanism in east-central Alaska and the Yukon (Dusel-Bacon et al., 2006) (Fig. 8). In south-central British Columbia (Fig. 8), a Late Devonian continental arc and backarc developed (Paradis et al., 2006).

Colpron et al. (2007) and Colpron and Nelson (2009) have proposed a direct link between the Antler orogeny and coeval tectonism of western Laurentia. They propose that a “Northwest Passage” opened in mid-Paleozoic time between Laurentia and Siberia, and a Scotia-style arc developed along the northern Laurentian margin in the Early Devonian (Fig. 10). The Alexander terrane, and other fragments such as the eastern Klamath and northern Sierran terranes, were transported from a Baltica origin to northwestern Laurentia through the Northwest Passage via the westward migration of the arc’s subduction zone (Fig. 10). By Middle Devonian time, a sinistral transform fault developed at the southern end of this passage and extended southward along western Laurentia. This system transported these terranes and fragments south along the margin. Colpron and Nelson (2009) note progressively younger deformation southward along the

Laurentian margin, from Alaska and the Yukon to Nevada, and suggest that this records the southward propagation of the transpressional system. They propose that this fault system could have provided the weakness along which Devonian subduction initiated.

#### b. **Roberts Mountains Allochthon Strata**

The RMA strata sampled (Figs. 9 and 11; Table 2) are arenite beds within units that are predominantly chert and argillite with some limestone and mafic volcanic rocks. Most contacts between and within units are structural, and the stratigraphic bases and tops of units are not known. The strata of the RMA are described briefly below, as evidence of their depositional environments.

<b>Sample Location</b>	<b>UTM: NAD 83</b>	<b>Easting</b>	<b>Northing</b>
Ordovician Snow Canyon Formation Snow Canyon; Independence Mtns		0579760	4585698 (UTM 11T)
Ordovician lower Vinini Formation Petes Summit; Toquima Range		0518089	4337111 (UTM 11S)
Ordovician McAfee Quartzite McAfee Peak; Independence Mtns		0590637	4599583 (UTM 11T)
Ordovician upper Vinini Formation Petes Summit; Toquima Range		0518089	4337111 (UTM 11S)
Silurian Elder Sandstone Elder Creek; Shoshone Range		0516196	4460270 (UTM 11T)
Devonian Slaven Chert Slaven Canyon; Shoshone Range		0519428	4479302 (UTM 11T)

Table 2: Locations of samples analyzed in this study referenced to UTM locations.

The Snow Canyon Formation and the McAfee Quartzite in the Independence Mountains (Figs. 9 and 11; Table 2) are the equivalent of the upper Vinini and Valmy formations, respectively (Holm-Denoma et al., 2011). Both units are Middle Ordovician



based on graptolite fauna (Churkin and Kay, 1967). The Snow Canyon Formation is predominantly chert with arenite, shale, and siltstone layers, and basaltic lavas with interbedded limestone (Churkin and Kay, 1967). The McAfee Quartzite is predominantly massive cliff-forming quartzite with intervals of shale and siltstone and bedded chert (Churkin and Kay, 1967). The arenite intervals in these formations are interpreted as turbidites (Miller and Larue, 1983).

The Vinini Formation (Figs. 9 and 11; Table 2) was first mapped in the Roberts Mountains by Merriam and Anderson (1942) along Vinini Creek. Merriam and Anderson (1942) recognized two informal units (upper and lower) based on lithology and graptolite fauna and described the extreme structural disruption of these rocks. In later work, Noble and Finney (1999) used precise radiolarian biostratigraphy to demonstrate a high degree of structural imbrication both within the Vinini Formation and within Devonian cherts. In the Toquima Range, near Petes Summit (Fig. 9), the Vinini Formation is divided into two informal units (upper and lower), which are mapped in depositional contact, and the extreme structural complexity and repetition of thrust slices is also mapped (McKee, 1976). We observed the depositional contact at Petes Summit, where the quartz arenite of the upper Vinini rests on shale of the lower Vinini. The lower Vinini Formation is predominantly quartz arenite, with siltstone, shale, chert, and limestone (Finney et al., 1993). The lower Vinini Formation is Upper Lower to Lower Middle Ordovician in age, based on graptolite and conodont fauna (Finney et al., 1993). The arenite intervals in the lower Vinini Formation are interpreted as turbidites (Finney et al., 1993). The upper Vinini Formation is predominantly shale and bedded chert, with some siltstone and arenite (Finney et al., 1993). The unit is Middle Middle to Upper Ordovician, based on

graptolites and conodonts (Finney et al., 1993). Graptolites and conodonts of the lower Vinini Formation are similar to those found in coeval Laurentian shelf carbonate deposits (Finney and Ethington, 1992; Finney, 1998). At Petes Summit, we observed low-angle cross lamination and hummocky cross stratification in the arenite of the upper Vinini Formation. We therefore interpret the upper Vinini as having been deposited in a high-energy environment at a depth above storm wave base on the continental shelf and probably at less than 100 m depth.

The Elder Sandstone (Figs. 9 and 11; Table 2) is predominantly fine-grained sandstone and siltstone, with some cherty shale and quartzite (Gilluly and Gates, 1965). Fossils are sparse in the unit; the age is Lower Silurian based on graptolites (Gilluly and Gates, 1965). The Elder Sandstone is interpreted as a turbidite deposit (Madrid, 1987).

The Slaven Chert (Figs. 9 and 11; Table 2) is predominantly black, bedded chert with shale beds and some limey sandstone and siltstone (Gilluly and Gates, 1965). The unit is Middle Devonian based on a variety of fossils (Gilluly and Gates, 1965). The arenite intervals in the Slaven Chert are interpreted as turbidites (Madrid, 1987).

#### **4. Methods**

Zircon grains from six arenite samples were analyzed for U-Pb ages and Hf isotope ratios (Figs. 9 and 11; Table 2). A small number of zircon grains from these samples were previously analyzed for U-Pb ages by Gehrels et al. (2000a), using ID-TIMS (Fig. 5). Zircon grains were separated and analyzed at the University of Arizona LaserChron facility using standard techniques described by Gehrels and Pecha (2014) to yield a best age distribution reflective of the true distribution of detrital zircon ages in each sample.

Approximately 200 randomly selected grains were analyzed in each sample for U-Pb ages. Approximately 50 of these grains were subsequently analyzed for Hf isotopes. Hf analyses were conducted on top of the pits left after U-Pb analysis, to ensure that Hf isotope data were collected from the same domain as the U-Pb age. Analyses were conducted by LA-ICPMS using the Photon Machines Anlyte G2 excimer laser connected to the Nu Plasma high-resolution inductively coupled plasma-mass spectrometer, using methods identical to those described by Gehrels and Pecha (2014).

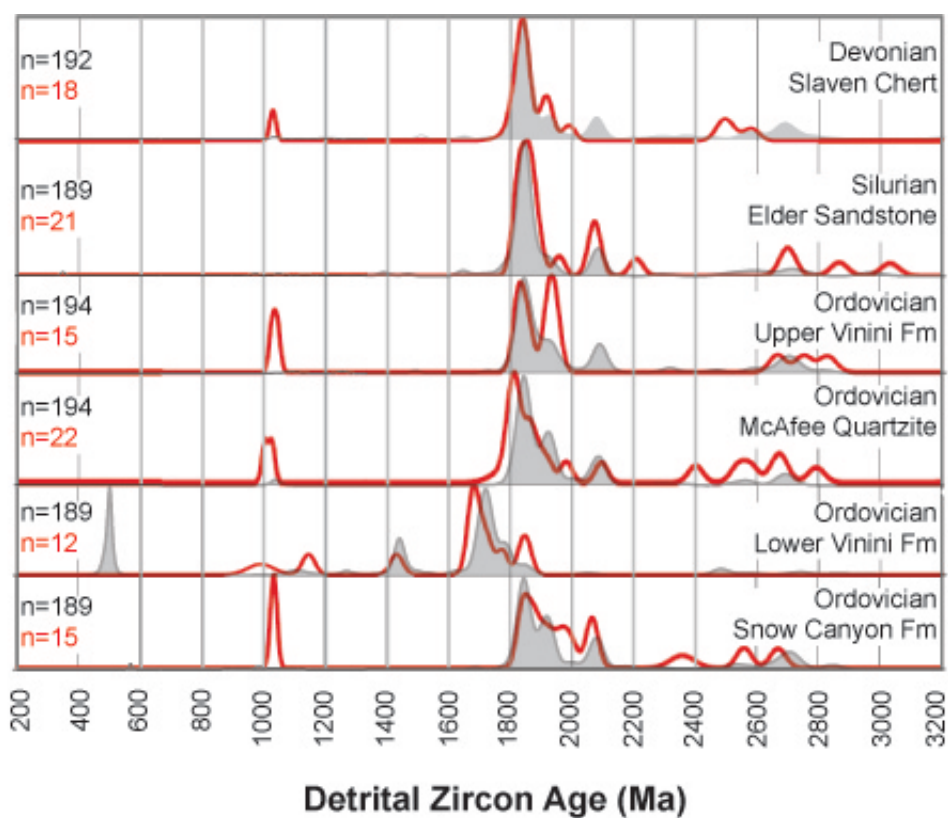


Figure 12. Normalized probability plots showing U-Pb ages of strata sampled. Red lines show the data from isotope-dilution thermal ionization mass spectrometry (Gehrels et al., 2000a); grayfilled curves are the data from laser-ablation inductively coupled plasma mass spectrometry (this study). Numbers of grains analyzed are shown (n = ).

### **a. Uranium-Lead Geochronology**

Analytical results are displayed graphically on normalized probability plots (Figs. 12 and 13), which allow visual comparison between zircon populations. U-Pb geochronology results are displayed in Figure 12, which contains both data from the original ID-TIMS analyses of these samples (Gehrels et al., 2000a) and the LA-ICPMS analyses of the current study, and in Figure 13, which displays the U-Pb results and Hf isotope analyses of the current study on the same chart. The essential U-Pb isotope information and ages are reported in Appendix C.

We compared detrital zircon age distributions both visually and statistically. Our initial appraisal was visual comparison of the probability plots. We also compared age distributions using the Kolmogorov-Smirnov (K-S) statistic (Guynn and Gehrels, 2006) (Table 3). The K-S statistic calculates whether a statistically significant difference exists between two distributions.  $P < 0.05$  indicates >95% probability that two U-Pb distributions are not the same. The K-S statistic is sensitive to proportions of ages present, and a low-P value may indicate that the proportions of age peaks are different, even though the ages are similar (Gehrels, 2012).

### **b. Hafnium Isotope Analysis**

Hafnium isotope data are shown on Hf-evolution diagrams (Fig. 13) that display epsilon Hf(t) ( $\epsilon\text{Hf}(t)$ ) values at the time of crystallization. The essential Hf isotope information is reported in Appendix D.

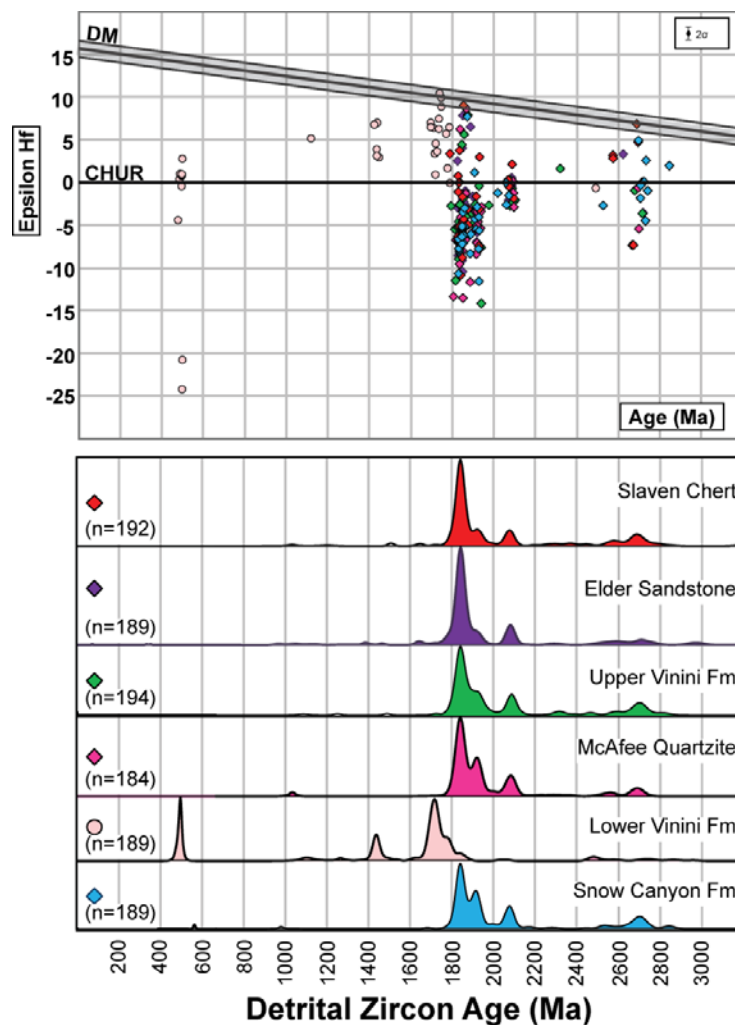


Figure 13. U-Pb ages and Hf isotope data for Roberts Mountains allochthon strata. U-Pb dates were run for all sample grains; approximately one-fourth of these grains were analyzed for hafnium isotopes. The upper graph shows  $\epsilon\text{Hf}(t)$  (epsilon Hf) values for each sample. The average measurement uncertainty for all hafnium analyses is shown in the upper right at the  $2\sigma$  of the values. Reference lines on the Hf plot are as follows: Depleted mantle (DM) is calculated using  $^{176}\text{Hf}/^{177}\text{Hf} = 0.283225$ ;  $^{176}\text{Lu}/^{177}\text{Hf} = 0.038513$  (Vervoort and Blichert-Toft, 1999); CHUR—chondritic uniform reservoir, is calculated using  $^{176}\text{Hf}/^{177}\text{Hf} = 0.282785$  and  $^{176}\text{Lu}/^{177}\text{Hf} = 0.0336$  (Bouvier et al., 2008).

	Slaven Chert	Elder Sandstone	Valmy Formation	Upper Vinini Formation	McAfee Quartzite	Snow Cyn Formation	Eureka Quartzite	Kinnikinic Quartzite	Mt Wilson Formation
Slaven Chert		0.130	0.156	0.402	0.103	0.069	0.064	0.239	0.010
Elder Sandstone	0.130		0.001	0.001	0.013	0.000	0.000	0.000	0.000
Valmy Formation	0.156	0.001		0.330	0.018	0.458	0.118	0.010	0.050
Upper Vinini Formation	0.402	0.001	0.330		0.203	0.872	0.748	0.307	0.348
McAfee Quartzite	0.103	0.013	0.018	0.203		0.318	0.433	0.990	0.020
Snow Canyon Formation	0.069	0.000	0.458	0.872	0.318		0.786	0.313	0.200
Eureka Quartzite	0.064	0.000	0.118	0.748	0.433	0.786		0.948	0.768
Kinnikinic Quartzite	0.239	0.000	0.010	0.307	0.990	0.313	0.948		0.104
Mt Wilson Formation	0.010	0.000	0.050	0.348	0.020	0.200	0.768	0.104	

Table 3: K-S statistical analysis results. The RMA strata are shown with green highlights, and the passive margin strata are shown with blue highlights. Comparisons between units with values greater than 0.05 are highlighted in yellow.  $P < 0.05$  indicates >95% probability that two U-Pb distributions are not the same.

### 5. Results: U-Pb Ages and Hf Isotope Ratios

Although the ID-TIMS data (Gehrels et al., 2000a) are similar to the new LA-ICPMS data (Fig. 12), there are variations in the proportions of age groups. The two studies used different grain-selection procedures. For the ID-TIMS study, zircon crystals were selected from color and morphology groups, without regard to the number of grains in each group. For our LA-ICPMS study, we attempted to select grains at random from the entire population of grains. This procedure resulted in a more representative age distribution because the grains are chosen randomly. The results and interpretations that follow are all based upon LA-ICPMS ages from our current study.

U-Pb geochronology and Hf isotope analyses reveal that the RMA strata are in two distinct groups. Five of the six samples (the Snow Canyon Formation, the McAfee Quartzite, the upper Vinini Formation, the Elder Sandstone, and the Slaven Chert) yield similar U-Pb age spectra, while the remaining sample, the lower Vinini Formation, yields significantly different U-Pb age spectra (Figs. 12 and 13). The Hf data from the five samples with similar U-Pb ages are similar, while the lower Vinini Formation sample, because of its different age spectra, yields significantly different Hf ratios (Fig. 13).

## **6. Provenance of the Roberts Mountains Allochthon**

To interpret provenance, we compared the data from our study to known U-Pb ages and Hf isotope data from Laurentian basement provinces and other sedimentary units.

### **a. Provenance of the Roberts Mountains Allochthon Exclusive of the Lower Vinini**

The detrital zircon age spectra of the Snow Canyon Formation, the McAfee Quartzite, the upper Vinini Formation, the Elder Sandstone, and the Slaven Chert are consistent with provenance in the Peace River Arch (PRA) region of western Canada (Fig. 14). The 1820–1960 Ma grains are similar in age to magmatic arcs in the PRA region, including the Fort Simpson, the Rimbey, the Ksituan, and the Great Bear arcs (Hoffman, 1989; Ross, 1991; Villeneuve et al., 1993) (Fig. 14). The 2060–2120 Ma grains are similar in age to accreted terranes in the PRA region, including the Buffalo Head and Chincaga terranes (Hoffman, 1989; Ross, 1991; Villeneuve et al., 1993) (Fig. 14). The 2650–2750 Ma grains are similar in age to Archean terranes in the PRA region, including the Nova and Hearne terranes (Hoffman, 1989; Ross, 1991; Villeneuve et al., 1993) (Fig. 14).

The Hf isotope data are consistent with provenance in the PRA region. The 1820–1960 Ma grains have a wide range of values, from juvenile and moderately juvenile through evolved ( $eHf(t) +10$  to  $-15$ ), similar to those of other units interpreted to originate in the PRA region (Gehrels and Pecha, 2014). The 2060–2120 Ma grains are more narrowly grouped, with moderately juvenile to evolved values of  $eHf(t) +3$  to  $-6$ , compatible with other units originating in the PRA region (Gehrels and Pecha, 2014).

The 2560–2750 Ma grains have juvenile, moderately juvenile, and evolved values of  $\epsilon_{\text{Hf}}(t)$  +6 to –15, also compatible with PRA origin (Gehrels and Pecha, 2014). The ages of basement terranes that comprise the PRA region (Fig. 7) are all represented in the age spectra of the RMA samples (exclusive of the lower Vinini Formation).

The detrital zircon U-Pb ages and Hf isotope data from these RMA strata are similar to selected passive margin strata and RMA strata analyzed in other studies (Fig. 15). The RMA strata sampled in this study (exclusive of the lower Vinini) have U-Pb age spectra similar to those of the Ordovician Valmy Formation of the RMA (Gehrels and Pecha, 2014), as well as the Eureka Quartzite and the Mount Wilson Formation (Gehrels and Pecha, 2014), and the Kinnikinic Quartzite (Barr, 2009), Ordovician units of the western Laurentian passive margin (Figs. 15 and 16). The K-S analyses of the RMA and the Ordovician passive margin units discussed above do not contradict our interpretation that the RMA strata have a common provenance with the Ordovician passive margin sandstones (Table 3). These RMA strata also show similar Hf isotope ratios to the Valmy Formation (Gehrels and Pecha, 2014) and to the Eureka Quartzite and the Mount Wilson Formation (Gehrels and Pecha, 2014) (Fig. 15).

The Peace River Arch region of western Canada is the source for the RMA units in this study, exclusive of the lower Vinini Formation, and for the Ordovician passive margin sandstones. The Peace River Arch region was an uplifted region from late Neoproterozoic through Middle Devonian time (Cant, 1988; Cant and O'Connell, 1988; Cecile et al., 1997). Igneous bodies in the PRA region have ages similar to the U-Pb ages of zircons in the RMA rocks sampled (Figs. 14 and 15). The U-Pb age spectra of the RMA rocks sampled are not consistent with derivation from the central Laurentian



craton; the Yavapai-Mazatzal terranes are 1.6–1.8 Ga and cannot serve as a source of the 1.8–2.0 Ga grains in the samples.

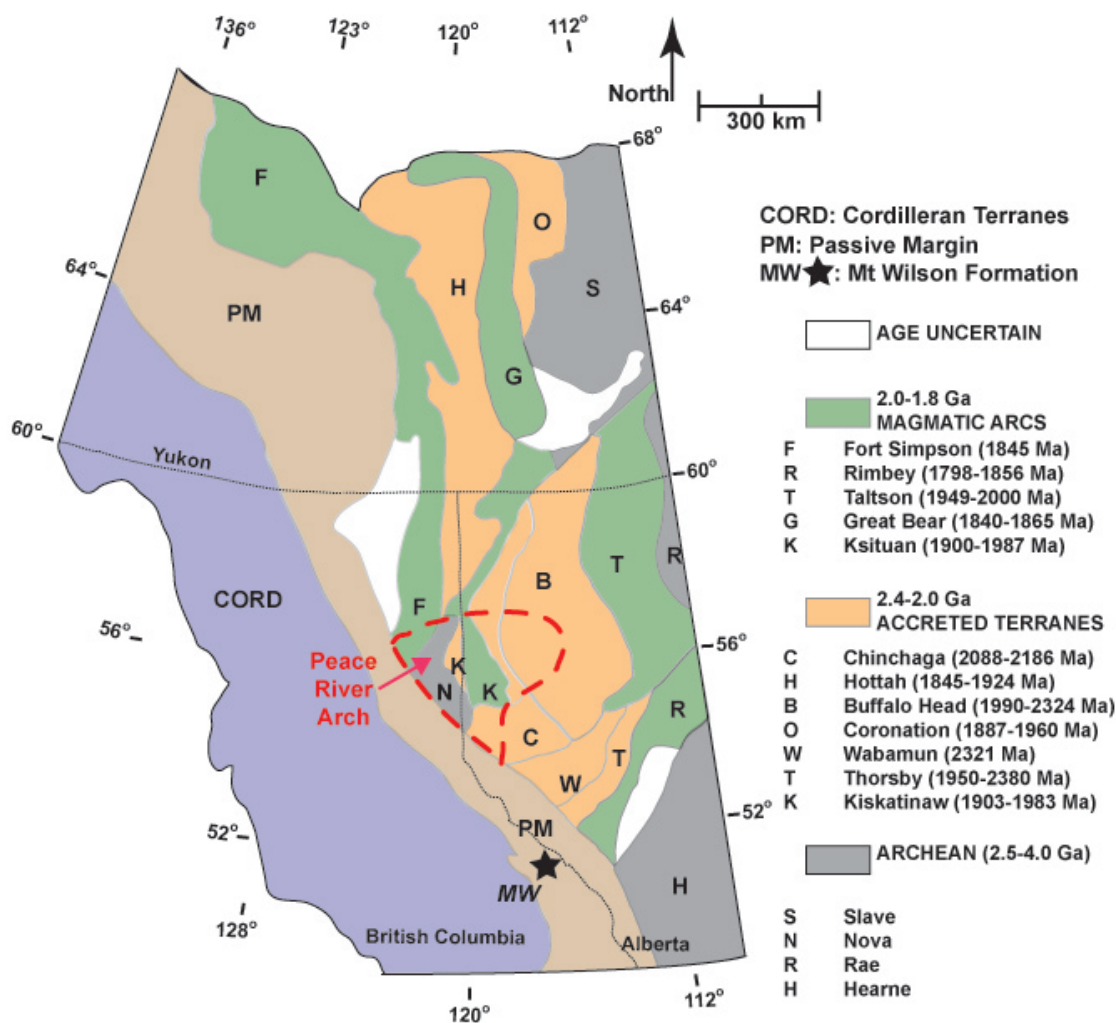


Figure 14. Map of western Canada showing the Cordilleran accreted terranes, the Cordilleran passive margin, and the basement provinces of the Canadian Shield. The Peace River Arch region is outlined by the red dashed line. The location of the Mount Wilson Formation sample (Gehrels and Pecha, 2014) is shown. The map is after Gehrels and Ross (1998); the basement provinces are compiled from Hoffman (1989), Ross (1991), and Villeneuve et al. (1993). WY Wyoming province; HE—Hearn province; SU—Superior province; CORD—Cordilleran; APP—Appalachian; OA—Ouachita-Marathon; MOJ—Mojavia.

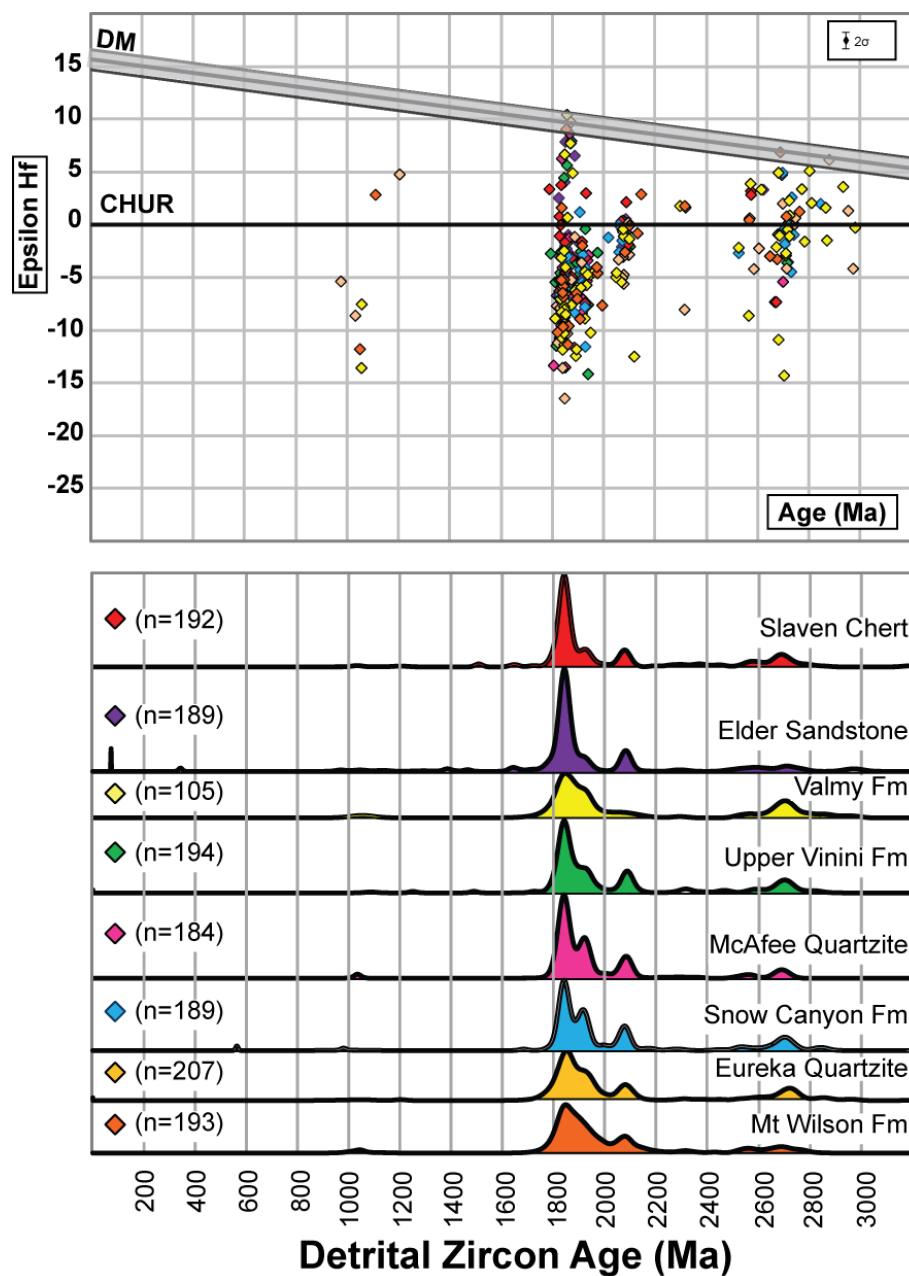


Figure 15. U-Pb ages and Hf isotope data of RMA and coeval passive margin strata. The data from the Mount Wilson Formation, the Eureka Quartzite, and the Valmy Formation are from Gehrels and Pecha (2014). Diagrams and symbols are as in Figure 13.

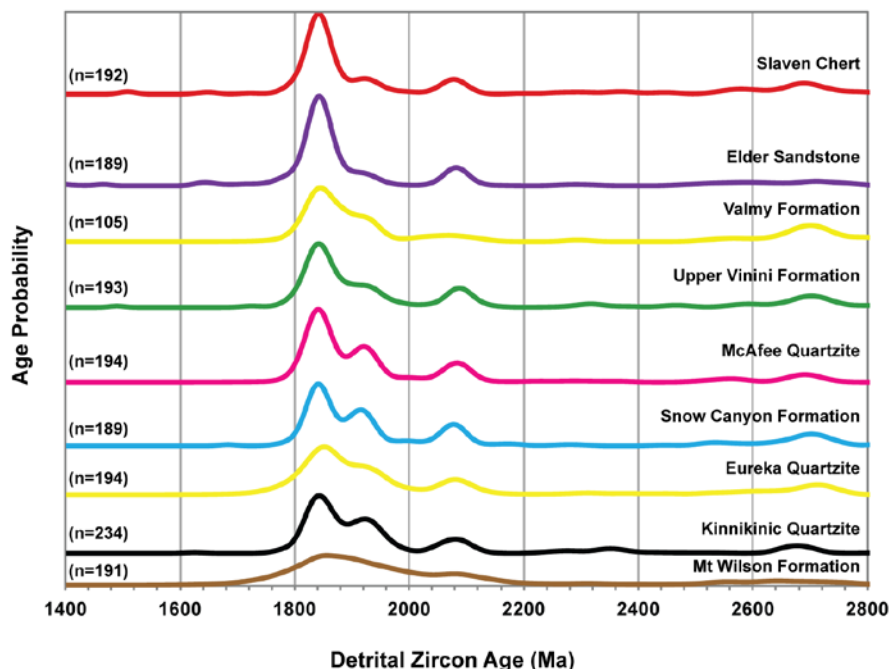


Figure 16. Normalized probability plot of Roberts Mountains allochthon (RMA) strata from this study, exclusive of the lower Vinini Formation. The plot includes the Valmy Formation of the RMA, analyzed by Gehrels and Pecha (2014). The plot also includes select Ordovician passive margin strata: the Mount Wilson Formation and Eureka Quartzite (Gehrels and Pecha, 2014) and the Kinnikinic Quartzite (Barr, 2009). No Hf isotope data are available for the Kinnikinic Quartzite, so only a normalized probability plot is shown.

### **b. Provenance of the Lower Vinini Formation**

The U-P age spectra of the lower Vinini Formation are consistent with provenance in north-central Laurentia. The 490–500 Ma grains are similar in age to plutonic suites in roof pendants and inliers within the Challis volcanic-plutonic complex and the Idaho batholith (Lund et al., 2010). The 1110–1120 Ma grains are consistent with the Grenville orogen; the 1420 Ma grains are consistent with the central Laurentian anorogenic granites; the 1660–1800 Ma grains are consistent with the Yavapai-Mazatzal terranes; and the 2470–2750 Ma grains are consistent with the Archean craton (Bickford et al., 1986; Hoffman, 1989; Ross, 1991; Anderson and Morrison, 1992; Bickford and

Anderson, 1993; Van Schmus et al., 1993) (Fig. 8). River systems traversing the north-central craton from east to west transported sediments from these crystalline bedrock sources—or from sediments recycled from them—and subsequently deposited them off the western Laurentian margin as the lower Vinini Formation.

The Hf isotope data of the lower Vinini grains are also consistent with origin in north-central Laurentia. The 490–500 Ma grains have mostly moderately juvenile to evolved values ( $\epsilon\text{Hf}(t)$  +3 to –5), with two grains highly evolved ( $\epsilon\text{Hf}(t)$  –20 to –25). The moderately juvenile to evolved grains are compatible with the plutonic suites in Idaho; however, the highly evolved grains are unlike any analyzed in these suites (Todt and Link, 2013). The 1110–1120 Ma grains have moderately juvenile values ( $\epsilon\text{Hf}(t)$  +4 to +6), similar to those of the Grenville orogen (Mueller et al., 2008; Bickford et al., 2010). The 1420 Ma grains have juvenile to moderately juvenile values ( $\epsilon\text{Hf}(t)$  +7 to +3), compatible with the anorogenic granitoids of the mid-Laurentian craton (Goodge and Vervoort, 2006). The 1660–1800 Ma grains have juvenile to moderately juvenile values ( $\epsilon\text{Hf}(t)$  +10–0), similar to the Yavapai-Mazatzal terranes (Bickford et al., 2008). The 2470–2750 Ma grains have moderately juvenile to evolved values ( $\epsilon\text{Hf}(t)$  +6 to –6), compatible with those in northern Greenland and Arctic Canada (Rohr et al., 2008, 2010) (Fig. 8).

The Early Cambrian uplift of the Transcontinental Arch altered the drainage patterns in western Laurentia; this change is recorded in the changing detrital zircon age patterns between upper Neoproterozoic and Lower Cambrian passive margin strata (Linde et al., 2014, and references cited therein) (Fig. 17). The uplift of the arch blocked the transport of Grenville-age grains and created, on the west flank of the arch itself, a new highland

and source of sand, consisting of Yavapai-Mazatzal basement rocks and sedimentary rocks recycled from this basement. In many older passive margin strata that predate the uplift of the arch, Grenville-age grains predominate (Fig. 17). These grains were transported by continent-spanning rivers that drained the central craton and Grenville orogenic terrane to the western Laurentian margin through the late Neoproterozoic (Rainbird et al., 1997, 2012). In many younger passive margin strata, deposited after the uplift of the arch, Yavapai-Mazatzal-age grains dominate (Fig. 17). Rivers originating in the central craton were blocked from flowing to the west by the uplifted arch, which blocked the transport of many Grenville-age grains (Amato and Mack, 2012; Gehrels and Pecha, 2014; Linde et al., 2014; Yonkee et al., 2014).

The detrital zircon U-Pb ages and Hf isotope data of the lower Vinini Formation resemble those of the younger, post-Transcontinental Arch uplift, passive margin strata, such as the Geersten Quartzite of Utah and the Osgood Mountains Quartzite of Nevada (Fig. 18). These are the only post-arch uplift passive margin data sets for which we have both U-Pb ages and Hf isotope data. The lower Vinini Formation U-Pb age spectra and Hf isotope ratios are similar to those of the younger passive margin strata. The provenance of the lower Vinini Formation is central Laurentian, shed from the western flanks of the Transcontinental Arch and the regions to the west of the arch, after the uplift of the arch (Fig 8).

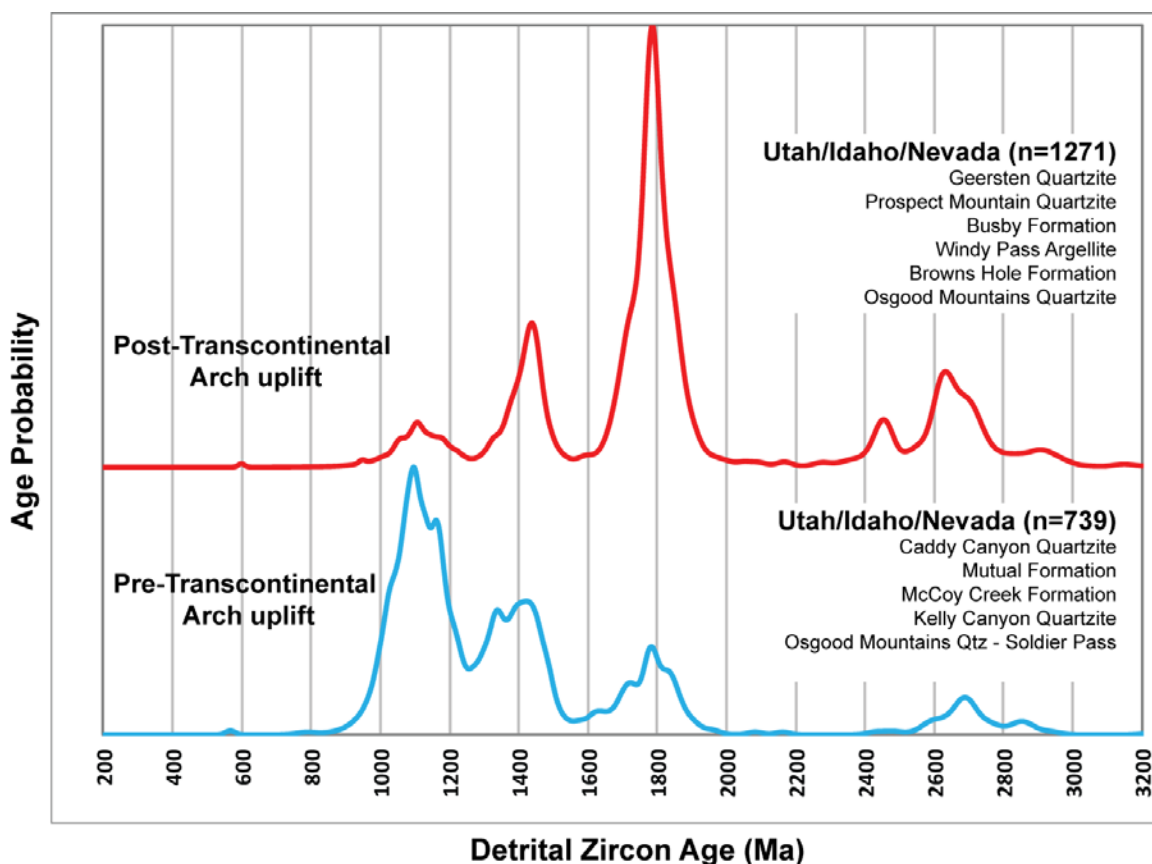


Figure 17. Compilation plots of units showing the distribution of detrital zircon ages in upper Neoproterozoic–Cambrian western Laurentian passive margin units (after Linde et al., 2014). The upper curve (red) is a compilation of ages of the relatively younger strata throughout the region. The lower curve (blue) is a compilation of ages of the relatively older strata throughout the region. The curves are normalized probability plots. The number of detrital zircon grains comprising each compilation is shown on the right. Osgood Mountains Quartzite (Linde et al., 2014); Kelley Canyon Quartzite, Caddy Canyon Quartzite, Brown’s Hole Formation, Geersten Canyon Quartzite, Mutual Formation, Prospect Mountain Quartzite, McCoy Creek Group, Busby Formation, and Windy Pass Argillite (Yonkee et al., 2014); Kelley Canyon Quartzite, Caddy Canyon Quartzite, Mutual Formation, and Prospect Mountain Quartzite (Lawton et al., 2010).

## **7. Discussion: Sedimentology and Paleogeographic Implications**

Sedimentological analyses provide a further constraint and suggest that the Ordovician passive margin sandstones are not the source of the RMA strata, but rather that these strata have a common source. Finney and Perry (1991) proposed that the Eureka

Quartzite (an extensive Ordovician passive margin unit) was the source of the sandstones in the younger sections of the Vinini and Valmy formations of the RMA. However, the grains of the Ordovician passive margin sandstones are more texturally mature than those of the RMA strata, whose grains are coarser, larger, and more poorly sorted (Ketner, 1966). The more mature shelf sands such as the Eureka Quartzite and Mount Wilson Formation could not be the source of the more immature RMA sandstones. The RMA and passive margin sandstones have similar U-Pb age spectra and Hf isotope ratios (Fig. 14) and share a common source in the PRA region.

The Ordovician passive margin sands and the RMA strata sampled have different depositional histories (Fig. 19). The Mount Wilson Formation was deposited in a nearshore to shelf environment immediately outboard of the Peace River Arch (Kent, 1994). Other Ordovician passive margin sandstones, now preserved as the Eureka Quartzite and the Kinnikinic Quartzite, were shed from the Peace River Arch, and subsequently moved southward along the western Laurentian margin via longshore transport to the depositional basin (Ketner, 1968) (Fig. 19B). The evidence for this transport is that grain size decreases and sorting improves in Ordovician arenites from near the PRA source (the Mount Wilson Formation) southward through Idaho (the Kinnikinic Quartzite) and into Nevada and California (the Eureka Quartzite) (Ketner, 1968). The texturally immature arenites of the RMA did not undergo the extensive reworking of this longshore transport. Sediments of the RMA strata, other than the lower Vinini Formation, were deposited as turbidites (Miller and Larue, 1983; Madrid, 1987; Finney et al., 1993) offshore of the Peace River Arch (Figs. 19A–19C). To reach their current geographic location, the RMA strata were *tectonically* transported south along the

western Laurentian margin, in Latest Devonian time (Fig. 19E). This is consistent with a sinistral transpressional fault system, as proposed by Colpron and Nelson (2009) (Figs. 19D and 19E). Subsequent shortening moved the RMA up onto the craton in the Antler orogeny of Latest Devonian–Earliest Mississippian time (Figs. 19E and 19F).

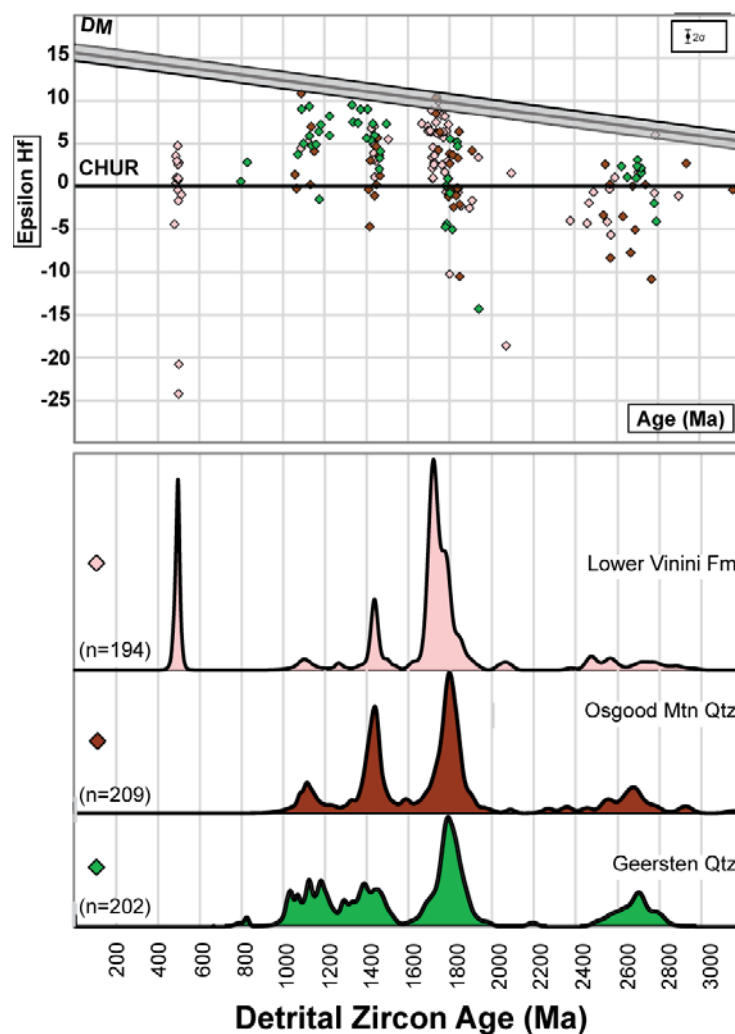


Figure 18. U-Pb ages and Hf isotope data for Roberts Mountains allochthon (RMA) and select Laurentian passive margin strata. The data from the Osgood Mountains Quartzite and the Geersten Canyon Quartzite are from Gehrels and Pecha (2014). The lower Vinini data includes those from this study ( $n = 189$ ) and from Gehrels and Pecha (2014) ( $n = 105$ ). Diagrams and symbols are as in Figure 13.



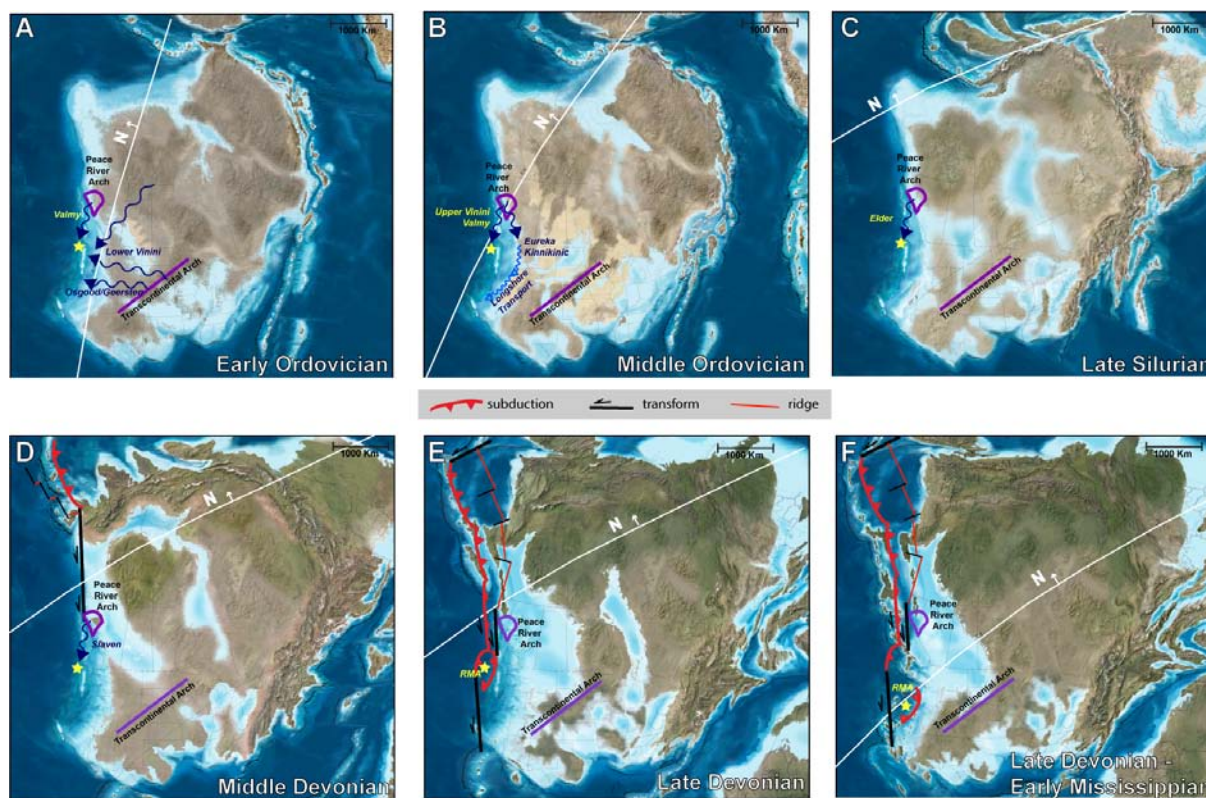


Figure 19. Paleogeographic maps of Laurentia from Middle Ordovician through Mississippian time (Blakey, 2013). Stars represent the depositional basin of Roberts Mountains allochthon (RMA) strata. White lines show the approximate position of the paleoequator. Blue wavy lines show approximate depositional pathways of units discussed. Transcontinental Arch (Sloss, 1988) and Peace River Arch (Ross, 1991) are superimposed. (A) Early Ordovician time. The lower Vinini is derived from the central craton and Transcontinental Arch; the Valmy Formation is derived from the Peace River Arch. (B) Middle Ordovician time. The upper Vinini and Valmy formations are shed from the Peace River Arch into an oceanic basin. The Eureka Quartzite is also derived from the Peace River Arch and transported via longshore current along the western Laurentian margin. (C) Late Silurian time. The Elder Sandstone is shed from the Peace River Arch region. (D) Middle Devonian time. The Slaven Chert is derived from the Peace River Arch. A Scotia-style arc has moved to the western margin of northern Laurentia, and a sinistral transpressional fault system has developed along the western margin. RMA strata are tectonically transported south along the margin by this fault system. (E) Late Devonian time. Subduction has initiated along much of the western margin of Laurentia, moving the RMA strata onto the craton. (F) Early Mississippian time. The Antler orogeny has uplifted the RMA strata into a highland on the western Laurentian margin.

## **8. Conclusions**

Sedimentological analyses provide a further constraint and suggest that the Ordovician passive margin sandstones are not the source of the RMA strata, but rather that these strata have a common source. Finney and Perry (1991) proposed that the Eureka Quartzite (an extensive Ordovician passive margin unit) was the source of the sandstones in the younger sections of the Vinini and Valmy formations of the RMA. However, the grains of the Ordovician passive margin sandstones are more texturally mature than those of the RMA strata, whose grains are coarser, larger, and more poorly sorted (Ketner, 1966). The more mature shelf sands such as the Eureka Quartzite and Mount Wilson Formation could not be the source of the more immature RMA sandstones. The RMA and passive margin sandstones have similar U-Pb age spectra and Hf isotope ratios (Fig. 14) and share a common source in the PRA region.

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### **Chapter 3**

Detrital zircon geochronology and HF isotope geochemistry of the Harmony Formation of Nevada: New insights into provenance, transport, and age

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#### **1. Abstract**

U-Pb geochronology and Hf-isotope geochemistry of detrital zircons provide new insights into the origin, age, and tectonic evolution of the enigmatic Harmony Formation, commonly considered a unit of the Roberts Mountains allochthon of north-central Nevada. Using laser-ablation inductively coupled plasma mass spectrometry, ten arenite samples were analyzed for U-Pb ages, and eight of these samples were further analyzed for Hf-isotope ratios. Three of the sampled arenite units have similar U-Pb age peaks and Hf-isotope ratios, including a 1.0-1.4 Ga peak with  $\epsilon_{\text{Hf}}$  values of +12 to -3 and a 2.5-2.7 Ga peak with  $\epsilon_{\text{Hf}}$  values of +7 to -5. The remaining seven samples differ significantly; each of these has age peaks of 1.7-1.9 Ga with  $\epsilon_{\text{Hf}}$  of +10 to -20 and age peaks of 2.3-2.7 Ga with  $\epsilon_{\text{Hf}}$  of +6 to -8. The data confirm the subdivision of the Harmony Formation into two petrofacies: quartzose (Harmony A) and feldspathic (Harmony B) (Gehrels et al., 2000a). The three samples that revealed the 1.0–1.4 and 2.5–2.7 Ga peaks are the

Harmony A, which originated in the central Laurentian craton. The other seven units sampled are the Harmony B, which was sourced from eastern Alberta-western Saskatchewan, north of the Harmony A source. We propose that all of the Harmony Formation strata were deposited near eastern Alberta and subsequently tectonically interleaved with the Roberts Mountains allochthon strata. The entire package was tectonically transported south along the Laurentian margin and then emplaced eastward onto the craton during the Late Devonian-Early Mississippian Antler orogeny.

## **2. Introduction**

Early Paleozoic time along the western Laurentian margin has commonly been interpreted as a quiescent interval (e.g., Poole et al., 1992; Dickinson, 2009, and references cited therein). The final Neoproterozoic rifting that separated the Rodinian supercontinent lasted from ca. 570-520 Ma and was followed by the deposition of passive margin sediments through mid-Devonian time (Poole et al., 1992; Dickinson, 2009; Yonkee et al., 2014). The quiescent interval came to an end with the Antler orogeny, during which the Roberts Mountains allochthon (RMA) was emplaced onto the western Laurentian margin (Dickinson, 2006). The RMA is an internally disrupted package of Cambrian–Devonian oceanic sediments, primarily composed of chert, argillite, quartzose turbidites, and some pillow basalts and volcanogenic debris flows (Doebrich, 1994; Dickinson, 2006; Dickinson, 2009) (Fig. 20).

The Harmony Formation of north-central Nevada presents an opportunity to investigate the tectonic evolution of western Laurentia. It is an enigmatic, fault-bounded,

feldspathic arenite mapped as a unit of the RMA (Madrid, 1987; Doebrich, 1994; Theodore, et al., 1994; Dickinson, 2006) (Fig. 20). Its age, provenance, and stratigraphic relationship with other RMA strata have long been controversial. It is primarily a texturally immature feldspathic arenite, which indicates minimal reworking and therefore minimal transport to the basin of deposition.

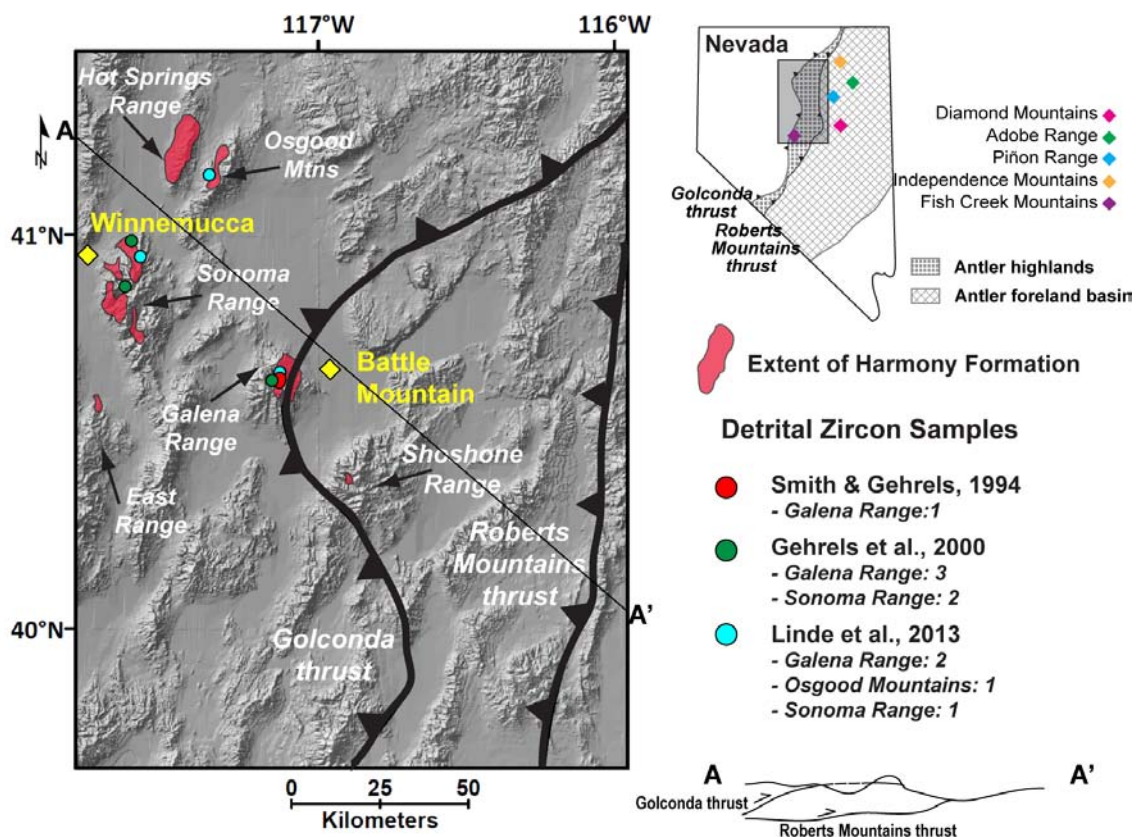


Figure 20. Map of north-central Nevada, showing sample locations (colored symbols), number of samples in each range, and the traces of the Roberts Mountains and Golconda thrusts. Some RMA rocks crop out to the west of the Golconda thrust in tectonic windows through the allochthon as shown in a notional, not-to-scale cross-section. The Antler orogenic highlands and Antler Foreland Basin are shown on the inset state map, with other Nevada locations mentioned in the text. Thrust traces are after Dickinson (2006); Antler highlands and basin are after Poole (1974). (Harmony Formation mapping: Ferguson et al., 1951; Roberts, 1964; Hotz and Willden, 1964; Gilluly, 1967; Madrid, 1987; Doebrich, et al., 1994; Theodore et al., 1994; Jones 1997a, 1997b). An inset map of Nevada shows where the larger scale map is located, as well as the location of other mountain ranges mentioned in the text. UTM coordinates of Harmony Formation samples are compiled in Table 4.

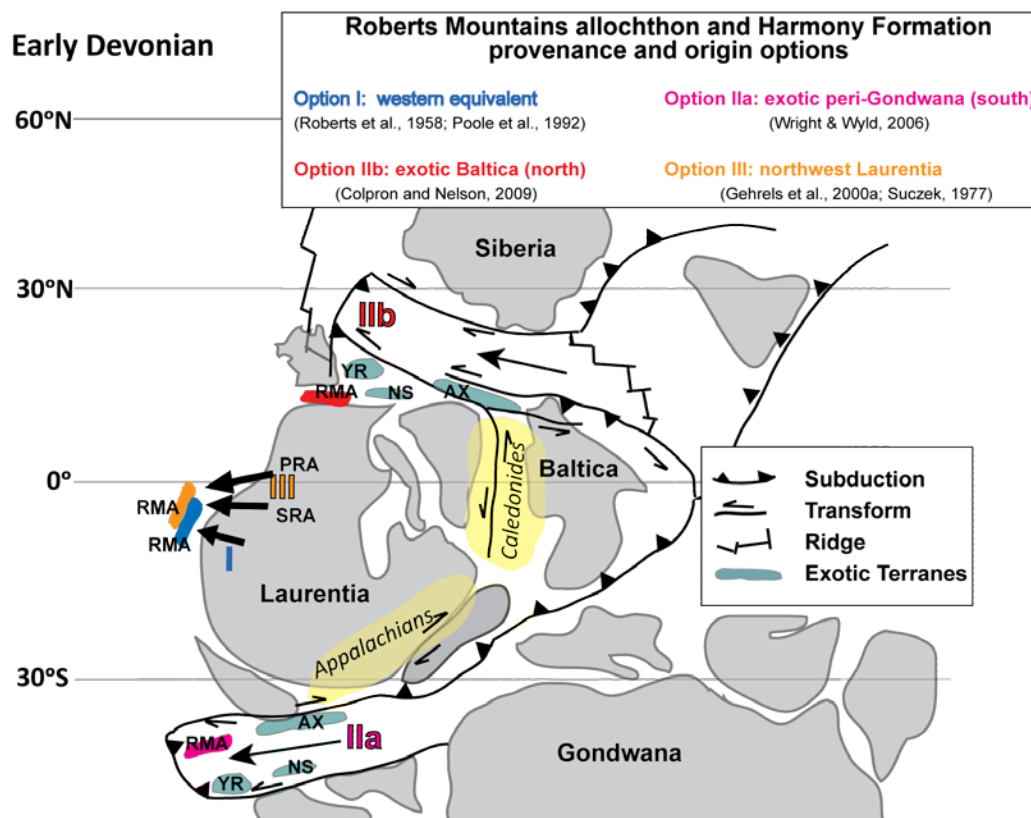


Figure 21: Contrasting tectonic models proposed to explain the source and tectonic transport of the Roberts Mountains allochthon and Harmony Formation, shown in Early Devonian time. Exotic terranes include the Alexander, Klamath, and Northern Sierran terranes. AX—Alexander Terrane; YR—Yreka Terrane; NS—Northern Sierra Terrane; RMA—Roberts Mountains allochthon; PRA—Peace River Arch; SRA—Salmon River Arch. Map is after Colpron and Nelson (2009).

Three general models have been invoked to explain the source, deposition, and transport of the RMA and the Harmony Formation (Fig. 21). The first model proposes that the oceanic sediments of the RMA and Harmony Formation are the coeval deep marine equivalent of the autochthonous passive margin strata of eastern Nevada (e.g., Roberts et al., 1958; Burchfiel and Davis, 1972; Poole et al., 1992) (Fig. 21, option I). This model is based upon the similar ages of the autochthonous and allochthonous strata, and the relative physical proximity of the proposed basin of deposition, directly offshore



the eventual location of the emplaced strata. The second model calls for an exotic, extra-Laurentian origin for the RMA and Harmony Formation and subsequent tectonic transport to a location offshore the emplacement. Two such models were proposed; one infers a peri-Gondwanan origin for the RMA and Harmony Formation with subsequent transport around the southern margin of Laurentia (Wright and Wyld, 2006) (Fig. 21, option IIa). The other model invokes origin in Baltica and subsequent tectonic transport around the northern margin of Laurentia (Colpron and Nelson, 2009) (Fig. 21, option IIb). These models are based on similarity of detrital-zircon ages and geologic history between the RMA and Harmony Formation and other terranes interpreted to be exotic and far traveled (Wright and Wyld, 2006; Colpron and Nelson, 2009). The third model calls for a distant, though still Laurentian, origin for the RMA and Harmony Formation and subsequent sedimentary transport to a location offshore western Laurentia with subsequent emplacement (Fig. 21, option III). Gehrels et al. (2000a) interpreted a source for the RMA in the Peace River Arch region of eastern Canada, based on similarity of detrital zircon U-Pb ages. Suczek (1977) suggested that the Harmony Formation originated in the Salmon River Arch region of Idaho, due to lithological similarity between the Harmony Formation and potential sedimentary sources (Fig. 22).

The age of the Harmony Formation has long been controversial. The earliest mappers interpreted the unit as Mississippian, based on Pennsylvanian limestone deposited unconformably on the Harmony Formation (Ferguson et al., 1951). Later workers interpreted the Harmony Formation as Cambrian, based on Cambrian trilobites found in the unit (Hotz and Willden, 1964). Jones (1997a) extracted a single conodont interpreted as Devonian and thus interpreted a Devonian age. Based on Jones (1997a) and

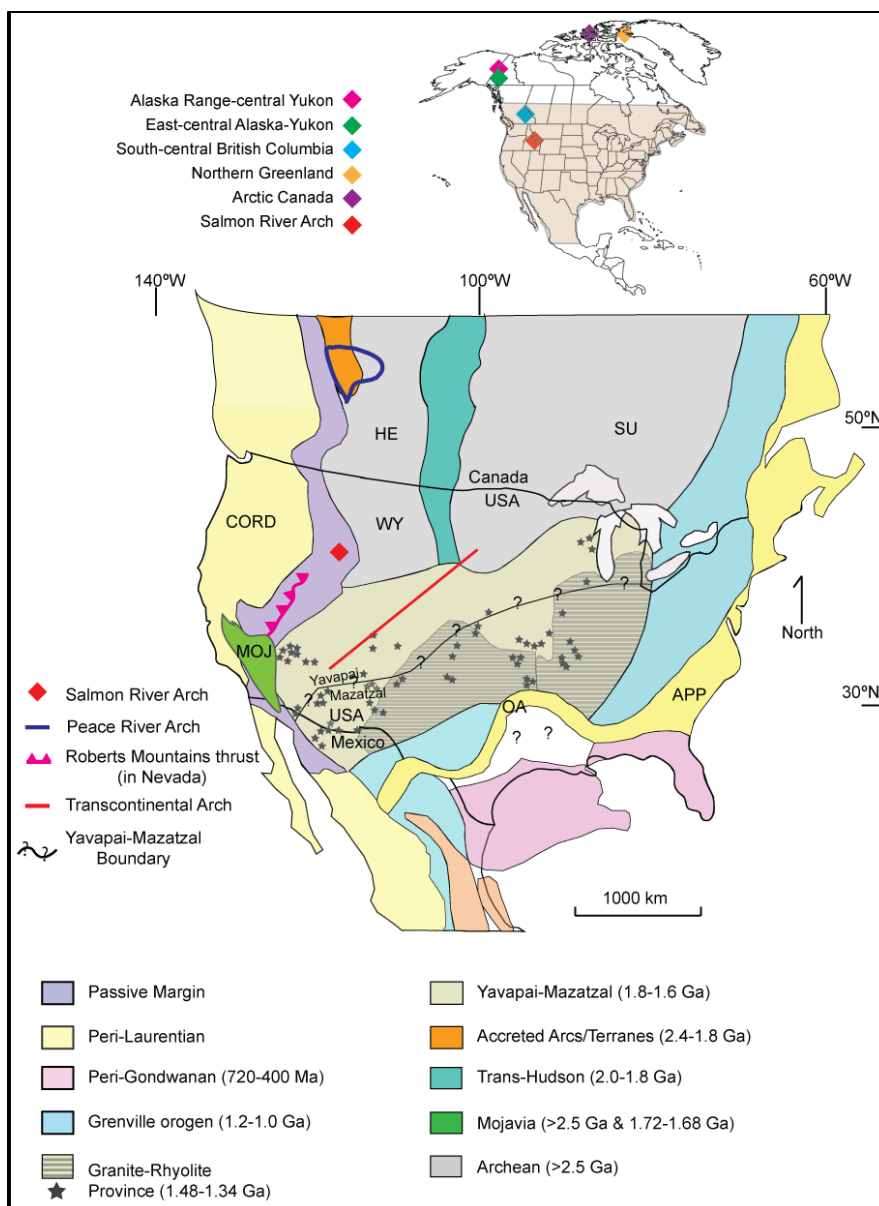


Figure 22: Locations of the main age provinces in North America that are potential source terranes for western Laurentian strata. The location of Transcontinental Arch is shown as a red line (Sloss, 1988); the Peace River Arch is shown as a blue line and the Salmon River Arch is a red diamond. The trace of the Roberts Mountains thrust in Nevada is shown. WY—Wyoming province; HE—Hearn province; SU—Superior province; CORD—Cordilleran; APP—Appalachian; OA—Ouachita-Marathon. Figure is after Gehrels et al. (2011) and compiled from Bickford et al. (1986), Hoffman (1989), Ross (1991), Burchfiel et al. (1992), Anderson and Morrison (1992), Bickford and Anderson (1993), Van Schmus et al. (1993), Villeneuve et al. (1993), Dickinson and Lawton (2001), and Dickinson and Gehrels (2009). An inset map of North America shows other locations referred to in the text.

interpreted geological relationships, Ketner et al. (2005) interpreted a Devonian-Mississippian age.

Analysis of detrital-zircon provenance is a useful method for understanding links between sedimentary units and source terranes (e.g. Fedo et al., 2003; Cawood, et al., 2007). To investigate the source, depositional history, and tectonic evolution of the Harmony Formation, we analyzed detrital zircons to obtain U-Pb ages and Hf-isotope ratios. U-Pb ages are important for identifying and then characterizing the provenance of sedimentary strata, and for comparison among sedimentary units (Gehrels, 2000; Fedo et al., 2003; Gehrels, 2012; Gehrels, 2014). Hf-isotope compositions can be used to determine the geochemical character of the magma in which the zircons crystallized (Bahlburg et al., 2011; Gehrels and Pecha, 2014). Using our new detrital-zircon data, we interpret that the Harmony Formation is from western Laurentia and not exotic. We propose that all of the Harmony Formation strata were deposited near eastern Alberta and subsequently tectonically interleaved with the remainder of the RMA strata. The entire package was tectonically transported south along the Laurentian margin and subsequently emplaced eastward onto the craton during the Late Devonian-Early Mississippian Antler orogeny.

### **3. Geologic Setting**

#### **a. Regional Tectonostratigraphic Framework**

The North American craton contains several Proterozoic and Archean crustal provinces that are geologically distinct source terranes for the upper Proterozoic and lower Paleozoic Laurentian strata (e.g., Gehrels et al., 2011 and references cited therein)

(Fig. 22). The Yavapai-Mazatzal Province (1.8-1.6 Ga) extends across central North America (Hoffman, 1989; Whitmeyer and Karlstrom, 2007) (Fig. 22). The granite-rhyolite province crops out within the Yavapai-Mazatzal Province (Bickford, et al., 1986; Anderson and Morrison, 1992). The Yavapai-Mazatzal Province is bounded on the north and northwest by the Trans-Hudson orogenic terrane (Whitmeyer and Karlstrom, 2007) (2.0-1.8 Ga) and Archean rocks (> 2.5 Ga) of the Wyoming and Superior provinces (Fig. 22), on the east and southeast by the terranes of the Grenville orogen (1.2-1.0 Ga) (Hoffman, 1989), and on the west by the Mojavia terrane (> 2.0-2.4 Ga with 1.7-1.8 Ga arcs) (Whitmeyer and Karlstrom, 2007; Nelson et al., 2011) (Fig. 22).

Detrital-zircon data demonstrate that sources for the upper Neoproterozoic-lower Paleozoic western Laurentian passive margin changed between upper Proterozoic and Lower Cambrian time (Linde et al., 2014a, and references cited therein). A significant sediment source for the upper Proterozoic Laurentian passive margin strata from the northwest U.S. to Sonora, Mexico was the 1.2–1.0 Ga Grenville orogen of southern and eastern North America (e.g., Rainbird et al., 1997; Lawton et al., 2010; Rainbird et al., 2012; Gehrels and Pecha, 2014; Yonkee et al., 2014; Linde et al., 2014a) (Fig. 22). In contrast, the 1.8–1.6 Ga Yavapai-Mazatzal and 1.48-1.34 Ga mid-continent granite-rhyolite provinces are the more predominant sediment sources of many strata higher in the passive-margin section (e.g., Lawton et al., 2010; Gehrels and Pecha, 2014; Linde et al., 2014a; Yonkee et al., 2014) (Fig. 22).

The Roberts Mountains allochthon is commonly interpreted as a sequence of Cambrian through Devonian oceanic sedimentary rocks emplaced structurally eastward onto the western Laurentian craton during the Late Devonian-Early Mississippian Antler

orogeny (Roberts et al., 1958; Poole et al., 1992). The Antler orogeny may be linked to tectonism that is inferred along the western margin of Laurentia from Late Devonian through Early Mississippian time; a model that explains this tectonism may also explain possible tectonic transport and emplacement of the RMA. This tectonism is recorded from Alaska to British Columbia and is younger to the south. Middle to Late Devonian continental arc magmatism occurred in the Alaska Range and central Yukon (Piercey et al., 2006) (Fig. 22). Upper Devonian – Lower Mississippian felsic igneous and meta-igneous rocks record bimodal volcanism in east-central Alaska and the Yukon (Dusel-Bacon et al., 2006) (Fig. 22), and in south-central British Columbia, a Late Devonian continental arc and backarc developed (Paradis et al., 2006) (Fig. 22).

Colpron and Nelson (2009) proposed a model that links the Antler orogeny and Devonian-Mississippian tectonism of the western Laurentian margin. They propose a "Northwest Passage" opened between Laurentia and Siberia in mid-Paleozoic time, within which an arc developed along the northern Laurentian margin in the Early Devonian (Colpron et al., 2007; Colpron and Nelson, 2009) (Fig. 21). The Alexander terrane, and other fragments such as the eastern Klamath and northern Sierran terranes commonly interpreted as exotic to Laurentia (e.g., Bazard et al., 1995; Gehrels et al., 1996; Grove et al., 2008; Colpron and Nelson, 2009; Beranek et al., 2013), were transported from their origin in Baltica to northwestern Laurentia through the "Northwest Passage" by means of the westward migration of the subduction zone (Fig. 21). By Middle Devonian time, a sinistral transpressional system developed at the southern end of this passage. The transform fault gradually propagated southward along western

Laurentia and transported the terranes and fragments south along the margin. Colpron and Nelson (2009) inferred that the progressively younger tectonism southward along the Laurentian margin records the southward propagation of the transpressional system. They propose that this fault system could have provided the weakness along which Devonian shortening initiated, resulting in the emplacement of the RMA (Colpron and Nelson, 2009).

RMA strata crop out in north-central Nevada between the Roberts Mountains thrust on the east and the Golconda thrust on the west, though some units are exposed in tectonic windows west of the Golconda thrust (Fig. 20). Strata of the allochthon structurally overlie coeval rocks of the western Laurentian passive margin (e.g., Schuchert, 1923; Kay, 1951; Roberts et al., 1958; Madrid, 1987). RMA strata are highly deformed, and include imbricated older over younger thrust sheets (Evans and Theodore, 1978; Oldow, 1984; Noble and Finney, 1999). The metamorphic grade of RMA rocks is generally greenschist facies or lower (Gehrels et al., 2000a). Sediments shed from the rising Antler highlands filled the Antler foreland basin, east of the Antler orogen and west of the Laurentian craton, between Devonian and Early Mississippian time (Poole, 1974; Trexler et al., 2003) (Fig. 20).

#### **b. The Harmony Formation**

The Harmony Formation is primarily a texturally immature feldspathic arenite; it is often interpreted as a part of the RMA, due to its structural imbrication with, and position above, RMA units (Madrid, 1987). The Harmony Formation was first described in Harmony Canyon of the Sonoma Range as a coarse micaceous and feldspathic sandstone

(Ferguson et al., 1951) (Figs. 20 and 23). Interbedded shale and limestone and graded beds are common (Ferguson et al., 1951; Roberts, 1964) (Fig. 23). Ferguson et al. (1951) estimated the thickness of the unit as up to 1524m (5,000 ft). The contact between the exposed base of the Harmony Formation and other units are faults (Fig. 24); therefore, the true thickness of the Harmony Formation is unknown (Ferguson et al., 1951; Ferguson et al., 1952; Roberts, 1964; Hotz and Willden, 1964).



Figure 23. The Harmony Formation in Little Cottonwood Canyon, Galena Range, Nevada. Left photo: the red bar shows a 1m trekking pole. Right photo: the scale totals 15cm. Photos by Gwen Linde.

The Harmony Formation has been interpreted as a turbidite deposit. Roberts (1964) observed the extensive graded bedding, fragments of marine fossils, and lack of cross-beds and ripple marks, and concluded that the Harmony was deposited by density currents. In comparison, Hotz and Willden (1964) suggested that the Harmony

Formation was deposited in an ocean basin by turbidity currents, based on its textural and mineralogical immaturity. Suczek (1977) documented features typical of turbidites, including sole marks, graded beds, plane beds, and interbedded marine shales. Suczek (1977) also documented facies A, B, and C of the classic Bouma (1963) sequence and interpreted the depositional environment of the Harmony Formation as turbidites deposited on inner and middle submarine fans.

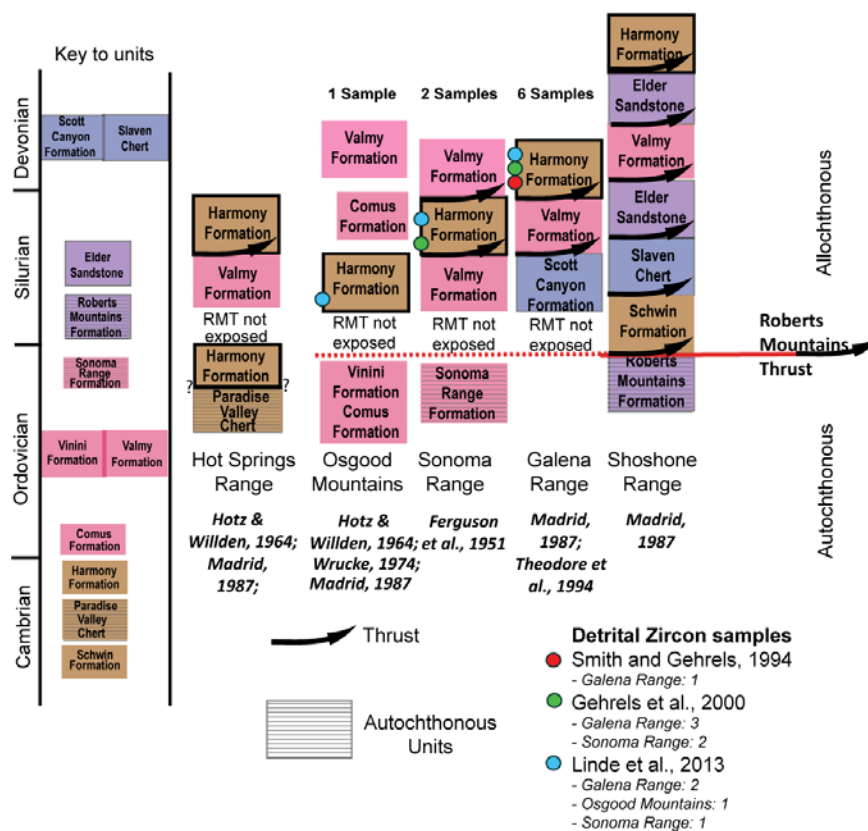


Figure 24. Tectonostratigraphic diagram of units of the Roberts Mountains allochthon in north-central Nevada mountain ranges in which the Harmony Formation crops out, showing locations of detrital zircon samples. Units are shown in their physical, structurally superimposed, and not chronostratigraphic order. Most units are internally disrupted with multiple imbricate thrusts not shown on this chart. Autochthonous units are shown with crosshatching. RMT—Roberts Mountains Thrust, as mapped by Ferguson et al. (1951); Hotz and Willden (1964); Wrucke (1974); Madrid (1987); and Theodore et al. (1994).



The Harmony Formation crops out in several ranges in north-central Nevada (Fig. 20). In the Hot Springs Range the Harmony Formation rests on the Cambrian Paradise Valley Chert with a contact that is interpreted as depositional (Hotz and Willden, 1964), structural (Madrid, 1987), or obscured (Jones, 1997a) (Figs. 20 and 24). In the other locations where the Harmony Formation is exposed, it is mapped as the top thrust in the RMA strata, except in the Sonoma Range where it is found both structurally above and below the Valmy Formation, which consists primarily of arenite, chert, and greenstone (Ferguson et al., 1951; Roberts, 1964; Hotz and Willden, 1964; Gilluly, 1967; Madrid, 1987) (Figs. 20 and 24). In the Sonoma Range, the Pennsylvanian-Permian Antler Peak Limestone was deposited on the Harmony Formation (Ferguson et al., 1951), and in the Galena Range, the Pennsylvanian Battle Formation, a conglomerate, was deposited on the Harmony Formation (Roberts, 1964). In both cases, the contact is an angular unconformity.

The Harmony Formation is inferred to have been exposed as part of the advancing Antler orogen, because clasts of the unit are found in Mississippian strata of the Antler foreland basin, which is interpreted to have filled from detritus shed from the Antler highlands (Dickinson, 2009). These Harmony Formation clasts are found in strata in the Diamond Mountains (Harbaugh, 1980), the Adobe Range (Ketner and Ross, 1990), the Independence Mountains, the Piñon Range, the East Range, and the Fish Creek Mountains (Ketner, 1998) (Fig. 20). Clasts and olistoliths of Cambrian trilobite-bearing Harmony Formation strata are in the Devonian Scott Canyon Formation in the Galena Range (Theodore, et al., 1994; Doebrich, 1994; this study) (Figs. 21 and 25). Because of the contained clasts and olistoliths of Harmony rocks, Doebrich (1994) speculated that

the Scott Canyon Formation was deposited in a volcanogenic debris flow sourced from a highland and/or the advancing allochthon that contained Harmony rocks.

The age of the Harmony Formation has long been controversial. The Harmony Formation was first mapped in the Sonoma (Ferguson et al., 1951) and Galena Ranges (Roberts 1951) and interpreted to be no younger than Mississippian because Pennsylvanian strata were deposited on the Harmony beds. Hotz and Willden (1964) interpreted the age of the Harmony Formation as Late Cambrian, based on trilobite faunas of Laurentian affinity that are re-sedimented in Harmony limestone lenses in the



Figure 25. The Devonian Scott Canyon Formation with Harmony Formation clast inclusion, in Galena Canyon, Galena Range, Nevada. The red arrow points to the Harmony Formation clast. Photo by Gwen Linde.

Hot Springs Range and the Osgood Mountains (Fig.20). The Harmony Formation is therefore no older than Cambrian. Doebrich et al. (1994) and Theodore (1994) interpreted the Harmony Formation as Cambrian based on both the trilobites and their field observation of clasts and olistoliths of Harmony Formation in the Devonian Scott Canyon Formation in the Galena Range (Figs. 20 and 25). Jones (1997a and 1997b) processed limestone from the Harmony Formation in the Hot Springs Range and reported a single conodont interpreted as Devonian. Later workers interpreted the Harmony Formation as Upper Devonian–Lower Mississippian, based on Jones, 1997a and 1997b (Ketner et al., 2005; Ketner, 2008). These studies interpreted the limestone beds which contain the Cambrian trilobites as resedimented *mélange* blocks (in the Osgood Mountains; Jones, 1997a and 1997b) or a separate unit which they mapped as "Cambrian Limestone" (in the Hot Springs Range; Jones, 1997a and 1997b).

Though it had been mapped as one lithostratigraphic formation (e.g. Roberts, 1951; Gilluly, 1967; Doebrich, 1994), the Harmony Formation comprises two distinct sub-units that can be distinguished by their detrital-zircon spectra and with less clarity, by composition (Fig. 26). Gehrels et al. (2000a) designated these as petrofacies, and called the generally more quartzose, more texturally mature "Harmony A" and the generally more feldspathic, more texturally immature "Harmony B" (Gehrels et al., 2000a). The Harmony B crops out everywhere the Harmony Formation is found, while the Harmony A has been described only in Little Cottonwood Canyon of the Galena Range (Fig. 20). Dickinson and Gehrels (2000) interpreted the contact between the Harmony A and B in Little Cottonwood Canyon as sharp, but concordant, and proposed that the two petrofacies recorded two successive submarine fan deposits.

The provenance of the Harmony Formation has also been much debated. Some workers proposed that the Harmony Formation was derived from a possible western land mass, because they could not find an eastern feldspathic granitoid source (Ferguson et al., 1951; Ketner, 1977). Several studies have interpreted the Salmon River Arch of western Idaho as the source of Harmony Formation sands, based on lithological similarities between potential granitic sources and the Harmony B (Rowell et al., 1979; Schweickert and Snyder, 1981; Suczek, 1977; Stewart and Suczek, 1977). Based on early U-Pb analyses of detrital zircons, Wallin (1990) and Smith and Gehrels (1994) thought the source was south of the present position of Harmony Formation outcrops. Based on more

<b>SAMPLE LOCATION UTM: NAD 83 (11N)</b>	<b>Easting</b>	<b>Northing</b>	<b>Sample from</b>	<b>Harmony A/B</b>	<b>Hf?</b>
Elbow Canyon <i>Sonoma Range</i>	441026	4514450	Gehrels et al., 2000	B	yes
Harmony Canyon <i>Sonoma Range</i>	446225	4533064	Gehrels et al., 2000	B	yes
Kluncy Canyon <i>Sonoma Range</i>	0447927	4531085	Linde et al., 2013	B	no
Gough's Canyon <i>Osgood Mountains</i>	0471723	4556266	Linde et al., 2013	B	no
Little Cottonwood Canyon #10 <i>Galena Range</i>	0490510	4495682	Linde et al., 2013	A	yes
Little Cottonwood Canyon #9 <i>Galena Range</i>	0491232	4495602	Linde et al., 2013	B	yes
Little Cottonwood Canyon #4 <i>Galena Range</i>	0490623	4494977	Gehrels et al., 2000	B	yes
Little Cottonwood Canyon #3 <i>Galena Range</i>	490459	4494916	Gehrels et al., 2000	B	yes
Little Cottonwood Canyon #2 <i>Galena Range</i>	490529	4495008	Gehrels et al., 2000	A	yes
Little Cottonwood Canyon #1 <i>Galena Range</i>	490600	4494792	Smith & Gehrels, 1994	A	yes

Table 4: Locations of samples analyzed in this study referenced to UTM locations.

extensive U-Pb analyses of detrital zircons (Fig. 26), Gehrels et al. (2000a), proposed that the Harmony A originated in central or southern Laurentia, and the Harmony B was derived from northern Laurentia (Fig.22). Later workers suggested that the Harmony Formation was an exotic accreted terrane, based on the inability to find any local source terranes (Ketner et al., 2005; Jones-Crafford 2008).

#### **4. Methods**

This study describes U-Pb and Hf-isotope zircon analyses from ten arenite samples of the Harmony Formation (Figs. 20 and 24; Table 4). Zircons from six of these samples were previously analyzed by ID-TIMS (Smith and Gehrels, 1994; Gehrels et al, 2000a) (Figs. 20 and 26; Table 4). This study reports ~200 additional U-Pb analyses and ~50 new Hf-isotope analyses of zircons from these same six samples by LA-ICPMS (Fig. 26). Moreover, ~100 LA-ICPMS U-Pb zircon analyses were performed from four additional arenite samples; these samples are previously described in Linde et al. (2013). On two of these samples, ~25 of the grains were further analyzed by LA-ICPMS for Hf-isotope ratios. The re-analysis of the original six arenite samples incorporated new methodology. Grains were selected randomly rather than by color or morphology, we increased the sample size to 200 grains for the U-Pb ages to obtain a more statistically significant sample size, and we incorporated Hf-isotope analyses. The sampling in the 2000s was accomplished to broaden the geographic scope into new drainages in the Sonoma Range and the Osgood Mountains and to test the finding by Gehrels et al. (2000a) of two distinct petrofacies in the Galena Range. A smaller number of grains were analyzed for these

samples in 2009; reduced cost and increased speed of analyses since that time has made analyses of greater numbers of grains more affordable and practical.

We separated and analyzed zircon at the University of Arizona LaserChron facility using standard techniques to yield a best age distribution reflective of the true distribution of detrital-zircon ages in each sample (Gehrels and Pecha, 2014). We selected zircon grains at random, choosing from all areas of the sample mount and not biasing selection by size, crystal shape, or color. We excluded grains with fractures, inclusions, or zonation. We positioned the analyses in the zircon cores to reduce the chance of analyzing overgrowths that might be compromised by Pb loss. Hf analyses were not conducted on every arenite sample due to cost constraints. Hf analyses were positioned on top of the U-Pb ablated pits to ensure that Hf-isotope data were collected from the same growth domain as the U-Pb ages. The data was collected over multiple trips. LA-ICPMS analyses in 2009 used a New Wave UP193 HE excimer laser, and in 2013-2014, the analyses were performed using a Photon Machines Analyte G2 excimer laser. In all analyses, the laser was connected to a Nu Plasma multi-collector high resolution ICPMS, using methods described in Gehrels and Pecha (2014). For both U-Pb and Hf analyses, we used a beam diameter of 35  $\mu\text{m}$ ; for a few very small zircon grains we used a beam diameter of 30  $\mu\text{m}$ .

Analytical results are displayed graphically on normalized-probability plots and Hf-evolution diagrams for visual comparison among zircon populations (Figs. 26, 27, 28, and 30). Hf-evolution diagrams display epsilon Hf ( $\epsilon\text{Hf}_{(t)}$ ) values at the time of zircon crystallization (Figs. 27, 28, and 30). For U-Pb analyses, measured ion intensities from the Nu HR ICPMS are imported into a data reduction program, "agecalc," which reduces

data, calculates ages, applies corrections and filters, and creates data tables, concordia diagrams, histograms, and normalized-probability plots (Gehrels and Pecha, 2014). For Hf analyses, a data-reduction program, "hfcalf," reduces data, calculates Hf ratios, applies corrections, and creates data tables and Hf-evolution charts (Gehrels and Pecha, 2014). We rejected U-Pb analyses for which uncertainties are greater than 10%, discordance is greater than 20%, and reverse discordance is greater than 5%. For Hf analyses, we applied a 2-sigma filter (Gehrels and Pecha, 2014).

In addition to visual evaluation, we also compared age distributions using the Kolmogorov-Smirnov (K-S) statistic (Guynn and Gehrels, 2006) (Table 5). The K-S calculates whether a statistically significant difference exists between two distributions.  $P$  (probability)  $<0.05$  indicates  $>95\%$  probability that two U-Pb distributions are not the same. The K-S statistic is sensitive to proportions of ages present, and a low  $P$  value may not indicate statistical difference between the ages of populations but rather that the proportions of age peaks are different (Gehrels, 2012).

## **5. Results: Uranium-Lead Ages and Hf-isotope ratios**

Although the ID-TIMS data are similar to our new LA-ICPMS data, there are variations in the proportions of age groups (Fig. 26). Different grain selection procedures were used in the two studies. In the ID-TIMS study, zircon crystals were selected based on color and morphology (Gehrels et al., 2000a). For the current LA-ICPMS study, we attempted to select grains randomly from the entire population of grains, to obtain a more representative age distribution. The results and interpretations that follow are based upon the LA-ICPMS ages.

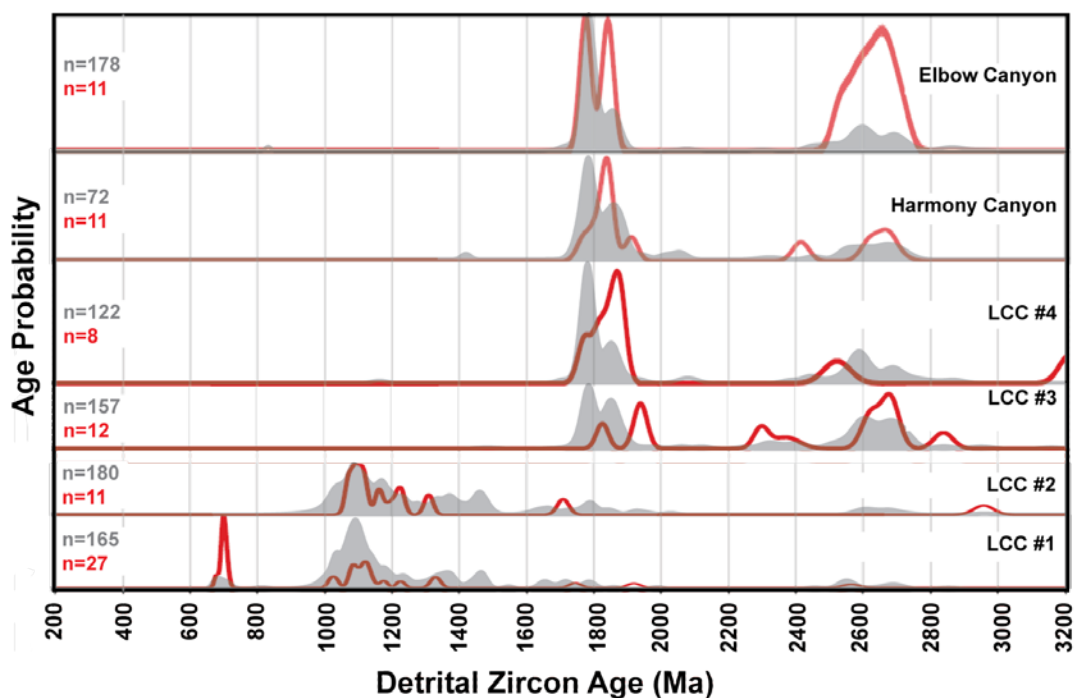


Figure 26. Normalized probability plots showing U-Pb ages of strata sampled. These data are analyses of the six samples collected and first analyzed in the 1990s and reanalyzed for this study. Red lines show the data from isotope-dilution thermal ionization mass spectrometry (Gehrels et al., 2000a); grey-filled curves are the data from laser-ablation inductively coupled plasma mass spectrometry (this study). Numbers of grains analyzed are shown.

The three Harmony A samples, LCC #1, #2, and #10, have similar U-Pb ages and Hf-isotope ratios and differ significantly from Harmony B (Figs. 26 and 27). In Harmony A samples, there is a major age peak at ca. 1.0-1.2 Ga, which includes 45-62% of the analyses from each sample (Figs. 26 and 27). These grains yielded  $\epsilon\text{Hf}_{(t)}$  ratios of +10 to -8 (Fig. 27). Smaller age populations from ca. 1.2-1.5 Ga, ca. 1.6-1.8 Ga, and ca. 2.5-2.8 Ga each comprise up to 15% of the zircons (Figs. 26 and 27). The age populations of 1.2-1.5 Ga yielded  $\epsilon\text{Hf}_{(t)}$  ratios of +12 to -1, the age populations of 1.6-1.8 Ga yielded  $\epsilon\text{Hf}_{(t)}$  ratios of +10 to -13, and the age populations of 2.5-2.8 Ga yielded  $\epsilon\text{Hf}_{(t)}$  ratios of +7 to -2



(Fig. 27). The LCC#1 sample also has a small ( $n = 7$ ) peak at ca. 685 Ma, consisting of 4% of the total analyses.

In comparison, the seven Harmony B samples, LCC #3, #4, and #9, and Kluncy, Gough's, Harmony, and Elbow Canyons, revealed a major age peak at ca. 1.7-1.9 Ga, consisting of 60-65% of the analyses (Figs. 26 and 27). These grains yielded  $\epsilon\text{Hf}(t)$  ratios of +1- to -20 (Fig. 27). The samples also yield smaller age populations at ca. 2.4-2.8 Ga, representative of 20% of the sample (Figs. 26 and 27). The age populations of 2.4-2.7 Ga yielded  $\epsilon\text{Hf}(t)$  ratios of +5 to -5 and the age populations of 2.7-2.8 Ga yielded  $\epsilon\text{Hf}(t)$  ratios of +7 to -9 (Fig. 27).

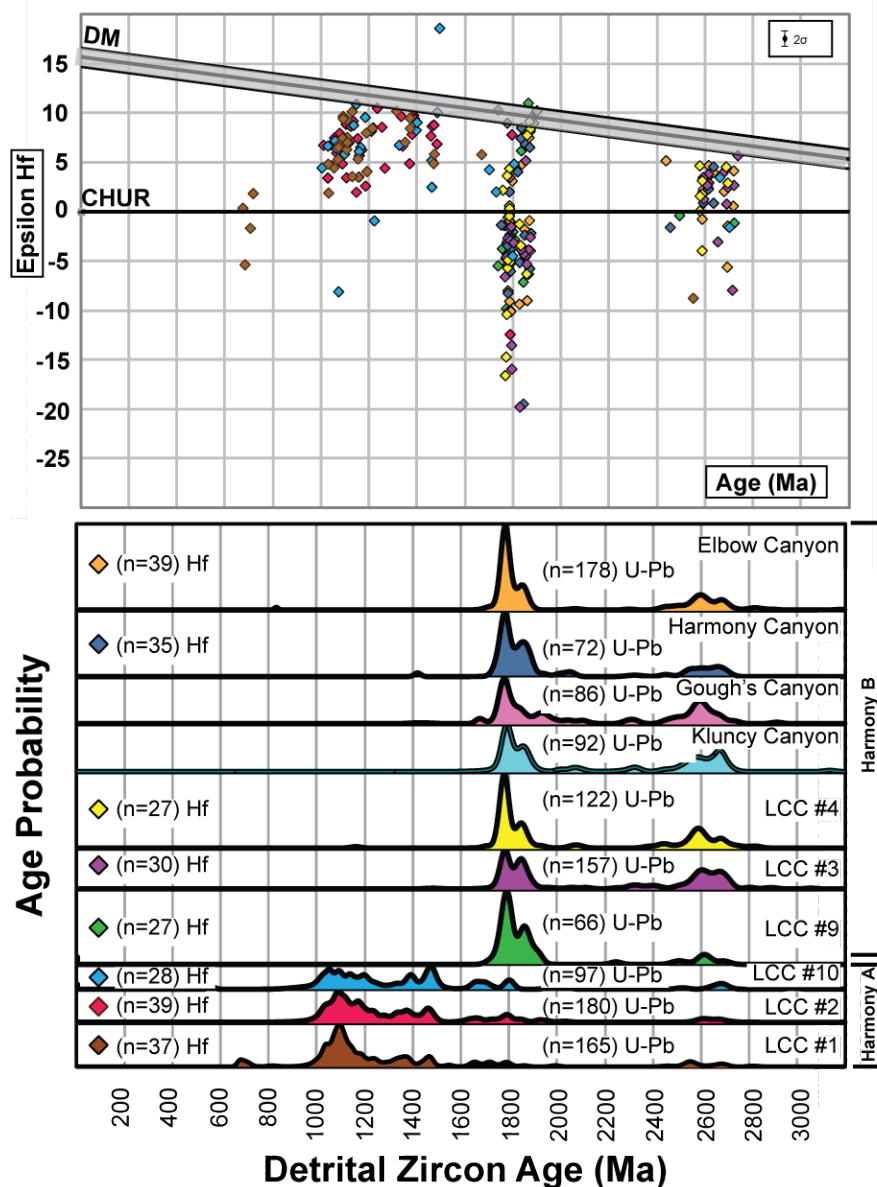


Figure 27. U-Pb ages and Hf isotope data for Harmony Formation. U-Pb dates were run for all sample grains; approximately  $\frac{1}{4}$  of these grains were analyzed for hafnium isotopes. The lower graph shows the normalized probability plots of U-Pb ages. The subdivision of Harmony A and B is indicated. The number of grains analyzed for U-Pb ages and Hf isotopes is shown. The upper graph shows Epsilon Hf values at the time of zircon crystallization ( $\epsilon\text{Hf}(t)$ ) vs. age for each sample. The average measurement uncertainty for all hafnium analyses on this chart in  $\pm$  epsilon units is shown in the upper right at the  $2\sigma$  level. Reference lines on the Hf plot are as follows: Depleted mantle (DM) is calculated using  $^{176}\text{Hf}/^{177}\text{Hf}=0.283225$ ;  $^{176}\text{Lu}/^{177}\text{Hf}=0.038513$  (Vervoort and Blichert-Toft, 1999); Chondritic uniform reservoir, CHUR, is calculated using  $^{176}\text{Hf}/^{177}\text{Hf}=0.282785$  and  $^{176}\text{Lu}/^{177}\text{Hf}=0.0336$  (Bouvier et al., 2008).

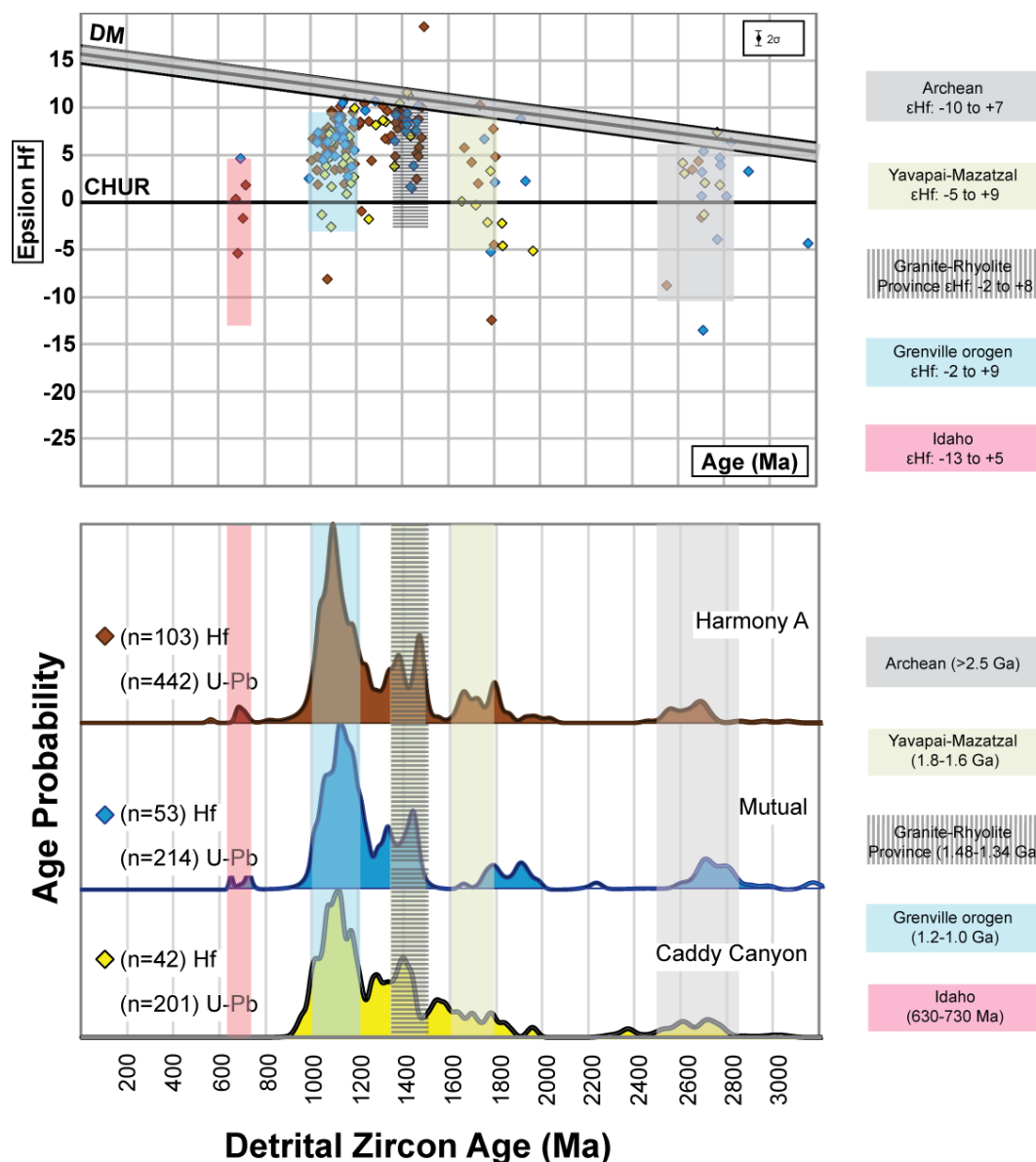


Figure 28. U-Pb ages and Hf isotope data for Harmony A samples and selected western Laurentian passive margin strata, showing the similarities between the U-Pb ages and Hf isotope analyses of the Harmony A and these passive margin strata. Data from the Neoproterozoic Mutual Formation and Caddy Canyon Quartzite are from Gehrels and Pecha (2014). Colored age bars that correspond to Laurentian basement terrane ages are superimposed over the U-Pb ages on the normalized probability plots. The ages are from references cited in Fig. 3. Colored Hf-isotope range bars that correspond to the same Laurentian terranes are shown on the Hf evolution diagram (Grenville: Bickford et al., 2010; Mueller et al., 2008. Granite-Rhyolite province: Goodge and Vervoort, 2006; Mueller et al., 2008. Yavapai-Mazatzal: Holm et al., 2013. Achaean: Rohr et al., 2008; Rohr et al., 2010. Idaho: Gaschnig, et al., 2013). Diagrams and symbols are as in Figure 27.

Table 5: K-S statistical analysis results. The Harmony A samples are shown with green highlights, central Laurentian passive margin strata are shown with blue highlights, Harmony B samples are shown with orange highlights, and southern British Columbia-northeastern Washington samples are shown with red highlights. Comparisons between units with values greater than 0.05 are highlighted in yellow.  $P < 0.05$  indicates >95% probability that two U-Pb distributions are not the same.

	LCC #1	LCC #2	LCC #10	Caddy Cyn Quartzite	Mutual Formation	LCC #9	LCC #3	LCC #4	Klunzy Canyon	Gough's Canyon	Harmony Canyon	Elbow Canyon	Horseshief Creek Group	Hamill Group	Addy Quartzite
LCC #1		0.077	0.010	0.013	0.007	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
LCC #2	0.077		0.196	0.838	0.357	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
LCC #10	0.010	0.196		0.588	0.227	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
Caddy Canyon Quartzite	0.013	0.838	0.588		0.689	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.000
Mutual Formation	0.007	0.357	0.227	0.689		0.000	0.000	0.033	0.000	0.000	0.745	0.090	0.041	0.061	0.006
LCC #9	0.000	0.000	0.000	0.000	0.000		0.000	0.018	0.942	0.199	0.001	0.000	0.009	0.000	0.000
LCC #3	0.000	0.000	0.000	0.000	0.000	0.033		0.018	0.085	0.154	0.376	0.982	0.028	0.005	0.003
LCC #4	0.000	0.000	0.000	0.000	0.000	0.000	0.942		0.085	0.154	0.376	0.982	0.028	0.005	0.003
Klunzy Canyon	0.000	0.000	0.000	0.000	0.000	0.000	0.000	0.085		0.722	0.024	0.011	0.139	0.004	0.000
Gough's Canyon	0.000	0.000	0.000	0.000	0.000	0.000	0.199	0.722	0.722		0.014	0.010	0.092	0.004	0.000
Harmony Canyon	0.000	0.000	0.000	0.000	0.000	0.000	0.154	0.024	0.024	0.014		0.010	0.147	0.065	0.010
Elbow Canyon	0.000	0.000	0.000	0.000	0.000	0.090	0.000	0.982	0.011	0.010	0.642		0.001	0.001	0.003
Horseshief Creek Group	0.000	0.000	0.000	0.000	0.000	0.041	0.009	0.028	0.139	0.092	0.147	0.001		0.000	0.000
Hamill Group	0.000	0.000	0.000	0.000	0.000	0.061	0.000	0.005	0.004	0.022	0.065	0.001	0.000		0.008
Addy Quartzite	0.000	0.000	0.000	0.000	0.000	0.006	0.000	0.003	0.000	0.000	0.010	0.003	0.000	0.008	

## **6. Discussion: Provenance and Age of the Harmony Formation**

To understand the tectonic evolution of the Harmony Formation, we must first understand its provenance. To interpret provenance, we compared the data from our study to known U-Pb ages and Hf-isotope data from other coeval sedimentary units, as well as Laurentian basement provinces.

### **a. Provenance of the Harmony A**

The Harmony A arenites are composed of detrital zircons that have potential sources in the Laurentian craton. We can account for these age spectra as follows: the 1.0-1.2 Ga grains are consistent with provenance in the Grenville orogen (Hoffman, 1989; Bickford and Anderson, 1993; Van Schmus et al., 1993) (Figs. 22 and 28), and the 1.3-1.5 Ga grains are similar to those with origin in the granite-rhyolite province of central Laurentia (Bickford et al., 1986; Hoffman, 1989; Anderson and Morrison, 1992; Bickford and Anderson, 1993; Van Schmus et al., 1993) (Figs. 22 and 28). The 1.6-1.8 Ga grains are consistent with provenance in the Yavapai-Mazatzal terranes and the 2.5-2.8 Ga grains are comparable with those originating in the Archean craton (Bickford et al., 1986; Hoffman, 1989; Ross, 1991; Van Schmus et al., 1993) (Figs. 22 and 28). The seven zircons from 673-716 Ma in sample LCC#1 sample have a potential source in Idaho. Several workers have reported zircon ages that fall within this range in igneous suites in Idaho:  $684 \pm 4$  Ma and  $685 \pm 7$  Ma (Lund et al., 2003),  $667 \pm 5$  Ma,  $717 \pm 4$  Ma, and  $709 \pm 5$  Ma (Fanning and Link, 2004), and 680-706 Ma (Durk et al., 2007). The U-Pb ages of the Harmony A are compatible with provenance in the central Laurentian craton. The

grains in the Harmony A arenites could have been derived from the igneous terranes or from recycled sediments originally derived from these terranes. The Hf-isotope ratios of the Harmony A samples are also similar to those in potential source terranes in central Laurentia. The 1.0-1.2 Ga grains have juvenile to intermediate values ( $\epsilon\text{Hf}_{(t)}+10$  to  $-8$ ), similar to those of the Grenville orogen (Bickford et al., 2010; Mueller et al., 2008), and the 1.4-1.5 Ga grains have juvenile to moderately juvenile values ( $\epsilon\text{Hf}_{(t)}+10$  to  $+2.5$ ), compatible with the granitoids of the mid-Laurentian craton (Goodge and Vervoort, 2006) (Figs. 22 and 28). The 1.6-1.8 Ga grains have juvenile to intermediate values ( $\epsilon\text{Hf}_{(t)}+10$  to  $-6$ ), similar to the Yavapai orogenic terrane (Holm et al., 2013). The 2.5-2.8 Ga grains have moderately juvenile to intermediate values ( $\epsilon\text{Hf}_{(t)}+7$  to  $-2$ ), compatible with those in northern Greenland and Arctic Canada of the Northeast Canadian shield (Rohr et al., 2008; Rohr et al., 2010) (Figs. 22 and 28). The 673-716 Ma grains have intermediate to evolved values ( $\epsilon\text{Hf}_{(t)}+1.9$  to  $-5.3$ ), within the range of 630-730 Ma zircons in the Idaho batholith, interpreted as inherited from the Windermere Supergroup volcanics (Gaschnig et al., 2013).

The Harmony A U-Pb ages and Hf-isotope ratios are similar to those of coeval passive margin sedimentary units of western Laurentia (Fig. 28). The upper Neoproterozoic-Lower Cambrian Mutual Formation and Caddy Canyon Quartzite are interpreted as originating in central Laurentia prior to the uplift of the Transcontinental Arch (Gehrels and Pecha, 2014; Linde et al., 2014a; Yonkee et al., 2014) (Fig. 28). The K-S analyses of these units support an interpretation of a common provenance for these units; the Mutual Formation and Caddy Canyon Quartzite have P-values of  $> 0.25$  with one another and Harmony A samples (Table 5). We interpret that the sediments comprising the Harmony

A, the Mutual Formation, and the Caddy Canyon Quartzite were derived from sources in the central Laurentian craton prior to the uplift of the Transcontinental Arch, and were transported to the western Laurentian margin by continent-spanning rivers.

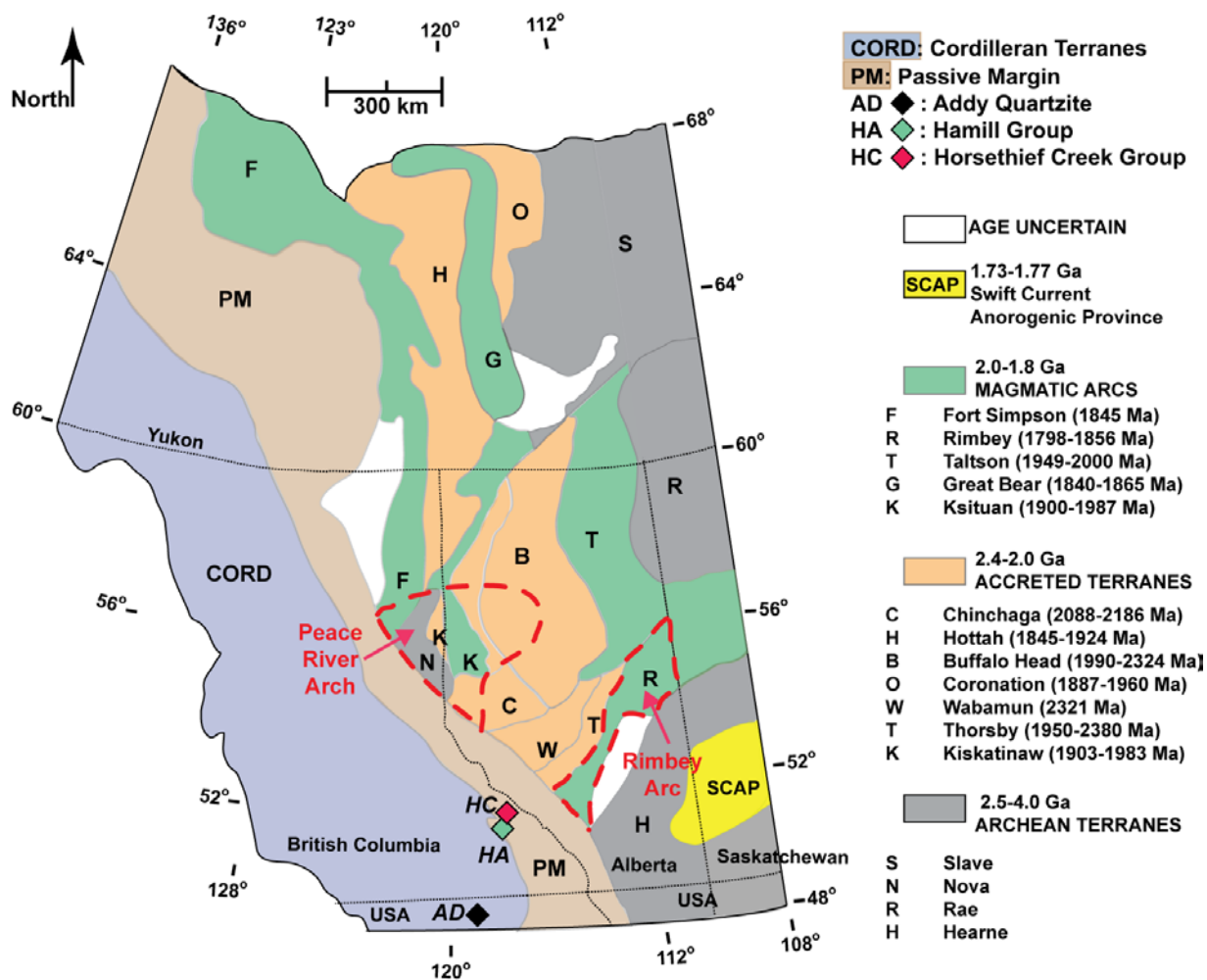


Figure 29. Map of western Canada showing the Cordilleran accreted terranes, the Cordilleran passive margin, and the basement provinces of the Canadian Shield. The Peace River Arch and Rimbey arc regions are outlined by red dashed lines. The Swift Current anorogenic province (SCAP) is highlighted in yellow. The location of the Hamill Group and Horsethief Creek Group samples (Gehrels and Pecha, 2014) are shown. Sample locations are shown by colored diamonds; Horsethief Creek Group (HC) and Hamill Group (HA) are from Gehrels and Pecha (2014); Addy Quartzite (AD) is from Linde et al. (2014b). Map is after Gehrels and Ross (1998); the basement provinces are compiled from Collerson et al. (1988), Hoffman (1989), Ross (1991), Villeneuve et al. (1993), and Peterson et al. (2015).

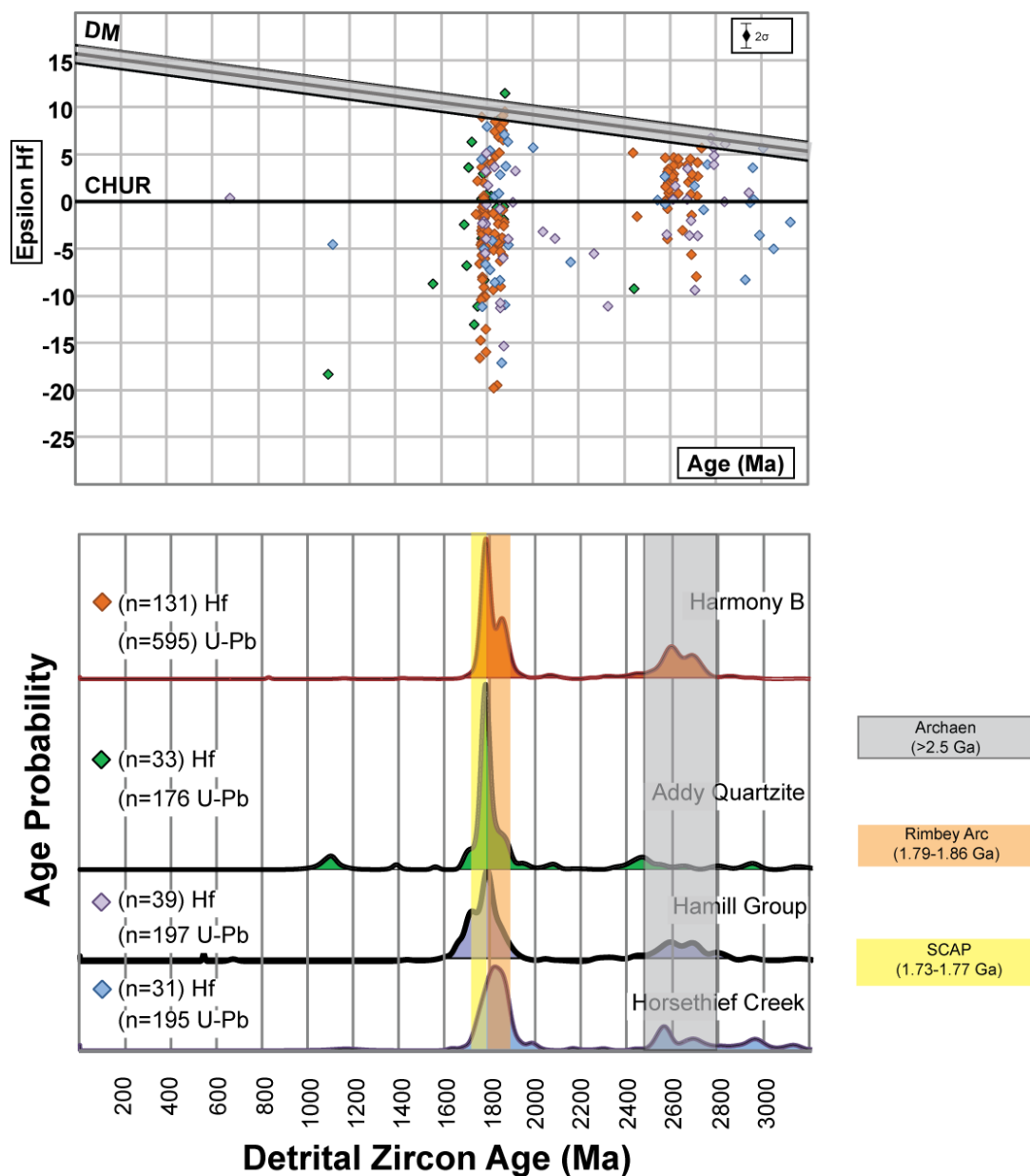


Figure 30. U-Pb ages and Hf isotope data for Harmony B samples and select Laurentian passive margin strata, showing the similarities between the U-Pb ages and Hf isotope analyses of the Harmony B and these passive margin strata. Colored age bars that correspond to Peace River Arch region basement terrane ages are superimposed over the U-Pb ages on the normalized probability plots. Data from the Horsethief Creek Group and the Hamill Group are from Gehrels and Pecha (2014). U-Pb analyses of the Addy Quartzite are from Linde et al. (2014b); Hf analyses of the Addy Quartzite are Linde's unpublished work. Diagrams and symbols are as in Figure 27.



### **b. Provenance of the Harmony B**

The Harmony B arenites are composed of detrital zircons that have potential sources in igneous terranes of eastern Alberta and western Saskatchewan. The Harmony B primary age populations of 1.75-1.9 Ga are within the zircon ages of two igneous provinces: the 1.73-1.77 Ga Swift Current anorogenic province of the Kivalliq Igneous Suite of western Saskatchewan (Collerson et al., 1988; Peterson et al., 2015) and the 1.79-1.86 Ga Rimbey arc of eastern Alberta (Villeneuve et al., 1993; Gehrels and Ross, 1998) (Figs. 29 and 30). The Harmony B age populations of ca. 2.4-2.7 Ga and 2.7-2.8 Ga are similar to those of Archean terranes in the same region (Figs. 29 and 30). Hf-isotope analyses have not been accomplished on the igneous provinces of eastern Alberta and western Saskatchewan; therefore, no direct comparison between the Hf ratios of the Harmony B samples and these potential source terranes is possible. However, the Harmony B grains have a range of U-Pb ages and  $\epsilon\text{Hf}_{(t)}$  values similar to those of the Horsethief Creek and Hamill Groups and the Addy Quartzite, passive margin units interpreted to have originated in eastern Alberta-western Saskatchewan (Linde et al., 2013; Gehrels and Pecha, 2014) (Fig. 30).

We considered other potential source terranes for the Harmony B: the Belt Supergroup, the Trans-Hudson orogen, and the Yavapai-Mazatzal terrane, and considered the possibility that the Harmony Formation shared provenance with the strata of the Roberts Mountains allochthon. The Belt Supergroup and Harmony B sediments share many detrital zircon age populations, but approximately 20% of the Belt Supergroup zircons are ca 1.4-1.7 Ga, younger than the youngest (ca. 1.75 Ga) Harmony B zircons (Link et al., 2013). If the Belt Supergroup was the source of the Harmony B, we would

expect to find some of these 1.4-1.7 Ga zircons in the Harmony B, but do not. The Trans-Hudson orogen and the Harmony B sediments also share many age populations; however, the 1.8-2.0 Ga Trans-Hudson orogen (Whitmeyer and Karlstrom, 2007) could not have supplied the 1.75-1.8 Ga zircons which comprise nearly 35% of the Harmony B. The Yavapai-Mazatzal terrane also shares many age populations with the Harmony B; however, Harmony B samples are missing grains of the age of the other major provinces in central Laurentia, including the Grenville orogenic terrane and the granite-rhyolite magmatic province; the later is located within the Yavapai-Mazatzal terrane (Figs. 22 and 30). It is not likely that river systems would flow across the Yavapai-Mazatzal terrane and transport only grains of that age, without also entraining Grenville or granite-rhyolite province aged grains. We therefore conclude that it is unlikely that the Harmony B originated in central Laurentia. The RMA strata and the Harmony B share some age populations, but dissimilarities preclude a common source. The RMA strata have major age populations from 1.80-1.95 Ga, with no grains < ca. 1.8 Ga (Linde et al., 2016); however, the Harmony B has a significant age population < ca. 1.8 Ga. We interpret that the Harmony B sediments did not originate in the same source as or from the RMA strata.

The detrital zircon U-Pb age spectra and Hf-isotope ratios of the Harmony B are consistent with origin in eastern Alberta and western Saskatchewan (Figs. 29 and 30). This region was exposed throughout Cambrian time and submerged from Ordovician through Mississippian time (Cant, 1988; Cant and O'Connell, 1988; Kent, 1994). We interpret that the provenance of the sediments of the Harmony B, the Horsethief Creek Group, the Hamill Group, and the Addy Quartzite is the Rimbey arc of eastern Alberta and the Swift Current anorogenic province of western Saskatchewan (Figs. 29 and 30).

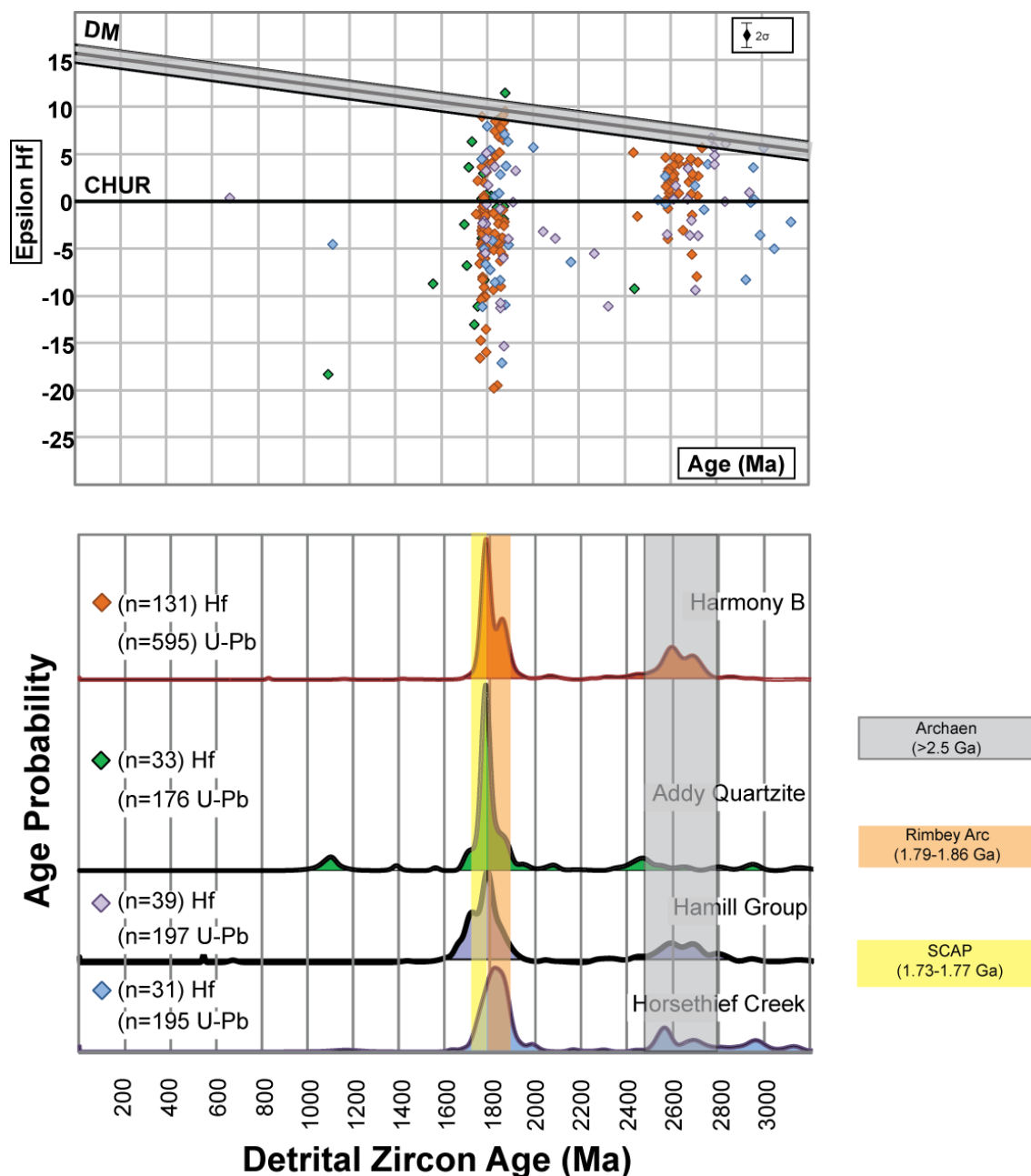


Figure 30. U-Pb ages and Hf isotope data for Harmony B samples and select Laurentian passive margin strata, showing the similarities between the U-Pb ages and Hf isotope analyses of the Harmony B and these passive margin strata. Colored age bars that correspond to Peace River Arch region basement terrane ages are superimposed over the U-Pb ages on the normalized probability plots. Data from the Horsethief Creek Group and the Hamill Group are from Gehrels and Pecha (2014). U-Pb analyses of the Addy Quartzite are from Linde et al. (2014b); Hf analyses of the Addy Quartzite are Linde's unpublished work. Diagrams and symbols are as in Figure 27.

### **c. Age of the Harmony Formation**

The depositional age of the Harmony Formation has been interpreted variously as Cambrian, Devonian, and Devonian-Mississippian. We considered several lines of evidence to interpret the age of the unit. 1) Late Cambrian fossils of North American affinity have been identified in Harmony Formation outcrops in two mountain ranges (Hotz and Willden, 1964). 2) Clasts and olistoliths of the Harmony B have been observed in the Devonian Scott Canyon Formation, near the mouth of Galena Canyon in the Galena Range (Doebrich, 1994; Theodore et al., 1994; Theodore pers. comm. 2015), and confirmed by our field work (Fig. 25). 3) By Mississippian time, the Harmony B strata were subaerally exposed on the craton as a part of the Antler highlands. We deduce this because clasts of the Harmony B were shed from the Antler highlands and transported eastward into the Antler foreland basin where they have been documented in many Mississippian Antler foreland basin units (Harbaugh, 1980; Ketner and Ross, 1990; Ketner, 1998) (Fig. 20). 4) Analysis of deposits interpreted as back- and forebulge deposits overlying the autochthonous passive margin indicates that the RMA was approaching the western Laurentian margin by early Late Devonian time (Goebel, 1991; Dickinson, 2000). 5) The Rimbey arc and Swift Current anorogenic province—the most likely sources of the zircons in the Harmony B—were submerged during the Devonian, not subaerially exposed (Cant, 1988; Cant and O'Connell, 1988; Kent, 1994) (Fig. 31). This region was the site of deposition of shallow shelf carbonates in the Devonian (Kent, 1994). 6) Though there are no fossils in the Harmony A, the similarity of Harmony A U-Pb detrital zircon age spectra and Hf-isotope ratios with those of late Neoproterozoic – Early Cambrian passive margin units, such as the Mutual Formation and the Caddy

Canyon Quartzite, strongly suggest deposition at the same time as these units and before the uplift of the Transcontinental Arch.

We also considered the evidence for a Devonian-Mississippian age of the Harmony Formation. The only datum supporting a Devonian or younger age for the Harmony Formation is a single conodont interpreted as Devonian (Jones, 1997a; 1997b). Ketner et al. (2005) cited this datum in arguing for a Late Devonian-Early Mississippian age for the Harmony Formation. Ketner et al. (2005) also cited additional sampling, but no sample numbers or photos are available or have been published. Weighing the unrepeated nature of the single occurrence against the other evidence presented, we conclude that the Harmony Formation is Cambrian.

## **7. Paleogeographic Implications**

The Harmony A and B are two distinct units with different provenance and depositional history. We prefer the interpretation that the sediments of the Harmony A originated in the central Laurentian craton, before the uplift of the Transcontinental Arch, and were deposited as turbidites in a basin off the western margin in late Neoproterozoic or Early Cambrian time (Fig. 31A). We propose that the sediments of the Harmony B originated in eastern Alberta-western Saskatchewan and were deposited as turbidites in a basin off the northwestern Laurentian margin in Late Cambrian time (Fig. 31B). It is likely that passive-margin arenites such as the Horsethief Creek Group, Hamill Group, and Addy Quartzite also originated in eastern Alberta – western Saskatchewan and were deposited in southern British Columbia and northeastern Washington (Figs. 31A and 31B). These units are the nearshore and shallow shelf equivalents of the deeper marine

Harmony B (Lindsey and Gaylord, 1992; Gehrels and Ross, 1998). From Ordovician through Devonian time, other RMA strata were deposited in the region (Linde et al., 2016).

We propose that some of the models considered above can now be excluded as reasonable explanations for the origin and transport of the Harmony Formation. The first model proposed that the RMA and the Harmony Formation were the deep marine equivalents of coeval allochthonous passive margin strata. The source of the Harmony A sediments is the central Laurentian craton, but the source of the Harmony B sediments is not. Additionally, the RMA strata, exclusive of the lower Vinini Formation, do not share age populations or a source with coeval sediments of the passive margin (Linde et al., 2016). The second model proposed that the RMA and Harmony Formation were extra-Laurentian and were tectonically transported to the western Laurentian margin along with known exotic terranes such as the Alexander terrane. The Harmony Formation has similar age populations and Hf-isotope ratios as western Laurentian terranes. While an exotic source might be possible, a Laurentian source is the simpler solution. The Harmony B also contains trilobite faunas of North American affinity, precluding an exotic origin (Hotz and Willden, 1964). The third model proposed a source in western Laurentia, in the Peace River Arch (PRA) for the RMA and the Salmon River Arch (SRA) for the Harmony Formation. The SRA can be ruled out, as there are no potential source terranes with appropriate age populations. However, the PRA region is a good candidate as a source for RMA sediments, and the Rimbey arc-Swift Current Anorogenic province to the southeast of the PRA is the likely source for Harmony B sediments.

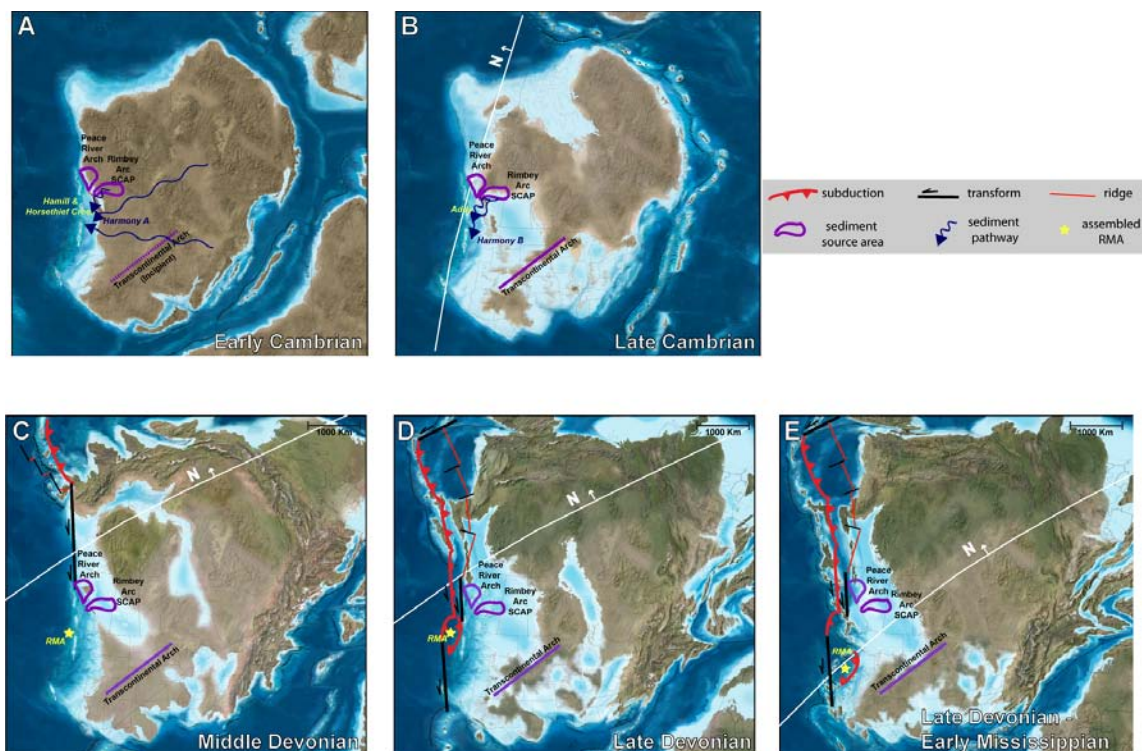


Figure 31: Paleogeographic maps of Laurentia for Early Cambrian through Mississippian time (Blakey, 2013). White lines show the approximate position of the paleoequator. Blue wavy lines show approximate sediment transport pathways of units discussed. The Transcontinental Arch (Sloss, 1988) and Peace River Arch (Ross, 1991) are superimposed. (A) Early Cambrian time. The Harmony A is derived from the central Laurentia, before the uplift of the Transcontinental Arch. The Horsethief Creek Group and Hamill Group are deposited near southeast British Columbia (B) Late Cambrian Time. The Transcontinental Arch has been uplifted. The Harmony B is shed from the Rimbey arc of eastern Alberta and the Swift Current Anorogenic Province of western Saskatchewan (C) Middle Devonian time. An arc has moved westward around the north edge of Laurentia, from northern Baltica to the western margin of northern Laurentia, and a sinistral transpressional fault system has developed along the western margin of Laurentia. The Harmony A and B, and remaining Roberts Mountains allochthon (RMA) strata, are tectonically transported south along the margin by this fault system. (D) Late Devonian time. Shortening and development of an accretionary prism has initiated along much of the western margin of Laurentia, moving the Harmony Formation and RMA strata onto the craton. (E) Early Mississippian time. The Antler orogeny has uplifted the accretionary prism of the Harmony Formation and RMA strata into a highland on the western Laurentian margin.

To reach their current position, we propose that the Harmony A and B were tectonically transported south along the western Laurentian margin in latest Devonian time (Fig. 31D) along with the other RMA strata (Linde et al., 2016). This is consistent with the sinistral transpressional fault system along the western Laurentian margin proposed by Colpron and Nelson (2009) (Figs 31C and 31D). Subsequent shortening emplaced the RMA up onto the western Laurentian craton in the Antler orogeny of latest Devonian-earliest Mississippian time (Fig. 12E). The Harmony Formation strata were imbricated with the rest of the RMA strata during the southward transport, or during emplacement onto the craton, or both.

## **8. Conclusions**

These U-Pb geochronology and Hf-isotope analyses of the Harmony Formation strata give new insight on its provenance, age, and tectonic history. We confirmed previous U-Pb detrital zircon geochronology that had proposed two distinct sub-units of the Harmony Formation, and we interpret different sources for these sub-units (Gehrels et al., 2000a; Gehrels, et al., 2000b). We also confirm that the Harmony Formation is distinct from the strata of the RMA. Our new data indicate that the provenance of the Harmony A was likely in central Laurentia, prior to the uplift of the Transcontinental Arch, and that the provenance of the Harmony B was probably in eastern Alberta-western Saskatchewan. We propose that the Harmony A and B were deposited off the western Laurentian margin of British Columbia in late Neoproterozoic-Cambrian time. The remainder of the RMA was deposited in the same general area, from Ordovician through Devonian time.



Our preferred model is that the Harmony A and B, and the other RMA strata, were tectonically transported in late Devonian time southward along the western Laurentian margin on a sinistral transpressional fault system. The entire RMA package, including the Harmony Formation, was subsequently emplaced eastward onto the craton during the Late Devonian-Early Mississippian Antler orogeny. The Harmony Formation strata were imbricated with the rest of the RMA strata during the tectonic transport to the south or during structural emplacement onto the craton.

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## **Conclusions**

### **1. Summary and Conclusions**

This dissertation investigated Neoproterozoic–Devonian units of the western Laurentian passive margin and Roberts Mountains allochthon (RMA) and determined new U-Pb detrital ages and Hf isotope zircon analyses. From these analyses, this study reports new interpretations of the early Paleozoic tectonics of western Laurentia.

The study reported in Chapter 1 used detrital zircon U-Pb geochronology to confirm that the Upper Neoproterozoic–Lower Cambrian Osgood Mountain Quartzite and the Upper Cambrian–Lower Ordovician Preble Formation in the Osgood Mountains of northern Nevada are units of the western Laurentian passive margin. This also found that U-Pb age populations of the detrital zircons shift with stratal age within the Osgood Mountain Quartzite, indicating a change in provenance within the Osgood Mountain Quartzite. This shift in U-Pb age populations and change in provenance is similar in coeval passive margin strata across the Great Basin. This was a result of the Late Neoproterozoic–Early Cambrian uplift of the Transcontinental Arch, which changed sedimentary transport patterns, and resulted in the observed shift in provenance. This study provided independent corroboration of the existence of the Transcontinental Arch and better precision for the timing at which the Arch uplifted. The study also recorded the impact of the uplifted Arch on continent-wide sediment dispersal patterns—the change in predominant source terranes—and confirmed the Arch as a sediment source for passive margin strata.

The study reported in Chapter 2 used detrital zircon U-Pb geochronology and Hf-isotope analyses to study Roberts Mountains allochthon strata. Zircon grains from six Ordovician to Devonian arenite samples were analyzed for U-Pb ages and some were further analyzed for Hf-isotope ratios. Five of the units have similar U-Pb age populations and Hf-isotope ratios, while those of the Ordovician lower Vinini Formation differ significantly. The lower Vinini Formation originated in the north-central Laurentian craton. The other five units, as well as Ordovician passive-margin sandstones of the western Laurentian margin, had a common source in the Peace River Arch region of western Canada. All of the RMA strata were deposited near the Peace River Arch region and then tectonically transported south along the Laurentian margin, where they were emplaced onto the craton during the Antler orogeny.

Chapter 3 used detrital zircon U-Pb geochronology and Hf-isotope analyses to investigate the Harmony Formation, an enigmatic unit in the RMA. Zircon grains from ten arenite samples were analyzed. Three of the arenite units (Harmony A) have similar U-Pb age peaks and Hf isotope ratios, whereas seven (Harmony B) differ significantly. The data confirmed the subdivision of the Harmony Formation into two petrofacies: quartzose (Harmony A) and feldspathic (Harmony B). Harmony A originated in the central Laurentian craton. Harmony B had a common source in eastern Alberta–western Saskatchewan, north of the source of the Harmony A. All of the Harmony Formation strata were deposited near eastern Alberta in Late Neoproterozoic through Cambrian time and subsequently imbricated tectonically with other Roberts Mountains allochthon strata. The entire package was tectonically transported south along the Laurentian margin and

then emplaced eastward onto the craton during the Late Devonian to Early Mississippian Antler orogeny.

These three studies demonstrated the utility of detrital zircon U-Pb geochronology and Hf isotope analyses to better understand difficult sedimentary and tectonic problems. The studies also provided new insights into the Early Paleozoic tectonic evolution of western Laurentia.

## **2. Recommendations for Future Work**

Future work could use detrital zircons to analyze units that are suspected to be, or are correlated with, Roberts Mountains allochthon strata. Several such units described below are good candidates for future work.

- Intriguing U-Pb analyses of metasediments in Sierran roof pendants are a starting point for further work. The May Lake pendant has some detrital-zircon ages quite similar to the Harmony B (Memeti et al., 2010); limited sampling should be expanded and Hf-isotope analyses added.
- The El Paso terrane of the southern Sierra/northern Mojave Desert region has been correlated with the Roberts Mountains allochthon, and limited detrital zircon geochronology shows that the ages are similar (Chapman et al., 2015). This initial sampling of the El Paso terrane should be expanded and Hf isotope analyses included.
- The Middle to Upper Ordovician Palmetto Formation of the Toiyabe, Toiyabe, and Monitor Ranges, and in Esmeralda county, Nevada, has been interpreted as an



allochthonous sequence of the Roberts Mountains allochthon (Ross, 1967; Kleinhampl and Ziony, 1985; McKee, 1968). The Palmetto Formation has also been interpreted as autochthonous and not part of the RMA (Stewart, 1980).

- Units of the Roberts Mountains allochthon in Northeastern Nevada (e.g., McFarlane, 1997; McFarlane, 2001) are ideal candidates for detrital zircon geochronology and Hf isotope studies, to determine if the same age and Hf isotope patterns found in this study are repeated.
- The Roberts Mountains thrust is interpreted in strata in Idaho (e.g., Coats, 1980; Link and Geslin, 1999), and detrital zircon geochronology and Hf isotope studies would determine if these rocks are Roberts Mountains allochthon strata.

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## **APPENDIX A**

### **Statistical Analysis of Osgood Mountain Quartzite Samples**

A visual scan of the Goughs Canyon and Golconda Mine relative probability graphs (Figs. 6 and 7) reveals similar age peaks, with similar numbers of grains, at ca. 1700 Ma. Both of these graphs also show peaks at ca. 2500 Ma and 2900 Ma, though with different numbers of grains forming the peaks. The K-S statistical test found very low correlation ( $P < 0.05$ ) between these two samples, when we compared the entire data set for each sample. However, the K-S test is very sensitive to proportions of ages, and as with these samples, will indicate no or low correlation, though a visual examination coupled with geologic understanding indicates the contrary—a high likelihood of common sources (Gehrels, 2012). Within the Osgood Mountain Quartzite, we used the K-S statistical test to compare the distinct Neoproterozoic and Archean grain subpopulations. The correlation between the Goughs Canyon and Golconda Mine Neoproterozoic subpopulations was 0.957, while the correlation between these two samples' Archean subpopulations was 0.174. This comparison of subpopulations allows us to account the different proportions of grains of similar ages that would otherwise indicate no correlation.

## APPENDIX B

### U-Pb geochronologic analyses of selected Osgood Mountain Quartzite and Preble Formation strata

Table notes are at the end of the appendix.

Isotope ratios											Apparent ages (Ma)								
Analysis	U	<sup>206</sup> Pb	U/Th	<sup>206</sup> Pb*	±	<sup>207</sup> Pb*	±	<sup>206</sup> Pb*	±	error	<sup>206</sup> Pb*	±	<sup>207</sup> Pb*	±	<sup>206</sup> Pb*	±	Best age	±	Conc
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)
<b>SAMPLE: Osgood Mountain Quartzite GC-03-COM</b>																			
<b>LOCATION: Goughs Canyon, Osgood Mountains 0468951 4554410 (NAD 83 UTM 11N)</b>																			
GC-03-COMB-2	39	5019	1.6	13.3082	7.0	1.9400	7.3	0.1873	2.3	0.31	1106.5	23.2	1095.0	49.1	1072.3	139.8	1072.3	139.8	103.2
GC03OEM-99	227	2831	1.4	12.8138	2.6	1.7529	3.3	0.1629	2.0	0.60	972.9	18.1	1028.2	21.4	1147.9	52.3	1147.9	52.3	84.8
GC03OEM-85	79	12349	1.6	12.7692	1.6	2.0042	3.4	0.1856	3.0	0.88	1097.6	30.1	1116.9	23.1	1154.8	32.8	1154.8	32.8	95.0
GC03OEM-88	116	60376	2.6	12.7043	1.5	2.1192	3.0	0.1953	2.6	0.87	1149.8	27.0	1155.1	20.4	1164.9	29.1	1164.9	29.1	98.7
GC03OEM-34	137	58888	3.0	12.6600	0.9	2.1557	2.2	0.1979	2.0	0.91	1164.2	21.0	1166.9	15.0	1171.9	17.7	1171.9	17.7	99.3
GC03OEM-37	38	6599	2.2	12.6355	9.2	2.1072	9.5	0.1991	2.2	0.23	1138.2	22.9	1151.2	65.5	1175.7	183.3	1175.7	183.3	96.8
GC03OEM-74	87	28506	2.3	12.6005	1.1	2.0615	3.6	0.1884	3.4	0.95	1112.7	34.5	1136.1	24.3	1181.2	21.9	1181.2	21.9	94.2
GC03OEM-14	137	5688	1.1	11.6434	1.2	2.7830	2.9	0.2350	2.6	0.91	1360.7	31.9	1351.0	21.3	1335.6	22.6	1335.6	22.6	101.9
GC03OEM-36	123	15944	1.2	11.5991	1.7	2.2045	3.5	0.1854	3.0	0.87	1096.7	30.6	1182.5	24.3	1343.0	33.0	1343.0	33.0	81.7
GC-03-COMB-15	69	12589	4.2	11.5051	1.4	2.8992	2.1	0.2419	1.6	0.76	1396.6	20.2	1381.7	16.0	1358.7	26.7	1358.7	26.7	102.8
GC-03-COMB-4	85	16572	1.1	11.4792	1.6	2.9036	2.7	0.2417	2.2	0.81	1395.7	27.4	1382.8	20.3	1363.0	30.1	1363.0	30.1	102.4
GC03OEM-40	91	72904	1.4	11.3947	1.5	2.8545	2.5	0.2359	2.0	0.79	1365.4	24.0	1370.0	18.5	1377.2	28.7	1377.2	28.7	99.1
GC03OEM-32	178	90171	2.1	11.3817	1.0	2.7877	1.7	0.2301	1.4	0.81	1335.1	16.4	1352.3	12.5	1379.4	18.8	1379.4	18.8	96.8
GC-03-COMB-17	58	10200	3.8	11.3592	2.0	2.9652	2.9	0.2443	2.0	0.71	1408.9	25.7	1398.7	21.8	1383.2	39.1	1383.2	39.1	101.9
GC03OEM-61	117	46256	2.7	11.3553	0.6	2.9245	1.9	0.2409	1.8	0.94	1391.1	22.8	1388.3	14.6	1383.9	12.4	1383.9	12.4	100.5
GC03OEM-54	123	43469	0.9	11.3012	1.2	2.9874	1.7	0.2449	1.3	0.75	1411.9	16.5	1404.4	13.3	1399.0	22.2	1399.0	22.2	101.4
GC03OEM-50	120	67158	1.8	11.2873	1.5	3.0835	3.4	0.2524	3.0	0.89	1451.0	38.8	1428.6	25.7	1395.4	29.2	1395.4	29.2	104.0
GC03OEM-53	78	23155	2.0	11.1918	1.1	3.0536	2.3	0.2479	2.1	0.88	1427.5	26.3	1421.1	17.8	1411.7	21.2	1411.7	21.2	101.1
GC03OEM-19	194	10573	1.1	11.1409	1.3	2.5264	3.1	0.2041	2.8	0.91	1197.5	30.7	1279.7	22.5	1420.4	24.4	1420.4	24.4	84.3
GC03OEM-13	261	16815	2.1	11.0933	0.4	2.9720	2.1	0.2391	2.0	0.98	1382.1	25.4	1400.5	15.7	1428.6	7.1	1428.6	7.1	96.7
GC03OEM-26	257	58196	1.9	11.0899	0.4	3.0205	2.4	0.2429	2.3	0.98	1402.0	29.5	1412.8	18.2	1429.2	8.5	1429.2	8.5	98.1
GC03OEM-17	328	140166	1.9	11.0829	0.3	3.0727	1.8	0.2470	1.8	0.98	1422.9	23.1	1425.9	14.1	1430.4	6.4	1430.4	6.4	99.5
GC-03-COMB-13	132	22326	5.1	11.0637	0.9	3.1578	2.9	0.2534	2.7	0.95	1455.9	35.7	1446.9	22.2	1433.7	17.0	1433.7	17.0	101.6
GC03OEM-11	133	33704	2.5	11.0538	0.4	3.1250	2.2	0.2505	2.2	0.98	1441.2	27.8	1438.9	16.8	1435.4	7.4	1435.4	7.4	100.4
GC-03-COMB-7	53	9842	2.5	11.0367	3.8	2.9531	4.2	0.2364	1.7	0.41	1367.9	20.9	1395.6	31.5	1438.3	72.4	1438.3	72.4	95.1
GC03OEM-78	258	71339	1.3	11.0227	0.8	3.1213	3.2	0.2495	3.1	0.97	1436.1	39.7	1438.0	24.6	1440.8	15.6	1440.8	15.6	99.7
GC03OEM-79	64	16660	3.2	11.0090	2.2	3.1663	6.2	0.2528	5.8	0.94	1453.0	75.9	1449.0	48.2	1443.1	42.0	1443.1	42.0	100.7
GC03OEM-56	77	25088	1.5	10.9914	1.6	3.1966	2.0	0.2548	1.2	0.61	1463.3	16.2	1456.3	15.8	1446.2	31.0	1446.2	31.0	101.2
GC03OEM-59	211	53488	1.7	10.9415	0.7	3.0655	1.9	0.2433	1.8	0.93	1403.7	22.1	1424.1	14.4	1454.8	13.1	1454.8	13.1	96.5
GC03OEM-60	205	46982	3.2	10.8966	0.7	3.1786	1.9	0.2512	1.7	0.92	1444.7	22.4	1452.0	14.6	1462.6	14.1	1462.6	14.1	98.8
GC03OEM-82	112	85617	2.4	10.8362	1.9	3.0915	3.6	0.2430	3.1	0.86	1402.1	38.9	1430.6	27.6	1473.2	35.2	1473.2	35.2	95.2
GC03OEM-6	34	22175	6.0	10.6891	4.3	3.2104	4.6	0.2489	1.4	0.31	1432.7	18.1	1459.7	35.3	1499.1	81.9	1499.1	81.9	95.6
GC-03-COMB-8	43	5216	0.9	9.7865	2.3	3.2668	4.1	0.2319	3.4	0.83	1344.3	41.1	1473.2	31.9	1664.1	42.9	1664.1	42.9	80.8
GC03OEM-76	305	8984	1.4	9.6557	0.8	3.7015	7.0	0.2592	7.0	0.99	1485.8	92.8	1571.7	56.3	1689.0	14.5	1689.0	14.5	88.0
GC03OEM-7	129	43009	4.0	9.6507	0.6	4.2113	2.4	0.2948	2.3	0.96	1665.3	33.9	1676.2	19.6	1689.9	11.6	1689.9	11.6	98.5
GC-03-COMB-18	49	10308	1.7	9.5933	1.5	4.2623	2.2	0.2966	1.6	0.73	1674.2	24.2	1686.1	18.4	1700.9	27.9	1700.9	27.9	98.4
GC03OEM-63	163	53730	3.6	9.5676	0.6	4.3672	2.3	0.3030	2.3	0.97	1706.4	34.2	1706.2	19.4	1705.9	10.5	1705.9	10.5	100.0
GC03OEM-72	191	108859	1.8	9.5480	0.5	4.4355	2.2	0.3071	2.1	0.97	1726.7	31.7	1719.0	17.9	1709.6	9.6	1709.6	9.6	101.0
GC-03-COMB-12	149	24932	2.4	9.5405	0.6	4.3513	1.0	0.3011	0.8	0.79	1696.7	12.4	1703.1	8.7	1711.1	11.8	1711.1	11.8	99.2
GC-03-COMB-10	34	7849	0.4	9.4460	1.7	4.4494	2.4	0.3048	1.8	0.72	1715.2	26.7	1721.6	20.3	1729.4	31.0	1729.4	31.0	99.2
GC03OEM-98	210	171831	2.6	9.4323	0.5	4.2438	2.1	0.2903	2.0	0.96	1643.1	29.2	1682.5	17.2	1732.0	10.1	1732.0	10.1	94.9
GC03OEM-52	225	47789	2.2	9.4147	0.6	4.6787	2.0	0.3195	2.0	0.96	1787.1	30.5	1763.4	17.1	1735.5	11.1	1735.5	11.1	103.0
GC03OEM-9	79	41900	1.5	9.2316	1.1	4.8444	2.0	0.3243	1.7	0.85	1810.9	26.7	1792.6	16.8	1771.4	19.2	1771.4	19.2	102.2
GC03OEM-24	230	85672	0.7	9.2312	0.4	4.6750	1.4	0.3130	1.4	0.96	1755.4	20.9	1762.8	11.8	1771.5	7.1	1771.5	7.1	99.1
GC03OEM-90	398	29867	2.1	9.2141	0.3	4.4033	2.2	0.2943	2.2	0.99	1662.8	31.9	1713.0	18.2	1774.9	5.3	1774.9	5.3	93.7
GC03OEM-30	94	60580	2.1	9.1982	0.9	4.7791	2.3	0.3188	2.1	0.92	1784.0	32.5	1781.2	19.1	1778.0	16.6	1778.0	16.6	100.3
GC03OEM-8	116	79078	2.0	9.1850	0.9	4.8060	2.1	0.3202	1.9	0.89	1790.5	28.9	1785.9	17.5	1780.6	17.3	1780.6	17.3	100.6
GC03OEM-31	274	226840	2.0	9.1826	0.4	4.8687	4.0	0.3242	3.9	1.00	1810.4	62.2	1796.9	33.4	1781.1	6.9	1781.1	6.9	101.6
GC-03-COMB-16	191	39374	2.0	9.1809	0.5	4.7562	1.9	0.3167	1.9	0.97	1773.6	29.0	1777.2	16.3	1781.5	9.2	1781.5	9.2	99.6
GC03OEM-91	70	31319	1.5	9.1776	1.0	4.8327	2.3	0.3217	2.1	0.91	1797.9	32.8	1790.6	19.4	1782.1	17.9	1782.1	17.9	100.9
GC03OEM-22	294	69402	3.9	9.1728	0.6	4.5721	3.7	0.3042	3.6	0.99	1712.0	54.7	1744.2	30.7	1783.0	10.4	1783.0	10.4	96.0
GC03OEM-46	278	77460	2.6	9.1616	0.8	4.6066	1.3	0.3061	1.0	0.76	1721.4	14.4	1750.5	10.5	1785.3	14.9	1785.3	14.9	96.4
GC03OEM-77	101	31274	1.4	9.1516	1.1	4.7773	3.3	0.3171	3.1	0.94	177								

## U-Pb geochronologic analyses of the Osgood Mountain Quartzite and Preble Formation

Isotope ratios											Apparent ages (Ma)									
Analysis	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc	
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	
<b>SAMPLE: Osgood Mountain Quartzite GC-03-COM LOCATION: Goughs Canyon, Osgood Mountains 0468951 4554410 (NAD 83 UTM 11N)</b>																				
GC03OEM-62	169	84340	2.7	9.0977	0.8	4.8988	2.5	0.3232	2.4	0.94	1805.5	37.4	1802.0	21.3	1798.0	15.4	1798.0	15.4	100.4	
GC03OEM-68	317	85489	3.1	9.0924	0.3	4.9351	3.8	0.3254	3.8	1.00	1816.2	59.4	1808.3	31.8	1799.1	6.1	1799.1	6.1	101.0	
GC03OEM-95	407	112318	3.5	9.0844	0.2	4.8573	1.3	0.3200	1.3	0.99	1789.9	19.8	1794.9	10.8	1800.7	3.7	1800.7	3.7	99.4	
GC03OEM-27	172	61107	2.9	9.0039	0.4	4.8839	1.3	0.3189	1.3	0.96	1784.5	19.6	1799.5	11.0	1816.9	6.5	1816.9	6.5	98.2	
GC03OEM-20	477	290763	2.8	8.9754	0.3	5.0916	2.9	0.3314	2.9	0.99	1845.4	46.9	1834.7	25.0	1822.6	5.7	1822.6	5.7	101.2	
GC03OEM-38	163	90863	6.0	8.8549	0.4	5.1698	1.5	0.3320	1.4	0.97	1848.1	23.2	1847.7	12.7	1847.1	6.8	1847.1	6.8	100.1	
GC03OEM-93	135	78186	2.3	8.8384	1.1	5.2158	1.8	0.3343	1.5	0.80	1859.4	23.7	1855.2	15.6	1850.5	19.7	1850.5	19.7	100.5	
GC03OEM-73	124	17769	2.9	8.7728	1.0	4.9347	3.9	0.3140	3.8	0.96	1760.2	58.0	1808.2	33.0	1863.9	18.9	1863.9	18.9	94.4	
GC03OEM-81	45	11611	6.8	8.7620	1.9	4.9462	3.3	0.3143	2.7	0.82	1761.9	41.8	1810.2	27.8	1866.2	33.7	1866.2	33.7	94.4	
GC03OEM-23	90	54496	7.3	8.6767	0.7	5.4207	2.0	0.3411	1.9	0.93	1892.1	31.1	1888.1	17.4	1883.8	13.3	1883.8	13.3	100.4	
GC03OEM-16	300	224669	2.4	6.3516	1.7	9.1419	2.5	0.4211	1.9	0.76	2265.6	36.8	2352.3	23.3	2428.3	28.2	2428.3	28.2	93.3	
GC03OEM-43	314	20734	1.9	6.3106	0.6	9.8417	2.2	0.4504	2.1	0.96	2397.2	42.4	2420.1	20.3	2439.3	10.4	2439.3	10.4	98.3	
GC-03-COMB-11	32	8231	0.8	6.2967	1.7	10.0435	3.0	0.4587	2.5	0.83	2433.7	51.3	2438.8	28.1	2443.0	28.5	2443.0	28.5	99.6	
GC03OEM-47	128	49917	2.9	6.1750	0.5	10.2161	2.5	0.4575	2.4	0.98	2428.7	48.6	2454.5	22.7	2476.0	8.9	2476.0	8.9	98.1	
GC03OEM-5	289	207482	2.3	6.1693	0.3	10.5546	1.5	0.4723	1.5	0.98	2493.4	30.9	2484.7	14.1	2477.6	4.7	2477.6	4.7	100.6	
GC03OEM-49	159	89144	2.4	6.1616	0.4	10.5095	3.1	0.4696	3.1	0.99	2482.0	62.8	2480.8	28.6	2479.7	7.2	2479.7	7.2	100.1	
GC03OEM-28	409	17300	1.9	5.7087	0.6	10.0832	4.4	0.4175	4.3	0.99	2249.0	81.8	2442.4	40.2	2607.7	10.0	2607.7	10.0	86.2	
GC03OEM-29	202	62420	2.6	5.5469	0.2	11.2570	2.3	0.4529	2.3	1.00	2408.0	45.7	2544.6	21.3	2655.5	3.3	2655.5	3.3	90.7	
GC-03-COMB-19	103	37380	2.2	5.4778	0.5	12.9928	1.7	0.5162	1.6	0.95	2683.0	35.8	2679.1	16.2	2676.2	8.8	2676.2	8.8	100.3	
GC03OEM-39	84	85566	2.0	5.9987	0.3	13.4200	2.2	0.5255	2.2	0.99	2722.3	48.1	2709.7	20.7	2700.2	5.6	2700.2	5.6	100.8	
GC03OEM-84	242	54946	1.5	5.9938	0.6	12.4797	1.8	0.4882	1.7	0.95	2562.9	35.3	2641.2	16.5	2701.7	9.1	2701.7	9.1	94.9	
GC-03-COMB-5	157	53140	1.3	4.7877	0.2	16.1005	1.1	0.5591	1.1	0.98	2862.8	24.8	2882.8	10.5	2896.8	3.5	2896.8	3.5	98.8	
<b>SAMPLE: Osgood Mountain Quartzite GOL-01-COM LOCATION: Golconda Mine, Edna Mountain 0462354 4531791 (NAD 83 UTM 11N)</b>																				
GOL-01COM-51	26	3924	2.4	13.6577	9.4	1.8193	10.0	0.1802	3.4	0.34	1068.1	33.8	1052.4	65.6	1020.0	190.5	1020.0	190.5	104.7	
GOL-01COM-46	45	6990	1.6	13.5514	5.3	1.8252	5.6	0.1794	1.8	0.31	1063.7	17.2	1054.6	36.8	1035.8	107.5	1035.8	107.5	102.7	
GOL-01COM-12	45	8099	1.1	13.3101	3.8	1.8621	4.4	0.1798	2.1	0.48	1065.6	20.5	1067.7	28.8	1072.0	77.0	1072.0	77.0	99.4	
GOL-01COM-72	86	10775	1.5	13.2350	3.3	1.9576	4.1	0.1879	2.3	0.57	1110.0	23.7	1101.1	27.3	1083.4	66.6	1083.4	66.6	102.5	
GOL-01COM-45	318	49259	2.5	13.0614	0.6	1.9557	2.0	0.1853	2.0	0.95	1095.7	19.7	1100.4	13.8	1109.8	12.5	1109.8	12.5	98.7	
GOL-01COM-95	143	33160	4.3	12.8660	1.7	2.0924	3.5	0.1952	3.1	0.88	1149.7	32.8	1146.3	24.4	1139.9	34.0	1139.9	34.0	100.9	
GOL-01COM-86	150	21622	3.4	12.8372	1.3	2.0823	3.5	0.1939	3.2	0.92	1142.3	33.8	1143.0	24.0	1144.3	26.5	1144.3	26.5	99.8	
GOL-01COM-10	77	6732	1.8	12.0043	3.1	2.3690	4.3	0.2063	2.9	0.69	1208.8	32.4	1233.3	30.4	1276.3	60.2	1276.3	60.2	94.7	
GOL-01COM-29	39	10467	1.9	11.6034	2.9	2.7822	3.2	0.2341	1.4	0.44	1356.2	17.3	1350.8	24.2	1342.2	56.2	1342.2	56.2	101.0	
GOL-01COM-16	47	11262	2.0	11.5760	2.7	2.7672	4.1	0.2323	3.0	0.74	1346.7	36.9	1346.7	30.4	1346.8	52.6	1346.8	52.6	100.0	
GOL-01COM-49	99	13817	1.7	11.4196	1.2	2.9033	3.9	0.2405	3.7	0.95	1389.1	45.8	1382.8	29.3	1373.0	24.0	1373.0	24.0	101.2	
GOL-01COM-57	26	5978	1.8	11.3486	3.1	2.9455	4.5	0.2424	3.3	0.73	1399.4	41.1	1399.7	33.9	1385.0	58.8	1385.0	58.8	101.0	
GOL-01COM-38	117	23756	2.0	11.3376	1.1	2.8013	2.9	0.2303	2.7	0.93	1336.3	32.7	1355.9	21.8	1386.9	20.6	1386.9	20.6	96.4	
GOL-01COM-42	103	25635	1.4	11.3318	2.1	2.8660	2.5	0.2355	1.4	0.56	1363.5	17.2	1373.0	18.9	1387.8	39.9	1387.8	39.9	98.2	
GOL-01COM-73	159	27780	2.2	11.2095	1.4	2.8957	7.1	0.2354	7.0	0.98	1362.8	85.6	1380.8	53.6	1408.7	26.5	1408.7	26.5	96.7	
GOL-01COM-47	72	13207	1.3	11.1517	2.2	2.9964	3.3	0.2423	2.5	0.75	1398.9	31.0	1406.7	25.1	1418.5	42.0	1418.5	42.0	98.6	
GOL-01COM-93	607	25552	1.4	11.1105	0.3	2.7594	1.5	0.2224	1.4	0.98	1294.3	16.9	1344.6	11.0	1425.6	6.0	1425.6	6.0	90.8	
GOL-01COM-40	429	78412	4.4	11.1026	0.4	3.0786	3.0	0.2479	3.0	0.99	1427.6	37.9	1427.4	22.9	1427.0	8.2	1427.0	8.2	100.0	
GOL-01COM-79	100	19216	1.7	11.0673	1.5	3.1435	2.4	0.2523	1.9	0.80	1450.5	25.1	1443.4	18.7	1433.0	28.1	1433.0	28.1	101.2	
GOL-01COM-31	94	15195	1.2	11.0484	1.5	3.1237	4.1	0.2503	3.8	0.93	1440.0	49.1	1438.5	31.6	1436.3	29.1	1436.3	29.1	100.3	
GOL-01COM-36	60	13627	3.0	11.0460	2.5	3.1292	3.8	0.2507	2.8	0.74	1442.1	36.1	1439.9	28.9	1436.7	47.9	1436.7	47.9	100.4	
GOL-01COM-53	186	42046	3.4	11.0139	1.1	3.1526	3.9	0.2518	3.7	0.96	1447.9	48.2	1445.6	29.8	1442.3	20.4	1442.3	20.4	100.4	
GOL-01COM-62	228	36186	11.1	10.9886	0.8	3.1181	2.6	0.2485	2.5	0.95	1430.8	31.8	1437.2	20.0	1446.7	15.1	1446.7	15.1	98.9	
GOL-01COM-70	181	40113	1.8	10.9321	1.1	3.1373	2.2	0.2487	1.9	0.86	1432.0	24.8	1441.9	17.3	1456.5	21.8	1456.5	21.8	98.3	
GOL-01COM-71	232	21681	1.6	10.9250	0.7	3.0729	1.3	0.2435	1.1	0.84	1404.8	13.5	1426.0	9.8	1457.7	13.2	1457.7	13.2	96.4	
GOL-01COM-80	84	21050	1.6	10.0944	2.0	3.7886	3.3	0.2774	2.6	0.78	1578.1	35.9	1590.3	26.3	1606.6	38.0	1606.6	38.0	98.2	
GOL-01COM-98	197	19607	2.5	9.7650	0.5	3.5698	0.6	0.2528	0.4	0.58	1453.0	4.7	1542.8	4.9	1668.2	9.4	1668.2	9.4	87.1	
GOL-01COM-64	388	45114	0.9	9.7465	0.4	3.3658	8.4	0.2379	8.3	1.00	1375.9	103.4	1496.5	65.5	1671.7	7.7	1671.7	7.7	82.3	
GOL-01COM-58	40	11025	3.1	9.7348	3.0	4.2986	4.2	0.3035	2.9	0.70	1708.6	44.2	1693.1	34.7	1673.9	55.8	1673.9	55.8	102.1	
GOL-01COM-25	120	14129	2.4	9.6798	1.5	4.0620	3.6	0.2852	3.3	0.91	1617.4	47.5	1646.7	29.7	1684.4	27.6	1684.4	27.6	96.0	
GOL-01COM-89	81	19013	2.9	9.5914	1.3	4.3599	3.6	0.3033	3.4	0.93	1707.6	50.6	1704.8	29.8	1701.3	23.6	1701.3	23.6	100.4	
GOL-01COM-77	106	35024	3.6	9.5661	1.2	4.3757	3.1	0.3036	2.8	0.93	1709.1	42.5	1707.8	25.3	1706.1	21.3	1706.1	21.3	100.2	
GOL-01COM-48	106	24655	2.6	9.5334	0.8	4.3393	2.1	0.3000	1.9	0.92	1691.5	28.6	1700.9	17.3	1712.4	15.3	1712.4	15.3	98.8	
GOL-01COM-55	194	33259	4.2	9.5260	0.9	4.3526	2.3	0.3007	2.2	0.92	1694.9	32.1	1703.4	19.4	1713.9	17.2	1713.9	17.2	98.9	
GOL-01COM-6	200	37784	3.5	9.5108	0.7	4.0840	4.5	0.2817	4.4	0.99	1600.0	63.0	1651.1	36.7	1716.8	13.0	1716.8	13.0	93.2	
GOL-01COM-61	295	71681	3.3	9.4933	0.2	4.2777	5.8	0.2945	5.8	1.00	1664.1	85.3	1689.1	47.9	1720.2	4.1	1720.2	4.1	96.7	
GOL-01COM-50	92	13088	2.5	9.4850	1.8	4.4356	2.8	0.3051	2.2	0.78	1716.7	33.2	1719.0	23.3	1721.8	32				



## U-Pb geochronologic analyses of the Osgood Mountain Quartzite and Preble Formation

Isotope ratios											Apparent ages (Ma)								
Analysis	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)
<b>SAMPLE: Osgood Mountain Quartzite GOL-01-COM LOCATION: Golconda Mine, Edna Mountain 0462354 4531791 (NAD 83 UTM 11N)</b>																			
GOL-01COM-99	1025	121561	3.2	9.1197	0.2	4.8393	0.8	0.3201	0.7	0.95	1790.1	11.5	1791.8	6.5	1793.6	4.5	1793.6	4.5	99.8
GOL-01COM-21	278	71135	2.8	9.1111	0.5	4.7199	1.4	0.3119	1.3	0.94	1750.0	19.5	1770.8	11.4	1795.4	8.7	1795.4	8.7	97.5
GOL-01COM-9	470	67337	2.0	9.0910	0.3	4.7369	1.9	0.3123	1.9	0.99	1752.1	28.6	1773.8	15.9	1799.4	5.9	1799.4	5.9	97.4
GOL-01COM-68	114	33334	0.9	9.0140	0.8	5.1006	2.4	0.3335	2.2	0.93	1855.1	35.4	1836.2	20.0	1814.8	15.4	1814.8	15.4	102.2
GOL-01COM-69	96	33488	3.9	8.9502	0.9	4.9969	2.4	0.3244	2.2	0.92	1811.0	34.2	1818.8	20.0	1827.7	17.0	1827.7	17.0	99.1
GOL-01COM-91	247	24013	7.9	8.9118	0.7	4.8931	9.4	0.3163	9.4	1.00	1771.4	145.2	1801.1	79.4	1835.5	13.6	1835.5	13.6	96.5
GOL-01COM-54	23	3297	207.7	8.9059	3.2	5.1609	4.0	0.3834	2.4	0.59	1854.6	38.0	1846.2	34.1	1836.7	58.7	1836.7	58.7	101.0
GOL-01COM-88	154	44121	2.2	8.7748	0.5	5.1831	2.2	0.3299	2.1	0.98	1837.7	38.6	1849.8	18.3	1863.5	8.3	1863.5	8.3	98.6
GOL-01COM-8	56	27209	10.4	8.6638	1.0	5.4374	2.6	0.3417	2.4	0.92	1894.7	39.8	1890.8	22.7	1886.5	18.9	1886.5	18.9	100.4
GOL-01COM-87	46	16005	0.7	6.3038	1.2	9.8794	2.1	0.4517	1.7	0.81	2402.7	34.6	2423.6	19.6	2441.1	21.1	2441.1	21.1	98.4
GOL-01COM-76	172	48370	2.7	6.2451	0.6	10.0210	3.2	0.4539	3.2	0.98	2412.5	63.8	2436.7	29.8	2457.0	10.3	2457.0	10.3	98.2
GOL-01COM-81	259	100137	1.0	5.8698	0.3	11.4287	2.9	0.4865	2.9	0.99	2555.7	61.7	2558.8	27.5	2561.2	5.7	2561.2	5.7	99.8
GOL-01COM-41	201	65965	1.1	5.6262	0.3	12.3932	2.4	0.5057	2.4	0.99	2638.3	51.5	2634.7	22.5	2631.9	4.7	2631.9	4.7	100.2
GOL-01COM-19	124	21437	1.0	5.6213	0.4	12.4133	1.3	0.5061	1.3	0.96	2639.9	27.5	2636.2	12.4	2633.3	6.4	2633.3	6.4	100.2
GOL-01COM-17	627	44578	6.8	5.5828	0.9	10.5440	2.0	0.4269	1.8	0.90	2291.9	34.0	2483.8	18.3	2644.7	14.5	2644.7	14.5	86.7
GOL-01COM-15	137	61474	1.1	5.5788	0.6	12.7109	2.9	0.5143	2.9	0.98	2674.9	63.1	2658.5	27.6	2645.9	9.4	2645.9	9.4	101.1
GOL-01COM-66	469	97806	2.4	5.4487	0.3	12.5770	3.0	0.4970	3.0	1.00	2601.0	65.0	2648.5	28.7	2685.0	4.2	2685.0	4.2	96.9
GOL-01COM-18	281	78230	2.3	5.4179	0.2	13.3935	1.4	0.5263	1.4	0.99	2725.8	31.6	2707.8	13.6	2694.4	3.6	2694.4	3.6	101.2
GOL-01COM-100	41	16993	2.3	5.3891	1.4	12.7942	2.3	0.5001	1.9	0.81	2614.1	41.0	2664.6	22.0	2703.2	22.4	2703.2	22.4	96.7
GOL-01COM-27	47	10243	2.5	5.2954	0.9	13.1530	2.4	0.5052	2.3	0.94	2635.9	49.1	2690.7	22.9	2732.1	14.1	2732.1	14.1	96.5
GOL-01COM-3	36	16902	3.0	5.2576	1.4	14.5856	2.3	0.5562	1.9	0.80	2850.8	43.3	2788.6	22.3	2743.9	23.1	2743.9	23.1	103.9
GOL-01COM-63	158	47721	0.9	5.1538	0.8	10.0810	3.0	0.3768	2.9	0.96	2061.4	50.7	2442.2	27.6	2776.6	13.5	2776.6	13.5	74.2
GOL-01COM-22	436	15206	0.5	4.9970	1.3	9.9882	7.9	0.3623	7.8	0.99	1993.3	134.4	2434.6	73.4	2827.2	21.0	2827.2	21.0	70.5
GOL-01COM-59	99	53930	5.6	4.8216	0.3	16.0083	2.2	0.5598	2.2	0.99	2865.8	51.5	2877.3	21.5	2885.3	4.8	2885.3	4.8	99.3
GOL-01COM-26	90	42028	1.5	4.8168	0.4	15.9893	1.9	0.5586	1.9	0.98	2860.8	43.7	2876.2	18.4	2886.9	6.1	2886.9	6.1	99.1
GOL-01COM-90	445	190198	1.9	4.8033	0.2	16.2707	2.0	0.5668	1.9	0.99	2894.7	45.2	2892.8	18.7	2891.5	3.6	2891.5	3.6	100.1
GOL-01COM-35	139	61928	1.7	4.8020	0.3	16.7150	1.6	0.5821	1.6	0.98	2957.5	37.6	2918.6	15.5	2891.9	5.2	2891.9	5.2	102.3
GOL-01COM-4	213	87181	1.3	4.7856	0.1	16.4807	1.8	0.5720	1.8	1.00	2916.1	41.8	2905.1	17.1	2897.5	2.4	2897.5	2.4	100.6
GOL-01COM-56	165	24965	1.4	4.7325	0.4	15.0154	0.8	0.5154	0.7	0.89	2679.5	16.0	2816.2	7.8	2915.6	6.0	2915.6	6.0	91.9
GOL-01COM-83	141	54482	2.5	4.5839	0.4	17.5685	2.2	0.5841	2.2	0.99	2965.4	51.3	2966.4	21.0	2967.1	5.9	2967.1	5.9	99.9
GOL-01COM-28	137	58113	2.7	4.5212	0.3	17.5811	1.8	0.5765	1.8	0.98	2934.5	41.3	2967.1	17.1	2989.3	5.1	2989.3	5.1	98.2
GOL-01COM-65	119	86632	1.5	3.6976	0.6	25.2440	2.8	0.6770	2.8	0.98	3332.7	72.3	3317.7	27.7	3308.6	9.2	3308.6	9.2	100.7
<b>SAMPLE: Osgood Mountain Quartzite SP-01-COM LOCATION: Soldier Pass, Osgood Mountains 0463750 4548029 (NAD 83 UTM 11N)</b>																			
SP-01COM-24	41	4488	1.9	13.6778	4.6	1.7684	5.2	0.1754	2.6	0.49	1041.9	24.6	1033.9	34.1	1017.0	92.9	1017.0	92.9	102.4
SP-01COM-82	62	12554	2.6	13.6665	3.1	1.8008	3.6	0.1785	1.7	0.48	1058.7	16.9	1045.7	23.4	1018.7	63.4	1018.7	63.4	103.9
SP-01COM-82	62	12554	2.6	13.6665	3.1	1.8008	3.6	0.1785	1.7	0.48	1058.7	16.9	1045.7	23.4	1018.7	63.4	1018.7	63.4	103.9
SP-01COM-84	92	15305	2.8	13.4667	2.7	1.8274	3.3	0.1785	2.0	0.59	1058.7	19.3	1055.4	21.9	1048.5	54.3	1048.5	54.3	101.0
SP-01COM-54	81	13785	2.4	13.4638	3.9	1.8647	5.5	0.1821	3.9	0.71	1078.4	38.8	1068.6	36.5	1048.9	78.5	1048.9	78.5	102.8
SP-01COM-65	64	6863	1.3	13.3954	2.5	1.7676	2.7	0.1717	1.2	0.45	1021.6	11.7	1033.7	17.8	1059.2	49.3	1059.2	49.3	96.5
SP-01COM-95	87	13398	2.3	13.3840	2.6	1.8458	3.6	0.1792	2.5	0.69	1062.5	24.0	1061.9	23.5	1060.9	52.2	1060.9	52.2	100.1
SP-01COM-70	122	26481	2.0	13.3639	1.8	1.8977	3.1	0.1839	2.4	0.80	1088.4	24.4	1080.3	20.3	1063.9	37.1	1063.9	37.1	102.3
SP-01COM-101	80	17510	2.7	13.3586	3.5	1.9207	4.3	0.1861	2.5	0.58	1100.1	25.6	1088.3	29.0	1064.7	71.1	1064.7	71.1	103.3
SP-01COM-98	88	16706	2.7	13.3409	3.5	1.8659	4.0	0.1805	1.8	0.46	1069.9	18.0	1069.1	26.4	1067.4	71.2	1067.4	71.2	100.2
SP-01COM-81	92	19617	2.7	13.3374	2.1	1.8916	3.2	0.1830	2.3	0.74	1083.2	23.3	1078.1	21.1	1067.9	43.1	1067.9	43.1	101.4
SP-01COM-93	115	14418	2.0	13.3334	3.1	1.9031	3.6	0.1840	1.8	0.50	1089.0	17.7	1082.2	23.7	1068.5	62.3	1068.5	62.3	101.9
SP-01COM-57	195	32750	2.6	13.3310	1.5	1.9747	4.0	0.1909	3.7	0.93	1126.4	38.1	1106.9	26.8	1068.9	29.8	1068.9	29.8	105.4
SP-01COM-72	181	33236	2.3	13.3267	2.1	1.8680	3.0	0.1806	2.2	0.72	1070.0	21.8	1069.8	20.1	1069.5	42.2	1069.5	42.2	100.0
SP-01COM-21	79	10158	2.4	13.3058	2.6	1.9438	4.5	0.1876	3.7	0.83	1108.3	38.1	1096.3	30.4	1072.7	51.4	1072.7	51.4	103.3
SP-01COM-8	135	20259	2.7	13.3052	1.2	1.9122	2.7	0.1845	2.4	0.90	1091.6	24.4	1085.3	18.1	1072.8	24.1	1072.8	24.1	101.8
SP-01COM-50	110	20289	2.0	13.2992	1.6	1.9350	2.4	0.1866	1.8	0.76	1103.1	18.5	1093.3	16.2	1073.7	31.7	1073.7	31.7	102.7
SP-01COM-86	208	33550	3.1	13.2880	1.1	1.8420	2.3	0.1775	2.0	0.88	1053.4	19.7	1060.6	15.2	1075.4	22.3	1075.4	22.3	98.0
SP-01COM-58	105	17841	2.4	13.2828	4.1	1.8981	4.6	0.1829	2.1	0.45	1082.6	20.7	1080.4	30.5	1076.2	82.2	1076.2	82.2	100.6
SP-01COM-39	59	22120	1.2	13.2599	4.0	1.8495	4.2	0.1779	1.5	0.35	1055.3	14.3	1063.3	27.8	1079.6	79.4	1079.6	79.4	97.8
SP-01COM-61	124	24044	2.0	13.2461	2.1	1.9353	3.9	0.1859	3.3	0.84	1099.2	32.9	1093.4	26.0	1081.7	42.3	1081.7	42.3	101.6
SP-01COM-41	76	18468	2.2	13.2360	3.1	1.9667	3.3	0.1907	1.0	0.29	1125.3	9.9	1111.0	22.0	1083.2	62.5	1083.2	62.5	103.9
SP-01COM-35	135	25863	1.2	13.1930	2.0	1.9453	2.8	0.1861	1.9	0.69	1100.4	19.5	1096.8	18.8	1089.7	40.6	1089.7	40.6	101.0
SP-01COM-5	106	19010	3.2	13.1887	1.8	1.9889	2.1	0.1902	0.9	0.46	1122.7	9.8	1111.7	13.9	1090.4	36.5	1090.4	36.5	103.0
SP-01COM-48	144	30952	2.4	13.1677	1.1	1.8919	2.0	0.1807	1.6	0.83	1070.7	16.1	1078.3	13.1	1093.6	22.1	1093.6	22.1	97.9
SP-01COM-76	127	32397	1.9	13.1664	2.9	1.9275	4.1	0.1841	2.9	0.71	1089.1	29.1	1090.7	27.4	1093.8	58.0	1093.8	58.0	99.6
SP-01COM-16	62	15291	2.8	13.1603	4.9	1.8337	5.5	0.1750	2.6	0.47	1039.7	24.6	1057.6	36.1	1094.				

## U-Pb geochronologic analyses of the Osgood Mountain Quartzite and Preble Formation

Isotope ratios													Apparent ages (Ma)						
Analysis	U (ppm)	206Pb/204Pb	U/Th	206Pb*/207Pb*	±	207Pb*/235U*	±	206Pb*/238U*	±	error corr.	206Pb*/238U*	±	207Pb*/235U*	±	206Pb*/207Pb*	±	Best age (Ma)	±	Conc (ppm)
<b>SAMPLE: Osgood Mountain Quartzite SP-01-COM LOCATION: Soldier Pass, Osgood Mountains 0463750 4548029 (NAD 83 UTM 11N)</b>																			
SP-01COM-18	106	23522	1.6	12.9363	2.1	2.0377	3.5	0.1912	2.8	0.81	1127.8	29.3	1128.2	23.9	1129.0	41.5	1129.0	41.5	99.9
SP-01COM-14	82	22652	2.0	12.9306	2.7	1.9805	3.5	0.1857	2.2	0.62	1098.2	21.9	1108.9	23.5	1129.9	54.5	1129.9	54.5	97.2
SP-01COM-62	424	57530	8.7	12.9266	0.6	2.0362	2.3	0.1909	2.2	0.96	1126.2	22.7	1127.7	15.6	1130.5	12.6	1130.5	12.6	99.6
SP-01COM-10	133	32834	3.4	12.9249	1.8	1.9608	2.8	0.1838	2.2	0.78	1087.7	22.4	1102.2	19.2	1130.7	35.2	1130.7	35.2	96.2
SP-01COM-30	150	24825	2.1	12.9204	2.0	1.9777	3.0	0.1853	2.3	0.76	1096.0	23.2	1107.9	20.3	1131.5	39.0	1131.5	39.0	96.9
SP-01COM-45	108	22766	3.2	12.9050	1.5	2.0024	3.8	0.1874	3.5	0.91	1107.4	35.2	1116.3	25.7	1133.8	30.8	1133.8	30.8	97.7
SP-01COM-34	328	41054	2.0	12.8879	0.8	2.1380	2.8	0.1998	2.7	0.96	1174.5	28.7	1161.2	19.4	1136.5	16.2	1136.5	16.2	103.3
SP-01COM-3	123	22900	2.3	12.8770	1.4	1.9750	2.0	0.1844	1.5	0.74	1091.2	15.0	1107.0	13.7	1138.2	27.4	1138.2	27.4	95.9
SP-01COM-19	120	24221	2.5	12.8273	2.0	1.9097	3.8	0.1777	3.2	0.85	1054.2	31.0	1084.5	25.1	1145.8	39.6	1145.8	39.6	92.0
SP-01COM-37	211	42890	5.0	12.7819	0.9	2.1304	2.0	0.1975	1.7	0.88	1161.8	18.3	1158.7	13.6	1152.8	18.7	1152.8	18.7	100.8
SP-01COM-78	199	6061	2.6	12.7549	3.8	2.0154	4.8	0.1864	3.0	0.62	1102.0	30.1	1120.7	32.8	1157.0	75.6	1157.0	75.6	95.2
SP-01COM-90	79	14112	1.5	12.6955	3.1	2.2832	4.3	0.2102	3.0	0.69	1230.0	33.2	1207.1	30.5	1166.3	62.4	1166.3	62.4	105.5
SP-01COM-59	80	32635	2.0	12.6526	2.4	2.0926	3.6	0.1920	2.7	0.75	1132.3	28.1	1146.4	24.9	1173.0	47.7	1173.0	47.7	96.5
SP-01COM-99	376	74969	4.5	12.5662	1.3	2.2287	5.2	0.2031	5.0	0.97	1192.1	54.5	1190.1	36.3	1186.6	25.8	1186.6	25.8	100.5
SP-01COM-60	63	9690	2.2	12.5195	2.2	2.3172	2.7	0.2104	1.5	0.58	1231.0	17.3	1217.6	19.0	1193.9	43.3	1193.9	43.3	103.1
SP-01COM-88	53	3684	1.2	12.4120	5.5	1.8531	9.0	0.1668	7.1	0.79	994.5	65.2	1064.5	59.4	1210.9	109.2	1210.9	109.2	82.1
SP-01COM-74	68	12974	2.0	12.3664	5.5	1.9849	16.1	0.1780	15.1	0.94	1056.2	147.5	1110.4	109.1	1218.1	107.5	1218.1	107.5	86.7
SP-01COM-79	100	13006	2.1	11.8504	1.3	2.6118	1.4	0.2245	0.6	0.42	1305.5	6.9	1304.0	10.3	1301.4	24.7	1301.4	24.7	100.3
SP-01COM-23	118	33956	1.9	11.7955	1.6	2.5846	3.3	0.2211	2.8	0.87	1287.8	33.2	1296.3	23.8	1310.5	30.6	1310.5	30.6	98.3
SP-01COM-47	64	11686	1.2	11.6103	2.1	2.7367	4.1	0.2304	3.6	0.87	1336.8	43.2	1338.5	30.7	1341.1	40.0	1341.1	40.0	99.7
SP-01COM-91	82	14759	2.1	11.5216	2.4	2.7454	3.9	0.2294	3.0	0.78	1331.4	36.5	1340.9	28.9	1355.9	46.6	1355.9	46.6	98.2
SP-01COM-25	70	18994	1.5	11.4371	2.0	2.7028	2.6	0.2242	1.5	0.60	1304.0	18.1	1329.2	18.9	1370.1	39.4	1370.1	39.4	95.2
SP-01COM-92	79	14776	1.3	11.4171	1.6	2.8655	3.3	0.2373	2.9	0.88	1372.5	36.1	1372.9	24.9	1373.4	29.9	1373.4	29.9	99.9
SP-01COM-96	65	11764	2.1	11.4133	3.6	2.7556	5.3	0.2281	3.9	0.73	1324.5	46.9	1343.6	39.8	1374.1	69.7	1374.1	69.7	96.4
SP-01COM-15	101	18535	2.5	11.3596	1.5	2.9901	3.5	0.2463	3.2	0.91	1419.6	40.8	1405.1	26.8	1383.1	28.4	1383.1	28.4	102.6
SP-01COM-52	109	51476	0.9	11.3049	1.2	2.9897	2.4	0.2451	2.1	0.88	1413.3	27.2	1405.0	18.6	1392.4	22.6	1392.4	22.6	101.5
SP-01COM-28	87	15887	0.9	10.8949	1.3	3.2267	2.4	0.2550	2.1	0.85	1464.0	27.1	1463.6	18.8	1462.9	23.9	1462.9	23.9	100.1
SP-01COM-94	195	27235	2.1	10.0082	4.5	3.2864	9.6	0.2386	8.5	0.88	1379.2	105.7	1477.8	75.1	1622.5	83.6	1622.5	83.6	85.0
SP-01COM-1	95	21643	1.4	9.8616	1.4	4.0563	3.0	0.2901	2.6	0.88	1642.1	37.8	1645.6	24.1	1650.0	25.8	1650.0	25.8	99.5
SP-01COM-64	203	38127	3.8	9.6270	0.4	4.2076	1.8	0.2938	1.7	0.97	1660.4	25.4	1675.5	14.6	1694.5	7.8	1694.5	7.8	98.0
SP-01COM-71	239	77679	3.9	9.5299	0.5	4.3175	4.9	0.2984	4.9	0.99	1683.4	72.7	1696.7	40.7	1713.1	9.4	1713.1	9.4	98.3
SP-01COM-77	225	97962	1.7	9.5244	0.6	4.4156	2.1	0.3050	2.0	0.95	1716.1	29.7	1715.3	17.1	1714.2	11.3	1714.2	11.3	100.1
SP-01COM-27	410	27054	4.3	9.4716	0.2	4.0280	2.5	0.2767	2.5	1.00	1574.7	34.7	1639.9	20.3	1724.4	4.4	1724.4	4.4	91.3
SP-01COM-38	268	56645	3.9	9.4428	0.5	4.5997	1.7	0.3150	1.6	0.95	1765.3	25.0	1749.2	14.2	1730.0	9.8	1730.0	9.8	102.0
SP-01COM-36	242	46769	3.2	9.4069	0.3	4.4365	1.2	0.3027	1.2	0.97	1704.6	17.7	1719.2	10.1	1737.0	5.7	1737.0	5.7	98.1
SP-01COM-7	108	22966	3.5	9.2424	0.9	4.8327	2.7	0.3239	2.6	0.95	1809.0	41.0	1790.6	23.1	1769.3	16.1	1769.3	16.1	102.2
SP-01COM-83	237	44995	4.5	9.1897	0.4	4.7346	2.3	0.3156	2.3	0.99	1768.0	35.8	1773.4	19.7	1779.7	6.8	1779.7	6.8	99.3
SP-01COM-4	346	75753	3.5	9.1351	0.4	4.7975	1.9	0.3179	1.8	0.97	1779.2	28.4	1784.5	15.7	1790.6	7.6	1790.6	7.6	99.4
SP-01COM-43	165	89167	4.2	9.0598	0.4	4.8364	1.9	0.3178	1.8	0.97	1778.9	28.5	1791.3	15.9	1805.6	7.9	1805.6	7.9	98.5
SP-01COM-49	574	27141	2.5	9.0038	0.2	4.1659	2.7	0.2720	2.7	1.00	1551.2	37.7	1667.3	22.5	1816.9	3.6	1816.9	3.6	85.4
SP-01COM-89	219	80385	1.7	8.9469	0.4	5.1578	2.6	0.3347	2.6	0.99	1861.1	41.5	1845.7	22.1	1828.4	7.4	1828.4	7.4	101.8
SP-01COM-87	108	17381	3.4	8.9142	0.6	5.1129	1.6	0.3306	1.4	0.92	1841.1	23.1	1838.3	13.4	1835.0	11.4	1835.0	11.4	100.3
SP-01COM-29	48	21377	2.9	8.7328	1.2	5.3921	3.1	0.3415	2.8	0.92	1894.0	46.0	1883.6	26.2	1872.2	21.8	1872.2	21.8	101.2
SP-01COM-11	36	19263	1.6	5.3524	0.9	13.1036	3.0	0.5087	2.9	0.96	2651.0	63.1	2687.1	28.6	2714.5	14.5	2714.5	14.5	97.7
<b>SAMPLE: Preble Formation GS-01-CP LOCATION: Garden Spring, Osgood Mountains 0470302 4548660 (NAD 83 UTM 11N)</b>																			
GS-01CP-25	99	10357	1.1	13.3338	3.1	1.6203	4.4	0.1567	3.2	0.72	938.4	28.0	978.1	27.9	1068.5	62.0	1068.5	62.0	87.8
GS-01CP-50	168	25130	1.7	13.3220	1.7	1.9273	3.4	0.1862	2.9	0.87	1100.9	29.6	1090.6	22.5	1070.2	33.3	1070.2	33.3	102.9
GS-01CP-63	153	37815	1.8	13.2444	1.2	1.9350	3.5	0.1859	3.3	0.94	1098.9	33.5	1093.3	23.6	1081.9	24.3	1081.9	24.3	101.6
GS-01CP-39	136	56330	1.5	13.0935	1.5	1.8890	4.1	0.1794	3.8	0.93	1063.6	37.5	1077.2	27.3	1104.9	30.4	1104.9	30.4	96.3
GS-01CP-86	64	8160	1.9	13.0511	3.3	2.0664	3.8	0.1956	2.0	0.52	1151.6	20.7	1137.7	26.1	1111.3	65.3	1111.3	65.3	103.6
GS-01CP-03	57	7458	2.0	13.0010	4.5	1.9581	5.4	0.1846	3.0	0.55	1092.2	29.7	1101.2	36.2	1119.1	89.6	1119.1	89.6	97.6
GS-01CP-19	119	21877	2.4	12.9944	1.3	1.9596	2.2	0.1847	1.7	0.80	1092.5	17.6	1101.7	14.7	1120.1	26.0	1120.1	26.0	97.5
GS-01CP-72	132	25545	2.0	12.9306	1.9	1.9306	2.6	0.1811	1.8	0.69	1072.7	17.9	1091.8	17.5	1129.9	37.7	1129.9	37.7	94.9
GS-01CP-59	37	4896	2.3	12.9228	3.7	2.0715	4.3	0.1942	2.3	0.52	1143.8	23.6	1139.4	29.7	1131.1	73.6	1131.1	73.6	101.1
GS-01CP-46	73	14997	1.1	11.8136	2.7	2.7489	2.9	0.2355	1.0	0.34	1363.4	12.0	1341.8	21.4	1307.5	52.5	1307.5	52.5	104.3
GS-01CP-13	60	10768	1.2	11.7710	4.7	2.6394	5.1	0.2253	1.9	0.38	1310.0	22.9	1311.7	37.4	1314.5	91.0	1314.5	91.0	99.7
GS-01CP-66	69	6183	1.7	11.6233	3.0	2.7596	3.3	0.2326	1.3	0.40	1348.3	15.7	1344.7	24.4	1338.9	58.1	1338.9	58.1	100.7
GS-01CP-47	114	32535	2.3	11.5984	0.8	2.8339	3.0	0.2384	2.9	0.96	1378.3	36.2	1364.5	22.7	1343.1	15.8	1343.1	15.8	102.6
GS-01CP-91	40	6655	1.1	11.5726	2.7	2.7108	3.4	0.2275	2.2	0.63	1321.5	25.7	1331.4	25.3	1347.4	51.2	1347.4	51.2	98.1
GS-01CP-24	54	18099	0.7	11.4864	1.9	2.8214	3.3	0.2350	2.7	0.82	1360.9	33.7	1361.2	25.1	1361.8	36.8	1		



## U-Pb geochronologic analyses of the Osgood Mountain Quartzite and Preble Formation

		Isotope ratios								Apparent ages (Ma)									
Analysis	U (ppm)	206Pb	U/Th	206Pb* (%)	±	207Pb* (%)	±	206Pb* (%)	±	error corr.	206Pb* 238U* (Ma)	±	207Pb* 235U (Ma)	±	206Pb* 207Pb* (Ma)	±	Best age (Ma)	±	Conc (%)
SAMPLE: Preble Formation GS-01-CP LOCATION: Garden Spring, Osgood Mountains 0470302 4548660 (NAD 83 UTM 11N)																			
GS-01CP-100	119	35401	1.4	11.0702	0.9	3.2529	1.3	0.2612	1.0	0.76	1495.8	13.4	1469.9	10.3	1432.6	16.6	1432.6	16.6	104.4
GS-01CP-36	185	25447	1.0	11.0463	0.7	2.9381	3.0	0.2354	2.9	0.97	1362.7	35.9	1391.8	22.8	1436.7	13.7	1436.7	13.7	94.8
GS-01CP-53	101	18801	2.2	11.0317	1.2	3.2046	3.0	0.2564	2.8	0.92	1471.4	36.6	1458.3	23.4	1439.2	22.3	1439.2	22.3	102.2
GS-01CP-51	85	19617	1.2	10.9874	1.6	3.2288	3.4	0.2573	3.1	0.89	1476.0	40.3	1464.1	26.7	1446.9	30.0	1446.9	30.0	102.0
GS-01CP-38	130	27153	1.6	10.9823	1.1	3.1621	3.7	0.2519	3.6	0.96	1448.1	46.4	1448.0	28.8	1447.7	20.5	1447.7	20.5	100.0
GS-01CP-08	133	31769	2.6	10.9637	1.1	3.2434	2.1	0.2579	1.8	0.86	1479.1	24.2	1467.6	16.4	1451.0	20.4	1451.0	20.4	101.9
GS-01CP-31	235	39440	1.2	10.9311	0.9	3.1570	2.0	0.2503	1.8	0.88	1440.0	22.8	1446.7	15.5	1456.6	17.9	1456.6	17.9	98.9
GS-01CP-94	234	88414	2.0	10.9244	1.2	3.1610	3.2	0.2504	2.9	0.93	1440.8	38.0	1447.7	24.4	1457.8	22.3	1457.8	22.3	98.8
GS-01CP-41	87	9900	2.3	10.9220	1.1	3.1119	3.8	0.2465	3.6	0.96	1420.4	45.8	1435.6	28.9	1458.2	20.8	1458.2	20.8	97.4
GS-01CP-65	183	37569	2.9	10.9151	0.7	3.1724	2.0	0.2511	1.8	0.93	1444.4	23.9	1450.5	15.2	1459.4	13.5	1459.4	13.5	99.0
GS-01CP-85	103	23975	1.6	10.9144	0.8	3.1469	2.2	0.2491	2.1	0.93	1433.9	26.4	1444.2	17.1	1459.5	15.8	1459.5	15.8	98.2
GS-01CP-10	83	34693	1.7	10.8904	2.0	3.1760	3.9	0.2509	3.3	0.85	1442.9	43.1	1451.3	30.1	1463.7	38.5	1463.7	38.5	98.6
GS-01CP-73	166	17043	1.5	10.8498	0.8	3.2494	2.7	0.2557	2.5	0.95	1467.8	33.4	1469.0	20.7	1470.8	15.5	1470.8	15.5	99.8
GS-01CP-67	111	6575	1.8	10.8075	3.2	3.2162	3.5	0.2521	1.5	0.42	1449.3	19.2	1461.1	27.0	1478.2	59.8	1478.2	59.8	98.0
GS-01CP-12	258	30683	4.0	9.9184	0.7	3.7924	2.0	0.2728	1.9	0.94	1555.0	25.9	1591.1	15.9	1639.3	12.1	1639.3	12.1	94.9
GS-01CP-02	336	15353	1.3	9.9010	0.6	2.9804	2.0	0.2140	1.9	0.96	1250.2	21.8	1402.6	15.2	1642.6	10.6	1642.6	10.6	76.1
GS-01CP-52	475	19821	1.7	9.8064	0.4	3.3968	2.2	0.2416	2.2	0.99	1395.0	27.1	1503.7	17.6	1660.4	7.1	1660.4	7.1	84.0
GS-01CP-16	110	17675	2.5	9.7623	0.9	4.0445	2.5	0.2864	2.4	0.93	1623.3	34.1	1643.2	20.7	1668.7	16.7	1668.7	16.7	97.3
GS-01CP-84	190	19260	2.6	9.7005	0.8	3.8262	1.9	0.2692	1.7	0.90	1536.7	23.2	1598.3	15.2	1680.4	15.3	1680.4	15.3	91.4
GS-01CP-89	247	118384	3.8	9.6313	0.5	4.2425	1.3	0.2963	1.2	0.93	1673.2	17.9	1682.3	10.7	1693.6	8.8	1693.6	8.8	98.8
GS-01CP-80	383	18114	3.2	9.6253	0.5	4.1416	3.3	0.2891	3.3	0.99	1637.1	47.1	1662.5	26.9	1694.8	9.1	1694.8	9.1	96.6
GS-01CP-42	203	21933	4.2	9.6161	0.7	4.0759	1.4	0.2843	1.3	0.89	1612.8	18.3	1649.5	11.8	1696.6	12.1	1696.6	12.1	95.1
GS-01CP-70	77	14166	1.7	9.5685	1.1	4.2266	4.5	0.2933	4.4	0.97	1658.1	64.1	1679.2	37.0	1705.7	19.5	1705.7	19.5	97.2
GS-01CP-07	245	88480	3.7	9.5542	0.4	4.2824	4.2	0.2967	4.2	0.99	1675.1	62.1	1690.0	34.9	1708.4	7.8	1708.4	7.8	98.1
GS-01CP-18	133	24880	3.7	9.5322	1.0	4.1195	1.7	0.2848	1.4	0.80	1615.5	19.9	1658.2	14.2	1712.7	19.1	1712.7	19.1	94.3
GS-01CP-96	311	11420	2.8	9.5205	0.5	4.0071	3.4	0.2767	3.3	0.99	1574.7	46.6	1635.6	27.5	1714.9	10.1	1714.9	10.1	91.8
GS-01CP-87	283	107158	2.7	9.5162	0.4	4.5231	1.2	0.3122	1.1	0.94	1751.4	17.6	1735.2	10.2	1715.8	7.9	1715.8	7.9	102.1
GS-01CP-76	116	12014	1.3	9.5149	2.5	4.4488	4.0	0.3070	3.2	0.79	1726.0	48.0	1721.5	33.4	1716.0	45.8	1716.0	45.8	100.6
GS-01CP-45	310	79524	2.1	9.5130	0.4	4.4290	3.3	0.3056	3.2	0.99	1718.9	48.8	1717.8	27.0	1716.4	6.5	1716.4	6.5	100.1
GS-01CP-28	469	59047	1.3	9.4960	0.2	4.2203	1.3	0.2907	1.3	0.98	1644.8	18.2	1678.0	10.5	1719.7	4.2	1719.7	4.2	95.6
GS-01CP-06	87	27765	1.1	9.4826	0.9	4.3558	2.7	0.2996	2.6	0.94	1689.2	38.4	1704.0	22.7	1722.3	17.4	1722.3	17.4	98.1
GS-01CP-04	159	50814	1.7	9.4757	0.5	4.4583	3.4	0.3064	3.3	0.99	1722.9	50.5	1723.2	28.0	1723.6	8.6	1723.6	8.6	100.0
GS-01CP-27	290	103177	2.3	9.4724	0.4	4.4854	2.1	0.3082	2.1	0.98	1731.6	31.4	1728.3	17.5	1724.2	7.4	1724.2	7.4	100.4
GS-01CP-61	430	48448	2.5	9.2261	0.5	4.8871	2.6	0.3270	2.5	0.98	1823.9	40.4	1800.0	21.8	1772.5	8.5	1772.5	8.5	102.9
GS-01CP-57	500	136252	3.3	9.1801	0.2	4.9260	2.0	0.3280	1.9	0.99	1828.5	30.9	1806.7	16.5	1781.6	4.1	1781.6	4.1	102.6
GS-01CP-95	139	33330	2.0	9.1781	0.6	4.7749	2.2	0.3178	2.2	0.96	1779.2	33.5	1780.5	18.9	1782.0	11.6	1782.0	11.6	99.8
GS-01CP-01	333	121268	2.2	9.1763	0.3	4.7919	3.1	0.3189	3.1	1.00	1784.4	47.7	1783.5	25.8	1782.4	5.4	1782.4	5.4	100.1
GS-01CP-78	305	68002	2.3	9.1515	0.5	4.7183	2.7	0.3132	2.7	0.98	1756.3	41.0	1770.5	22.7	1787.3	8.7	1787.3	8.7	98.3
GS-01CP-69	267	64348	2.2	9.1435	0.2	4.7588	2.8	0.3156	2.8	1.00	1768.1	43.9	1777.7	23.9	1788.9	4.1	1788.9	4.1	98.8
GS-01CP-88	555	120816	3.5	9.1154	0.5	4.9925	4.1	0.3301	4.1	0.99	1838.7	65.5	1818.1	34.9	1794.5	9.0	1794.5	9.0	102.5
GS-01CP-20	199	62338	3.6	9.0830	0.5	4.7635	3.3	0.3138	3.3	0.99	1759.4	50.1	1778.5	27.6	1801.0	8.7	1801.0	8.7	97.7
GS-01CP-55	380	46290	1.9	9.8906	0.5	4.6418	1.5	0.3023	1.4	0.94	1702.9	21.6	1756.8	12.9	1821.6	9.7	1821.6	9.7	93.5
GS-01CP-81	208	85830	0.8	8.9501	0.5	5.0251	2.1	0.3262	2.1	0.97	1819.9	32.5	1823.6	17.8	1827.7	8.6	1827.7	8.6	99.6
GS-01CP-83	535	99813	4.1	8.9097	0.3	5.0028	1.1	0.3233	1.1	0.96	1805.7	17.1	1819.8	9.6	1835.9	5.9	1835.9	5.9	98.4
GS-01CP-75	142	18385	2.1	8.8476	0.8	4.9499	3.4	0.3176	3.3	0.97	1778.1	51.7	1810.8	28.9	1848.6	14.8	1848.6	14.8	96.2
GS-01CP-17	120	29448	3.4	8.8410	0.5	5.0998	2.6	0.3270	2.6	0.98	1823.8	40.6	1836.1	22.0	1849.9	8.2	1849.9	8.2	98.6
GS-01CP-99	285	24331	1.7	8.7883	0.3	4.4771	1.1	0.2954	1.0	0.95	1618.3	14.3	1726.7	8.8	1860.8	6.2	1860.8	6.2	87.0
GS-01CP-92	269	68081	4.0	8.6789	0.5	5.2985	1.6	0.3335	1.5	0.94	1855.4	23.4	1868.6	13.3	1883.3	9.9	1883.3	9.9	98.5
GS-01CP-23	478	23962	2.5	8.6886	2.1	7.3711	4.0	0.3683	3.4	0.85	2021.2	58.4	2157.5	35.4	2289.8	35.8	2289.8	35.8	88.3
GS-01CP-32	617	49511	3.0	6.1105	3.0	9.6403	5.7	0.4272	4.8	0.85	2293.2	92.5	2401.0	52.2	2493.7	51.1	2493.7	51.1	92.0
GS-01CP-56	409	113804	1.5	5.7766	0.1	11.8377	2.1	0.4960	2.1	1.00	2596.4	45.0	2591.7	19.8	2587.9	2.5	2587.9	2.5	100.3
GS-01CP-62	445	24465	1.2	5.7061	0.3	9.8035	2.5	0.4057	2.4	0.99	2195.3	45.3	2416.5	22.6	2608.4	5.8	2608.4	5.8	84.2
GS-01CP-26	447	21996	1.9	5.6219	0.7	9.6921	4.8	0.3952	4.8	0.99	2146.8	86.8	2406.0	44.3	2633.1	12.2	2633.1	12.2	81.5
GS-01CP-34	237	65637	1.0	5.5164	0.2	11.8885	3.5	0.4476	3.5	1.00	2384.7	68.8	2539.0	32.2	2664.6	3.0	2664.6	3.0	89.5
GS-01CP-44	65	36641	1.4	5.4257	0.5	12.9959	2.2	0.5114	2.1	0.97	2662.6	45.6	2679.4	20.3	2692.0	8.6	2692.0	8.6	98.9
GS-01CP-14	200	38071	1.2	4.7923	0.4	14.7942	3.8	0.5142	3.8	0.99	2674.6	83.7	2802.1	36.6	2895.2	7.1	2895.2	7.1	92.4
GS-01CP-09	161	42489	1.6	4.5683	0.3	17.4065	2.9	0.5767	2.9	1.00	2935.3	68.9	2957.5	28.2	2972.6	4.4	2972.6	4.4	98.7
GS-01CP-49	104	54313	1.6	3.7424	0.4	24.0357	2.2	0.6524	2.1	0.98	3237.5	54.0	3269.8	21.1	3289.7	6.4	3289.7	6.4	98.4
SAMPLE: Preble Formation GOL-02-CP LOCATION: Golconda Mine, Edna Mountain 0465271 4533874 (NAD 83 UTM 11N)																			
GOL-02CP-54	128	8820	3.3	13.6763	1.8	1.5939	2.6	0.1581	1.9	0.72	946.2	16.3	967.8	16.1	1017.2	36.2	1017.2	36.2	



## U-Pb geochronologic analyses of the Osgood Mountain Quartzite and Preble Formation

Analysis	Isotope ratios										Apparent ages (Ma)								
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	207Pb*	±	206Pb*	±	Best age	±	Conc	
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(%)	
SAMPLE: Preble Formation GOL-02-CP LOCATION: Golconda Mine, Edna Mountain 0465271 4533874 (NAD 83 UTM 11N)																			
GOL-02CP-98	60	13554	1.9	11.5077	3.3	2.7890	4.3	0.2328	2.7	0.63	1349.0	33.3	1352.6	32.2	1358.2	64.2	1358.2	64.2	99.3
GOL-02CP-71	94	13760	2.9	11.4941	2.2	2.8500	3.0	0.2376	2.0	0.68	1374.1	25.3	1368.8	22.5	1360.5	42.0	1360.5	42.0	101.0
GOL-02CP-70	82	13267	2.1	11.4669	2.8	2.8581	3.8	0.2377	2.5	0.66	1374.7	31.0	1370.9	28.4	1365.1	54.5	1365.1	54.5	100.7
GOL-02CP-19	87	13061	0.8	11.4572	1.5	2.7783	2.6	0.2309	2.2	0.83	1339.0	26.7	1349.7	19.8	1366.7	28.1	1366.7	28.1	98.0
GOL-02CP-6	96	14881	2.3	11.4096	1.6	2.8485	2.3	0.2357	1.6	0.72	1364.4	19.9	1368.4	17.0	1374.7	30.3	1374.7	30.3	99.2
GOL-02CP-13	95	14093	2.0	11.4007	1.8	2.7702	5.6	0.2291	5.3	0.95	1329.6	63.8	1347.5	41.9	1376.2	34.8	1376.2	34.8	96.6
GOL-02CP-16	184	39901	2.8	11.3737	1.4	2.9025	3.1	0.2394	2.7	0.90	1383.7	34.1	1382.6	23.1	1380.8	26.1	1380.8	26.1	100.2
GOL-02CP-66	81	13331	1.3	11.3435	1.7	2.8970	3.1	0.2383	2.6	0.84	1378.0	32.2	1381.1	23.2	1385.9	31.8	1385.9	31.8	99.4
GOL-02CP-63	66	10675	2.4	11.3133	4.2	2.7702	4.6	0.2273	1.8	0.39	1320.3	21.2	1347.5	34.3	1391.0	81.5	1391.0	81.5	94.9
GOL-02CP-30	13	3078	2.4	11.3112	11.7	3.0118	12.0	0.2471	2.7	0.23	1423.4	34.8	1410.6	91.9	1391.3	225.4	1391.3	225.4	102.3
GOL-02CP-88	106	18521	2.3	11.2712	1.7	2.8732	2.5	0.2349	1.8	0.73	1360.0	22.6	1374.9	19.1	1398.1	33.2	1398.1	33.2	97.3
GOL-02CP-17	84	4318	1.3	11.2024	4.4	2.8828	5.6	0.2342	3.5	0.62	1356.6	42.5	1377.4	42.2	1409.9	83.9	1409.9	83.9	96.2
GOL-02CP-75	187	24157	1.1	11.1826	0.9	2.9993	2.2	0.2433	2.0	0.91	1403.6	25.6	1407.4	16.9	1413.3	17.6	1413.3	17.6	99.3
GOL-02CP-48	77	9608	0.9	11.1472	1.6	3.0849	4.2	0.2494	3.8	0.92	1435.4	49.5	1428.9	32.0	1419.3	31.0	1419.3	31.0	101.1
GOL-02CP-56	98	15044	0.9	11.1324	1.4	3.0539	2.4	0.2466	2.0	0.83	1420.8	25.5	1421.2	18.5	1421.9	26.1	1421.9	26.1	99.9
GOL-02CP-58	133	19875	2.9	11.1057	1.1	3.1100	2.6	0.2505	2.4	0.91	1441.0	30.3	1435.2	19.9	1426.4	20.6	1426.4	20.6	101.0
GOL-02CP-11	131	22755	1.8	11.1034	0.7	3.1260	1.4	0.2517	1.3	0.88	1447.4	16.4	1439.1	11.1	1426.8	13.1	1426.8	13.1	101.4
GOL-02CP-67	340	7541	1.4	11.0749	1.1	2.7976	3.5	0.2247	3.3	0.95	1306.7	39.4	1354.9	26.3	1431.7	21.3	1431.7	21.3	91.3
GOL-02CP-40	98	20386	1.2	11.0659	1.5	3.1236	2.3	0.2507	1.7	0.76	1442.1	22.3	1438.5	17.5	1433.3	28.4	1433.3	28.4	100.6
GOL-02CP-49	121	16321	1.5	11.0557	1.1	3.1738	4.0	0.2545	3.9	0.96	1461.6	51.0	1450.8	31.2	1435.1	20.5	1435.1	20.5	101.8
GOL-02CP-8	62	13504	2.3	11.0449	1.6	3.2916	2.3	0.2637	1.7	0.73	1508.6	22.4	1479.1	17.8	1436.9	29.8	1436.9	29.8	105.0
GOL-02CP-45	96	14724	2.8	11.0413	1.9	3.1164	2.9	0.2496	2.2	0.76	1436.2	28.6	1436.8	22.4	1437.5	35.9	1437.5	35.9	99.9
GOL-02CP-22	156	24532	2.3	11.0354	1.3	3.0671	2.1	0.2455	1.7	0.81	1415.1	22.0	1424.5	16.5	1438.6	24.3	1438.6	24.3	98.4
GOL-02CP-91	102	16431	1.7	11.0284	1.8	3.1329	2.1	0.2506	1.1	0.52	1441.5	14.3	1440.8	16.4	1439.8	34.5	1439.8	34.5	100.1
GOL-02CP-9	288	48627	1.5	11.0271	0.6	3.1570	3.3	0.2525	3.3	0.99	1451.3	42.4	1446.7	25.5	1440.0	10.7	1440.0	10.7	100.8
GOL-02CP-39	114	15557	1.7	11.0203	2.2	3.2222	2.7	0.2575	1.6	0.58	1477.2	20.9	1462.5	21.1	1441.2	42.1	1441.2	42.1	102.5
GOL-02CP-28	402	75182	1.3	10.9893	0.5	3.2052	3.5	0.2555	3.5	0.99	1466.6	45.3	1458.4	27.0	1446.5	9.7	1446.5	9.7	101.4
GOL-02CP-64	121	19495	1.6	10.9667	1.3	3.1578	3.7	0.2512	3.4	0.93	1444.5	44.5	1446.9	28.4	1450.4	25.0	1450.4	25.0	99.6
GOL-02CP-18	230	42276	1.8	10.9364	0.5	3.1297	2.4	0.2482	2.3	0.97	1429.4	29.6	1440.0	18.3	1455.7	10.4	1455.7	10.4	98.2
GOL-02CP-44	215	33775	2.5	10.9144	1.1	3.2384	3.8	0.2563	3.7	0.95	1471.1	48.3	1466.4	29.9	1459.5	21.8	1459.5	21.8	100.8
GOL-02CP-27	90	17093	0.9	10.9079	2.1	3.2022	3.6	0.2533	2.9	0.81	1455.6	37.6	1457.7	27.6	1460.7	40.1	1460.7	40.1	99.7
GOL-02CP-65	125	15211	1.6	10.8753	1.3	3.2250	3.3	0.2544	3.0	0.91	1461.0	39.2	1463.2	25.4	1466.4	25.3	1466.4	25.3	99.6
GOL-02CP-89	131	16339	1.3	10.8738	1.2	3.2286	2.1	0.2546	1.7	0.82	1462.3	22.7	1464.0	16.4	1466.6	23.1	1466.6	23.1	99.7
GOL-02CP-94	114	15076	1.7	10.8728	2.3	3.2507	3.1	0.2563	2.0	0.65	1471.1	26.6	1469.4	24.0	1466.8	44.5	1466.8	44.5	100.3
GOL-02CP-25	199	29410	2.8	10.8526	1.1	3.2369	2.1	0.2548	1.8	0.86	1463.1	23.7	1466.0	16.3	1470.3	20.1	1470.3	20.1	99.5
GOL-02CP-83	118	25794	2.4	10.8389	1.0	3.8563	3.0	0.2808	2.8	0.94	1595.3	39.8	1604.6	24.2	1616.8	19.1	1616.8	19.1	98.7
GOL-02CP-43	45	9704	1.2	9.7928	2.0	4.2478	2.5	0.3017	1.5	0.59	1699.7	22.0	1683.3	20.4	1662.9	37.0	1662.9	37.0	102.2
GOL-02CP-87	103	18643	2.3	9.6623	0.8	4.3084	1.2	0.3019	1.0	0.77	1700.8	14.3	1695.0	10.3	1687.7	14.7	1687.7	14.7	100.8
GOL-02CP-5	170	34880	1.8	9.6617	0.8	4.3796	1.4	0.3069	1.1	0.81	1725.4	17.4	1708.5	11.7	1687.8	15.3	1687.8	15.3	102.2
GOL-02CP-36	83	15655	3.1	9.6526	1.1	4.4956	3.2	0.3147	3.0	0.94	1763.9	46.6	1730.2	26.8	1689.6	20.9	1689.6	20.9	104.4
GOL-02CP-81	68	15520	3.4	9.6019	1.7	4.3154	2.6	0.3005	2.0	0.77	1693.9	29.5	1696.3	21.3	1699.3	30.6	1699.3	30.6	99.7
GOL-02CP-1	232	21177	2.3	9.5847	0.4	4.2648	1.8	0.2965	1.7	0.97	1673.8	25.1	1686.6	14.4	1702.6	8.1	1702.6	8.1	98.3
GOL-02CP-100	322	15116	2.5	9.5826	1.3	3.4675	4.9	0.2410	4.7	0.97	1391.9	59.3	1519.9	38.7	1703.0	23.2	1703.0	23.2	81.7
GOL-02CP-12	282	29515	3.2	9.5706	0.5	4.4041	2.9	0.3057	2.8	0.99	1719.5	42.7	1713.1	23.7	1705.3	8.9	1705.3	8.9	100.8
GOL-02CP-10	290	69506	1.9	9.5656	0.7	4.5368	2.6	0.3147	2.5	0.97	1764.0	39.1	1737.7	21.8	1706.2	12.1	1706.2	12.1	103.4
GOL-02CP-72	469	49364	4.3	9.5655	0.5	4.1149	8.4	0.2855	8.4	1.00	1618.8	119.8	1657.3	68.6	1706.3	9.6	1706.3	9.6	94.9
GOL-02CP-65	137	17659	1.0	9.5603	1.4	4.3318	2.7	0.3004	2.3	0.86	1693.1	34.9	1699.4	22.5	1707.3	25.8	1707.3	25.8	99.2
GOL-02CP-53	113	13025	3.5	9.5462	1.1	3.9456	2.1	0.2732	1.8	0.86	1556.9	64.7	1623.1	16.9	1710.0	19.9	1710.0	19.9	101.0
GOL-02CP-55	162	32982	3.2	9.5421	0.9	4.4099	4.4	0.3052	4.3	0.98	1717.0	24.9	1714.2	36.4	1710.8	16.5	1710.8	16.5	100.4
GOL-02CP-69	118	21234	1.7	9.5339	1.0	4.4538	2.7	0.3080	2.5	0.92	1730.7	38.2	1722.4	22.6	1712.4	19.2	1712.4	19.2	101.1
GOL-02CP-90	224	52317	1.5	9.5246	1.0	4.4689	1.9	0.3087	1.6	0.84	1734.4	24.0	1725.2	15.5	1714.1	18.5	1714.1	18.5	101.2
GOL-02CP-26	216	46815	4.7	9.5222	0.5	4.4045	3.1	0.3042	3.1	0.99	1712.0	46.8	1713.2	26.1	1714.6	9.0	1714.6	9.0	99.8
GOL-02CP-84	152	34159	1.8	9.5138	0.6	4.4714	2.1	0.3085	2.0	0.95	1733.5	30.2	1725.7	17.3	1716.2	11.8	1716.2	11.8	101.0
GOL-02CP-4	153	27283	2.1	9.4951	0.9	4.5074	1.7	0.3104	1.4	0.86	1742.7	21.8	1732.3	13.8	1719.9	15.8	1719.9	15.8	101.3
GOL-02CP-61	132	33675	1.7	9.4436	1.0	4.3538	2.1	0.2982	1.9	0.88	1682.4	27.7	1703.6	17.5	1729.8	18.2	1729.8	18.2	97.3
GOL-02CP-95	96	17639	2.7	9.4422	1.0	4.5463	2.4	0.3113	2.2	0.90	1747.3	33.7	1739.5	20.2	1730.1	19.0	1730.1	19.0	101.0
GOL-02CP-79	120	25839	2.9	9.4175	0.8	4.5515	3.4	0.3109	3.3	0.97	1745.0	51.1	1740.4	28.7	1734.9	15.3	1734.9	15.3	100.6
GOL-02CP-24	169	38576	2.3	9.4134	0.6	4.4807	1.9	0.3059	1.8	0.95	1720.5	27.5	1727.4	15.9	1735.7	11.0	1735.7	11.0	99.1
GOL-02CP-62	108	20485	1.4	9.3606	1.7	4.3343	3.5	0.2943	3.0	0.86	1662.8	44.1	1699.9	28.7	1746.0	32.0	1746.0	32.0	95.2
GOL-02CP-51	105	23012																	

## U-Pb geochronologic analyses of the Osgood Mountain Quartzite and Preble Formation

Analysis	Isotope ratios								Apparent ages (Ma)										
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc
	(ppm)	204Pb		(%)	(%)	(%)	(%)	(%)	(%)	corr.	(Ma)	(Ma)	(Ma)	(Ma)	(Ma)	(Ma)	(Ma)	(%)	
<b>SAMPLE: Preble Formation GOL-02-CP LOCATION: Golconda Mine, Edna Mountain 0465271 4533874 (NAD 83 UTM 11N)</b>																			
GOL-02CP-31	174	46215	4.4	5.3994	0.4	13.0579	4.1	5.1113	4.0	0.99	2662.4	88.0	2683.8	36.3	2700.0	7.4	2700.0	7.4	98.6
GOL-02CP-47	49	12873	0.5	5.3403	0.6	12.9992	2.4	5.0353	2.4	0.97	2628.7	51.2	2679.6	23.0	2718.2	9.2	2718.2	9.2	96.7
GOL-02CP-34	134	30130	2.1	5.1674	0.4	14.2616	0.9	5.5345	0.8	0.90	2760.3	17.4	2767.3	8.2	2772.3	6.3	2772.3	6.3	99.6
<b>SAMPLE: Preble Formation GC-01-CP LOCATION: Goughs Canyon, Osgood Mountains 0468513 4554509 (NAD 83 UTM 11N)</b>																			
GC-01CP-99	67	7864	0.9	13.3853	4.2	1.9362	4.5	0.1880	1.7	0.39	1110.4	17.8	1093.7	30.2	1060.7	83.8	1060.7	83.8	104.7
GC-01CP-42	374	7922	2.8	13.3562	1.0	1.6731	2.2	0.1621	1.9	0.88	968.3	17.1	998.4	13.8	1065.1	21.1	1065.1	21.1	90.9
GC-01CP-35	126	12881	3.2	13.3228	1.3	1.8839	2.8	0.1820	2.4	0.88	1078.1	24.2	1075.4	18.5	1070.1	26.9	1070.1	26.9	100.7
GC-01CP-1	82	6225	1.1	13.3174	1.8	1.8699	2.0	0.1806	0.8	0.43	1070.3	8.3	1070.5	13.0	1070.9	35.7	1070.9	35.7	99.9
GC-01CP-54	92	9626	1.9	13.2099	2.0	1.9456	2.3	0.1864	1.1	0.50	1101.9	11.5	1096.9	15.3	1087.2	39.5	1087.2	39.5	101.4
GC-01CP-39	254	33601	2.8	13.1546	1.7	1.9956	2.8	0.1904	2.3	0.81	1123.5	23.5	1114.0	19.1	1095.6	33.4	1095.6	33.4	102.5
GC-01CP-70	241	26616	2.3	13.1206	1.3	1.9557	3.3	0.1861	3.0	0.92	1100.2	30.8	1100.4	22.2	1100.7	25.6	1100.7	25.6	100.0
GC-01CP-41	66	7593	1.5	12.8284	6.1	2.1988	6.5	0.2046	2.3	0.35	1199.9	24.8	1180.7	45.3	1145.7	120.9	1145.7	120.9	104.7
GC-01CP-33	86	7658	1.5	11.5874	2.1	2.7403	4.4	0.2303	3.9	0.88	1336.0	46.8	1339.5	32.8	1344.9	40.7	1344.9	40.7	99.3
GC-01CP-14	94	10628	3.6	11.5536	1.4	2.7866	2.4	0.2335	1.9	0.81	1352.8	23.1	1352.0	17.6	1350.6	26.8	1350.6	26.8	100.2
GC-01CP-36	100	12913	1.4	11.5343	2.3	2.8680	4.3	0.2399	3.6	0.84	1386.3	45.1	1373.5	32.2	1353.8	44.2	1353.8	44.2	102.4
GC-01CP-10	106	9003	1.8	11.4787	1.0	2.8828	3.1	0.2400	2.9	0.94	1386.7	36.6	1377.4	23.5	1363.1	20.0	1363.1	20.0	101.7
GC-01CP-76	96	12962	2.7	11.4763	2.4	2.9075	3.3	0.2420	2.2	0.68	1397.1	27.8	1383.9	24.6	1363.5	46.2	1363.5	46.2	102.5
GC-01CP-89	153	15603	1.7	11.4563	1.0	2.7335	5.1	0.2271	5.0	0.98	1319.4	60.0	1337.6	38.2	1366.8	19.8	1366.8	19.8	96.5
GC-01CP-62	165	28978	2.9	11.4391	1.3	2.7267	1.8	0.2262	1.3	0.71	1314.6	15.3	1335.7	13.4	1369.7	24.5	1369.7	24.5	96.0
GC-01CP-31	84	4346	1.9	11.4310	3.6	2.6653	3.7	0.2210	0.8	0.21	1287.0	9.0	1318.9	27.2	1371.1	69.4	1371.1	69.4	93.9
GC-01CP-59	109	12825	2.6	11.4108	1.0	2.9356	1.4	0.2429	1.0	0.72	1402.0	13.1	1391.1	10.9	1374.5	19.3	1374.5	19.3	102.0
GC-01CP-50	115	8845	1.2	11.3600	1.9	2.7880	2.9	0.2297	2.2	0.76	1332.9	26.6	1352.3	21.8	1383.1	36.4	1383.1	36.4	96.4
GC-01CP-26	115	14427	2.7	11.3268	1.4	2.7298	2.1	0.2243	1.6	0.75	1304.3	18.9	1336.6	15.9	1388.7	27.3	1388.7	27.3	93.9
GC-01CP-43	557	14007	4.2	11.2330	0.5	2.3215	3.2	0.1891	3.1	0.99	1116.7	32.2	1218.9	22.6	1404.6	9.7	1404.6	9.7	79.5
GC-01CP-24	145	2814	3.8	11.2240	3.8	2.4936	6.9	0.2030	5.8	0.83	1191.4	62.8	1270.2	50.3	1406.2	73.4	1406.2	73.4	84.7
GC-01CP-65	83	9995	2.1	11.1861	1.6	2.8973	2.4	0.2351	1.8	0.73	1361.0	21.6	1381.2	18.1	1412.6	31.1	1412.6	31.1	96.3
GC-01CP-21	93	10864	1.2	11.1224	2.8	3.1118	4.1	0.2510	3.0	0.73	1443.7	38.3	1435.6	31.3	1423.6	53.3	1423.6	53.3	101.4
GC-01CP-29	166	26944	1.4	11.1110	2.0	2.8894	2.7	0.2328	1.8	0.68	1349.4	22.5	1379.2	20.6	1425.5	38.4	1425.5	38.4	94.7
GC-01CP-64	464	18751	3.7	11.0575	0.5	2.7383	1.6	0.2196	1.5	0.95	1279.8	17.5	1338.9	11.8	1434.7	9.1	1434.7	9.1	89.2
GC-01CP-38	211	18810	2.7	11.0570	1.3	3.0858	2.2	0.2475	1.8	0.81	1425.3	22.8	1429.2	16.8	1434.8	24.4	1434.8	24.4	99.3
GC-01CP-67	98	11381	1.6	11.0432	1.5	3.1627	2.3	0.2533	1.7	0.76	1455.5	22.7	1448.1	17.7	1437.2	28.3	1437.2	28.3	101.3
GC-01CP-72	97	14360	1.7	11.0299	2.3	3.2547	2.4	0.2604	0.9	0.38	1491.7	12.2	1470.3	18.9	1439.5	42.9	1439.5	42.9	103.6
GC-01CP-94	255	37775	2.2	11.0263	0.8	3.2311	2.6	0.2584	2.5	0.95	1481.6	32.8	1464.6	20.2	1440.1	15.1	1440.1	15.1	102.9
GC-01CP-71	233	32667	3.2	11.0141	0.4	3.0656	1.8	0.2449	1.7	0.97	1412.1	21.7	1424.1	13.4	1442.2	7.8	1442.2	7.8	97.9
GC-01CP-34	108	19907	1.2	10.9687	1.0	3.0783	1.9	0.2449	1.6	0.85	1412.0	20.6	1427.3	14.6	1450.1	18.8	1450.1	18.8	97.4
GC-01CP-90	55	7827	1.5	10.9655	3.7	3.1771	4.2	0.2527	2.0	0.48	1452.2	26.4	1451.6	32.5	1450.7	70.2	1450.7	70.2	100.1
GC-01CP-44	124	20276	1.3	10.9306	1.5	3.1568	2.8	0.2503	2.3	0.85	1439.8	30.2	1446.7	21.3	1456.7	27.8	1456.7	27.8	98.8
GC-01CP-5	54	8473	2.0	10.8660	2.7	3.1817	3.6	0.2507	2.4	0.66	1442.3	30.8	1452.7	28.1	1468.0	52.2	1468.0	52.2	98.3
GC-01CP-17	75	11068	2.4	10.8580	2.1	3.0557	3.8	0.2406	3.1	0.83	1390.0	38.8	1421.7	28.7	1469.4	40.0	1469.4	40.0	94.6
GC-01CP-5	72	18148	1.6	10.8498	2.0	3.0802	3.1	0.2424	2.3	0.75	1399.1	29.3	1427.8	23.7	1470.8	38.4	1470.8	38.4	95.1
GC-01CP-92	265	10446	4.0	10.8561	1.3	2.9013	5.3	0.2200	5.2	0.97	1282.0	60.2	1382.3	40.3	1540.6	24.0	1540.6	24.0	83.2
GC-01CP-51	52	6556	1.9	9.9760	2.6	3.9491	5.0	0.2857	4.3	0.85	1620.2	61.4	1623.8	40.7	1628.5	48.7	1628.5	48.7	99.5
GC-01CP-22	51	8637	2.1	9.9705	2.3	4.0989	3.3	0.2964	2.5	0.74	1673.5	36.2	1654.1	27.3	1629.6	42.1	1629.6	42.1	102.7
GC-01CP-20	318	20120	3.3	9.9617	1.0	3.0599	22.0	0.2211	21.9	1.00	1287.5	256.2	1422.7	169.7	1631.2	18.9	1631.2	18.9	78.9
GC-01CP-19	539	12548	2.9	9.8920	0.5	3.1051	4.8	0.2228	4.8	1.00	1296.5	56.4	1433.9	37.1	1644.2	8.7	1644.2	8.7	78.9
GC-01CP-55	340	9070	2.2	9.8518	0.5	3.0216	9.1	0.2159	9.1	1.00	1260.2	103.8	1413.1	69.4	1651.8	9.3	1651.8	9.3	76.3
GC-01CP-13	161	18614	2.2	9.8037	0.5	3.8386	1.0	0.2729	0.9	0.88	1555.7	12.8	1600.9	8.4	1660.9	9.0	1660.9	9.0	93.7
GC-01CP-40	120	21740	4.9	9.7596	1.3	4.1653	2.8	0.2948	2.4	0.87	1665.6	35.5	1667.2	22.7	1669.2	24.8	1669.2	24.8	99.8
GC-01CP-85	167	7494	3.6	9.6775	1.7	3.4897	2.7	0.2449	2.1	0.77	1412.3	26.0	1524.9	21.0	1684.8	31.2	1684.8	31.2	83.8
GC-01CP-60	69	10401	2.1	9.6750	1.4	4.2454	3.1	0.2979	2.7	0.88	1680.9	40.4	1682.9	25.3	1685.3	26.5	1685.3	26.5	99.7
GC-01CP-73	134	16995	5.1	9.6560	1.0	3.7830	3.4	0.2649	3.2	0.96	1515.0	43.2	1589.1	26.9	1688.9	18.3	1688.9	18.3	89.7
GC-01CP-56	81	14357	4.6	9.6522	1.9	4.3534	2.9	0.3048	2.2	0.75	1714.8	33.1	1703.5	24.1	1689.6	35.4	1689.6	35.4	101.5
GC-01CP-9	118	18378	1.6	9.6459	1.0	4.3890	1.6	0.3070	1.3	0.78	1726.2	19.4	1710.3	13.5	1690.8	18.8	1690.8	18.8	102.1
GC-01CP-57	443	6826	3.2	9.6453	0.8	3.8147	1.1	0.2669	0.8	0.73	1524.8	11.0	1595.8	8.9	1690.9	13.9	1690.9	13.9	90.2
GC-01CP-27	248	31181																	

## U-Pb geochronologic analyses of the Osgood Mountain Quartzite and Preble Formation

Isotope ratios											Apparent ages (Ma)								
Analysis	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)
SAMPLE: Preble Formation GC-01-CP LOCATION: Goughs Canyon, Osgood Mountains 0468513 4554509 (NAD 83 UTM 11N)																			
GC-01CP-47	92	15847	4.7	9.0776	1.2	4.9457	2.5	0.3256	2.2	0.87	1817.1	34.3	1810.1	21.0	1802.1	22.2	1802.1	22.2	100.8
GC-01CP-7	102	15599	9.9	8.8955	1.1	5.0375	2.2	0.3250	1.9	0.86	1814.1	29.5	1825.6	18.4	1838.8	20.1	1838.8	20.1	98.7
GC-01CP-63	218	41428	10.7	8.8852	0.5	5.0193	2.9	0.3234	2.8	0.98	1806.6	44.4	1822.6	24.3	1840.9	9.3	1840.9	9.3	98.1
GC-01CP-86	96	13285	6.2	8.8626	1.1	5.1343	2.7	0.3300	2.4	0.91	1838.5	39.1	1841.8	22.9	1845.5	20.6	1845.5	20.6	99.6
GC-01CP-6	51	9705	1.8	8.6385	2.5	5.4723	3.8	0.3428	2.8	0.74	1900.4	46.1	1896.3	32.4	1891.7	45.4	1891.7	45.4	100.5
GC-01CP-77	222	20999	1.5	5.6427	2.0	10.7180	2.6	0.4386	1.7	0.65	2344.5	32.9	2499.0	24.0	2627.0	32.7	2627.0	32.7	89.2
GC-01CP-52	208	25549	1.8	5.6363	0.5	10.9378	1.2	0.4471	1.1	0.91	2382.4	22.3	2517.9	11.4	2628.9	8.3	2628.9	8.3	90.6
GC-01CP-4	125	52373	1.2	5.5033	0.5	13.0332	2.0	0.5202	2.0	0.97	2700.0	43.9	2682.1	19.3	2668.5	7.8	2668.5	7.8	101.2
GC-01CP-98	212	69019	1.2	5.5013	0.4	12.6590	1.5	0.5051	1.4	0.96	2635.6	30.6	2654.6	13.9	2669.1	6.8	2669.1	6.8	98.7
GC-01CP-75	154	50100	1.2	5.4968	0.4	12.6644	3.1	0.5049	3.1	0.99	2634.8	66.3	2655.0	29.1	2670.5	6.7	2670.5	6.7	98.7
GC-01CP-58	290	27914	2.8	5.4287	0.9	10.5648	3.3	0.4160	3.2	0.96	2242.1	61.0	2485.6	31.0	2691.1	14.7	2691.1	14.7	83.3
GC-01CP-18	105	20755	1.1	5.3618	0.4	12.4128	4.2	0.4827	4.1	1.00	2539.0	87.0	2636.1	39.1	2711.6	6.3	2711.6	6.3	93.6
GC-01CP-80	106	39471	2.3	5.3240	0.4	14.2995	1.0	0.5522	0.8	0.88	2834.1	19.2	2769.8	9.0	2723.2	7.4	2723.2	7.4	104.1
GC-01CP-93	359	44185	2.3	5.2510	0.3	13.4855	1.5	0.5136	1.5	0.98	2671.9	32.7	2714.3	14.3	2746.0	4.4	2746.0	4.4	97.3

### Notes:

1. Analyses with >10% uncertainty (1-sigma) in 206Pb/238U age are not included.
2. Analyses with >10% uncertainty (1-sigma) in 206Pb/207Pb age are not included, unless 206Pb/238U age is <500 Ma.
3. Best age is determined from 206Pb/238U age for analyses with 206Pb/238U age <1000 Ma and from 206Pb/207Pb.
4. Concordance is based on 206Pb/238U age / 206Pb/207Pb age. Value is not reported for 206Pb/238U ages <500
5. Analyses with 206Pb/238U age > 500 Ma and with >30% discordance (<70% concordance) are not included.
6. Analyses with 206Pb/238U age > 500 Ma and with >5% reverse discordance (<105% concordance) are not included.
7. All uncertainties are reported at the 1-sigma level, and include only measurement errors.
8. Systematic errors are as follows (at 2-sigma level): [sample 1: 2.5% (206Pb/238U) & 1.4% (206Pb/207Pb)] These values are reported on cells U1 and W1 of NUagecalc.
9. Analyses conducted by LA-MC-ICPMS, as described by Gehrels et al. (2008).
10. U concentration and U/Th are calibrated relative to Sri Lanka zircon standard and are accurate to ~20%.
11. Common Pb correction is from measured 204Pb with common Pb composition interpreted from Stacey and Kramers (1975).
12. Common Pb composition assigned uncertainties of 1.5 for 206Pb/204Pb, 0.3 for 207Pb/204Pb, and 2.0 for 208Pb/204Pb.
13. U/Pb and 206Pb/207Pb fractionation is calibrated relative to fragments of a large Sri Lanka zircon of  $563.5 \pm 3.2$  Ma (2-sigma).
14. U decay constants and composition as follows:  $238\text{U} = 9.8485 \times 10^{-10}$ ,  $235\text{U} = 1.55125 \times 10^{-10}$ ,  $238\text{U}/235\text{U} = 137.88$ .
15. Weighted mean and concordia plots determined with Isoplot (Ludwig, 2008).



**APPENDIX C**

**U-Pb geochronologic analyses of  
selected Roberts Mountains allochthon strata**

Table notes are at the end of the appendix.

	Isotope ratios										Apparent ages (Ma)									
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc.	Discord.
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)
<b>Sample: Snow Canyon Formation. Location: Snow Canyon, Independence Mountains; 0579760 4585698 (NAD 83 UTM 11T)</b>																				
SNOW-CYN-171	530	103043	4.4	16.9299	0.8	0.7446	1.2	0.0914	0.9	0.76	564.0	5.1	565.1	5.4	569.6	17.5	564.0	5.1	99.0	1.0
SNOW-CYN-182	221	138508	1.1	13.9234	0.6	1.6320	1.2	0.1648	1.0	0.87	983.4	9.3	982.6	7.4	980.9	11.6	980.9	11.6	100.3	-0.3
SNOW-CYN-75	40	18372	1.4	13.8037	3.3	1.7128	3.6	0.1715	1.5	0.41	1020.2	14.0	1013.3	23.0	998.4	66.2	998.4	66.2	102.2	-2.2
SNOW-CYN-82	92	61584	1.0	9.6861	0.8	4.2147	1.1	0.2961	0.8	0.73	1671.9	12.1	1676.9	9.3	1683.2	14.4	1683.2	14.4	99.3	0.7
SNOW-CYN-118	228	22936	4.3	9.1736	0.3	4.1335	3.1	0.2750	3.1	0.99	1566.2	42.8	1660.9	25.3	1782.9	5.7	1782.9	5.7	87.8	12.2
SNOW-CYN-61	38	34156	0.7	9.1307	1.7	4.9244	3.1	0.3261	2.6	0.85	1819.5	41.5	1806.4	26.2	1791.4	30.2	1791.4	30.2	101.6	-1.6
SNOW-CYN-98	59	33444	0.8	9.0275	0.6	4.8745	1.2	0.3192	1.0	0.85	1785.6	15.4	1797.9	9.8	1812.1	10.9	1812.1	10.9	98.5	1.5
SNOW-CYN-87	55	60278	1.1	9.0027	1.0	4.8071	1.7	0.3139	1.4	0.81	1759.7	21.2	1786.1	14.3	1817.1	18.0	1817.1	18.0	96.8	3.2
SNOW-CYN-63	29	74802	0.7	9.9658	1.3	4.9828	2.2	0.3240	1.8	0.82	1809.3	28.8	1816.4	18.8	1824.6	23.2	1824.6	23.2	99.2	0.8
SNOW-CYN-9	62	32646	0.6	8.9568	0.9	5.1270	2.4	0.3331	2.2	0.93	1853.2	35.9	1840.6	20.3	1826.4	15.6	1826.4	15.6	101.5	-1.5
SNOW-CYN-30	53	42708	0.9	8.9538	0.8	4.9671	1.2	0.3226	0.9	0.76	1802.2	14.9	1813.7	10.5	1827.0	14.6	1827.0	14.6	98.6	1.4
SNOW-CYN-196	64	59693	0.7	8.9475	1.0	5.0527	1.6	0.3279	1.3	0.80	1828.1	20.8	1828.2	13.9	1828.3	17.9	1828.3	17.9	100.0	0.0
SNOW-CYN-113	61	77813	0.6	8.9453	0.6	5.0860	1.3	0.3300	1.1	0.88	1838.2	18.3	1833.8	11.0	1828.7	11.2	1828.7	11.2	100.5	-0.5
SNOW-CYN-38	40	62208	0.8	8.9424	1.5	4.9646	1.7	0.3220	0.8	0.47	1799.4	12.2	1813.3	14.1	1829.3	26.8	1829.3	26.8	98.4	1.6
SNOW-CYN-32	257	74199	1.0	8.9331	0.3	4.9875	0.8	0.3231	0.8	0.93	1805.0	12.0	1817.2	6.9	1831.2	5.3	1831.2	5.3	98.6	1.4
SNOW-CYN-27	38	25786	0.7	8.9315	1.4	4.9938	1.7	0.3235	0.9	0.55	1806.7	14.4	1818.3	14.1	1831.5	25.3	1831.5	25.3	98.6	1.4
SNOW-CYN-14	41	40359	0.9	8.9305	1.7	5.1028	2.3	0.3305	1.5	0.68	1840.8	24.6	1836.6	19.2	1831.7	30.2	1831.7	30.2	100.5	-0.5
SNOW-CYN-141	42	34914	1.2	8.9287	1.0	5.1209	1.9	0.3316	1.6	0.85	1846.2	26.3	1839.6	16.4	1832.1	18.5	1832.1	18.5	100.8	-0.8
SNOW-CYN-36	22	21333	1.4	8.9287	2.3	4.9182	2.8	0.3185	1.5	0.53	1782.3	22.9	1805.4	23.3	1832.1	42.3	1832.1	42.3	97.3	2.7
SNOW-CYN-121	59	43360	1.3	8.9282	1.1	5.0217	1.4	0.3252	0.9	0.65	1815.0	14.4	1823.0	11.8	1832.2	19.1	1832.2	19.1	99.1	0.9
SNOW-CYN-144	44	51830	1.5	8.9251	1.0	5.0984	2.2	0.3300	1.9	0.88	1838.5	31.1	1835.8	18.7	1832.8	18.7	1832.8	18.7	100.3	-0.3
SNOW-CYN-70	244	279260	0.9	8.9240	0.2	5.1293	1.4	0.3320	1.4	0.98	1848.0	21.7	1841.0	11.7	1833.0	4.3	1833.0	4.3	100.8	-0.8
SNOW-CYN-77	69	122422	1.0	8.9221	0.8	4.9495	1.2	0.3203	1.0	0.79	1791.1	15.3	1810.7	10.4	1833.4	13.7	1833.4	13.7	97.7	2.3
SNOW-CYN-22	28	27707	1.2	8.9195	2.0	5.0542	2.5	0.3270	1.4	0.58	1823.6	23.0	1828.5	21.1	1833.9	36.7	1833.9	36.7	99.4	0.6
SNOW-CYN-179	41	29085	0.8	8.9153	0.9	5.1042	1.6	0.3300	1.3	0.83	1838.6	21.2	1836.8	13.6	1834.8	16.3	1834.8	16.3	100.2	-0.2
SNOW-CYN-78	35	49548	0.8	8.9141	1.0	5.0933	1.6	0.3293	1.3	0.80	1834.9	20.7	1835.0	13.7	1835.0	17.7	1835.0	17.7	100.0	0.0
SNOW-CYN-142	35	24716	1.0	8.9133	1.4	5.1169	1.9	0.3308	1.3	0.67	1842.2	20.2	1838.9	15.9	1835.2	25.2	1835.2	25.2	100.4	-0.4
SNOW-CYN-126	138	84479	0.7	8.9119	0.4	4.9794	1.0	0.3218	0.9	0.91	1798.7	14.5	1815.8	8.5	1835.5	7.5	1835.5	7.5	98.0	2.0
SNOW-CYN-191	51	36556	1.7	8.9117	1.0	5.0967	1.3	0.3294	0.9	0.66	1835.6	13.6	1835.6	10.9	1835.9	17.4	1835.5	17.4	100.0	0.0
SNOW-CYN-169	79	42642	1.1	8.9099	1.0	4.7323	2.1	0.3058	1.9	0.89	1720.0	28.7	1773.0	17.9	1835.9	17.9	1835.9	17.9	93.7	6.3
SNOW-CYN-132	51	53246	1.6	8.9013	1.0	5.0986	1.2	0.3292	0.7	0.55	1834.3	10.6	1835.9	10.3	1837.6	18.3	1837.6	18.3	99.8	0.2
SNOW-CYN-118	32	36028	0.7	8.9007	1.1	5.0476	1.8	0.3258	1.5	0.81	1818.2	23.1	1827.4	15.3	1837.8	19.2	1837.8	19.2	98.9	1.1
SNOW-CYN-151	79	97830	1.2	8.8996	0.5	5.1611	1.1	0.3331	1.0	0.90	1853.5	15.4	1846.2	9.0	1838.0	8.4	1838.0	8.4	100.8	-0.8
SNOW-CYN-135	115	155050	1.3	8.8940	0.4	5.1009	1.7	0.3290	1.7	0.97	1833.7	26.8	1836.3	14.6	1839.1	7.0	1839.1	7.0	99.7	0.3
SNOW-CYN-164	111	106505	0.9	8.8935	0.4	5.1512	0.9	0.3323	0.9	0.91	1849.3	13.8	1844.6	8.0	1839.2	6.9	1839.2	6.9	100.5	-0.5
SNOW-CYN-10	63	56602	1.7	8.8918	0.7	5.1925	1.3	0.3349	1.1	0.86	1861.9	18.1	1851.4	11.0	1839.6	11.9	1839.6	11.9	101.2	-1.2
SNOW-CYN-170	77	174559	1.1	8.8910	0.5	5.0152	1.1	0.3234	1.0	0.88	1806.3	15.6	1821.9	9.5	1839.7	9.6	1839.7	9.6	98.2	1.8
SNOW-CYN-172	20	30578	1.4	8.8891	1.8	5.1124	2.6	0.3296	1.8	0.71	1836.4	29.4	1838.2	21.9	1840.1	32.8	1840.1	32.8	99.8	0.2
SNOW-CYN-107	36	78308	1.1	8.8882	0.5	5.0762	1.2	0.3272	1.1	0.90	1824.9	17.7	1832.1	10.5	1840.3	9.8	1840.3	9.8	99.2	0.8
SNOW-CYN-89	43	63599	1.0	8.8864	0.9	4.9870	1.2	0.3214	0.8	0.65	1796.6	11.9	1817.1	9.9	1840.7	16.1	1840.7	16.1	97.6	2.4
SNOW-CYN-31	89	73950	1.2	8.8863	0.5	5.1486	4.2	0.3318	4.2	0.99	1847.2	67.3	1844.2	35.9	1840.7	9.4	1840.7	9.4	100.4	-0.4
SNOW-CYN-116	112	142250	1.7	8.8851	0.4	5.0449	1.9	0.3251	1.8	0.98	1814.6	29.0	1826.9	15.9	1840.9	6.9	1840.9	6.9	98.6	1.4
SNOW-CYN-200	115	109959	1.2	8.8846	0.5	5.0839	1.2	0.3276	1.1	0.92	1826.7	17.7	1833.4	10.3	1841.0	8.8	1841.0	8.8	99.2	0.8
SNOW-CYN-166	21	23371	0.8	8.8836	1.7	5.0960	2.1	0.3283	1.3	0.60	1830.3	20.2	1835.4	17.8	1841.3	30.2	1841.3	30.2	99.4	0.6
SNOW-CYN-185	117	105425	1.0	8.8826	0.4	5.1330	1.1	0.3307	1.0	0.93	1841.7	16.0	1841.6	9.1	1841.5	7.3	1841.5	7.3	100.0	0.0
SNOW-CYN-67	159	213593	2.3	8.8809	0.4	5.1590	1.3	0.3323	1.2	0.95	1849.5	19.3	1845.9	10.7	1841.8	6.9	1841.8	6.9	100.4	-0.4
SNOW-CYN-147	97	119246	1.5	8.8803	0.7	5.1142	1.1	0.3294	0.8	0.73	1835.4	12.5	1838.5	9.1	1841.9	13.1	1841.9	13.1	99.6	0.4
SNOW-CYN-58	106	142558	1.4	8.8795	0.2	5.0922	0.8	0.3279	0.8	0.96	1828.4	12.8	1834.8	7.1	1842.1	4.3	1842.1	4.3	99.3	0.7
SNOW-CYN-153	120	161045	1.5	8.8774	0.6	5.1887	1.9	0.3341	1.8	0.96	1858.1	29.4	1850.8	16.2	1842.5	10.0	1842.5	10.0	100.8	-0.8
SNOW-CYN-162	73	160747	1.9	8.8767	0.6	5.2203	1.3	0.3361	1.1	0.88	1867.8	18.3	1855.9	10.9	1842.7	10.8	1842.7	10.8	101.4	-1.4
SNOW-CYN-45	145	153135	1.4	8.8719	0.4	5.1497	1.0	0.3314	1.0	0.93	1845.0	15.6	1844.3	8.9	1843.6	7.0	1843.6	7.0	100.1	-0.1
SNOW-CYN-97	51	75488</																		

## U-Pb geochronologic analyses of selected Roberts Mountains allochthon strata

	Isotope ratios										Apparent ages (Ma)									
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc.	Discord.
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)
<b>Sample: Snow Canyon Formation. Location: Snow Canyon, Independence Mountains; 0579760 4585698 (NAD 83 UTM 11T)</b>																				
SNOW-CYN-117	28	46056	0.4	8.8209	0.8	5.0624	2.1	0.3239	1.9	0.93	1808.6	30.7	1829.8	17.8	1854.1	14.4	1854.1	14.4	97.5	2.5
SNOW-CYN-43	63	96847	1.6	8.8191	0.8	5.1893	1.5	0.3319	1.2	0.82	1847.7	19.3	1850.9	12.5	1854.4	15.2	1854.4	15.2	99.6	0.4
SNOW-CYN-190	96	79453	1.0	8.8049	0.5	5.2493	1.2	0.3352	1.0	0.88	1863.6	16.7	1860.7	10.0	1857.3	9.8	1857.3	9.8	100.3	-0.3
SNOW-CYN-129	207	322649	4.7	8.8048	0.4	5.1065	1.5	0.3261	1.5	0.96	1819.4	23.1	1837.2	12.9	1857.4	7.4	1857.4	7.4	98.0	2.0
SNOW-CYN-7	90	95659	2.1	8.8007	0.7	5.2804	1.3	0.3370	1.1	0.86	1872.4	18.1	1865.7	11.0	1858.2	11.8	1858.2	11.8	100.8	-0.8
SNOW-CYN-15	143	160011	1.5	8.7858	0.3	5.2873	1.1	0.3369	1.0	0.96	1871.8	16.7	1866.8	9.1	1861.3	5.3	1861.3	5.3	100.6	-0.6
SNOW-CYN-198	84	67128	2.8	8.7442	0.7	5.3116	1.5	0.3369	1.3	0.87	1871.5	21.4	1870.7	12.9	1869.8	13.3	1869.8	13.3	100.1	-0.1
SNOW-CYN-173	101	187975	1.5	8.7147	1.9	5.3476	9.5	0.3380	9.3	0.98	1877.0	151.0	1876.5	81.1	1875.9	33.7	1875.9	33.7	100.1	-0.1
SNOW-CYN-49	155	236451	1.5	8.7019	0.3	5.3935	1.5	0.3404	1.5	0.98	1888.6	24.7	1883.8	13.3	1878.6	6.1	1878.6	6.1	100.5	-0.5
SNOW-CYN-161	20	14401	0.7	8.8899	3.0	5.3032	3.4	0.3342	1.7	0.49	1858.9	27.1	1869.4	29.1	1881.1	53.4	1881.1	53.4	98.8	1.2
SNOW-CYN-184	33	14196	0.8	8.8771	1.0	5.3435	1.4	0.3363	0.9	0.67	1868.7	15.1	1875.9	11.9	1883.7	18.6	1883.7	18.6	99.2	0.8
SNOW-CYN-193	47	51193	1.8	8.8649	1.1	5.3317	1.5	0.3351	1.0	0.88	1862.9	16.2	1874.0	12.5	1886.2	19.2	1886.2	19.2	98.8	1.2
SNOW-CYN-85	67	48040	0.6	8.6331	0.8	5.4444	1.5	0.3409	1.3	0.84	1891.0	20.9	1891.9	13.1	1892.9	15.0	1892.9	15.0	99.9	0.1
SNOW-CYN-163	50	19275	0.7	8.6250	1.0	5.4224	2.2	0.3392	2.0	0.90	1882.8	32.3	1888.4	18.9	1894.6	17.3	1894.6	17.3	99.4	0.6
SNOW-CYN-131	51	97967	0.7	8.6054	0.7	5.4474	1.0	0.3400	0.7	0.73	1886.6	11.5	1892.3	8.2	1898.6	11.8	1898.6	11.8	99.4	0.6
SNOW-CYN-17	41	19804	1.8	8.6010	1.4	5.5060	1.7	0.3435	0.9	0.53	1903.3	14.5	1901.5	14.4	1899.6	25.5	1899.6	25.5	100.2	-0.2
SNOW-CYN-44	263	143112	3.3	8.5768	0.2	5.5014	1.1	0.3422	1.1	0.99	1897.3	18.6	1900.8	9.8	1904.6	2.8	1904.6	2.8	99.6	0.4
SNOW-CYN-119	41	49974	1.2	8.5763	0.7	5.4997	1.2	0.3421	1.0	0.80	1896.7	15.7	1900.5	10.2	1904.7	12.7	1904.7	12.7	99.6	0.4
SNOW-CYN-66	46	64039	0.9	8.5736	0.9	5.5461	1.7	0.3449	1.5	0.85	1910.0	24.3	1907.8	14.9	1905.3	16.6	1905.3	16.6	100.2	-0.2
SNOW-CYN-24	24	26300	1.0	8.5711	1.8	5.5083	2.4	0.3424	1.6	0.67	1898.3	26.7	1901.9	20.9	1905.8	32.1	1905.8	32.1	99.6	0.4
SNOW-CYN-93	174	270312	1.3	8.5637	0.3	5.3319	1.1	0.3312	1.0	0.96	1844.0	16.4	1874.0	9.2	1907.4	5.5	1907.4	5.5	96.7	3.3
SNOW-CYN-197	73	17621	0.9	8.5621	0.7	5.2005	0.9	0.3229	0.6	0.66	1804.1	9.5	1852.7	7.8	1907.7	12.3	1907.7	12.3	94.6	5.4
SNOW-CYN-81	53	72976	0.8	8.5612	0.8	5.5731	2.4	0.3460	2.2	0.94	1915.7	36.6	1911.9	20.3	1907.9	15.0	1907.9	15.0	100.4	-0.4
SNOW-CYN-5	51	120427	0.9	8.5565	0.8	5.5541	2.1	0.3447	1.9	0.93	1909.1	32.2	1909.0	18.0	1908.9	13.9	1908.9	13.9	100.0	0.0
SNOW-CYN-73	64	43049	1.8	8.5523	0.8	5.4394	1.3	0.3374	1.1	0.81	1874.1	17.8	1891.1	11.6	1909.8	14.1	1909.8	14.1	98.1	1.9
SNOW-CYN-155	25	48093	0.9	8.5510	1.6	5.5469	2.2	0.3440	1.5	0.67	1905.9	24.1	1907.9	18.6	1910.0	28.7	1910.0	28.7	99.8	0.2
SNOW-CYN-120	52	68253	1.0	8.5493	0.8	5.4796	1.6	0.3398	1.3	0.87	1885.5	22.0	1897.4	13.4	1910.4	13.9	1910.4	13.9	98.7	1.3
SNOW-CYN-94	31	21928	0.7	8.5442	1.8	5.5028	2.3	0.3410	1.4	0.63	1891.5	23.7	1901.0	19.6	1911.5	31.8	1911.5	31.8	99.0	1.0
SNOW-CYN-12	178	265794	0.8	8.5365	0.2	5.6221	0.9	0.3481	0.9	0.98	1925.4	14.2	1919.5	7.5	1913.1	2.7	1913.1	2.7	100.6	-0.6
SNOW-CYN-139	153	165091	0.9	8.5273	0.3	5.4565	0.7	0.3375	0.6	0.92	1874.4	10.1	1893.8	5.8	1915.0	4.6	1915.0	4.6	97.9	2.1
SNOW-CYN-84	137	86779	1.4	8.5210	0.4	4.8362	1.7	0.2989	1.7	0.98	1685.7	24.6	1791.2	14.3	1916.4	6.5	1916.4	6.5	88.0	12.0
SNOW-CYN-42	78	96370	1.0	8.5178	0.3	5.5566	1.1	0.3403	1.1	0.96	1902.4	17.4	1909.4	9.5	1917.0	5.5	1917.0	5.5	99.2	0.8
SNOW-CYN-71	77	73803	1.3	8.5136	0.6	5.5132	1.2	0.3434	1.1	0.88	1888.7	17.6	1902.7	10.5	1919.0	10.3	1919.0	10.3	98.5	1.5
SNOW-CYN-52	102	104843	0.7	8.5122	0.6	5.6229	1.6	0.3471	1.5	0.92	1920.9	24.5	1919.6	13.8	1918.2	11.0	1918.2	11.0	100.1	-0.1
SNOW-CYN-91	163	95427	1.0	8.5103	0.3	5.5469	1.0	0.3424	1.0	0.96	1898.1	16.2	1907.9	8.8	1918.6	4.8	1918.6	4.8	98.9	1.1
SNOW-CYN-82	89	172417	1.1	8.4982	0.5	5.7182	1.8	0.3524	1.7	0.96	1946.2	28.7	1934.1	15.5	1921.2	9.5	1921.2	9.5	101.3	-1.3
SNOW-CYN-176	21	16687	0.9	8.4967	1.2	5.5232	1.9	0.3404	1.5	0.78	1888.4	24.5	1904.2	16.6	1921.5	21.8	1921.5	21.8	98.3	1.7
SNOW-CYN-56	84	70209	1.5	8.4947	0.6	5.6662	1.4	0.3491	1.3	0.89	1930.3	20.9	1926.2	12.1	1921.9	11.2	1921.9	11.2	100.4	-0.4
SNOW-CYN-167	133	11624	2.1	8.4863	0.3	5.5587	3.1	0.3421	3.1	1.00	1896.9	50.6	1909.7	26.6	1923.7	5.2	1923.7	5.2	98.6	1.4
SNOW-CYN-125	45	97694	1.2	8.4827	0.8	5.5844	1.4	0.3436	1.2	0.83	1903.8	19.1	1913.7	12.1	1924.4	14.1	1924.4	14.1	98.9	1.1
SNOW-CYN-60	68	36343	0.5	8.4825	0.5	5.5416	1.6	0.3409	1.5	0.96	1891.1	24.4	1907.1	13.4	1924.5	8.2	1924.5	8.2	98.3	1.7
SNOW-CYN-160	37	69009	0.8	8.4800	0.9	5.5747	1.6	0.3429	1.3	0.82	1900.4	21.8	1912.2	13.9	1925.0	16.7	1925.0	16.7	98.7	1.3
SNOW-CYN-8	21	22966	0.8	8.4799	2.5	5.5882	2.8	0.3437	1.2	0.45	1904.4	20.4	1914.3	23.8	1925.0	44.3	1925.0	44.3	98.9	1.1
SNOW-CYN-199	51	33912	1.2	8.4780	0.8	5.2437	2.1	0.3224	1.9	0.92	1801.6	30.2	1859.7	17.9	1925.4	14.9	1925.4	14.9	93.6	6.4
SNOW-CYN-51	68	129535	1.1	8.4745	0.7	5.6710	1.5	0.3486	1.4	0.90	1927.7	22.8	1927.0	13.2	1926.2	12.1	1926.2	12.1	100.1	-0.1
SNOW-CYN-16	61	65104	1.3	8.4678	1.0	5.6441	2.3	0.3466	2.1	0.91	1918.5	35.3	1922.9	20.1	1927.6	17.2	1927.6	17.2	99.5	0.5
SNOW-CYN-150	80	101032	1.0	8.4644	0.4	5.7295	1.2	0.3517	1.1	0.94	1942.9	18.4	1935.8	10.1	1928.3	6.9	1928.3	6.9	100.8	-0.8
SNOW-CYN-171	101	114861	1.0	8.4521	0.3	5.6969	1.1	0.3492	1.1	0.95	1930.9	17.9	1930.9	9.7	1930.9	6.1	1930.9	6.1	100.0	0.0
SNOW-CYN-130	185	400942	2.5	8.4329	0.3	5.5471	1.1	0.3393	1.0	0.95	1883.1	17.1	1907.9	9.5	1935.0	6.1	1935.0	6.1	97.3	2.7
SNOW-CYN-158	296	152000	3.1	8.4303	0.1	5.8503	1.6	0.3577	1.6	1.00	1971.3	26.3	1953.9	13.5	1935.5	2.4	1935.5	2.4	101.8	-1.8
SNOW-CYN-105	160	165134	1.1	8.4184	0.3	5.7095	1.4	0.3486	1.4	0.97	1927.9	23.2	1932.8	12.3	1938.0	5.7	1938.0	5.7	99.5	0.5
SNOW-CYN-76	42	42216	1.2	8.4155	1.0	5.6969	1.4	0.3477	0.9	0.69	1923.7	15.6	1930.9	11.8	1938.7	17.8	1938.7	17.8	99.2	0.8
SNOW-CYN-111	24	16024	0.4	8.3638	1.9	5.7827	2.8	0.3508	2.0	0.72	1938.3	33.4	1943.8	24.0	1949.7	34.4	1949.7	34.4	99.4	0.6
SNOW-CYN-20	140	189553	1.0	8.2253	0.2	6.0233	1.6	0.3593	1.6	0.99	1979.0	27.7	1979.2	14.3	1979.5	3.6	1979.5	3.6	100.0	0.0
SNOW-CYN-101	84	24275	1.6	8.1606	0.5	6.0538	0.9	0.3583	0.7	0.81	1974.1	12.4	1983.6	7.9	1993.5	9.5	1993.5	9.5	99.0	1.0
SNOW-CYN-145	112	168637	1.1	8.1305	0.6	6.2179	1.6	0.3667	1.5	0.93	2013.7	26.1	2007.0	14.1	2000.1	10.2	2000.1	10.2	100.7	-0.7
SNOW-CYN-175	29	26832	1.3	8.1279	1.3	6.0692	2.8	0.3578	2.5	0.89	1971.6	42.2	1985.8	24.4	2000.7	23.1	2000.7	23.1	98.5	1.5
SNOW-CYN-137	29	19947	1.0	8.0567	1.3	6.2843														



## U-Pb geochronologic analyses of selected Roberts Mountains allochthon strata

	Isotope ratios										Apparent ages (Ma)									
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc.	Discord.
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)
<b>Sample: Snow Canyon Formation. Location: Snow Canyon, Independence Mountains; 0579760 4585698 (NAD 83 UTM 11T)</b>																				
SNOW-CYN-187	22	6820	1.0	7.3431	1.4	7.4295	3.0	0.3957	2.6	0.88	2149.1	48.3	2164.5	26.9	2179.2	25.0	2179.2	25.0	98.6	1.4
SNOW-CYN-33	78	97913	1.0	7.3383	0.6	7.4187	2.1	0.3948	2.0	0.96	2145.2	37.1	2163.2	18.9	2180.3	9.8	2180.3	9.8	98.4	1.6
SNOW-CYN-50	48	44805	1.3	6.9949	0.5	8.2642	1.2	0.4193	1.1	0.91	2257.1	21.3	2260.4	11.2	2263.4	9.0	2263.4	9.0	99.7	0.3
SNOW-CYN-100	22	50899	1.4	6.8541	1.0	8.4369	2.5	0.4194	2.3	0.92	2257.8	43.6	2279.2	22.6	2298.4	16.6	2298.4	16.6	98.2	1.8
SNOW-CYN-177	52	61826	0.8	6.2895	0.4	10.1238	1.6	0.4618	1.5	0.97	2447.5	30.8	2446.1	14.4	2445.0	6.4	2445.0	6.4	100.1	-0.1
SNOW-CYN-69	56	63570	1.6	6.0084	0.3	11.0800	0.9	0.4827	0.9	0.93	2538.9	18.3	2529.9	8.7	2522.6	5.6	2522.6	5.6	100.6	-0.6
SNOW-CYN-46	276	284199	1.9	6.0006	0.6	11.0102	1.7	0.4792	1.6	0.94	2523.6	33.4	2524.0	15.8	2524.3	9.8	2524.3	9.8	100.0	0.0
SNOW-CYN-4	137	142690	1.1	5.9847	0.4	9.6794	4.6	0.4201	4.5	1.00	2261.1	86.7	2404.7	42.0	2528.7	6.3	2528.7	6.3	89.4	10.6
SNOW-CYN-6	32	61997	1.3	5.9371	0.3	11.3035	1.3	0.4867	1.2	0.96	2556.5	25.9	2548.5	11.9	2542.1	5.8	2542.1	5.8	100.6	-0.6
SNOW-CYN-96	65	69104	1.3	5.8166	0.4	11.5934	1.1	0.4891	1.0	0.94	2566.7	22.2	2572.1	10.5	2576.4	6.6	2576.4	6.6	99.6	0.4
SNOW-CYN-124	31	54605	1.4	5.8076	0.5	11.3933	1.2	0.4799	1.1	0.93	2526.8	23.6	2555.9	11.4	2579.0	7.7	2579.0	7.7	98.0	2.0
SNOW-CYN-178	42	11830	0.6	5.6477	0.8	11.9939	3.0	0.4913	2.9	0.96	2576.2	61.4	2603.9	28.1	2625.5	13.2	2625.5	13.2	98.1	1.9
SNOW-CYN-21	31	94305	1.2	5.6435	0.7	11.0815	1.6	0.4536	1.5	0.91	2411.1	29.7	2530.0	15.1	2626.8	11.0	2626.8	11.0	91.8	8.2
SNOW-CYN-72	101	95097	1.4	5.6658	0.3	12.2453	0.8	0.4943	0.8	0.92	2589.3	16.0	2623.4	7.7	2649.8	5.4	2649.8	5.4	97.7	2.3
SNOW-CYN-134	63	87523	0.6	5.5329	0.4	12.5376	1.3	0.5031	1.3	0.95	2627.2	27.5	2645.6	12.7	2659.6	7.2	2659.6	7.2	98.8	1.2
SNOW-CYN-54	101	158945	0.9	5.5320	0.1	12.5754	1.3	0.5045	1.3	1.00	2633.3	27.8	2648.4	12.2	2659.9	2.0	2659.9	2.0	99.0	1.0
SNOW-CYN-114	127	179368	1.8	5.4575	0.4	12.8436	1.8	0.5084	1.7	0.98	2649.6	37.4	2668.2	16.6	2682.4	6.3	2682.4	6.3	98.8	1.2
SNOW-CYN-156	27	49994	0.6	5.4306	0.9	13.0700	1.5	0.5148	1.2	0.78	2677.0	25.8	2684.7	14.2	2690.5	15.5	2690.5	15.5	99.5	0.5
SNOW-CYN-86	25	36023	1.2	5.4221	1.0	12.9442	1.3	0.5090	0.8	0.65	2652.5	17.8	2675.6	11.9	2693.1	15.9	2693.1	15.9	98.5	1.5
SNOW-CYN-25	115	205883	2.4	5.4168	0.2	13.2727	1.3	0.5214	1.3	0.99	2705.2	28.7	2699.2	12.4	2694.7	3.1	2694.7	3.1	100.4	-0.4
SNOW-CYN-65	57	55824	0.7	5.4016	0.5	13.2284	1.6	0.5182	1.6	0.96	2691.7	34.3	2696.1	15.4	2699.4	7.9	2699.4	7.9	99.7	0.3
SNOW-CYN-154	160	212749	1.7	5.3971	0.2	13.4276	1.0	0.5256	1.0	0.98	2722.9	21.9	2710.2	9.5	2700.7	3.4	2700.7	3.4	100.8	-0.8
SNOW-CYN-127	67	106148	0.4	5.3965	0.3	12.9554	1.0	0.5071	0.9	0.95	2644.1	20.0	2676.4	9.1	2700.9	5.0	2700.9	5.0	97.9	2.1
SNOW-CYN-104	91	192420	1.0	5.3937	0.2	13.6249	2.4	0.5330	2.4	1.00	2754.0	52.8	2724.0	22.4	2701.8	3.9	2701.8	3.9	101.9	-1.9
SNOW-CYN-13	116	242807	1.2	5.3870	0.2	13.3988	1.3	0.5235	1.3	0.99	2714.0	27.8	2708.2	12.0	2703.8	3.4	2703.8	3.4	100.4	-0.4
SNOW-CYN-41	58	98319	0.4	5.3824	0.4	13.2352	1.0	0.5167	0.9	0.89	2685.0	18.7	2696.6	9.1	2705.2	7.3	2705.2	7.3	99.3	0.7
SNOW-CYN-95	122	200709	0.4	5.3772	0.2	13.0938	1.1	0.5106	1.0	0.99	2659.4	22.7	2686.4	9.9	2706.8	2.5	2706.8	2.5	98.2	1.8
SNOW-CYN-112	22	31369	0.6	5.3767	0.7	13.3170	1.9	0.5193	1.8	0.93	2696.2	40.1	2702.4	18.4	2707.0	11.5	2707.0	11.5	99.6	0.4
SNOW-CYN-3	134	267317	1.1	5.3708	0.2	13.5939	1.2	0.5295	1.2	0.99	2739.4	27.2	2721.8	11.7	2708.8	3.0	2708.8	3.0	101.1	-1.1
SNOW-CYN-165	69	52575	0.7	5.3288	0.2	13.6884	1.9	0.5290	1.9	1.00	2737.4	42.3	2728.4	18.0	2721.7	2.7	2721.7	2.7	100.6	-0.6
SNOW-CYN-34	67	77982	0.9	5.3074	0.2	13.5090	0.8	0.5200	0.7	0.96	2699.2	15.8	2715.9	7.1	2728.4	3.7	2728.4	3.7	98.9	1.1
SNOW-CYN-192	85	137048	1.0	5.2992	0.3	13.7156	0.9	0.5271	0.9	0.95	2729.4	20.0	2730.3	8.9	2730.9	4.7	2730.9	4.7	99.9	0.1
SNOW-CYN-26	203	367278	0.8	5.2736	0.1	13.8384	0.5	0.5293	0.5	0.96	2738.5	10.2	2738.7	4.5	2738.9	2.3	2738.9	2.3	100.0	0.0
SNOW-CYN-152	50	103476	0.6	5.2493	0.4	14.0412	2.0	0.5346	2.0	0.98	2760.7	44.7	2752.5	19.4	2746.5	7.4	2746.5	7.4	100.5	-0.5
SNOW-CYN-23	34	83229	1.2	4.9896	0.5	15.1884	3.0	0.5496	2.9	0.98	2823.6	67.0	2827.1	28.4	2829.6	8.8	2829.6	8.8	99.8	0.2
SNOW-CYN-123	63	216756	1.3	4.9483	0.3	15.2936	0.8	0.5489	0.8	0.92	2820.5	17.2	2833.7	7.8	2843.1	5.1	2843.1	5.1	99.2	0.8
SNOW-CYN-157	85	108993	60.7	4.9314	0.2	15.7985	1.0	0.5651	0.9	0.97	2887.5	21.7	2864.7	9.1	2848.7	3.7	2848.7	3.7	101.4	-1.4
SNOW-CYN-128	33	44308	0.7	4.9293	0.2	15.0765	0.7	0.5390	0.6	0.95	2779.2	14.0	2820.1	6.2	2849.4	3.5	2849.4	3.5	97.5	2.5
SNOW-CYN-183	48	296306	1.3	3.2593	0.2	30.3488	1.4	0.7174	1.4	0.99	3486.4	36.6	3498.2	13.6	3504.9	3.4	3504.9	3.4	99.5	0.5
<b>Sample: McAfee Quartzite. Location: McAfee Creek, Independence Mountains; 0590637 4599583 (NAD 83 UTM 11T)</b>																				
McAfee-97SF1-70	148	54632	1.8	13.6151	1.3	1.7270	1.8	0.1705	1.2	0.66	1015.1	11.1	1018.6	11.5	1026.3	26.9	1026.3	26.9	98.9	1.1
McAfee-97SF1-8	365	133119	1.1	13.5575	0.4	1.7968	0.6	0.1767	0.4	0.66	1048.8	3.7	1044.3	3.8	1034.9	8.7	1034.9	8.7	101.3	-1.3
McAfee-97SF1-22	256	49041	1.2	13.5242	0.9	1.7693	1.3	0.1735	1.0	0.76	1031.6	9.4	1034.2	8.4	1039.8	17.2	1039.8	17.2	99.2	0.8
McAfee-97SF1-123	22	18308	0.7	9.4113	3.4	4.6400	3.8	0.3167	1.8	0.46	1773.6	27.3	1756.5	32.0	1736.1	62.3	1736.1	62.3	102.2	-2.2
McAfee-97SF1-196	100	59887	14.5	9.0724	0.6	4.8917	0.9	0.3219	0.7	0.74	1798.9	10.7	1800.8	7.7	1803.1	11.2	1803.1	11.2	99.8	0.2
McAfee-97SF1-118	125	79313	1.6	9.0669	0.4	5.0574	0.6	0.3326	0.5	0.80	1850.8	8.3	1829.0	5.5	1804.2	7.1	1804.2	7.1	102.6	-2.6
McAfee-97SF1-39	26	16134	0.4	9.0410	1.6	5.2194	2.3	0.3422	1.6	0.72	1897.4	26.7	1855.8	19.2	1809.4	28.4	1809.4	28.4	104.9	-4.9
McAfee-97SF1-132	39	20055	1.5	9.0232	1.7	5.1813	2.1	0.3391	1.2	0.56	1882.2	19.1	1849.5	17.9	1813.0	31.7	1813.0	31.7	103.8	-3.8
McAfee-97SF1-138	60	28368	0.8	9.0139	1.8	5.0199	3.3	0.3282	2.7	0.84	1829.5	43.6	1822.7	27.6	1814.8	32.4	1814.8	32.4	100.8	-0.8
McAfee-97SF1-105	41	18612	0.9	8.9972	1.6	5.1451	2.5	0.3357	2.0	0.78	1866.1	31.9	1843.6	21.5	1818.2	28.6	1818.2	28.6	102.6	-2.6
McAfee-97SF1-37	55	26707	0.4	8.9772	1.3	5.1850	2.0	0.3376	1.4	0.73	1875.1	23.3	1850.2	16.7	1822.3	24.2	1822.3	24.2	102.9	-2.9
McAfee-97SF1-88	47	19442	1.3	8.9765	1.4	5.0823	1.6	0.3309	0.9	0.54	1842.6	14.0	1833.1	13.8	1822.4	24.9	1822.4	24.9	101.1	-1.1
McAfee-97SF1-69	38	26642	0.6	8.9759	1.5	5.0202	2.0	0.3268	1.2	0.63	1822.9	19.6	1822.7	16.6	1822.5	27.7	1822.5	27.7	100.0	0.0
McAfee-97SF1-55	40	25808	0.6	8.9653	0.9	5.1478	1.3	0.3347	1.0	0.75	1861.2	15.7	1844.0	11.0	1824.7	15.7	1824.7	15.7	102.0	-2.0
McAfee-97SF1-67	193	230751	1.8	8.9651	0.4	5.0628	0.8	0.3292	0.7	0.84	1834.5	10.4	1829.9	6.6	1824.7	7.6	1824.7	7.6	100.5	-0.5
McAfee-97SF1-46	77	24989	1.1	8.9645	0.7	5.1457	1.1	0.3346	0.9	0.78	1860.4	14.0	1843.7	9.5	1824.8	12.7	1824.8	12.7	102.0	-2.0
McAfee-97SF1-73	46	32729	1.5	8.9609	1.1	4.9831	1.6	0.3239	1.1	0.72	1808.5	17.9	1816.5	13.4	1825.6	20.1	1825.6	20.1	99.1	0.9
McAfee-97SF1-59	51	29910	1.1	8.9555	1.0	5.1569	1.5	0.												

## U-Pb geochronologic analyses of selected Roberts Mountains allochthon strata

	Isotope ratios										Apparent ages (Ma)									
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc.	Discord.
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)
<b>Sample: McAfee Quartzite, Location: McAfee Creek, Independence Mountains; 0590637 4599583 (NAD 83 UTM 11T)</b>																				
McAfee-97SF1-87	112	85370	0.9	8.8939	0.6	5.1227	1.4	0.3304	1.3	0.89	1840.5	20.5	1839.9	12.2	1839.1	11.7	1839.1	11.7	100.1	-0.1
McAfee-97SF1-26	64	47470	0.8	8.8926	1.3	5.1972	2.6	0.3352	2.2	0.85	1863.5	35.7	1852.2	22.0	1839.4	24.3	1839.4	24.3	101.3	-1.3
McAfee-97SF1-127	31	28519	0.6	8.8922	1.4	5.3107	1.6	0.3425	0.7	0.45	1898.7	11.7	1870.6	13.6	1839.5	25.8	1839.5	25.8	103.2	-3.2
McAfee-97SF1-104	41	39172	0.6	8.8915	1.2	5.3890	2.7	0.3475	2.4	0.90	1922.7	39.7	1883.1	22.8	1839.6	21.4	1839.6	21.4	104.5	-4.5
McAfee-97SF1-180	58	47141	1.0	8.8888	0.9	5.1972	1.1	0.3350	0.6	0.53	1862.8	9.0	1852.2	9.0	1840.2	16.3	1840.2	16.3	101.2	-1.2
McAfee-97SF1-54	58	68284	1.0	8.8887	1.2	5.1392	1.4	0.3313	0.7	0.49	1844.7	10.8	1842.6	11.7	1840.2	21.7	1840.2	21.7	100.2	-0.2
McAfee-97SF1-80	54	24511	0.9	8.8875	0.8	5.2569	1.5	0.3389	1.2	0.82	1881.1	19.7	1861.9	12.5	1840.5	15.1	1840.5	15.1	102.2	-2.2
McAfee-97SF1-191	221	112477	1.4	8.8869	0.3	5.2600	0.5	0.3390	0.4	0.83	1882.0	7.0	1862.4	4.4	1840.6	5.1	1840.6	5.1	102.3	-2.3
McAfee-97SF1-140	143	92852	1.1	8.8855	0.9	5.3309	1.0	0.3435	0.6	0.56	1903.7	9.7	1873.8	8.9	1840.9	15.7	1840.9	15.7	103.4	-3.4
McAfee-97SF1-151	70	37149	0.9	8.8840	1.0	5.1835	1.3	0.3340	0.8	0.61	1857.7	12.6	1849.9	11.0	1841.2	18.5	1841.2	18.5	100.9	-0.9
McAfee-97SF1-130	82	30133	0.9	8.8816	0.7	5.1574	1.0	0.3322	0.7	0.71	1849.1	11.4	1845.6	8.5	1841.7	12.7	1841.7	12.7	100.4	-0.4
McAfee-97SF1-152	208	172847	1.4	8.8792	0.3	5.2894	0.6	0.3406	0.5	0.88	1889.7	8.7	1867.1	5.2	1842.1	5.3	1842.1	5.3	102.6	-2.6
McAfee-97SF1-65	112	62750	0.9	8.8774	0.5	5.2189	1.3	0.3360	1.2	0.92	1867.5	19.3	1855.7	11.0	1842.5	9.3	1842.5	9.3	101.4	-1.4
McAfee-97SF1-30	64	45044	0.9	8.8749	1.0	5.3672	1.5	0.3455	1.1	0.73	1912.9	17.7	1879.6	12.6	1843.0	18.2	1843.0	18.2	103.8	-3.8
McAfee-97SF1-129	106	109727	0.9	8.8734	0.5	5.2001	1.6	0.3347	1.5	0.94	1860.9	23.8	1852.6	13.3	1843.3	9.3	1843.3	9.3	101.0	-1.0
McAfee-97SF1-41	32	15171	0.8	8.8724	1.5	5.1290	1.7	0.3300	0.8	0.46	1838.6	12.6	1840.9	14.5	1843.5	27.5	1843.5	27.5	99.7	0.3
McAfee-97SF1-142	56	33815	1.0	8.8717	1.2	5.2151	2.2	0.3356	1.8	0.83	1865.3	29.9	1855.1	19.0	1843.7	22.5	1843.7	22.5	101.2	-1.2
McAfee-97SF1-2	48	36525	1.0	8.8696	1.5	5.2986	2.8	0.3408	2.3	0.84	1890.8	38.3	1868.6	23.7	1844.1	27.2	1844.1	27.2	102.5	-2.5
McAfee-97SF1-181	71	47991	1.2	8.8685	0.6	5.2031	0.8	0.3347	0.6	0.73	1880.9	9.9	1853.1	7.2	1844.3	10.4	1844.3	10.4	100.9	-0.9
McAfee-97SF1-95	43	24711	1.5	8.8684	1.2	5.2165	1.4	0.3355	0.7	0.52	1865.1	11.8	1855.3	11.9	1844.4	21.5	1844.4	21.5	101.1	-1.1
McAfee-97SF1-76	23	18357	0.5	8.8658	1.6	5.1964	2.9	0.3341	2.4	0.83	1858.4	39.1	1852.0	24.9	1844.9	29.8	1844.9	29.8	100.7	-0.7
McAfee-97SF1-85	46	33107	1.0	8.8658	1.4	5.1648	1.9	0.3321	1.3	0.66	1848.5	20.3	1846.8	16.3	1844.9	26.2	1844.9	26.2	100.2	-0.2
McAfee-97SF1-9	41	29173	0.8	8.8614	1.3	5.1313	1.8	0.3298	1.2	0.69	1837.3	19.5	1841.3	15.0	1845.8	23.1	1845.8	23.1	99.5	0.5
McAfee-97SF1-7	147	73596	1.8	8.8605	0.5	5.2193	0.6	0.3354	0.4	0.60	1864.5	5.9	1855.8	5.1	1846.0	8.7	1846.0	8.7	101.0	-1.0
McAfee-97SF1-75	55	45823	0.7	8.8590	1.0	5.1183	1.2	0.3289	0.7	0.61	1832.9	11.8	1839.2	10.3	1846.3	17.4	1846.3	17.4	99.3	0.7
McAfee-97SF1-21	29	38684	0.8	8.8559	1.6	5.1432	2.1	0.3303	1.3	0.63	1840.0	21.0	1843.3	17.8	1846.9	29.5	1846.9	29.5	99.6	0.4
McAfee-97SF1-15	141	135061	0.9	8.8552	0.4	5.2298	0.8	0.3359	0.7	0.86	1866.8	11.5	1857.5	7.0	1847.0	7.6	1847.0	7.6	101.1	-1.1
McAfee-97SF1-119	31	25016	0.8	8.8529	1.4	5.1270	1.5	0.3292	0.6	0.37	1834.5	9.0	1840.6	12.9	1847.5	25.4	1847.5	25.4	99.3	0.7
McAfee-97SF1-185	172	119121	1.4	8.8499	0.5	5.3336	0.6	0.3423	0.4	0.64	1897.9	6.5	1874.3	5.3	1848.1	8.5	1848.1	8.5	102.7	-2.7
McAfee-97SF1-179	148	175958	1.4	8.8479	0.4	5.2897	1.0	0.3394	1.0	0.94	1884.0	15.8	1867.2	8.8	1848.5	6.4	1848.5	6.4	101.9	-1.9
McAfee-97SF1-149	69	48868	0.8	8.8473	0.7	5.2416	1.3	0.3363	1.0	0.83	1869.0	17.0	1859.0	10.7	1848.7	12.6	1848.7	12.6	101.1	-1.1
McAfee-97SF1-170	350	260291	2.6	8.8434	0.2	5.3104	0.7	0.3406	0.7	0.95	1889.5	11.6	1870.5	6.4	1849.5	4.1	1849.5	4.1	102.2	-2.2
McAfee-97SF1-158	130	52699	0.7	8.8409	0.3	5.3466	0.6	0.3428	0.6	0.90	1900.3	9.2	1876.4	5.3	1850.0	4.9	1850.0	4.9	102.7	-2.7
McAfee-97SF1-121	95	17877	0.5	8.8303	0.6	5.2411	0.7	0.3357	0.4	0.57	1865.7	6.9	1859.3	6.4	1852.1	11.1	1852.1	11.1	100.7	-0.7
McAfee-97SF1-101	58	27385	0.9	8.8271	1.0	5.3236	2.3	0.3408	2.1	0.91	1890.6	34.7	1872.7	19.9	1852.8	17.2	1852.8	17.2	102.0	-2.0
McAfee-97SF1-114	122	65540	1.4	8.8252	0.8	5.2906	1.0	0.3386	0.7	0.63	1880.1	10.7	1867.3	8.9	1853.2	14.7	1853.2	14.7	101.5	-1.5
McAfee-97SF1-31	37	11541	1.0	8.8242	3.3	5.1687	3.4	0.3308	0.8	0.24	1842.2	13.1	1847.5	29.0	1853.4	59.7	1853.4	59.7	99.4	0.6
McAfee-97SF1-43	62	41289	0.8	8.8219	0.8	5.2456	1.1	0.3356	0.8	0.73	1865.6	13.1	1860.1	9.4	1853.9	13.6	1853.9	13.6	100.6	-0.6
McAfee-97SF1-134	42	50744	0.9	8.8111	1.3	5.2838	1.7	0.3377	1.1	0.65	1875.4	18.5	1866.3	14.9	1856.1	23.9	1856.1	23.9	101.0	-1.0
McAfee-97SF1-174	51	51900	0.6	8.8078	1.4	5.2234	1.7	0.3337	1.0	0.58	1856.2	16.0	1856.4	14.6	1856.7	25.2	1856.7	25.2	100.0	0.0
McAfee-97SF1-156	27	31386	0.7	8.8017	1.9	5.3535	2.1	0.3417	0.9	0.44	1895.1	14.8	1877.5	17.7	1858.0	33.5	1858.0	33.5	102.0	-2.0
McAfee-97SF1-11	42	55871	1.1	8.7904	1.0	5.1715	1.5	0.3297	1.1	0.73	1836.9	17.0	1847.9	12.4	1860.3	18.0	1860.3	18.0	98.7	1.3
McAfee-97SF1-97	58	9808	1.0	8.7848	2.4	5.1379	2.5	0.3274	0.5	0.22	1825.5	8.6	1842.4	23.1	1861.5	44.0	1861.5	44.0	98.1	1.9
McAfee-97SF1-148	133	60532	0.9	8.7816	0.4	5.3854	0.9	0.3430	0.8	0.88	1901.1	12.4	1882.5	7.4	1862.1	7.5	1862.1	7.5	102.1	-2.1
McAfee-97SF1-186	167	76118	0.5	8.7801	0.4	5.2625	0.8	0.3351	0.7	0.85	1863.1	11.0	1862.8	6.8	1862.4	7.7	1862.4	7.7	100.0	0.0
McAfee-97SF1-188	209	102704	2.1	8.7798	0.2	5.2888	1.3	0.3368	1.3	0.99	1871.1	21.1	1867.1	11.2	1862.5	3.6	1862.5	3.6	100.5	-0.5
McAfee-97SF1-135	71	27589	2.7	8.7796	0.9	5.4060	1.3	0.3442	0.9	0.74	1907.0	15.6	1885.8	10.9	1862.5	15.5	1862.5	15.5	102.4	-2.4
McAfee-97SF1-198	70	28156	1.8	8.7636	0.8	5.3948	1.0	0.3429	0.7	0.65	1900.6	11.0	1884.0	8.8	1865.8	14.1	1865.8	14.1	101.9	-1.9
McAfee-97SF1-74	138	29020	1.4	8.7609	1.3	5.1648	1.5	0.3282	0.8	0.51	1829.5	12.4	1846.8	12.9	1866.4	23.5	1866.4	23.5	98.0	2.0
McAfee-97SF1-68	130	27610	7.9	8.7590	0.5	5.3800	0.9	0.3418	0.8	0.82	1895.2	12.4	1881.7	7.9	1866.8	9.6	1866.8	9.6	101.5	-1.5
McAfee-97SF1-173	78	94259	0.9	8.7368	0.8	5.3305	1.1	0.3378	0.7	0.66	1875.9	11.4	1873.8	9.1	1871.3	14.4	1871.3	14.4	100.2	-0.2
McAfee-97SF1-168	23	13340	0.6	8.7152	2.7	5.5348	2.9	0.3498	1.1	0.38	1933.9	18.7	1906.0	25.0	1875.8	48.5	1875.8	48.5	103.1	-3.1
McAfee-97SF1-200	58	6394	1.6	8.7032	3.5	5.1737	3.6	0.3266	0.9	0.25	1821.8	14.1	1848.3	30.6	1878.3	62.8	1878.3	62.8	97.0	3.0
McAfee-97SF1-84	288	215283	2.2	8.6776	0.2	5.5389	1.4	0.3486	1.4	0.99	1927.9	23.5	1906.7	12.2	1883.6	3.1	1883.6	3.1	102.3	-2.3
McAfee-97SF1-82	51	46336	0.5	8.6595	1.0	5.6460	1.8	0.3546	1.5	0.84	1956.5	25.1	1923.2	15.2	1887.4	17.1	1887.4	17.1	103.7	-3.7
McAfee-97SF1-137	81	56270	1.0	8.6204	0.6	5.5716	0.7	0.3483	0.4	0.56	1926.7	6.3	1911.7	5.8	1895.5	10.0	1895.5	10.0	101.6	-1.6
McAfee-97SF1-50	20	15094	0.8																	



## U-Pb geochronologic analyses of selected Roberts Mountains allochthon strata

	Isotope ratios										Apparent ages (Ma)									
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc.	Discord
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)
<b>Sample: McAfee Quartzite. Location: McAfee Creek, Independence Mountains; 050637 4599583 (NAD 83 UTM 11T)</b>																				
McAfee-97SF1-125	172	82627	0.5	8.4678	0.3	5.8646	0.5	0.3602	0.4	0.82	1983.0	7.5	1956.0	4.6	1927.6	5.4	1927.6	5.4	102.9	-2.9
McAfee-97SF1-187	148	81397	0.6	8.4575	0.4	5.8662	0.6	0.3488	0.5	0.76	1928.8	7.6	1929.3	5.2	1929.7	7.0	1929.7	7.0	100.0	0.0
McAfee-97SF1-163	35	26329	1.4	8.4530	1.4	5.8284	2.5	0.3573	2.1	0.84	1969.5	35.7	1950.6	21.8	1930.7	24.5	1930.7	24.5	102.0	-2.0
McAfee-97SF1-100	198	173972	1.1	8.4528	0.4	5.7481	0.9	0.3524	0.8	0.91	1946.0	14.0	1938.6	7.9	1930.8	6.7	1930.8	6.7	100.8	-0.8
McAfee-97SF1-106	214	69591	1.4	8.4463	0.4	5.5636	1.6	0.3408	1.6	0.97	1890.6	26.1	1910.5	14.1	1932.1	6.8	1932.1	6.8	97.8	2.2
McAfee-97SF1-20	32	92931	0.3	8.4431	1.9	5.7755	2.4	0.3537	1.5	0.62	1952.1	25.1	1942.7	20.9	1932.8	34.0	1932.8	34.0	101.0	-1.0
McAfee-97SF1-103	177	90205	1.1	8.4409	0.4	5.8803	1.8	0.3600	1.8	0.98	1982.1	30.2	1958.3	15.7	1933.3	6.3	1933.3	6.3	102.5	-2.5
McAfee-97SF1-172	121	73267	1.9	8.4322	0.4	5.8590	1.0	0.3583	0.9	0.93	1974.2	16.1	1955.2	8.9	1935.1	6.9	1935.1	6.9	102.0	-2.0
McAfee-97SF1-98	47	19878	0.6	8.4262	1.1	5.7621	1.8	0.3521	1.4	0.78	1944.8	23.8	1940.7	15.7	1936.4	20.3	1936.4	20.3	100.4	-0.4
McAfee-97SF1-116	234	162236	1.5	8.4261	0.2	5.7291	0.6	0.3501	0.5	0.94	1935.1	8.9	1935.8	4.9	1936.4	3.3	1936.4	3.3	99.9	0.1
McAfee-97SF1-169	268	102378	1.2	8.4152	0.3	5.8655	0.6	0.3580	0.5	0.86	1972.6	9.0	1956.2	5.3	1938.7	5.6	1938.7	5.6	101.7	-1.7
McAfee-97SF1-155	65	48475	1.1	8.4013	0.7	5.8356	0.9	0.3556	0.6	0.63	1961.1	9.6	1951.7	7.8	1941.7	12.5	1941.7	12.5	101.0	-1.0
McAfee-97SF1-160	207	139323	0.6	8.3877	0.3	5.9517	0.5	0.3621	0.4	0.83	1991.9	7.6	1968.8	4.6	1944.6	5.3	1944.6	5.3	102.4	-2.4
McAfee-97SF1-143	31	20466	0.8	8.3693	1.3	5.8189	1.7	0.3532	1.1	0.66	1949.9	19.3	1949.2	15.1	1948.5	23.5	1948.5	23.5	100.1	-0.1
McAfee-97SF1-150	27	21294	1.1	8.1785	2.4	6.2373	2.6	0.3700	1.1	0.42	2029.3	19.1	2009.7	23.1	1989.6	42.7	1989.6	42.7	102.0	-2.0
McAfee-97SF1-79	166	190358	2.4	8.1658	1.0	6.1503	1.6	0.3642	1.3	0.80	2002.3	22.2	1997.4	14.0	1992.4	17.0	1992.4	17.0	100.5	-0.5
McAfee-97SF1-193	29	12644	1.2	8.1512	1.4	6.2774	1.9	0.3711	1.3	0.68	2034.6	22.5	2015.3	16.6	1995.6	24.5	1995.6	24.5	102.0	-2.0
McAfee-97SF1-190	49	56619	1.0	8.1233	1.1	6.2513	1.7	0.3683	1.4	0.78	2021.4	23.6	2011.7	15.3	2001.7	19.5	2001.7	19.5	101.0	-1.0
McAfee-97SF1-188	37	29947	1.2	8.0755	0.9	6.4587	4.0	0.3783	3.9	0.97	2068.2	69.1	2040.3	35.2	2012.1	16.1	2012.1	16.1	102.8	-2.8
McAfee-97SF1-61	28	21027	0.7	7.8837	1.6	6.7797	3.3	0.3876	2.9	0.88	2111.9	52.2	2083.1	29.2	2054.7	27.9	2054.7	27.9	102.8	-2.8
McAfee-97SF1-124	220	148907	1.5	7.8661	0.4	6.7456	0.7	0.3848	0.6	0.83	2098.8	10.3	2078.6	6.1	2058.6	6.8	2058.6	6.8	102.0	-2.0
McAfee-97SF1-3	131	122598	0.7	7.8628	0.4	6.6220	0.6	0.3776	0.5	0.83	2065.2	9.4	2062.3	5.6	2059.3	6.2	2059.3	6.2	100.3	-0.3
McAfee-97SF1-136	32	21139	1.0	7.8429	1.2	6.7060	1.7	0.3815	1.2	0.73	2083.1	22.2	2073.4	15.2	2063.8	20.8	2063.8	20.8	100.9	-0.9
McAfee-97SF1-17	34	48125	1.0	7.8174	1.3	6.7366	1.8	0.3819	1.2	0.68	2085.4	21.7	2077.4	16.0	2069.6	23.5	2069.6	23.5	100.8	-0.8
McAfee-97SF1-45	74	82299	1.6	7.8143	0.5	6.7425	0.9	0.3821	0.8	0.83	2086.2	13.8	2078.2	8.3	2070.3	9.3	2070.3	9.3	100.8	-0.8
McAfee-97SF1-44	32	23564	1.3	7.7902	1.4	6.7853	2.5	0.3834	2.1	0.84	2092.0	37.2	2083.8	22.0	2075.2	24.0	2075.2	24.0	100.8	-0.8
McAfee-97SF1-153	102	52342	1.6	7.7808	0.3	6.7569	0.5	0.3813	0.4	0.75	2082.4	6.9	2080.1	4.6	2077.8	6.1	2077.8	6.1	100.2	-0.2
McAfee-97SF1-159	92	59209	2.0	7.7772	0.3	6.8379	0.8	0.3857	0.7	0.91	2102.8	13.4	2090.6	7.3	2078.6	6.1	2078.6	6.1	101.2	-1.2
McAfee-97SF1-27	99	56650	1.9	7.7715	0.5	6.9050	1.7	0.3892	1.6	0.96	2119.2	28.6	2099.3	14.7	2079.9	8.6	2079.9	8.6	101.9	-1.9
McAfee-97SF1-128	35	28440	0.6	7.7635	1.3	6.7911	2.1	0.3824	1.6	0.78	2087.4	28.7	2084.6	18.2	2081.7	22.5	2081.7	22.5	100.3	-0.3
McAfee-97SF1-66	44	18261	1.3	7.7627	1.4	6.7588	1.7	0.3805	1.0	0.60	2078.7	18.2	2080.3	15.1	2081.9	23.9	2081.9	23.9	99.8	0.2
McAfee-97SF1-51	116	117902	1.5	7.7619	0.3	6.6816	0.7	0.3761	0.6	0.86	2058.2	10.5	2070.2	6.1	2082.1	6.1	2082.1	6.1	99.9	1.1
McAfee-97SF1-131	43	19453	1.3	7.7593	0.8	6.8678	1.3	0.3865	1.0	0.76	2106.6	17.4	2094.5	11.3	2082.7	14.5	2082.7	14.5	101.1	-1.1
McAfee-97SF1-96	37	41735	0.8	7.7519	1.3	6.7094	1.5	0.3772	0.7	0.47	2063.3	12.5	2073.9	13.2	2084.4	23.2	2084.4	23.2	99.0	-1.0
McAfee-97SF1-92	77	54214	1.0	7.7501	0.4	6.9409	0.5	0.3901	0.4	0.68	2123.5	6.6	2103.9	4.8	2084.8	7.0	2084.8	7.0	101.9	-1.9
McAfee-97SF1-157	96	56520	1.1	7.7247	0.7	7.0197	1.0	0.3933	0.7	0.72	2138.0	12.5	2113.9	8.5	2090.6	11.7	2090.6	11.7	102.3	-2.3
McAfee-97SF1-144	191	227075	3.0	7.7211	0.4	6.9669	0.6	0.3913	0.4	0.71	2128.7	7.7	2109.8	5.3	2091.4	7.4	2091.4	7.4	101.8	-1.8
McAfee-97SF1-112	96	58312	1.5	7.7181	0.4	6.9698	1.0	0.3901	0.9	0.91	2123.5	15.7	2107.6	8.5	2092.1	6.8	2092.1	6.8	101.5	-1.5
McAfee-97SF1-161	122	43647	2.6	7.7062	0.7	6.9036	1.8	0.3858	1.7	0.93	2103.5	30.8	2099.1	16.4	2094.8	11.8	2094.8	11.8	100.4	-0.4
McAfee-97SF1-117	70	60328	0.7	7.7054	0.9	6.9660	1.0	0.3893	0.4	0.43	2119.6	7.7	2107.1	8.8	2095.0	15.6	2095.0	15.6	101.2	-1.2
McAfee-97SF1-102	158	182899	2.0	7.7030	0.3	6.9974	1.1	0.3909	1.1	0.97	2127.1	19.5	2111.1	9.9	2095.5	4.5	2095.5	4.5	101.5	-1.5
McAfee-97SF1-58	61	67169	1.2	7.7000	1.0	6.9792	1.2	0.3898	0.7	0.59	2121.7	13.0	2108.8	10.8	2096.2	17.1	2096.2	17.1	101.2	-1.2
McAfee-97SF1-77	14	14773	0.7	7.6688	2.7	6.8338	4.4	0.3801	3.4	0.79	2076.7	61.2	2090.1	38.7	2103.3	47.1	2103.3	47.1	98.7	-1.3
McAfee-97SF1-192	141	38567	1.5	7.6612	0.8	7.0802	1.2	0.3934	0.9	0.74	2138.6	16.8	2121.6	11.1	2105.1	14.6	2105.1	14.6	101.6	-1.6
McAfee-97SF1-146	68	40942	0.9	7.4085	2.3	7.8361	4.6	0.4210	4.1	0.87	2265.2	77.5	2212.4	41.9	2163.7	39.6	2163.7	39.6	104.7	-4.7
McAfee-97SF1-177	42	26578	1.1	7.1380	1.3	7.7105	2.2	0.3992	1.7	0.78	2165.2	31.1	2197.8	19.4	2228.4	23.2	2228.4	23.2	97.2	-2.8
McAfee-97SF1-184	130	122276	0.6	6.9412	0.3	6.8582	0.7	0.4329	0.6	0.91	2318.6	11.8	2296.4	6.1	2276.7	4.9	2276.7	4.9	101.8	-1.8
McAfee-97SF1-193	78	104977	1.4	6.7845	0.7	8.8264	0.9	0.4343	0.6	0.63	2325.1	11.4	2320.2	8.5	2315.9	12.4	2315.9	12.4	100.4	-0.4
McAfee-97SF1-14	199	163185	2.2	6.5797	0.3	8.8274	0.8	0.4213	0.7	0.90	2266.2	13.0	2320.3	6.9	2368.3	5.5	2368.3	5.5	95.7	4.3
McAfee-97SF1-120	93	98992	1.0	6.0245	1.0	10.3590	1.1	0.4526	0.6	0.55	2406.9	12.6	2467.4	10.6	2517.6	16.1	2517.6	16.1	95.6	4.4
McAfee-97SF1-178	19	12587	1.4	5.9880	2.2	11.0919	2.8	0.4815	1.7	0.60	2534.0	35.3	2530.9	26.1	2528.4	37.5	2528.4	37.5	100.2	-2.2
McAfee-97SF1-28	164	187219	0.6	5.9014	0.4	11.3514	1.7	0.4859	1.7	0.98	2552.7	35.6	2552.4	16.1	2552.2	6.0	2552.2	6.0	100.0	0.0
McAfee-97SF1-6	33	40639	0.6	5.8551	0.8	11.3933	1.1	0.4838	0.8	0.71	2543.9	17.0	2555.9	10.7	2565.4	13.4	2565.4	13.4	99.2	0.8
McAfee-97SF1-33	148	110900	1.8	5.8451	0.2	11.8935	2.4	0.5042	2.4	1.00	2631.8	51.8	2596.1	22.6	2568.3	4.0	2568.3	4.0	102.5	-2.5
McAfee-97SF1-78	340	59900	1.9	5.8064	0.1	11.6908	1.0	0.4923	1.0	0.99	2580.7	20.6	2580.0	9.1	2579.3	1.7	2579.3	1.7	100.1	-0.1
McAfee-97SF1-80	33	38888	3.0	5.5220	0.5	12.9720	1.8	0.5195	1.7	0.95	2697.2	37.0	2677.6	16.6	2662.9	8.7	2662.9	8.7	101.3	-1.3
McAfee-97SF1-1																				



## U-Pb geochronologic analyses of selected Roberts Mountains allochthon strata

	Isotope ratios										Apparent ages (Ma)										
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc.	Discord.	
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)	
<b>Sample: lower Vinini Formation. Location: Petes Summit, Toquima Range: 0518089 4337111 (NAD 83 UTM 11T)</b>																					
LOW VININI 97SF12-190	359	140836	1.0	17.4780	1.1	0.6299	2.4	0.0799	2.2	0.89	495.2	10.3	496.0	9.6	499.8	24.8	495.2	10.3	99.1	0.9	
LOW VININI 97SF12-50	39	9890	0.5	18.4533	7.2	0.5967	8.2	0.0799	3.8	0.47	495.3	18.3	475.2	31.0	379.0	162.5	495.3	18.3	130.7	-30.7	
LOW VININI 97SF12-142	38	16456	2.0	17.6961	10.5	0.6223	10.6	0.0799	1.7	0.16	495.3	8.2	491.3	41.3	472.4	231.8	495.3	8.2	104.8	-4.8	
LOW VININI 97SF12-10	98	21025	0.5	17.4708	3.6	0.6309	3.9	0.0799	1.6	0.40	495.7	7.5	496.6	15.4	500.7	79.4	495.7	7.5	99.0	1.0	
LOW VININI 97SF12-130	214	73022	0.3	17.3277	1.2	0.6371	1.8	0.0801	1.4	0.77	496.5	6.7	500.5	7.2	518.8	25.6	496.5	6.7	95.7	4.3	
LOW VININI 97SF12-179	565	144200	0.7	17.4762	0.7	0.6319	1.3	0.0801	1.1	0.85	496.6	5.4	497.3	5.2	500.1	15.3	496.6	5.4	99.3	0.7	
LOW VININI 97SF12-152	211	114161	0.7	17.3700	1.3	0.6365	3.3	0.0802	3.0	0.92	497.3	14.6	500.2	13.1	513.4	28.2	497.3	14.6	96.8	3.2	
LOW VININI 97SF12-35	207	191128	0.7	17.4577	2.1	0.6334	2.4	0.0802	1.1	0.48	497.3	5.5	498.2	9.3	502.4	45.7	497.3	5.5	99.0	1.0	
LOW VININI 97SF12-106	319	110139	1.1	17.4367	1.4	0.6348	2.5	0.0803	2.1	0.83	497.8	9.8	499.1	9.7	505.0	29.8	497.8	9.8	98.6	1.4	
LOW VININI 97SF12-166	52	14553	0.8	17.4661	3.7	0.6342	4.8	0.0803	3.0	0.62	498.1	14.2	498.7	18.7	501.3	81.9	498.7	14.2	99.4	0.6	
LOW VININI 97SF12-33	65	35323	0.8	16.8624	6.0	0.6571	6.1	0.0804	0.8	0.14	498.3	4.0	512.8	24.4	578.3	130.4	498.3	4.0	86.2	13.8	
LOW VININI 97SF12-191	125	58994	0.7	17.2369	1.5	0.6438	2.9	0.0805	2.4	0.85	499.0	11.6	504.6	11.4	530.3	33.4	499.0	11.6	94.1	5.9	
LOW VININI 97SF12-25	125	66771	1.6	17.6443	2.2	0.6292	2.5	0.0805	1.2	0.47	499.2	5.6	495.6	9.7	478.9	48.5	499.2	5.6	104.2	-4.2	
LOW VININI 97SF12-36	153	8697	1.3	17.2246	4.2	0.6468	4.9	0.0808	2.6	0.53	500.9	12.4	506.5	19.5	531.9	91.3	500.9	12.4	94.2	5.8	
LOW VININI 97SF12-93	60	28077	0.8	17.2241	5.9	0.6475	7.2	0.0809	4.0	0.56	501.4	19.5	507.0	28.7	532.0	130.1	501.4	19.5	94.3	5.7	
LOW VININI 97SF12-117	173	51629	1.1	17.4399	2.4	0.6396	2.9	0.0809	1.6	0.56	501.5	7.9	502.0	11.6	504.6	53.8	501.5	7.9	99.4	0.6	
LOW VININI 97SF12-7	594	134712	0.5	17.5449	0.6	0.6375	1.6	0.0811	1.5	0.92	502.8	7.1	500.7	6.3	491.4	13.8	502.8	7.1	102.3	-2.3	
LOW VININI 97SF12-157	20	10328	1.6	13.1682	6.4	1.8758	7.3	0.1792	3.5	0.47	1062.3	33.8	1072.6	48.5	1093.5	129.3	1093.5	33.8	97.1	2.9	
LOW VININI 97SF12-129	42	39781	1.2	13.1076	3.4	2.0169	4.7	0.1917	3.2	0.69	1130.8	33.5	1121.2	31.7	1102.8	67.3	1102.8	33.5	102.5	-2.5	
LOW VININI 97SF12-45	122	97560	1.2	13.0154	1.0	2.0010	1.7	0.1889	1.4	0.81	1115.4	14.2	1115.9	11.6	1116.9	19.9	1116.9	14.2	99.9	0.1	
LOW VININI 97SF12-187	60	48933	0.9	13.0102	2.4	1.9910	4.5	0.1879	3.7	0.84	1109.8	38.0	1112.5	30.2	1117.7	48.8	1117.7	38.0	98.8	9.3	0.7
LOW VININI 97SF12-126	73	44396	1.2	13.0060	2.2	2.0029	2.8	0.1889	1.8	0.63	1115.6	18.2	1116.5	19.2	1118.3	44.1	1118.3	18.2	99.8	0.2	
LOW VININI 97SF12-177	43	28465	1.5	12.9205	3.7	2.0120	4.0	0.1885	1.4	0.35	1113.5	14.0	1119.6	26.9	1131.4	74.2	1131.4	14.0	98.4	1.6	
LOW VININI 97SF12-189	110	101253	1.6	12.1600	1.8	2.3227	2.0	0.2048	0.9	0.45	1201.3	9.9	1219.3	14.1	1251.2	34.7	1251.2	9.9	96.0	4.0	
LOW VININI 97SF12-193	128	85570	1.6	12.0755	0.6	2.4356	2.0	0.2133	2.0	0.96	1246.4	22.2	1253.2	14.7	1264.8	10.9	1264.8	10.9	98.5	1.5	
LOW VININI 97SF12-192	97	63131	2.1	11.9392	1.2	2.5747	1.7	0.2229	1.2	0.69	1297.4	13.7	1293.5	12.3	1286.9	23.6	1286.9	13.7	100.8	-0.8	
LOW VININI 97SF12-111	64	36049	1.4	11.4601	1.6	2.8730	3.6	0.2388	3.3	0.90	1380.4	40.8	1374.9	27.5	1366.2	30.8	1366.2	40.8	101.0	-1.0	
LOW VININI 97SF12-61	75	50496	0.9	11.1687	1.9	3.1573	2.3	0.2558	1.4	0.58	1468.1	17.8	1446.8	18.0	1415.6	36.2	1415.6	17.8	103.7	-3.7	
LOW VININI 97SF12-2	158	111192	1.2	11.1327	0.6	3.0652	0.9	0.2475	0.7	0.75	1425.5	6.4	1424.0	6.8	1421.8	11.3	1421.8	11.3	100.3	-0.3	
LOW VININI 97SF12-63	66	39255	1.5	11.1125	0.9	3.1780	4.7	0.2561	4.6	0.98	1470.0	80.5	1451.8	36.1	1425.3	16.3	1425.3	80.5	103.1	-3.1	
LOW VININI 97SF12-55	134	199794	1.8	11.0726	0.8	3.1441	1.6	0.2525	1.4	0.86	1451.3	18.2	1443.6	12.5	1432.1	15.8	1432.1	18.2	101.3	-1.3	
LOW VININI 97SF12-101	61	56680	0.8	11.0656	1.6	3.0774	2.7	0.2470	2.2	0.81	1422.9	27.7	1427.1	20.5	1433.4	29.7	1433.4	27.7	99.3	0.7	
LOW VININI 97SF12-121	99	108276	1.4	11.0628	0.7	3.1419	2.1	0.2521	2.0	0.95	1449.3	26.2	1443.0	16.4	1433.8	12.5	1433.8	12.5	101.1	-1.1	
LOW VININI 97SF12-69	143	11317	1.2	11.0543	0.7	2.9727	6.3	0.2383	6.3	0.99	1378.0	77.7	1400.7	47.9	1435.3	12.5	1435.3	77.7	96.0	4.0	
LOW VININI 97SF12-27	149	173305	1.2	11.0536	0.6	3.1318	1.1	0.2511	1.0	0.87	1444.0	13.0	1440.5	8.8	1435.4	10.7	1435.4	13.0	100.6	-0.6	
LOW VININI 97SF12-48	252	86501	0.8	11.0503	0.4	3.1161	2.3	0.2497	2.2	0.99	1437.1	28.9	1436.7	17.5	1436.0	7.5	1436.0	28.9	100.1	-1.1	
LOW VININI 97SF12-168	83	62578	0.9	11.0488	0.9	3.0775	1.6	0.2466	1.3	0.82	1421.0	16.4	1427.1	12.0	1436.2	17.2	1436.2	16.4	98.9	-1.1	
LOW VININI 97SF12-32	107	97581	1.2	11.0430	0.5	3.1211	1.1	0.2500	1.0	0.89	1438.3	12.5	1437.9	8.4	1437.2	9.3	1437.2	12.5	100.1	-0.1	
LOW VININI 97SF12-183	267	183766	1.6	11.0402	0.3	3.1989	1.7	0.2561	1.6	0.98	1470.1	21.5	1456.9	12.9	1437.7	6.6	1437.7	21.5	102.3	-2.3	
LOW VININI 97SF12-R81	47	57046	1.2	11.0387	2.1	3.1189	3.7	0.2497	3.1	0.83	1436.9	39.5	1437.4	28.4	1438.0	39.1	1438.0	39.5	99.9	0.1	
LOW VININI 97SF12-134	109	117842	1.2	11.0362	0.5	3.1392	1.4	0.2513	1.3	0.92	1445.0	16.3	1442.4	10.6	1438.4	10.4	1438.4	16.3	100.5	-0.5	
LOW VININI 97SF12-17	89	98704	1.0	11.0322	0.7	3.1131	3.0	0.2491	2.9	0.97	1433.8	37.1	1435.9	22.9	1439.1	14.0	1439.1	37.1	99.6	0.4	
LOW VININI 97SF12-12	163	206911	2.0	11.0321	0.3	3.1687	1.1	0.2535	1.1	0.95	1456.7	13.9	1449.6	8.7	1439.1	6.7	1439.1	13.9	101.2	-1.2	
LOW VININI 97SF12-173	131	113400	1.5	11.0269	0.8	3.1387	2.0	0.2510	1.8	0.91	1443.7	23.0	1442.2	15.1	1440.0	15.7	1440.0	23.0	100.3	-0.3	
LOW VININI 97SF12-180	144	114950	0.8	11.0037	0.7	3.1353	1.8	0.2502	1.7	0.92	1439.6	21.3	1441.4	13.8	1444.0	13.3	1444.0	21.3	99.7	0.3	
LOW VININI 97SF12-87	130	106569	1.3	11.0013	0.5	3.1620	1.5	0.2523	1.4	0.94	1450.3	18.2	1447.9	11.5	1444.5	9.5	1444.5	18.2	100.4	-0.4	
LOW VININI 97SF12-150	62	108335	1.5	11.0000	1.3	3.1422	2.0	0.2507	1.5	0.74	1442.0	18.9	1443.1	15.1	1447.7	24.9	1443.1	18.9	99.8	0.2	
LOW VININI 97SF12-90	42	18328	0.7	10.9930	2.8	2.8573	5.3	0.2278	4.5	0.85	1323.0	53.5	1370.7	39.8	1445.9	53.6	1445.9	53.5	91.5	8.5	
LOW VININI 97SF12-56	62	48211	0.9	10.9241	1.1	3.2460	1.9	0.2572	1.6	0.83	1475.4	20.6	1468.2	14.7	1457.9	20.2	1457.9	20.6	101.2	-1.2	
LOW VININI 97SF12-89	72	29370	0.5	10.9006	1.1	3.1895	2.0	0.2522	1.7	0.85	1449.6	22.3	1454.6	15.6	1462.0	20.0	1462.0	22.3	99.2	0.8	
LOW VININI 97SF12-181	81	94981	0.7	10.7807	1.2	3.2842	2.0	0.2568	1.6	0.81	1473.4	21.4	1477.3	15.6	1482.9	22.2	1482.9	21.4	99.4	0.6	
LOW VININI 97SF12-119	472	441019	1.8	10.6830	0.2	3.3801	1.5	0.2619	1.5	0.99	1499.5	19.9	1499.8	11.8	1500.2	4.5	1500.2	19.9	100.0	0.0	
LOW VININI 97SF12-103	117	25158	1.4	10.6723	0.8	3.4708	1.2	0.2636	0.8	0.70	1508.3	11.0	1520.6	9.2	1537.7	15.8	1537.7	11.0	98.1	1.9	
LOW VININI 97SF12-5	42	38927	1.0	10.6734	1.4	4.0212	2.4	0.2938	1.9	0.80	1660.4	28.3	1638.5	19.6	1610.4	26.8	1610.4	28.3	103.1	-3.1	
LOW VININI 97SF12-135	131	109421	1.4	10.6473	0.6	3.9296	1.8	0.2863	1.7	0.94	1623.3	23.8	1619.8	14.2	1615.3	11.1	1615.3	23.8	100.5	-0.5	
LOW VININI 97SF12-116	294	330298	7.9	9.9426	0.2	4.0194	0.6	0.2898	0.6	0.96	1640.7	8.1	1638.1	4.8	1634.8						

## U-Pb geochronological analyses of selected Roberts Mountains allochthon strata

		Isotope ratios										Apparent ages (Ma)									
		U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc.	Discord.
		(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)
<b>Sample: lower Vinini Formation. Location: Petes Summit, Toquima Range; 0518089 4337111 (NAD 83 UTM 11T)</b>																					
LOW VININI 97SF12-44	112	81278	2.5	9.5312	0.5	4.4437	1.8	0.3072	1.7	0.96	1728.8	25.6	1720.5	14.6	1712.9	9.1	1712.9	9.1	100.8	-0.8	
LOW VININI 97SF12-43	236	261283	1.6	9.5301	0.4	4.2598	3.2	0.2944	3.2	0.99	1663.7	47.0	1685.6	26.6	1713.1	7.0	1713.1	7.0	97.1	2.9	
LOW VININI 97SF12-154	220	316270	1.5	9.5289	0.3	4.4350	2.0	0.3065	1.9	0.98	1723.5	29.4	1718.9	16.4	1713.3	6.3	1713.3	6.3	100.6	-0.6	
LOW VININI 97SF12-77	187	223570	1.9	9.5282	0.3	4.3675	2.4	0.3018	2.4	0.99	1700.3	36.1	1706.2	20.1	1713.5	4.7	1713.5	4.7	99.2	0.8	
LOW VININI 97SF12-113	84	83469	1.3	9.5258	0.5	4.3577	4.1	0.3101	4.0	0.99	1696.6	60.2	1704.4	33.7	1713.9	10.0	1713.9	10.0	99.0	1.0	
LOW VININI 97SF12-78	153	466926	1.9	9.5191	0.6	4.4938	3.0	0.3012	3.0	0.98	1741.9	45.3	1729.8	25.1	1715.2	10.2	1715.2	10.2	101.6	-1.6	
LOW VININI 97SF12-109	155	88596	3.1	9.5185	0.5	4.5485	1.8	0.3140	1.7	0.96	1760.4	26.7	1739.9	15.1	1715.3	9.7	1715.3	9.7	102.6	-2.6	
LOW VININI 97SF12-23	126	226755	2.0	9.5150	0.7	4.4484	1.9	0.3070	1.8	0.92	1725.8	26.5	1721.4	15.8	1716.0	13.6	1716.0	13.6	100.6	-0.6	
LOW VININI 97SF12-107	138	155516	2.2	9.5096	0.5	4.3644	2.0	0.3100	2.0	0.97	1696.3	29.3	1705.6	16.6	1717.0	8.5	1717.0	8.5	98.8	1.2	
LOW VININI 97SF12-159	104	98128	1.3	9.5094	0.8	4.5069	2.9	0.3018	2.8	0.96	1744.8	42.2	1732.3	23.9	1717.1	15.3	1717.1	15.3	101.6	-1.6	
LOW VININI 97SF12-8	131	212812	0.9	9.5075	0.4	4.4582	2.2	0.3074	2.1	0.98	1728.0	32.3	1723.2	18.0	1717.4	7.6	1717.4	7.6	100.6	-0.6	
LOW VININI 97SF12-75	265	486407	2.8	9.5009	0.2	4.4980	1.1	0.3099	1.1	0.99	1740.4	16.2	1730.6	8.9	1718.7	3.3	1718.7	3.3	101.3	-1.3	
LOW VININI 97SF12-R82	153	399910	2.4	9.4953	0.3	4.5315	2.3	0.3121	2.3	0.99	1750.9	35.1	1736.8	19.2	1719.8	5.5	1719.8	5.5	101.8	-1.8	
LOW VININI 97SF12-18	121	128301	1.9	9.4950	0.4	4.4698	1.8	0.3078	1.7	0.98	1729.9	26.3	1725.4	14.7	1719.9	6.6	1719.9	6.6	100.6	-0.6	
LOW VININI 97SF12-175	155	180329	2.0	9.4932	0.4	4.4681	1.7	0.3076	1.7	0.98	1729.1	25.1	1725.1	14.1	1720.2	6.9	1720.2	6.9	100.5	-0.5	
LOW VININI 97SF12-102	123	105962	2.4	9.4903	0.4	4.4650	4.1	0.3073	4.1	1.00	1727.5	62.6	1724.5	34.4	1720.8	6.9	1720.8	6.9	100.4	-0.4	
LOW VININI 97SF12-125	200	245331	2.7	9.4902	0.3	4.4033	2.9	0.3031	2.9	1.00	1706.6	43.0	1713.0	23.8	1720.8	5.0	1720.8	5.0	99.2	0.8	
LOW VININI 97SF12-16	97	118590	1.8	9.4880	0.7	4.4611	1.4	0.3070	1.2	0.87	1725.8	18.1	1723.8	11.4	1721.2	12.4	1721.2	12.4	100.3	-0.3	
LOW VININI 97SF12-128	226	294688	1.6	9.4835	0.4	4.4540	2.3	0.3064	2.3	0.99	1722.7	34.8	1722.5	19.3	1722.1	6.5	1722.1	6.5	100.0	0.0	
LOW VININI 97SF12-149	224	158788	1.2	9.4781	0.4	4.5286	1.8	0.3113	1.8	0.98	1747.1	27.4	1736.2	15.2	1723.7	7.0	1723.7	7.0	101.4	-1.4	
LOW VININI 97SF12-58	147	178452	0.8	9.4771	0.6	4.4994	1.6	0.3093	1.5	0.93	1737.1	23.3	1730.9	13.7	1723.3	11.0	1723.3	11.0	100.8	-0.8	
LOW VININI 97SF12-138	115	192707	1.0	9.4738	0.7	4.4816	2.6	0.3079	2.5	0.96	1730.5	37.4	1727.6	21.4	1724.0	13.8	1724.0	13.8	100.4	-0.4	
LOW VININI 97SF12-194	174	369932	2.9	9.4714	0.5	4.4958	2.5	0.3088	2.5	0.98	1734.9	37.6	1730.2	20.9	1724.4	8.8	1724.4	8.8	100.6	-0.6	
LOW VININI 97SF12-198	70	68060	1.6	9.4710	0.9	4.5447	2.7	0.3122	2.6	0.94	1751.4	39.6	1739.2	22.8	1724.5	16.6	1724.5	16.6	101.6	-1.6	
LOW VININI 97SF12-66	287	207517	1.1	9.4674	0.2	4.4787	1.2	0.3075	1.2	0.98	1728.5	18.3	1727.0	10.2	1725.2	4.2	1725.2	4.2	100.2	-0.2	
LOW VININI 97SF12-163	188	134321	1.2	9.4661	0.3	4.4049	2.6	0.3024	2.6	0.99	1703.3	38.7	1713.3	21.6	1725.5	5.5	1725.5	5.5	98.7	1.3	
LOW VININI 97SF12-91	84	71940	1.0	9.4616	1.0	4.4819	3.3	0.3076	3.1	0.95	1728.7	47.1	1727.6	27.1	1726.3	18.1	1726.3	18.1	100.1	-0.1	
LOW VININI 97SF12-161	212	84543	1.1	9.4575	0.3	4.2385	1.5	0.2907	1.4	0.98	1645.2	20.8	1681.5	12.0	1727.1	4.8	1727.1	4.8	95.3	4.7	
LOW VININI 97SF12-13	93	154926	1.0	9.4549	0.3	4.4849	1.5	0.3075	1.5	0.98	1728.6	22.9	1728.2	12.8	1727.6	5.8	1727.6	5.8	100.1	-0.1	
LOW VININI 97SF12-141	149	145193	1.4	9.4387	0.6	4.5178	2.0	0.3093	1.9	0.95	1737.1	28.6	1734.2	16.4	1730.8	10.9	1730.8	10.9	100.4	-0.4	
LOW VININI 97SF12-100	86	117345	1.7	9.4377	0.7	4.5437	2.5	0.3110	2.5	0.97	1745.7	37.5	1739.0	21.1	1731.0	12.0	1731.0	12.0	100.8	-0.8	
LOW VININI 97SF12-R85	92	52081	1.2	9.4334	0.6	4.5009	1.7	0.3079	1.6	0.93	1730.6	24.3	1731.1	14.3	1731.8	11.4	1731.8	11.4	99.9	0.1	
LOW VININI 97SF12-3	181	304251	1.3	9.4292	0.5	4.5644	1.6	0.3121	1.5	0.95	1751.3	23.4	1742.8	13.3	1732.6	8.8	1732.6	8.8	101.1	-1.1	
LOW VININI 97SF12-80	78	76788	1.4	9.4284	1.0	4.5995	2.2	0.3145	2.0	0.90	1762.9	31.0	1749.2	18.6	1732.8	17.9	1732.8	17.9	101.7	-1.7	
LOW VININI 97SF12-9	74	71185	2.3	9.4278	0.9	4.5588	1.7	0.3117	1.5	0.85	1749.1	22.4	1741.8	14.3	1732.9	16.6	1732.9	16.6	100.9	-0.9	
LOW VININI 97SF12-164	64	56466	3.3	9.4108	1.1	4.4894	2.1	0.3064	1.8	0.86	1723.0	27.8	1729.0	17.7	1736.2	19.9	1736.2	19.9	99.2	0.8	
LOW VININI 97SF12-184	76	55717	2.0	9.4098	0.7	4.5488	2.4	0.3104	2.3	0.95	1742.8	34.7	1739.9	19.9	1736.4	13.5	1736.4	13.5	100.4	-0.4	
LOW VININI 97SF12-110	61	8771	1.3	9.3783	1.9	4.4996	5.5	0.3060	5.1	0.94	1721.2	77.6	1730.9	45.4	1742.6	33.9	1742.6	33.9	98.8	1.2	
LOW VININI 97SF12-37	221	184461	2.6	9.3718	0.3	4.6291	1.3	0.3146	1.3	0.98	1763.5	20.2	1754.5	11.2	1743.8	5.1	1743.8	5.1	101.1	-1.1	
LOW VININI 97SF12-67	99	73291	2.5	9.3703	0.5	4.5933	2.8	0.3122	2.7	0.98	1751.3	41.6	1748.0	23.1	1744.1	9.8	1744.1	9.8	100.4	-0.4	
LOW VININI 97SF12-26	163	114168	1.5	9.3689	0.3	4.4190	1.8	0.3003	1.8	0.99	1692.6	26.1	1715.9	14.7	1744.4	5.1	1744.4	5.1	97.0	3.0	
LOW VININI 97SF12-73	97	47754	3.2	9.3328	0.6	4.6284	2.6	0.3133	2.5	0.98	1756.9	38.5	1754.4	21.3	1751.5	10.4	1751.5	10.4	100.3	-0.3	
LOW VININI 97SF12-86	155	325229	0.8	9.3239	0.3	4.6496	1.2	0.3144	1.1	0.96	1762.4	17.1	1758.2	9.7	1753.2	6.1	1753.2	6.1	100.5	-0.5	
LOW VININI 97SF12-97	72	57679	2.1	9.3156	0.9	4.5912	2.5	0.3102	2.3	0.92	1741.7	34.8	1747.7	20.5	1754.8	17.2	1754.8	17.2	99.2	0.8	
LOW VININI 97SF12-34	156	229809	1.5	9.2652	0.6	4.8500	2.1	0.3259	2.1	0.97	1818.5	32.6	1793.6	18.0	1764.8	10.1	1764.8	10.1	103.0	-3.0	
LOW VININI 97SF12-74	103	80501	2.3	9.2586	0.7	4.7989	1.3	0.3222	1.1	0.84	1800.7	16.7	1784.7	10.6	1766.1	12.4	1766.1	12.4	102.0	-2.0	
LOW VININI 97SF12-174	73	81806	2.5	9.2494	1.0	4.6910	1.8	0.3147	1.5	0.84	1763.7	23.5	1765.6	15.2	1767.9	18.0	1767.9	18.0	99.8	0.2	
LOW VININI 97SF12-143	132	143079	1.4	9.2473	0.4	4.7128	1.4	0.3161	1.3	0.97	1770.5	20.5	1769.5	11.5	1768.3	6.4	1768.3	6.4	100.1	-0.1	
LOW VININI 97SF12-98	277	64839	2.6	9.2271	0.2	4.1950	1.2	0.2807	1.2	0.99	1595.0	16.6	1673.0	9.7	1772.3	2.8	1772.3	2.8	90.0	10.0	
LOW VININI 97SF12-39	217	291915	2.7	9.2168	0.4	4.7067	1.4	0.3146	1.4	0.96	1763.4	20.9	1768.4	11.7	1774.3	6.8	1774.3	6.8	99.4	0.6	
LOW VININI 97SF12-65	194	243517	2.9	9.2115																	



## U-Pb geochronologic analyses of selected Roberts Mountains allochthon strata

		Isotope ratios								Apparent ages (Ma)								Best age	Conc.	Discord.	
		U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±				
		(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)
<b>Sample: lower Vinini Formation. Location: Petes Summit, Toquima Range: 0518089 4337111 (NAD 83 UTM 11T)</b>																					
LOW VININI 97SF12-11	93	158753	2.2	5.2788	0.5	14.2477	2.5	0.5455	2.4	0.98	2806.3	55.2	2766.3	23.5	2737.3	7.9	2737.3	7.9	102.5	-2.5	
LOW VININI 97SF12-64	33	72078	0.5	5.2747	0.4	13.8462	2.7	0.5297	2.6	0.99	2740.2	58.7	2739.2	25.2	2738.6	6.1	2738.6	6.1	100.1	-0.1	
LOW VININI 97SF12-29	67	218455	1.2	5.1698	1.6	13.8605	4.9	0.5197	4.6	0.94	2897.9	102.3	2740.2	46.6	2771.6	26.6	2771.6	26.6	97.3	2.7	
LOW VININI 97SF12-182	49	97648	0.6	4.8298	0.4	16.1139	1.6	0.5645	1.6	0.97	2885.0	36.9	2883.6	15.7	2882.6	6.9	2882.6	6.9	100.1	-0.1	
LOW VININI 97SF12-40	139	189158	1.7	4.6013	0.2	17.4251	3.3	0.5815	3.3	1.00	2954.9	77.2	2958.5	31.4	2961.0	3.6	2961.0	3.6	99.8	0.2	
LOW VININI 97SF12-99	94	250515	0.9	3.8593	0.2	23.2126	1.4	0.6497	1.4	0.99	3227.1	35.9	3235.9	14.0	3241.3	3.8	3241.3	3.8	99.6	0.4	
<b>Sample: upper Vinini Formation. Location: Petes Summit, Toquima Range: 0518089 4337111 (NAD 83 UTM 11T)</b>																					
VININI-97SF-11-52	153	45733	1.5	13.2079	1.2	1.9055	2.1	0.1825	1.7	0.81	1080.8	16.8	1083.0	13.8	1087.5	24.0	1087.5	24.0	99.4	0.6	
VININI-97SF-11-191	164	95852	3.0	12.1570	0.8	2.4261	1.6	0.2139	1.4	0.87	1249.6	15.4	1250.4	11.2	1251.7	15.2	1251.7	15.2	99.8	0.2	
VININI-97SF-11-137	233	61061	1.5	10.7459	0.5	3.3328	1.2	0.2597	1.1	0.90	1488.6	14.7	1488.8	9.6	1489.1	10.1	1489.1	10.1	100.0	0.0	
VININI-97SF-11-40	95	35185	1.8	9.4755	1.0	4.5559	2.3	0.3131	2.1	0.91	1755.9	32.6	1741.2	19.4	1723.6	17.6	1723.6	17.6	101.9	-1.9	
VININI-97SF-11-74	196	89887	4.4	9.1612	0.3	4.7560	1.2	0.3160	1.1	0.96	1770.2	17.2	1777.2	9.7	1785.4	5.9	1785.4	5.9	99.2	0.8	
VININI-97SF-11-138	208	81623	2.8	9.1314	0.3	4.8870	1.1	0.3237	1.1	0.97	1807.5	17.7	1800.9	9.7	1791.3	4.7	1791.3	4.7	100.9	-0.9	
VININI-97SF-11-36	18	7005	1.8	9.1163	3.4	4.9391	10.6	0.3266	10.0	0.95	1821.7	158.6	1809.0	89.3	1794.3	61.8	1794.3	61.8	101.5	-1.5	
VININI-97SF-11-29	68	25326	1.1	9.0951	0.6	4.7530	3.1	0.3135	3.1	0.98	1758.1	47.4	1776.6	26.3	1798.5	10.7	1798.5	10.7	97.7	2.3	
VININI-97SF-11-93	153	109995	4.3	9.0487	0.2	4.9292	1.5	0.3235	1.4	0.99	1806.8	22.8	1807.3	12.4	1807.8	4.5	1807.8	4.5	99.9	0.1	
VININI-97SF-11-192	42	14707	1.3	9.0413	0.7	5.1360	1.4	0.3368	1.2	0.86	1871.2	19.6	1842.1	11.9	1809.3	13.1	1809.3	13.1	103.4	-3.4	
VININI-97SF-11-76	15	18239	0.9	9.0405	3.2	5.1142	3.5	0.3353	1.3	0.37	1864.2	20.7	1838.5	29.5	1809.5	58.8	1809.5	58.8	103.0	-3.0	
VININI-97SF-11-25	118	29729	7.5	9.0201	0.4	4.8775	1.1	0.3191	1.0	0.90	1785.3	14.9	1798.4	8.9	1813.6	8.2	1813.6	8.2	98.4	1.6	
VININI-97SF-11-54	25	23149	0.9	9.0059	2.0	5.1032	2.8	0.3333	2.0	0.72	1854.5	32.8	1836.6	24.0	1816.5	35.5	1816.5	35.5	102.1	-2.1	
VININI-97SF-11-56	106	40476	0.8	8.9972	0.6	5.1093	1.6	0.3364	1.5	0.92	1854.8	23.7	1837.7	13.5	1818.2	11.2	1818.2	11.2	102.0	-2.0	
VININI-97SF-11-42	39	16785	0.9	8.9801	1.7	5.0146	1.9	0.3266	1.0	0.51	1821.9	15.7	1821.8	16.5	1821.7	30.5	1821.7	30.5	100.0	0.0	
VININI-97SF-11-78	115	88956	1.0	8.9787	0.6	4.7961	0.9	0.3123	0.6	0.69	1752.1	9.4	1784.2	7.5	1821.9	11.7	1821.9	11.7	96.2	3.8	
VININI-97SF-11-187	182	67811	1.5	8.9739	0.5	5.2032	2.5	0.3366	2.5	0.98	1880.2	40.2	1853.1	21.4	1822.9	9.5	1822.9	9.5	103.1	-3.1	
VININI-97SF-11-100	180	57040	0.9	8.9688	0.3	5.0708	1.1	0.3298	1.1	0.97	1837.6	17.6	1831.2	9.6	1824.0	5.2	1824.0	5.2	100.7	-0.7	
VININI-97SF-11-157	271	123010	1.0	8.9588	0.2	5.0341	0.8	0.3271	0.8	0.97	1824.3	12.9	1825.1	7.1	1826.0	3.6	1826.0	3.6	99.9	0.1	
VININI-97SF-11-67	145	140883	2.1	8.9526	0.5	5.0379	1.6	0.3271	1.5	0.95	1824.4	24.1	1825.7	13.6	1827.2	9.4	1827.2	9.4	99.8	0.2	
VININI-97SF-11-164	59	26870	0.7	8.9480	0.9	5.1394	1.2	0.3335	0.9	0.71	1855.5	14.1	1842.6	10.4	1828.2	15.6	1828.2	15.6	101.5	-1.5	
VININI-97SF-11-27	59	25440	1.3	8.9476	0.8	4.9817	2.5	0.3233	2.4	0.95	1805.7	37.4	1816.2	21.1	1828.2	14.1	1828.2	14.1	98.8	1.2	
VININI-97SF-11-17	81	43137	2.4	8.9467	0.6	5.1746	1.4	0.3356	1.3	0.92	1866.3	20.5	1848.4	11.7	1828.0	10.0	1828.0	10.0	102.1	-2.1	
VININI-97SF-11-4	140	20590	2.3	8.9343	0.3	4.7053	2.5	0.3049	2.4	0.99	1715.5	36.9	1768.2	20.7	1830.9	5.8	1830.9	5.8	93.7	6.3	
VININI-97SF-11-85	80	53383	0.5	8.9333	0.7	5.1722	1.3	0.3351	1.1	0.85	1863.1	18.3	1848.1	11.4	1831.1	12.9	1831.1	12.9	101.7	-1.7	
VININI-97SF-11-190	114	125413	0.9	8.9229	0.6	5.2880	1.2	0.3422	1.1	0.86	1897.3	17.6	1866.9	10.6	1833.3	11.4	1833.3	11.4	103.5	-3.5	
VININI-97SF-11-112	132	59165	2.5	8.9211	0.3	5.2343	0.9	0.3387	0.9	0.95	1880.3	14.2	1858.2	7.8	1833.6	5.2	1833.6	5.2	102.5	-2.5	
VININI-97SF-11-180	80	115104	1.4	8.9208	0.6	5.1822	1.0	0.3353	0.8	0.81	1864.0	12.7	1849.7	8.2	1833.7	10.3	1833.7	10.3	101.7	-1.7	
VININI-97SF-11-182	229	129017	1.7	8.9198	0.2	5.1606	1.6	0.3339	1.6	0.99	1857.0	26.2	1846.1	14.0	1833.9	4.2	1833.9	4.2	101.3	-1.3	
VININI-97SF-11-173	96	28135	1.3	8.9184	0.6	4.9813	2.0	0.3222	1.9	0.96	1800.5	29.9	1816.2	16.8	1834.2	10.6	1834.2	10.6	98.2	1.8	
VININI-97SF-11-99	36	28280	0.8	8.9162	1.5	5.1436	2.1	0.3326	1.5	0.71	1851.1	23.7	1843.3	17.7	1834.6	26.8	1834.6	26.8	100.9	-0.9	
VININI-97SF-11-1	74	25752	1.2	8.9161	0.9	5.2119	3.0	0.3370	2.9	0.96	1872.4	46.7	1854.6	25.6	1834.6	15.8	1834.6	15.8	102.1	-2.1	
VININI-97SF-11-34	34	24461	0.9	8.9159	1.6	5.1908	2.4	0.3357	1.8	0.75	1865.8	28.9	1851.1	20.3	1834.7	28.7	1834.7	28.7	101.7	-1.7	
VININI-97SF-11-53	37	25202	1.0	8.9154	1.8	5.1984	2.8	0.3361	2.1	0.76	1868.0	34.4	1852.4	23.8	1834.8	33.1	1834.8	33.1	101.8	-1.8	
VININI-97SF-11-19	73	29432	1.4	8.9078	0.6	5.1612	1.5	0.3334	1.4	0.91	1855.0	21.8	1846.2	12.6	1836.3	11.0	1836.3	11.0	101.0	-1.0	
VININI-97SF-11-96	40	12582	0.6	8.9066	1.1	5.2902	3.3	0.3417	3.1	0.94	1895.0	50.3	1867.3	27.9	1836.2	20.7	1836.2	20.7	103.2	-3.2	
VININI-97SF-11-142	36	16147	0.7	8.9061	1.3	5.1591	1.7	0.3332	1.1	0.66	1854.1	18.1	1845.9	14.4	1836.7	22.9	1836.7	22.9	100.9	-0.9	
VININI-97SF-11-24	53	39772	1.2	8.9057	1.4	5.1759	1.7	0.3343	1.0	0.59	1859.2	16.0	1848.7	14.3	1836.8	24.7	1836.8	24.7	101.2	-1.2	
VININI-97SF-11-44	177	117286	1.1	8.9051	0.5	5.0228	1.0	0.3244	0.9	0.90	1811.2	14.8	1823.2	8.8	1836.9	8.4	1836.9	8.4	98.6	1.4	
VININI-97SF-11-59	113	53068	1.7	8.9033	0.6	5.0543	3.3	0.3264	3.2	0.98	1820.8	50.9	1828.5	27.7	1837.2	11.3	1837.2	11.3	99.1	0.9	
VININI-97SF-11-95	54	44905	1.6	8.9024	1.0	5.1617	1.9	0.3333	1.7	0.87	1854.2	26.7	1846.3	16.2	1837.4	17.2	1837.4	17.2	100.9	-0.9	
VININI-97SF-11-123	49	20002	1.4	8.9014	1.3	5.0465	1.9	0.3258	1.5	0.76	1818.0	23.3	1827.2	16.5	1837.6	23.1	1837.6	23.1	98.9	1.1	
VININI-97SF-11-153	95	37674	2.7	8.8976	0.5	5.1286	1.3	0.3310	1.2	0.92	1843.0	19.1	1840.9	11.0	1838.4	9.3	1838.4	9.3	100.3	-0.3	
VININI-97SF-11-35	159	68874	1.7	8.8969	0.3	5.2401	1.0	0.3381	0.9	0.96	1877.6	15.1	1859.2	8.3	1838.5	5.1	1838.5	5.1	102.1	-2.1	
VININI-97SF-11-3	106	33251	2.1	8.8942	0.4	5.2864	1.0	0.3410	0.9	0.90	1891.5	14.0	1866.7	8.1	1839.1	7.6	1839.1	7.6	102.9	-2.9	
VININI-97SF-11-79	95	106846	1.2	8.8942	1.0	5.2908	2.1	0.3413	1.8	0.88	1892.9	30.1	1867.4	17.8	1839.1	17.6	1839.1	17.6	102.9	-2.9	
VININI-97SF-11-127	109	25081	1.6	8.8908	0.5	5.1389	1.2	0.3314	1.1	0.91	1845.0	17.1	1842.6	9.9	1839.8	8.7	1839.8	8.7	100.3	-0.3	
VININI-97SF-11-105	36	23115	1.0	8.8894	1.7	5.2789	2.1	0.3403	1.2	0.59	1888.3	20.4	1865.5	18.0	1840.1	30.7	1840.1	30.7	102.6	-2.6	
VININI-97SF-11-177	139	81084	1.5	8.8892	0.4	5.1166	1.4	0.3299	1.3	0.96	1837.7	21.0	1838.9	11.6	1840.1	7.0	1840.1	7.0	99.9	0.1	
VININI-97SF-11-128	53	38166	0.7	8.8878	0.9	5.0888	1.3	0.3280													

## U-Pb geochronologic analyses of selected Roberts Mountains allochthon strata

	Isotope ratios										Apparent ages (Ma)									
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc.	Discord.
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)
<b>Sample: upper Vinini Formation. Location: Petes Summit, Toquima Range; 0518089 4337111 (NAD 83 UTM 11T)</b>																				
VININI-97SF-11-58	195	96148	0.9	8.7987	0.5	5.1796	1.6	0.3305	1.5	0.94	1840.9	23.5	1849.3	13.2	1858.6	9.3	1858.6	9.3	99.0	1.0
VININI-97SF-11-146	55	40013	1.6	8.7971	0.8	5.2271	1.8	0.3335	1.6	0.89	1855.4	25.3	1857.1	15.0	1858.9	14.5	1858.9	14.5	99.8	0.2
VININI-97SF-11-134	242	301056	2.5	8.7901	0.2	5.2515	0.9	0.3348	0.9	0.97	1861.6	14.4	1861.0	7.9	1860.4	4.2	1860.4	4.2	100.1	-0.1
VININI-97SF-11-167	74	27962	2.8	8.7875	0.4	4.9894	1.3	0.3180	1.2	0.95	1779.9	18.8	1817.5	10.7	1860.9	7.2	1860.9	7.2	95.6	4.4
VININI-97SF-11-132	94	22789	1.2	8.7866	0.6	5.0952	1.7	0.3247	1.6	0.93	1812.6	24.6	1835.3	14.2	1861.1	11.1	1861.1	11.1	97.4	2.6
VININI-97SF-11-8	132	82264	1.6	8.7745	0.6	5.2939	1.0	0.3369	0.9	0.84	1871.7	13.8	1867.9	8.7	1863.6	10.0	1863.6	10.0	100.4	-0.4
VININI-97SF-11-47	44	12818	0.8	8.7729	1.1	5.1821	1.8	0.3297	1.4	0.79	1837.0	22.9	1849.7	15.4	1863.9	19.9	1863.9	19.9	98.6	1.4
VININI-97SF-11-9	112	57066	2.6	8.7339	0.5	5.4446	1.0	0.3449	0.9	0.85	1910.1	14.7	1891.9	8.9	1872.0	9.8	1872.0	9.8	102.0	-2.0
VININI-97SF-11-140	47	18779	0.6	8.7206	1.5	5.4564	2.2	0.3451	1.7	0.74	1911.2	27.3	1893.8	19.0	1874.7	26.7	1874.7	26.7	101.9	-1.9
VININI-97SF-11-10	37	16789	1.2	8.7196	0.8	5.3984	2.2	0.3414	2.0	0.92	1893.4	32.6	1884.6	18.5	1874.9	15.3	1874.9	15.3	101.0	-1.0
VININI-97SF-11-158	106	54672	2.1	8.7155	0.5	5.3007	2.0	0.3351	2.0	0.97	1862.9	32.2	1869.0	17.5	1875.8	8.4	1875.8	8.4	99.3	0.7
VININI-97SF-11-121	220	71623	1.7	8.6691	0.4	5.2383	1.1	0.3294	1.0	0.92	1835.2	16.0	1858.9	9.3	1885.4	7.7	1885.4	7.7	97.3	2.7
VININI-97SF-11-88	88	26408	0.9	8.6669	0.7	5.4691	1.6	0.3438	1.4	0.89	1904.8	23.1	1895.8	13.6	1885.8	13.2	1885.8	13.2	101.0	-1.0
VININI-97SF-11-5	79	24795	1.2	8.6432	0.9	5.5642	1.6	0.3488	1.3	0.83	1928.9	22.4	1910.6	13.9	1890.8	16.3	1890.8	16.3	102.0	-2.0
VININI-97SF-11-144	63	23221	1.0	8.6421	0.7	5.4626	2.9	0.3424	2.8	0.97	1898.1	45.8	1894.7	24.5	1891.0	11.8	1891.0	11.8	100.0	-0.4
VININI-97SF-11-48	96	37594	1.9	8.6322	0.6	5.5456	1.1	0.3472	1.0	0.87	1921.2	16.1	1907.7	9.6	1893.1	10.1	1893.1	10.1	101.5	-1.5
VININI-97SF-11-150	172	95937	1.0	8.6252	0.4	5.5077	1.5	0.3445	1.5	0.96	1908.5	24.1	1901.8	13.0	1894.5	7.6	1894.5	7.6	100.7	-0.7
VININI-97SF-11-124	215	75870	1.6	8.6170	1.0	5.1652	2.5	0.3228	2.3	0.92	1803.4	36.3	1846.9	21.3	1896.2	17.5	1896.2	17.5	95.1	4.9
VININI-97SF-11-149	43	25262	0.9	8.6166	0.8	5.5322	1.5	0.3457	1.3	0.86	1914.2	22.1	1905.6	13.3	1896.3	14.1	1896.3	14.1	100.9	-0.9
VININI-97SF-11-193	200	226678	2.4	8.6105	0.2	5.5836	1.6	0.3487	1.6	0.99	1928.4	26.9	1913.6	14.1	1897.6	4.1	1897.6	4.1	101.6	-1.6
VININI-97SF-11-31	53	27944	0.7	8.6018	1.0	5.6715	2.6	0.3538	2.4	0.92	1952.8	40.2	1927.0	22.4	1899.4	18.4	1899.4	18.4	102.8	-2.8
VININI-97SF-11-32	56	66327	1.5	8.5947	0.8	5.4539	1.5	0.3400	1.3	0.84	1886.5	21.1	1893.4	13.3	1900.9	15.2	1900.9	15.2	99.2	0.8
VININI-97SF-11-133	19	18778	0.7	8.5709	2.8	5.4682	3.3	0.3399	1.8	0.53	1886.3	28.9	1895.6	28.5	1905.9	50.6	1905.9	50.6	99.0	1.0
VININI-97SF-11-2	75	54375	1.1	8.5587	0.6	5.6308	1.2	0.3495	1.0	0.85	1932.3	16.9	1920.8	10.3	1908.4	11.2	1908.4	11.2	101.3	-1.3
VININI-97SF-11-77	49	62223	1.1	8.5346	1.2	5.6287	1.5	0.3484	0.9	0.61	1927.0	15.2	1920.5	12.8	1913.5	21.1	1913.5	21.1	100.7	-0.7
VININI-97SF-11-71	47	80247	1.1	8.5337	1.7	5.6273	4.5	0.3463	4.2	0.93	1926.4	69.4	1920.3	38.9	1913.7	30.6	1913.7	30.6	100.7	-0.7
VININI-97SF-11-86	221	217719	1.7	8.5158	0.2	5.4175	1.8	0.3346	1.8	0.99	1860.6	29.2	1887.6	15.6	1917.4	3.8	1917.4	3.8	97.0	3.0
VININI-97SF-11-57	50	26794	0.6	8.5120	0.9	5.6866	1.6	0.3511	1.4	0.84	1939.7	23.0	1929.3	14.1	1918.2	16.0	1918.2	16.0	101.1	-1.1
VININI-97SF-11-97	37	24331	1.3	8.5046	1.3	5.6365	1.7	0.3477	1.1	0.63	1923.4	18.1	1921.7	14.9	1919.8	24.0	1919.8	24.0	100.2	-0.2
VININI-97SF-11-199	44	24016	2.1	8.4738	1.6	5.8044	2.2	0.3567	1.5	0.69	1966.6	25.7	1947.1	19.1	1926.3	28.7	1926.3	28.7	102.1	-2.1
VININI-97SF-11-41	28	19005	1.1	8.4723	1.4	5.5113	1.7	0.3387	0.9	0.53	1880.2	14.5	1902.4	14.4	1926.6	25.5	1926.6	25.5	97.6	2.4
VININI-97SF-11-156	86	35825	0.9	8.4715	0.4	5.5706	1.6	0.3423	1.6	0.97	1897.6	26.0	1911.6	14.0	1926.8	6.8	1926.8	6.8	98.5	1.5
VININI-97SF-11-197	40	69540	0.6	8.4589	1.3	5.7741	2.0	0.3542	1.5	0.76	1954.8	25.5	1942.5	17.2	1929.5	23.2	1929.5	23.2	101.3	-1.3
VININI-97SF-11-172	536	38226	3.1	8.4583	0.2	5.0091	1.2	0.3073	1.2	0.98	1727.3	18.0	1820.9	10.2	1929.6	3.9	1929.6	3.9	89.5	10.5
VININI-97SF-11-115	82	34522	1.0	8.4560	0.6	5.6348	1.8	0.3456	1.7	0.94	1913.4	28.6	1921.4	15.8	1930.1	10.8	1930.1	10.8	99.1	0.9
VININI-97SF-11-183	54	34688	1.2	8.4553	1.0	5.7332	1.3	0.3516	0.8	0.64	1942.1	14.1	1936.4	11.4	1930.2	18.2	1930.2	18.2	100.6	-0.6
VININI-97SF-11-101	292	164616	0.6	8.4505	0.3	5.7603	1.1	0.3530	1.1	0.97	1949.1	17.9	1940.5	9.5	1931.2	4.5	1931.2	4.5	100.9	-0.9
VININI-97SF-11-166	224	164058	0.6	8.4412	0.3	5.8299	0.9	0.3569	0.8	0.94	1967.5	14.0	1950.9	7.6	1933.2	5.5	1933.2	5.5	101.8	-1.8
VININI-97SF-11-122	94	27302	1.3	8.4402	0.7	5.6000	1.8	0.3428	1.7	0.92	1900.1	27.8	1916.1	15.9	1933.4	13.3	1933.4	13.3	98.3	1.7
VININI-97SF-11-70	42	32447	0.6	8.4369	1.0	5.8650	1.8	0.3589	1.5	0.84	1976.9	25.6	1966.1	15.6	1934.1	17.7	1934.1	17.7	102.2	-2.2
VININI-97SF-11-170	144	47086	0.7	8.4256	0.3	5.7505	1.6	0.3514	1.6	0.98	1941.3	27.0	1939.0	14.2	1936.5	5.5	1936.5	5.5	100.2	-0.2
VININI-97SF-11-11	64	5075	1.2	8.4222	1.1	5.8381	1.7	0.3566	1.3	0.78	1966.1	22.8	1952.1	15.0	1937.2	19.3	1937.2	19.3	101.5	-1.5
VININI-97SF-11-7	134	65089	0.7	8.4214	0.2	5.8099	1.0	0.3549	1.0	0.97	1957.7	16.7	1947.9	8.8	1937.4	4.1	1937.4	4.1	101.0	-1.0
VININI-97SF-11-65	27	8307	0.3	8.4210	2.3	5.8671	2.4	0.3583	0.8	0.33	1974.3	13.7	1956.4	20.9	1937.5	40.6	1937.5	40.6	101.9	-1.9
VININI-97SF-11-81	101	53354	0.4	8.3538	0.6	5.9342	1.0	0.3595	0.8	0.79	1980.0	14.1	1966.3	9.1	1951.8	11.5	1951.8	11.5	101.4	-1.4
VININI-97SF-11-110	130	60871	0.8	8.3056	0.4	5.9535	1.8	0.3586	1.7	0.97	1975.7	29.0	1969.1	15.3	1962.2	7.6	1962.2	7.6	100.7	-0.7
VININI-97SF-11-160	29	11398	1.3	8.2504	1.2	5.9962	1.9	0.3588	1.4	0.77	1976.5	24.3	1975.3	16.1	1974.1	21.0	1974.1	21.0	100.1	-0.1
VININI-97SF-11-14	237	205331	2.3	8.2273	0.5	5.9906	2.9	0.3575	2.8	0.98	1970.1	47.5	1974.5	24.8	1979.0	9.7	1979.0	9.7	99.5	0.5
VININI-97SF-11-38	45	10213	0.8	8.2272	4.2	5.8646	4.7	0.3499	2.1	0.45	1934.3	35.2	1956.0	40.7	1979.1	74.6	1979.1	74.6	97.7	2.3
VININI-97SF-11-114	81	25313	0.5	8.2147	0.5	6.0538	1.1	0.3607	0.9	0.87	1985.4	16.0	1983.6	9.4	1981.8	9.7	1981.8	9.7	100.2	-0.2
VININI-97SF-11-154	61	37663	1.4	8.0477	1.1	6.2700	1.7	0.3660	1.3	0.78	2010.4	22.4	2014.3	14.7	2018.2	18.7	2018.2	18.7	99.6	0.4
VININI-97SF-11-94	266	178919	4.8	7.9887	0.8	6.2291	1.7	0.3609	1.5	0.88	1986.5	26.3	2008.5	15.3	2031.3	14.4	2031.3	14.4	97.8	2.2
VININI-97SF-11-84	62	30123	1.4	7.8706	0.9	6.4274	4.3	0.3669	4.2	0.98	2014.8	73.0	2036.0	37.9	2057.6	15.8	2057.6	15.8	97.9	2.1
VININI-97SF-11-12	79	38142	1.3	7.7992	0.5	6.7935	0.9	0.3843	0.7	0.84	2096.2	12.8	2084.9	7.5	2073.7	8.2	2073.7	8.2	101.1	-1.1
VININI-97SF-11-179	36	25677	1.1	7.7867	1.2	6.8150	2.2	0.3849	1.8	0.84	2099.0	32.8	2087.7	19.3	2076.5	20.6	2076.5	20.6	101.1	-1.1
VININI-97SF-11-30	266	91671	2.0	7.7866	0.3	6.6047	1.9	0.3730	1.8	0.99	2043.5	32.1	2060.0	16.4	2076.5	5.0	2076.5	5.0	98.4	



### U-Pb geochronologic analyses of selected Roberts Mountains allochthon strata

	Isotope ratios										Apparent ages (Ma)									
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc.	Discord.
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)
<b>Sample: upper Vinini Formation. Location: Petes Summit, Toquima Range; 0518089 4337111 (NAD 83 UTM 11T)</b>																				
VININI-97SF-11-106	33	25913	1.3	6.2191	1.6	9.7716	4.0	0.4408	3.7	0.92	2354.0	72.9	2413.5	37.0	2464.0	26.4	2464.0	26.4	95.5	4.5
VININI-97SF-11-171	67	25273	1.2	6.1697	0.5	10.3846	2.2	0.4647	2.1	0.97	2460.2	43.8	2469.7	20.5	2477.5	9.3	2477.5	9.3	99.3	0.7
VININI-97SF-11-119	302	236598	1.9	5.8782	0.2	11.3351	1.3	0.4832	1.3	0.99	2541.4	26.6	2551.1	12.0	2558.8	3.6	2558.8	3.6	99.3	0.7
VININI-97SF-11-83	39	38140	1.3	5.8084	0.8	11.7178	1.7	0.4936	1.5	0.89	2586.3	32.5	2582.1	16.0	2578.8	12.9	2578.8	12.9	100.3	-0.3
VININI-97SF-11-64	78	56320	1.4	5.7655	1.0	11.4819	2.3	0.4801	2.1	0.90	2527.8	44.2	2563.1	21.9	2591.2	16.9	2591.2	16.9	97.6	2.4
VININI-97SF-11-116	132	68890	2.2	5.7549	0.3	11.5848	0.6	0.4835	0.5	0.83	2542.6	10.5	2571.5	5.6	2594.2	5.5	2594.2	5.5	98.0	2.0
VININI-97SF-11-43	97	90131	1.9	5.7124	0.3	12.3047	1.0	0.5098	0.9	0.94	2655.7	20.3	2627.9	9.3	2606.6	5.5	2606.6	5.5	101.9	-1.9
VININI-97SF-11-80	41	63401	0.7	5.5871	0.6	12.3834	1.4	0.5018	1.3	0.91	2621.5	27.6	2633.9	13.2	2643.5	9.6	2643.5	9.6	99.2	0.8
VININI-97SF-11-92	57	111053	1.3	5.5806	0.3	12.5648	0.9	0.5086	0.8	0.93	2650.5	18.5	2647.6	8.6	2645.4	5.6	2645.4	5.6	100.2	-0.2
VININI-97SF-11-200	60	100513	1.1	5.4764	0.6	13.4425	1.0	0.5339	0.8	0.81	2757.9	18.8	2711.3	9.8	2676.6	10.1	2676.6	10.1	103.0	-3.0
VININI-97SF-11-72	64	54439	0.5	5.4749	0.5	12.9087	1.1	0.5126	1.0	0.89	2667.6	21.7	2673.0	10.5	2677.1	8.3	2677.1	8.3	99.6	0.4
VININI-97SF-11-188	168	80782	1.1	5.4709	0.1	13.2225	1.2	0.5247	1.2	0.99	2718.9	27.3	2695.7	11.7	2678.3	2.4	2678.3	2.4	101.5	-1.5
VININI-97SF-11-69	48	52485	1.5	5.4691	0.4	13.2482	1.7	0.5255	1.7	0.97	2722.5	37.6	2697.5	16.4	2678.8	6.7	2678.8	6.7	101.6	-1.6
VININI-97SF-11-46	165	39003	0.6	5.4482	0.2	12.4556	1.9	0.4922	1.9	1.00	2580.1	40.0	2639.4	17.8	2685.2	3.0	2685.2	3.0	96.1	3.9
VININI-97SF-11-62	177	113543	0.9	5.4240	0.2	13.4588	4.0	0.5294	4.0	1.00	2739.1	88.4	2712.4	37.5	2692.5	2.8	2692.5	2.8	101.7	-1.7
VININI-97SF-11-102	26	21015	0.5	5.4048	0.8	13.3199	1.7	0.5221	1.5	0.90	2708.2	33.8	2702.6	16.1	2698.4	12.4	2698.4	12.4	100.4	-0.4
VININI-97SF-11-174	19	21930	1.2	5.3960	1.0	13.2173	2.3	0.5173	2.1	0.91	2687.6	46.4	2695.3	21.9	2701.1	16.0	2701.1	16.0	99.5	0.5
VININI-97SF-11-108	88	36668	0.6	5.3944	0.3	13.4664	2.3	0.5269	2.3	0.99	2728.2	51.1	2712.9	21.9	2701.6	4.4	2701.6	4.4	101.0	-1.0
VININI-97SF-11-73	130	109070	0.7	5.3937	0.2	12.8656	1.2	0.5033	1.2	0.98	2627.9	25.3	2669.9	11.2	2701.8	3.6	2701.8	3.6	97.3	2.7
VININI-97SF-11-75	112	64953	0.9	5.3782	0.3	13.3490	0.7	0.5207	0.6	0.92	2702.1	14.0	2704.7	6.5	2706.5	4.5	2706.5	4.5	99.8	0.2
VININI-97SF-11-90	79	143842	1.1	5.3658	0.4	13.9583	1.3	0.5432	1.2	0.95	2796.8	27.3	2746.9	12.0	2710.3	6.4	2710.3	6.4	103.2	-3.2
VININI-97SF-11-194	45	40729	0.6	5.3557	0.5	14.1007	2.1	0.5471	2.0	0.97	2815.7	45.8	2756.5	19.6	2713.4	8.0	2713.4	8.0	103.8	-3.8
VININI-97SF-11-22	92	74917	0.8	5.3557	0.2	13.6992	1.0	0.5327	0.9	0.97	2750.4	20.9	2729.1	9.1	2713.4	3.6	2713.4	3.6	101.4	-1.4
VININI-97SF-11-120	41	19469	0.4	5.3465	0.3	13.5105	1.0	0.5239	0.9	0.94	2715.7	19.9	2716.0	9.0	2716.3	5.3	2716.3	5.3	100.0	0.0
VININI-97SF-11-126	38	21471	0.3	5.3380	0.6	13.3728	1.1	0.5177	0.9	0.86	2689.5	20.6	2706.3	10.3	2718.9	9.1	2718.9	9.1	98.9	1.1
VININI-97SF-11-147	30	30860	0.4	5.3048	0.7	13.8093	1.6	0.5313	1.5	0.90	2746.9	32.9	2736.7	15.5	2729.2	11.7	2729.2	11.7	100.7	-0.7
VININI-97SF-11-185	37	19455	1.0	5.3014	0.6	13.6739	1.2	0.5258	1.1	0.88	2723.5	24.1	2727.4	11.7	2730.2	9.7	2730.2	9.7	99.8	0.2
VININI-97SF-11-89	48	38408	1.6	5.2898	0.5	13.9833	0.7	0.5365	0.5	0.73	2768.7	11.9	2748.6	6.8	2733.8	8.1	2733.8	8.1	101.3	-1.3
VININI-97SF-11-152	40	9737	0.8	5.1804	0.6	14.4418	1.3	0.5426	1.1	0.87	2794.4	25.7	2779.2	12.4	2768.2	10.6	2768.2	10.6	100.9	-0.9
VININI-97SF-11-82	82	28437	1.6	5.0348	0.4	11.7699	3.6	0.4298	3.6	0.99	2304.8	70.1	2586.3	34.1	2814.9	6.6	2814.9	6.6	81.9	18.1
VININI-97SF-11-103	40	8209	0.7	5.0250	0.6	15.3149	1.2	0.5581	1.1	0.87	2859.0	24.8	2835.0	11.7	2818.0	9.8	2818.0	9.8	101.5	-1.5
VININI-97SF-11-104	97	65054	3.3	4.9992	2.3	13.6211	3.9	0.4939	3.1	0.80	2587.4	65.8	2723.7	36.7	2826.5	38.2	2826.5	38.2	91.5	8.5
VININI-97SF-11-155	98	118308	0.8	3.2308	0.4	31.5737	2.7	0.7398	2.6	0.99	3570.0	72.3	3537.1	26.3	3518.5	6.3	3518.5	6.3	101.5	-1.5
<b>Sample: Elder Sandstone. Location: Elder Creek, Shoshone Range; 0516196 4460270 (NAD 83 UTM 11T)</b>																				
ELDRSS-143 <>	463	19005	7.1	21.5974	8.6	0.0727	8.9	0.0114	2.5	0.27	73.0	1.8	71.3	6.2	13.4	207.1	73.0	1.8	NA	NA
ELDRSS-29 <>	1260	12089	3.2	18.0411	0.7	0.4201	2.4	0.0550	2.2	0.95	344.9	7.5	356.1	7.1	429.5	16.7	344.9	7.5	NA	NA
ELDRSS-23 <>	256	54686	0.9	14.0057	1.0	1.5854	1.1	0.1610	0.5	0.50	962.6	4.9	966.9	6.8	968.9	19.4	968.9	19.4	99.4	0.6
ELDRSS-197 <>	56	26051	0.7	13.4985	2.7	1.7931	2.8	0.1755	1.0	0.35	1042.6	9.4	1043.0	18.4	1043.7	53.5	1043.7	53.5	99.9	0.1
ELDRSS-74 <>	132	56241	2.0	13.4913	1.3	1.7383	1.8	0.1701	1.2	0.69	1012.6	11.5	1022.8	11.4	1044.8	25.8	1044.8	25.8	96.9	3.1
ELDRSS-22 <>	136	85797	1.1	12.8970	1.3	2.0618	1.4	0.1929	0.5	0.34	1136.8	5.0	1136.2	9.7	1135.0	26.7	1135.0	26.7	100.2	-0.2
ELDRSS-99 <>	50	13222	1.8	11.9004	2.3	2.4903	4.9	0.2149	4.3	0.88	1255.2	49.2	1269.2	35.5	1293.2	45.5	1293.2	45.5	97.0	3.0
ELDRSS-185 <>	222	191878	2.0	11.3531	0.5	2.8591	0.8	0.2354	0.7	0.78	1362.8	8.1	1371.2	6.4	1384.2	10.2	1384.2	10.2	98.5	1.5
ELDRSS-155 <>	74	65602	2.1	11.2670	1.4	2.9634	1.5	0.2422	0.6	0.41	1397.9	7.9	1398.3	11.6	1398.8	26.6	1398.8	26.6	99.9	0.1
ELDRSS-87 <>	187	172967	0.9	10.8788	0.6	3.2648	1.3	0.2576	1.1	0.87	1477.5	14.7	1472.9	9.9	1465.8	11.7	1465.8	11.7	100.8	-0.8
ELDRSS-183 <>	143	136954	1.6	9.9179	0.7	4.0429	1.2	0.2908	0.9	0.80	1645.6	13.7	1642.9	9.6	1639.4	13.0	1639.4	13.0	100.4	-0.4
ELDRSS-184 <>	292	185849	5.2	9.9007	0.4	3.9798	1.3	0.2858	1.2	0.96	1620.4	17.9	1630.1	10.5	1642.6	6.5	1642.6	6.5	98.6	1.4
ELDRSS-145 <>	112	6588	0.7	9.8350	1.4	3.7989	3.2	0.2710	2.9	0.90	1545.7	39.3	1592.5	25.6	1655.0	25.7	1655.0	25.7	93.4	6.6
ELDRSS-109 <>	38	32248	1.3	9.6908	1.7	4.2513	2.1	0.2988	1.2	0.58	1685.3	18.1	1684.0	17.3	1682.3	31.7	1682.3	31.7	100.2	-0.2
ELDRSS-104 <>	614	87858	1.5	9.5047	0.2	3.6418	0.7	0.2510	0.6	0.95	1443.9	8.3	1558.7	5.4	1718.0	3.7	1718.0	3.7	84.0	16.0
ELDRSS-63 <>	337	228687	2.9	9.3265	0.4	4.5882	0.9	0.3104	0.8	0.91	1742.4	12.4	1747.1	7.4	1752.7	6.6	1752.7	6.6	99.4	0.6
ELDRSS-94 <>	325	445712	2.4	9.2080	0.2	4.7777	0.8	0.3191	0.8	0.97	1785.2	12.4	1781.0	6.9	1776.1	3.5	1776.1	3.5	100.5	-0.5
ELDRSS-3 <>	236	224617	2.7	9.1980	0.4	4.8018														

## U-Pb geochronologic analyses of selected Roberts Mountains allochthon strata

	Isotope ratios								Apparent ages (Ma)											
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc.	Discord.
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)
<b>Sample: Elder Sandstone. Location: Elder Creek, Shoshone Range; 0516196 4460270 (NAD 83 UTM 11T)</b>																				
ELDRSS-49 <>	217	218333	1.5	8.9118	0.4	5.1873	1.2	0.3353	1.1	0.95	1863.9	18.5	1850.5	10.3	1835.5	7.0	1835.5	7.0	101.5	-1.5
ELDRSS-64 <>	91	114121	1.2	8.9110	0.5	5.1990	1.1	0.3360	1.0	0.87	1867.4	16.0	1852.6	9.6	1835.7	9.9	1835.7	9.9	101.7	-1.7
ELDRSS-85 <>	334	120849	2.0	8.9086	0.2	4.7108	2.6	0.3044	2.6	1.00	1713.0	39.6	1769.2	22.1	1836.2	4.5	1836.2	4.5	93.3	6.7
ELDRSS-200 <>	111	83082	1.7	8.9086	0.5	5.1175	1.0	0.3306	0.9	0.88	1841.5	14.1	1839.0	8.5	1836.2	8.5	1836.2	8.5	100.3	-0.3
ELDRSS-88 <>	104	83527	1.1	8.9065	0.6	5.1581	0.8	0.3332	0.5	0.64	1853.8	8.5	1845.7	7.0	1836.6	11.4	1836.6	11.4	100.9	-0.9
ELDRSS-80 <>	76	99241	1.2	8.9050	1.3	5.1131	1.6	0.3302	0.9	0.56	1839.5	14.3	1838.3	13.7	1836.9	24.3	1836.9	24.3	100.1	-0.1
ELDRSS-75 <>	60	59563	1.6	8.9042	0.8	5.1059	1.3	0.3297	1.0	0.79	1837.1	16.7	1837.1	11.2	1837.1	14.5	1837.1	14.5	100.0	0.0
ELDRSS-98 <>	34	21414	0.7	8.9038	2.3	5.1372	2.7	0.3317	1.3	0.50	1846.8	21.6	1842.3	23.0	1837.1	42.6	1837.1	42.6	100.5	-0.5
ELDRSS-45 <>	155	142861	1.7	8.8976	0.5	5.1479	0.8	0.3322	0.6	0.77	1849.1	9.4	1844.1	6.4	1838.4	8.7	1838.4	8.7	100.6	-0.6
ELDRSS-150 <>	76	50979	1.3	8.8973	0.7	5.2407	1.5	0.3382	1.3	0.88	1877.9	21.0	1859.3	12.5	1838.5	12.7	1838.5	12.7	102.1	-2.1
ELDRSS-127 <>	157	120824	1.7	8.8954	0.6	5.174	2.1	0.2914	2.0	0.96	1648.7	29.0	1734.2	17.2	1839.9	10.4	1839.9	10.4	89.7	10.3
ELDRSS-51 <>	58	56474	1.9	8.8940	0.9	5.1072	1.3	0.3294	0.9	0.70	1835.7	14.1	1837.3	10.7	1839.1	16.3	1839.1	16.3	99.8	0.2
ELDRSS-6 <>	104	179003	1.2	8.8938	0.9	5.0954	1.3	0.3287	1.0	0.73	1832.0	15.4	1835.3	11.2	1839.2	16.2	1839.2	16.2	99.6	0.4
ELDRSS-100 <>	63	44668	3.2	8.8912	1.4	4.4198	1.9	0.2850	1.4	0.71	1616.5	19.8	1716.1	16.1	1839.7	24.6	1839.7	24.6	87.9	12.1
ELDRSS-138 <>	65	71616	1.6	8.8901	1.9	5.1041	2.1	0.3291	0.9	0.43	1834.0	14.6	1836.6	17.9	1839.9	34.5	1839.9	34.5	99.7	0.3
ELDRSS-108 <>	82	66583	1.0	8.8901	0.6	5.1680	1.1	0.3369	0.9	0.83	1853.3	14.5	1847.0	9.3	1839.9	11.1	1839.9	11.1	100.7	-0.7
ELDRSS-70 <>	55	56099	1.0	8.8900	1.0	4.5934	3.2	0.2962	3.0	0.95	1672.3	44.0	1748.1	26.3	1840.0	18.1	1840.0	18.1	90.9	9.1
ELDRSS-44 <>	105	114428	3.5	8.8891	0.7	5.2185	1.4	0.3363	1.2	0.85	1868.9	19.2	1855.3	11.8	1840.1	13.0	1840.1	13.0	101.6	-1.6
ELDRSS-118 <>	145	115477	2.0	8.8889	0.7	5.1279	1.1	0.3306	0.9	0.78	1841.2	14.1	1840.7	9.6	1840.2	12.7	1840.2	12.7	100.1	-0.1
ELDRSS-82 <>	235	480780	1.2	8.8848	0.4	5.1698	0.7	0.3331	0.5	0.76	1853.6	8.2	1847.7	5.7	1841.0	8.0	1841.0	8.0	100.7	-0.7
ELDRSS-25 <>	59	69246	1.4	8.8838	1.6	5.0809	1.8	0.3274	0.7	0.42	1825.6	11.8	1832.9	15.1	1841.2	29.2	1841.2	29.2	99.2	0.8
ELDRSS-144 <>	41	46781	0.8	8.8797	0.8	5.1500	1.1	0.3317	0.8	0.68	1846.5	12.6	1844.4	9.7	1842.0	15.2	1842.0	15.2	100.2	-0.2
ELDRSS-153 <>	63	60767	0.6	8.8721	0.5	5.1636	0.9	0.3323	0.8	0.85	1843.6	12.2	1846.6	7.6	1843.6	8.6	1843.6	8.6	100.3	-0.3
ELDRSS-112 <>	110	198208	2.1	8.8704	0.5	5.1722	1.1	0.3328	1.0	0.90	1851.7	16.5	1848.1	9.7	1843.9	9.1	1843.9	9.1	100.4	-0.4
ELDRSS-78 <>	90	258213	1.5	8.8701	0.6	5.2328	1.6	0.3366	1.4	0.91	1870.5	23.3	1858.0	13.4	1844.0	11.5	1844.0	11.5	101.4	-1.4
ELDRSS-149 <>	43	46516	1.2	8.8698	1.0	5.1303	1.3	0.3300	0.8	0.60	1838.5	12.5	1841.1	11.1	1844.1	19.0	1844.1	19.0	99.7	0.3
ELDRSS-34 <>	115	205282	1.3	8.8686	0.5	5.2001	0.8	0.3345	0.6	0.81	1860.0	10.0	1852.6	6.5	1844.3	8.2	1844.3	8.2	100.9	-0.9
ELDRSS-163 <>	155	114274	2.1	8.8682	0.4	4.6673	3.4	0.3002	3.4	0.99	1692.3	50.6	1761.4	28.7	1844.4	8.0	1844.4	8.0	91.8	8.2
ELDRSS-156 <>	206	425547	1.6	8.8681	0.4	5.2459	0.8	0.3374	0.7	0.87	1874.2	11.5	1860.1	6.9	1844.4	7.2	1844.4	7.2	101.6	-1.6
ELDRSS-28 <>	123	180092	2.1	8.8653	0.4	5.2032	1.8	0.3346	1.8	0.97	1860.4	28.5	1853.1	15.5	1845.0	7.6	1845.0	7.6	100.8	-0.8
ELDRSS-175 <>	81	65325	0.8	8.8650	0.7	5.1849	1.0	0.3334	0.8	0.78	1854.7	13.1	1850.1	8.9	1845.0	11.9	1845.0	11.9	100.5	-0.5
ELDRSS-151 <>	84	65452	1.5	8.8634	0.6	5.1969	0.8	0.3341	0.5	0.66	1858.1	8.4	1852.1	6.7	1845.4	10.8	1845.4	10.8	100.7	-0.7
ELDRSS-111 <>	99	92275	1.6	8.8631	0.6	5.1977	0.9	0.3341	0.7	0.74	1858.3	10.9	1852.2	7.7	1845.4	11.0	1845.4	11.0	100.7	-0.7
ELDRSS-177 <>	118	123861	1.8	8.8628	0.6	5.1150	0.9	0.3288	0.7	0.73	1832.5	10.4	1838.6	7.6	1845.5	10.9	1845.5	10.9	99.3	0.7
ELDRSS-72 <>	91	118382	1.6	8.8602	0.5	5.1936	0.9	0.3337	0.7	0.81	1856.5	11.8	1851.6	7.7	1846.9	9.7	1846.9	9.7	100.6	-0.6
ELDRSS-171 <>	42	53500	1.4	8.8600	1.5	5.2359	1.8	0.3364	1.0	0.57	1869.6	16.4	1858.5	15.1	1846.1	26.4	1846.1	26.4	101.3	-1.3
ELDRSS-162 <>	30	24123	0.9	8.8594	1.4	5.1953	2.4	0.3338	1.9	0.80	1856.9	31.3	1851.8	20.6	1846.2	26.0	1846.2	26.0	100.6	-0.6
ELDRSS-117 <>	143	201862	1.9	8.8593	0.3	5.1883	0.8	0.3334	0.8	0.92	1854.7	12.5	1850.7	7.2	1846.2	5.9	1846.2	5.9	100.5	-0.5
ELDRSS-126 <>	151	358113	1.7	8.8578	0.6	4.9248	1.3	0.3164	1.2	0.90	1772.0	18.2	1806.5	11.0	1846.5	10.5	1846.5	10.5	96.0	4.0
ELDRSS-16 <>	107	217485	1.8	8.8570	0.5	5.2303	0.8	0.3360	0.5	0.70	1867.3	8.6	1857.6	6.5	1846.7	9.7	1846.7	9.7	101.1	-1.1
ELDRSS-101 <>	117	107829	1.6	8.8569	0.4	5.1895	0.6	0.3334	0.4	0.68	1854.6	6.4	1850.9	5.0	1846.7	7.7	1846.7	7.7	100.4	-0.4
ELDRSS-102 <>	223	324482	1.8	8.8534	0.5	5.2102	0.9	0.3345	0.7	0.78	1860.4	11.0	1854.3	7.4	1847.4	9.8	1847.4	9.8	100.7	-0.7
ELDRSS-134 <>	118	101378	1.2	8.8525	0.5	5.1805	0.8	0.3326	0.7	0.80	1851.0	10.8	1849.4	7.1	1847.6	9.0	1847.6	9.0	100.2	-0.2
ELDRSS-165 <>	53	40051	1.3	8.8520	0.8	5.2337	1.1	0.3360	0.7	0.65	1867.4	11.4	1858.1	9.2	1847.7	14.7	1847.7	14.7	101.1	-1.1
ELDRSS-122 <>	199	153850	2.0	8.8479	0.3	5.2277	0.6	0.3355	0.6	0.91	1864.8	8.9	1857.1	5.2	1848.5	4.6	1848.5	4.6	100.9	-0.9
ELDRSS-30 <>	325	25733	3.0	8.8478	0.3	4.2523	3.3	0.2729	3.2	1.00	1555.3	44.8	1684.2	26.7	1848.6	5.4	1848.6	5.4	84.1	15.9
ELDRSS-103 <>	36	24389	1.0	8.8478	2.2	5.2694	2.5	0.3381	1.2	0.49	1877.7	20.0	1863.9	21.2	1848.6	39.1	1848.6	39.1	101.6	-1.6
ELDRSS-18 <>	31	36073	0.7	8.8459	1.3	5.2648	1.8	0.3378	1.3	0.68	1875.9	20.6	1863.2	15.8	1848.9	24.4	1848.9	24.4	101.5	-1.5
ELDRSS-107 <>	69	52402	2.2	8.8459	0.8	5.1842	1.4	0.3328	1.2	0.81	1851.0	18.6	1850.0	12.1	1848.9	14.9	1848.9	14.9	100.1	-0.1
ELDRSS-115 <>	235	306229	1.6	8.8457	0.2	5.2118	0.5	0.3344	0.4	0.90	1859.5	7.1	1854.5	4.2	1849.0	4.0	1849.0	4.0	100.6	-0.6
ELDRSS-33 <>	95	110969	2.0	8.8441	0.9	5.1608	1.5	0.3310	1.2	0.82	1843.4	19.2	1846.2	12.5	1849.3	15.4	1849.3	15.4	99.7	0.3
ELDRSS-13 <>	218	124615	1.7	8.8391	0.3	5.3010	1.5	0.3398	1.5	0.98	1885.9	24.0	1869.0	12.7	1850.3	4.9	1850.3	4.9	101.9	-1.9
ELDRSS-164 <>	298	194707	1.8	8.8355	0.2	5.0537	1.1	0.3238	1.0	0.98	1808.5	16.6	1828.4	9.1	1851.1	3.8	1851.1	3.8	97.7	2.3
ELDRSS-21 <>	30	30954	1.9	8.8344	2.1	5.2980	2.7	0.3359	1.7	0.64	1884.0	27.9	1868.5	22.9	1851.3	37.3	1851.3	37.3	101.8	-1.8
ELDRSS-116 <>	74	209506	0.8	8.8337	0.5	5.2547	1.1	0.3367	1.0	0.91	1870.6	16.5	1861.5	9.5	1851.4	8.3	1851.4	8.3	101.0	-1.0
ELDRSS-167 <>	219	190824	1.2	8.8313	0.3	5.2310	0.5	0.3351	0.4	0.81	1862.8	6.4	1857.7	4.1	1851.9	5.2	1851.9	5.2	100.6	-0.6
ELDRSS-159 <>	284	301976	1.9	8.8299	0.3	5.2690	0.6	0.												



## U-Pb geochronologic analyses of selected Roberts Mountains allochthon strata

	Isotope ratios													Apparent ages (Ma)							Conc.	Discord
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±				
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)		
<b>Sample: Elder Sandstone. Location: Elder Creek, Shoshone Range; 0516196 4460270 (NAD 83 UTM 11T)</b>																						
ELDRSS-41 <>	29	35676	0.9	8.5345	1.3	5.6243	1.4	0.3481	0.7	0.46	1925.7	10.9	1919.8	12.2	1913.5	22.5	1913.5	22.5	100.6	-0.6		
ELDRSS-92 <>	99	177855	0.7	8.5163	0.7	4.6904	1.6	0.2897	1.5	0.89	1640.1	21.2	1765.5	13.7	1917.3	13.1	1917.3	13.1	85.5	14.5		
ELDRSS-180 <>	56	36132	0.7	8.5119	1.1	5.6435	1.2	0.3484	0.5	0.40	1926.9	8.1	1922.8	10.4	1918.3	19.9	1918.3	19.9	100.5	-0.5		
ELDRSS-24 <>	112	283830	1.2	8.4649	0.8	5.7724	1.1	0.3544	0.7	0.67	1955.5	12.6	1942.3	9.7	1928.2	15.0	1928.2	15.0	101.4	-1.4		
ELDRSS-7 <>	34	45222	0.2	8.4617	1.3	5.7095	1.4	0.3504	0.5	0.37	1936.5	8.9	1932.8	12.5	1928.9	24.0	1928.9	24.0	100.4	-0.4		
ELDRSS-8 <>	109	181067	1.7	8.4584	0.7	5.7392	1.1	0.3521	0.8	0.76	1944.5	14.0	1937.3	9.6	1929.6	13.0	1929.6	13.0	100.8	-0.8		
ELDRSS-106 <>	42	136724	1.0	8.4175	1.2	5.7160	1.6	0.3490	1.0	0.62	1929.6	16.6	1933.8	13.8	1938.2	22.2	1938.2	22.2	99.6	0.4		
ELDRSS-132 <>	90	97520	1.2	8.3935	0.6	5.8141	0.7	0.3539	0.4	0.62	1953.4	7.5	1948.5	6.3	1943.4	10.1	1943.4	10.1	100.5	-0.5		
ELDRSS-53 <>	77	19906	1.0	8.2721	2.2	5.9103	2.3	0.3546	0.7	0.32	1956.5	12.4	1962.8	19.9	1969.4	38.7	1969.4	38.7	99.3	0.7		
ELDRSS-81 <>	41	44088	0.6	7.8282	1.3	6.7012	1.6	0.3805	1.0	0.58	2078.5	17.0	2072.8	14.5	2067.1	23.6	2067.1	23.6	100.5	-0.5		
ELDRSS-66 <>	84	112732	1.0	7.8206	0.7	6.6936	1.6	0.3797	1.5	0.91	2074.7	26.4	2071.8	14.5	2068.9	12.2	2068.9	12.2	100.3	-0.3		
ELDRSS-65 <>	37	7116	0.6	7.7917	0.9	6.7475	1.2	0.3813	0.8	0.67	2082.4	14.0	2078.9	10.4	2075.4	15.4	2075.4	15.4	100.3	-0.3		
ELDRSS-182 <>	56	82181	1.3	7.7877	1.0	6.7177	1.5	0.3794	1.1	0.73	2073.6	19.4	2075.0	13.2	2076.3	17.8	2076.3	17.8	99.9	0.1		
ELDRSS-172 <>	205	286259	2.3	7.7793	0.3	6.8295	0.7	0.3853	0.6	0.91	2101.1	10.7	2089.6	5.8	2078.2	4.7	2078.2	4.7	101.1	-1.1		
ELDRSS-57 <>	68	74519	1.4	7.7793	0.6	6.8074	1.2	0.3841	1.0	0.87	2095.3	17.9	2086.7	10.2	2078.2	10.1	2078.2	10.1	100.8	-0.8		
ELDRSS-190 <>	42	49227	1.3	7.7785	1.0	6.7742	1.3	0.3821	0.9	0.68	2085.9	15.9	2082.4	11.7	2078.8	17.1	2078.8	17.1	100.3	-0.3		
ELDRSS-1 <>	60	45885	1.0	7.7751	0.8	6.7537	1.3	0.3808	1.0	0.76	2080.2	17.6	2079.7	11.4	2079.1	14.7	2079.1	14.7	100.1	-0.1		
ELDRSS-188 <>	167	194472	1.5	7.7703	0.3	6.7408	0.7	0.3799	0.6	0.89	2075.7	10.8	2078.0	6.1	2080.2	5.5	2080.2	5.5	99.8	0.2		
ELDRSS-3 <>	16	23988	0.5	7.7679	1.7	6.6596	2.0	0.3752	1.0	0.53	2053.8	18.4	2067.3	17.5	2080.7	29.5	2080.7	29.5	98.7	1.3		
ELDRSS-37 <>	250	240027	2.3	7.7546	0.2	6.7902	1.0	0.3819	1.0	0.97	2085.1	17.7	2084.4	9.0	2083.8	4.2	2083.8	4.2	100.1	-0.1		
ELDRSS-38 <>	55	77482	1.6	7.7546	0.6	6.8097	1.4	0.3830	1.3	0.90	2090.2	23.0	2087.0	12.6	2083.8	10.8	2083.8	10.8	100.3	-0.3		
ELDRSS-11 <>	122	84593	2.1	7.7534	0.5	6.7676	1.0	0.3806	0.8	0.85	2078.9	15.0	2081.5	8.8	2084.0	9.2	2084.0	9.2	99.8	0.2		
ELDRSS-166 <>	108	187062	1.4	7.7458	0.3	6.8787	0.8	0.3864	0.7	0.90	2106.2	12.7	2095.9	6.9	2085.8	5.9	2085.8	5.9	101.0	-1.0		
ELDRSS-157 <>	78	71436	1.6	7.7360	0.9	6.8850	1.1	0.3863	0.6	0.59	2105.6	11.5	2096.7	9.7	2088.0	15.6	2088.0	15.6	100.8	-0.8		
ELDRSS-27 <>	114	112866	1.4	7.7357	0.5	6.9254	0.6	0.3885	0.4	0.66	2116.1	7.4	2101.9	5.5	2088.1	8.1	2088.1	8.1	101.3	-1.3		
ELDRSS-146 <>	106	79515	0.5	7.7339	0.8	6.7903	1.1	0.3809	0.8	0.68	2080.4	13.8	2084.5	10.1	2088.5	14.6	2088.5	14.6	99.6	0.4		
ELDRSS-114 <>	16	20153	0.5	7.7282	3.4	6.8791	3.8	0.3855	1.6	0.41	2101.8	27.9	2096.0	33.4	2090.2	60.3	2090.2	60.3	100.6	-0.6		
ELDRSS-178 <>	62	75401	1.4	7.7219	1.0	6.8626	1.4	0.3843	0.9	0.69	2096.5	17.0	2093.8	12.2	2091.2	17.5	2091.2	17.5	100.3	-0.3		
ELDRSS-123 <>	87	141464	1.0	7.7107	0.4	7.1694	3.1	0.4009	3.1	0.99	2173.3	57.5	2132.7	28.0	2093.8	6.3	2093.8	6.3	103.8	-3.8		
ELDRSS-195 <>	19	35465	0.8	7.6614	2.4	6.8129	2.8	0.3786	1.4	0.50	2069.6	25.1	2087.4	25.0	2105.0	43.0	2105.0	43.0	98.3	1.7		
ELDRSS-39 <>	253	624938	1.9	6.9476	0.2	6.3585	1.4	0.4212	1.4	0.99	2265.8	25.9	2270.7	12.4	2275.1	3.6	2275.1	3.6	99.6	0.4		
ELDRSS-46 <>	83	74014	1.7	6.7976	0.4	8.7993	1.1	0.4338	1.0	0.91	2322.9	19.2	2317.4	9.9	2312.6	7.7	2312.6	7.7	100.4	-0.4		
ELDRSS-54 <>	92	342415	2.0	6.1576	0.4	10.6530	0.8	0.4758	0.7	0.90	2508.7	14.9	2493.3	7.4	2480.8	5.9	2480.8	5.9	101.1	-1.1		
ELDRSS-147 <>	238	480675	1.2	6.0331	0.2	10.5102	0.6	0.4599	0.6	0.96	2439.1	12.2	2480.8	5.8	2515.2	3.1	2515.2	3.1	97.0	3.0		
ELDRSS-55 <>	86	109099	1.8	5.9692	0.5	11.1532	1.1	0.4829	1.0	0.92	2539.7	22.0	2536.0	10.7	2533.1	7.7	2533.1	7.7	100.3	-0.3		
ELDRSS-35 <>	90	69078	1.5	5.9336	0.4	11.3082	1.7	0.4866	1.7	0.98	2556.1	36.0	2548.9	16.3	2543.1	6.2	2543.1	6.2	100.5	-0.5		
ELDRSS-5 <>	107	85853	1.5	5.8319	0.5	11.6983	1.4	0.4948	1.3	0.95	2591.4	28.6	2580.6	13.2	2572.0	7.6	2572.0	7.6	100.8	-0.8		
ELDRSS-10 <>	94	186598	1.7	5.7951	0.4	11.7280	1.0	0.4929	0.9	0.92	2583.3	19.3	2582.9	9.2	2582.6	6.5	2582.6	6.5	100.0	0.0		
ELDRSS-160 <>	99	174317	1.6	5.7770	0.3	11.9143	0.7	0.4992	0.6	0.88	2610.3	13.3	2597.7	6.6	2587.8	5.5	2587.8	5.5	100.9	-0.9		
ELDRSS-26 <>	97	233706	1.3	5.7139	0.3	12.1106	1.4	0.5019	1.3	0.98	2621.8	28.9	2613.0	12.9	2606.2	4.8	2606.2	4.8	100.6	-0.6		
ELDRSS-62 <>	36	124177	1.3	5.6648	0.6	12.2916	0.9	0.5050	0.6	0.68	2635.2	13.1	2626.9	8.3	2620.5	10.8	2620.5	10.8	100.6	-0.6		
ELDRSS-137 <>	29	29832	1.0	5.5998	0.6	12.1245	1.5	0.4924	1.4	0.91	2581.1	29.6	2614.1	14.3	2639.7	10.5	2639.7	10.5	97.8	2.2		
ELDRSS-105 <>	111	101319	1.2	5.5549	0.3	12.7736	1.0	0.5146	0.9	0.94	2676.3	20.7	2663.1	9.5	2653.1	5.7	2653.1	5.7	100.9	-0.9		
ELDRSS-71 <>	166	205889	0.7	5.4167	0.3	13.3403	0.7	0.5241	0.7	1.02	2716.5	14.9	2704.0	6.9	2694.7	4.8	2694.7	4.8	100.8	-0.8		
ELDRSS-140 <>	112	102665	1.8	5.3951	0.4	14.0636	6.4	0.5503	6.4	0.90	2826.4	145.7	2754.0	60.6	2701.3	6.7	2701.3	6.7	104.6	-4.6		
ELDRSS-67 <>	37	54762	0.8	5.3887	0.5	13.6905	1.1	0.5351	1.0	0.89	2762.7	22.3	2728.5	10.5	2703.3	8.2	2703.3	8.2	102.2	-2.2		
ELDRSS-131 <>	63	363538	1.1	5.3569	0.4	14.3233	0.8	0.5215	0.7	0.87	2705.6	14.5	2709.9	7.1	2713.1	6.2	2713.1	6.2	99.7	0.3		
ELDRSS-95 <>	87	132426	0.7	5.3448	0.3	13.5930	0.8	0.5269	0.7	0.92	2728.5	16.2	2721.8	7.5	2716.8	5.0	2716.8	5.0	100.4	-0.4		
ELDRSS161 <>	57	27381	0.9	5.3413	0.4	13.7374	0.9	0.5322	0.8	0.89	2750.6	17.6	2731.8	8.4	2717.9	6.0	2717.9	6.0	101.2	-1.2		
ELDRSS-77 <>	32	41093	2.2	5.2395	0.7	14.3539	3.8	0.5455	3.7	0.98	2806.3	85.1	2773.4	36.1	2749.5	10.9	2749.5	10.9	102.1	-2.1		
ELDRSS-20 <>	78	83592	0.7	5.1814	0.3	14.0466	4.4	0.5279	4.4	1.00	2732.4	98.1	2752.9	41.8</								

## U-Pb geochronologic analyses of selected Roberts Mountains allochthon strata

	Isotope ratios									Apparent ages (Ma)										
	U (ppm)	<sup>206</sup> Pb 204Pb	U/Th	<sup>206</sup> Pb* 207Pb*	± (%)	<sup>207</sup> Pb* 235U*	± (%)	<sup>206</sup> Pb* 238U	± (%)	error corr.	<sup>206</sup> Pb* 238U*	± (Ma)	<sup>207</sup> Pb* 235U	± (Ma)	<sup>206</sup> Pb* 207Pb*	± (Ma)	Best age (Ma)	± (Ma)	Conc. (%)	Discord. (%)
<b>Sample: Slaven Chert. Location: Slaven Canyon, Shoshone Range: 0519428 4479302 (NAD 83 UTM 11T)</b>																				
SLAVEN RMA-142	117	109625	1.1	8.9285	0.5	5.1101	1.2	0.3309	1.1	0.91	1842.8	17.1	1837.8	10.0	1832.1	9.1	1832.1	9.1	100.6	-0.6
SLAVEN RMA-186	191	58078	1.2	8.9271	0.3	5.0303	1.8	0.3257	1.7	0.99	1817.5	27.5	1824.4	15.0	1832.4	5.5	1832.4	5.5	99.2	0.8
SLAVEN RMA-131	31	61577	1.5	8.9217	1.5	5.0472	2.0	0.3266	1.4	0.69	1821.8	22.1	1827.3	17.1	1833.5	26.4	1833.5	26.4	99.4	0.6
SLAVEN RMA-171	54	61627	1.6	8.9207	1.0	5.2060	2.5	0.3368	2.3	0.91	1871.4	37.0	1853.6	21.3	1833.7	18.8	1833.7	18.8	102.1	-2.1
SLAVEN RMA-12	86	122215	1.4	8.9190	0.9	5.0822	2.4	0.3287	2.3	0.93	1832.3	36.3	1833.1	20.8	1834.0	16.4	1834.0	16.4	99.9	0.1
SLAVEN RMA-22	128	86826	1.1	8.9172	0.5	5.1138	1.5	0.3307	1.4	0.95	1841.9	22.7	1838.4	12.7	1834.4	8.4	1834.4	8.4	100.4	-0.4
SLAVEN RMA-28	55	59183	1.2	8.9161	0.8	4.8488	1.8	0.3136	1.5	0.88	1758.2	23.6	1793.4	14.8	1834.6	15.3	1834.6	15.3	95.8	4.2
SLAVEN RMA-124	125	111124	1.4	8.9146	0.4	4.6678	2.9	0.3018	2.8	0.99	1700.2	42.5	1761.5	24.0	1834.9	7.5	1834.9	7.5	92.7	7.3
SLAVEN RMA-46	76	120714	1.7	8.9144	0.4	4.9994	1.1	0.3232	1.1	0.94	1805.5	17.0	1819.2	9.7	1835.0	7.1	1835.0	7.1	98.4	1.6
SLAVEN RMA-16	44	29740	1.3	8.9117	1.6	5.1475	2.3	0.3227	1.6	0.69	1851.5	25.5	1844.0	19.4	1835.5	29.7	1835.5	29.7	100.9	-0.9
SLAVEN RMA-52	166	172336	1.4	8.9055	0.3	4.9763	2.3	0.3214	2.3	0.99	1796.6	35.4	1815.3	19.3	1836.8	5.9	1836.8	5.9	97.8	2.2
SLAVEN RMA-195	257	205593	1.7	8.9024	0.3	4.8982	2.3	0.3163	2.3	0.99	1771.4	35.4	1801.9	19.4	1837.4	4.6	1837.4	4.6	96.4	3.6
SLAVEN RMA-158	200	131219	1.6	8.9022	0.3	4.9000	2.4	0.3164	2.4	0.99	1772.0	37.0	1802.3	20.3	1837.5	5.7	1837.5	5.7	96.4	3.6
SLAVEN RMA-102	307	167140	2.2	8.9015	0.2	4.9225	1.7	0.3178	1.7	0.99	1779.0	26.1	1806.1	14.3	1837.6	3.1	1837.6	3.1	96.8	3.2
SLAVEN RMA-197	154	109996	1.7	8.9003	0.3	5.0744	1.8	0.3276	1.8	0.99	1826.5	29.0	1831.8	15.6	1837.9	4.9	1837.9	4.9	99.4	0.6
SLAVEN RMA-93	94	114605	1.5	8.9002	0.6	5.0491	2.2	0.3259	2.1	0.97	1818.6	33.4	1827.6	18.5	1837.9	10.0	1837.9	10.0	99.0	1.0
SLAVEN RMA-122	85	169822	1.2	8.8996	0.8	5.0800	1.5	0.3229	1.2	0.86	1828.2	19.8	1832.8	12.4	1837.9	13.7	1837.9	13.7	99.5	0.5
SLAVEN RMA-112	286	239066	4.9	8.8996	0.3	5.1671	0.9	0.3335	0.8	0.95	1855.4	13.5	1847.2	7.5	1838.0	5.2	1838.0	5.2	100.9	-0.9
SLAVEN RMA-193	71	73587	2.5	8.8988	0.9	5.1146	1.3	0.3301	0.9	0.71	1838.8	15.2	1838.5	11.3	1838.8	17.0	1838.8	17.0	100.0	0.0
SLAVEN RMA-127	44	41862	1.0	8.8983	1.0	5.1307	3.0	0.3311	2.9	0.95	1843.8	45.9	1841.2	25.7	1838.3	17.8	1838.3	17.8	100.3	-0.3
SLAVEN RMA-73	67	134699	1.3	8.8980	0.5	5.0011	1.4	0.3227	1.4	0.94	1803.1	21.5	1819.5	12.2	1838.3	8.6	1838.3	8.6	98.1	1.9
SLAVEN RMA-92	68	65527	1.2	8.8968	0.6	5.0815	1.2	0.3279	1.0	0.84	1828.1	15.4	1833.0	9.8	1838.6	11.5	1838.6	11.5	99.4	0.6
SLAVEN RMA-62	124	133404	1.8	8.8942	0.4	5.0354	2.2	0.3248	2.2	0.99	1813.2	34.5	1825.3	18.8	1839.1	6.9	1839.1	6.9	98.6	1.4
SLAVEN RMA-143	65	125684	1.6	8.8936	1.1	5.0114	1.9	0.3232	1.6	0.83	1805.6	24.9	1821.2	16.2	1839.2	19.4	1839.2	19.4	98.2	1.8
SLAVEN RMA-56	66	67339	1.2	8.8925	1.0	4.9855	2.2	0.3215	2.0	0.90	1797.2	31.0	1816.9	18.7	1839.4	17.8	1839.4	17.8	97.7	2.3
SLAVEN RMA-94	122	66177	1.4	8.8912	0.4	5.1196	1.3	0.3301	1.2	0.96	1839.1	19.6	1839.4	10.8	1839.7	6.4	1839.7	6.4	100.0	0.0
SLAVEN RMA-105	79	57645	1.2	8.8888	0.8	4.9791	1.7	0.3210	1.5	0.89	1794.6	23.9	1815.8	14.4	1840.2	13.8	1840.2	13.8	97.5	2.5
SLAVEN RMA-60	182	217075	1.7	8.8886	0.3	4.4069	1.9	0.2841	1.9	0.99	1612.0	26.5	1713.6	15.6	1840.2	5.6	1840.2	5.6	87.6	12.4
SLAVEN RMA-6	56	54293	1.8	8.8874	1.1	5.1284	2.5	0.3306	2.3	0.90	1841.1	36.8	1840.8	21.6	1840.5	19.8	1840.5	19.8	100.0	0.0
SLAVEN RMA-150	138	141996	1.7	8.8862	0.4	5.1879	2.0	0.3344	2.0	0.98	1859.5	31.6	1850.6	17.0	1840.7	7.4	1840.7	7.4	101.0	-1.0
SLAVEN RMA-80	111	181988	2.1	8.8858	0.4	5.0233	0.8	0.3237	0.7	0.87	1807.9	11.6	1823.3	7.1	1840.7	7.5	1840.7	7.5	98.2	1.8
SLAVEN RMA-75	195	141169	1.6	8.8853	0.3	4.7846	3.0	0.3063	3.0	1.00	1732.5	45.4	1762.2	25.2	1840.9	5.3	1840.9	5.3	94.1	5.9
SLAVEN RMA-32	101	93368	1.7	8.8824	0.5	5.1485	2.5	0.3317	2.4	0.98	1846.5	39.2	1844.2	21.2	1841.5	9.4	1841.5	9.4	100.3	-0.3
SLAVEN RMA-31	74	100498	0.4	8.8815	0.9	4.9953	1.8	0.3218	1.6	0.86	1798.4	24.6	1818.5	15.5	1841.7	17.0	1841.7	17.0	97.6	2.4
SLAVEN RMA-121	125	148722	2.3	8.8814	0.3	5.1081	0.8	0.3290	0.8	0.93	1833.7	12.5	1837.5	7.2	1841.7	5.7	1841.7	5.7	99.6	0.4
SLAVEN RMA-151	103	44551	1.6	8.8810	0.5	5.0733	1.5	0.3268	1.5	0.95	1822.7	23.2	1831.6	13.1	1841.8	8.8	1841.8	8.8	99.0	1.0
SLAVEN RMA-98	91	77128	1.4	8.8803	1.0	5.1371	1.7	0.3309	1.4	0.80	1842.6	21.8	1842.3	14.4	1841.9	18.5	1841.9	18.5	100.0	0.0
SLAVEN RMA-87	74	55717	0.7	8.8791	0.6	5.0334	1.3	0.3241	1.1	0.89	1809.9	17.9	1825.0	10.8	1842.2	10.4	1842.2	10.4	98.2	1.8
SLAVEN RMA-96	127	203826	1.3	8.8777	0.5	5.1169	2.2	0.3295	2.1	0.97	1835.8	34.0	1838.9	18.5	1842.5	8.8	1842.5	8.8	99.6	0.4
SLAVEN RMA-42	104	173022	1.3	8.8771	0.4	5.0378	2.3	0.3243	2.2	0.98	1810.9	35.0	1825.7	19.1	1842.6	7.2	1842.6	7.2	98.3	1.7
SLAVEN RMA-188	36	17569	0.8	8.8770	2.3	5.0687	2.5	0.3263	0.9	0.86	1820.6	14.4	1830.9	21.2	1842.6	42.1	1842.6	42.1	98.8	1.2
SLAVEN RMA-8	71	60018	1.4	8.8758	0.7	5.1571	2.1	0.3320	2.0	0.95	1848.0	32.2	1845.6	18.0	1842.8	12.5	1842.8	12.5	100.3	-0.3
SLAVEN RMA-68	180	132934	2.0	8.8757	0.4	5.0572	2.7	0.3255	2.7	0.99	1816.8	42.1	1829.0	22.8	1842.9	7.3	1842.9	7.3	98.6	1.4
SLAVEN RMA-132	86	64621	1.8	8.8710	0.4	5.1180	1.3	0.3293	1.2	0.94	1834.9	19.3	1839.1	10.9	1843.8	7.6	1843.8	7.6	99.5	0.5
SLAVEN RMA-111	39	46488	1.0	8.8684	0.9	5.1458	1.4	0.3310	1.0	0.73	1843.1	15.8	1843.7	11.5	1844.4	16.8	1844.4	16.8	99.9	0.1
SLAVEN RMA-66	103	195797	1.5	8.8678	0.4	5.1058	1.2	0.3284	1.1	0.94	1830.5	17.6	1837.1	10.0	1844.5	7.0	1844.5	7.0	99.2	0.8
SLAVEN RMA-71	70	92108	0.8	8.8673	1.1	4.9940	1.9	0.3212	1.5	0.81	1795.4	23.8	1818.3	15.9	1844.6	20.1	1844.6	20.1	97.3	2.7
SLAVEN RMA-145	26	14743	1.8	8.8662	1.5	5.1792	2.6	0.3330	2.2	0.82	1853.1	34.8	1849.2	22.3	1844.8	26.9	1844.8	26.9	100.4	-0.4
SLAVEN RMA-126	48	35931	1.5	8.8627	0.8	5.1862	2.0	0.3334	1.8	0.92	1854.6	29.1	1850.4	16.7	1845.5	13.9	1845.5	13.9	100.5	-0.5
SLAVEN RMA-53	59	88254	1.0	8.8626	0.6	4.9532	1.1	0.3184	0.9	0.84	1781.8	14.1	1811.4	9.1	1845.5	10.6	1845.5	10.6	96.5	3.5
SLAVEN RMA-108	64	67789	1.3	8.8616	0.5	5.0687	1.5	0.3256	1.4	0.95	1817.2	22.8	1830.6	12.9	1845.7	9.0	1845.7	9.0	98.5	1.5
SLAVEN RMA-140	49	59245	1.3	8.8592	1.8	5.0364	1.9	0.3236	0.5	0.25	1807.3	7.4	1825.5	15.8	1846.2	32.7	1846.2	32.7	97.9	2.1
SLAVEN RMA-13	30	5021	0.8	8.8573	2.4	5.1960	3.2	0.3338	2.2	0.67	1856.7	35.0	1852.0	27.4	1846.6	43.0	1846.6	43.0	100.5	-0.5
SLAVEN RMA-7	75	120099	1.1	8.8572	0.6	5.0739	2.0	0.3259	1.9	0.95	1818.7	29.6	1831.7	16.7	1846.6	11.6	1846.6	11.6	98.5	1.5
SLAVEN RMA-1	38	99655	1.6	8.8524	0.9	5.1692	2.2	0.3319	2.0	0.92	1847.5	32.5	1847.6	18.8	1847.6	16.1	1847.6	16.1	100.0	0.0
SLAVEN RMA-20	41	46129	1.4	8.8468	1.2	5.1637	3.4	0.3313	3.2	0.94	1844.8	50.7	1846.6	28.7	1848.8	21.3	1848.8	21.3	99.8	0.2
SLAVEN RMA-67	63	87611	1.5	8.8434	0.7	5.1530	3.8	0.3305	3.8	0.99	1840.8	60.2	1844.9	32.5	1849.5	11.8	1849.5	11.8	99.5	0.5
SLAVEN RMA-110	34	2047	0.2																	



### U-Pb geochronologic analyses of selected Roberts Mountains allochthon strata

	Isotope ratios										Apparent ages (Ma)									
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc.	Discord
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)
<b>Sample: Slaven Chert. Location: Slaven Canyon, Shoshone Range: 0519428 4479302 (NAD 83 UTM 11T)</b>																				
SLAVEN RMA-47	38	23566	0.5	8.5290	0.6	5.3984	3.4	0.3339	3.3	0.98	1857.4	53.4	1884.6	28.9	1914.7	11.2	1914.7	11.2	97.0	3.0
SLAVEN RMA-125	392	88729	4.1	8.5172	0.2	5.2718	4.6	0.3257	4.6	1.00	1817.3	72.5	1864.3	39.1	1917.1	4.0	1917.1	4.0	94.8	5.2
SLAVEN RMA-174	154	193641	0.4	8.5114	0.4	4.7403	2.2	0.2926	2.2	0.99	1654.6	31.9	1774.4	18.6	1918.4	6.8	1918.4	6.8	86.3	13.7
SLAVEN RMA-138	53	100317	1.5	8.5067	0.9	5.4901	2.5	0.3387	2.4	0.93	1880.5	38.7	1899.0	21.8	1919.4	16.2	1919.4	16.2	98.0	2.0
SLAVEN RMA-37	58	74694	1.0	8.5025	0.9	5.4795	4.5	0.3379	4.4	0.98	1876.6	72.1	1897.4	38.8	1920.2	15.7	1920.2	15.7	97.7	2.3
SLAVEN RMA-25	63	72016	1.0	8.4971	0.5	5.6448	1.7	0.3479	1.7	0.95	1924.4	27.8	1923.0	15.1	1921.4	9.4	1921.4	9.4	100.2	-0.2
SLAVEN RMA-109	45	46541	1.2	8.4793	0.8	5.6046	1.4	0.3447	1.2	0.85	1909.1	20.1	1916.8	12.4	1925.1	13.5	1925.1	13.5	99.2	0.8
SLAVEN RMA-156	89	71249	1.0	8.4691	0.4	5.6200	1.0	0.3452	0.9	0.90	1911.7	14.6	1919.2	8.4	1927.3	7.7	1927.3	7.7	99.2	0.8
SLAVEN RMA-18	130	163380	1.3	8.4610	0.3	5.6023	2.5	0.3438	2.4	0.99	1904.8	40.3	1916.4	21.2	1929.0	5.8	1929.0	5.8	98.7	1.3
SLAVEN RMA-119	64	67698	0.7	8.4166	0.4	5.7846	2.2	0.3531	2.1	0.98	1949.4	35.4	1944.1	18.6	1938.4	8.0	1938.4	8.0	100.6	-0.6
SLAVEN RMA-30	187	167062	1.4	8.4066	0.3	5.6973	1.7	0.3474	1.7	0.98	1922.0	28.4	1931.0	15.0	1940.5	5.9	1940.5	5.9	99.0	1.0
SLAVEN RMA-70	50	60609	0.5	8.3967	1.0	5.6820	1.9	0.3460	1.6	0.84	1915.6	26.3	1928.6	16.2	1942.7	18.2	1942.7	18.2	98.6	1.4
SLAVEN RMA-106	37	49729	1.2	8.3881	0.8	5.5876	1.2	0.3399	0.9	0.74	1886.3	14.5	1914.2	10.4	1944.5	14.6	1944.5	14.6	97.0	3.0
SLAVEN RMA-21	102	137776	1.3	8.3844	0.5	5.8493	3.1	0.3557	3.1	0.99	1961.7	52.2	1953.7	27.1	1943.3	8.9	1943.3	8.9	100.8	-0.8
SLAVEN RMA-81	62	66867	0.5	8.3192	0.5	5.6688	1.1	0.3420	0.9	0.87	1896.2	15.1	1926.6	9.2	1959.2	9.5	1959.2	9.5	96.8	3.2
SLAVEN RMA-163	99	106591	0.8	8.2100	0.7	4.7029	4.2	0.2800	4.1	0.99	1591.5	58.4	1767.8	35.2	1982.8	12.7	1982.8	12.7	80.3	19.7
SLAVEN RMA-134	81	90875	1.9	8.1417	0.6	6.0998	2.7	0.3602	2.6	0.97	1983.1	44.3	1990.2	23.2	1997.6	10.7	1997.6	10.7	90.3	0.7
SLAVEN RMA-104	31	32749	0.9	8.0968	1.6	6.0550	2.5	0.3556	1.9	0.77	1961.1	32.5	1983.8	21.7	2007.5	28.2	2007.5	28.2	97.7	2.3
SLAVEN RMA-72	92	82821	2.2	7.9056	0.6	6.4055	2.0	0.3673	1.9	0.95	2016.5	32.6	2033.0	17.4	2049.8	10.6	2049.8	10.6	98.4	1.6
SLAVEN RMA-91	101	73045	0.9	7.8573	0.4	6.5378	4.8	0.3726	4.8	1.00	2041.5	83.5	2051.0	42.2	2060.6	7.6	2060.6	7.6	99.1	0.9
SLAVEN RMA-59	24	25752	0.9	7.8338	1.5	6.5396	1.9	0.3716	1.2	0.84	2036.7	21.0	2051.2	16.6	2065.9	25.7	2065.9	25.7	98.6	1.4
SLAVEN RMA-107	41	78567	1.3	7.8322	1.3	6.5547	1.5	0.3723	0.8	0.52	2040.4	13.7	2053.3	13.2	2066.2	22.6	2066.2	22.6	98.7	1.3
SLAVEN RMA-39	65	100705	1.2	7.8145	0.7	6.6290	1.6	0.3757	1.4	0.90	2056.2	25.1	2063.2	13.9	2070.2	11.9	2070.2	11.9	99.3	0.7
SLAVEN RMA-45	56	53937	0.8	7.7924	0.7	6.5721	1.5	0.3714	1.4	0.91	2036.1	24.5	2055.6	13.7	2075.2	11.5	2075.2	11.5	98.1	1.9
SLAVEN RMA-183	43	107679	0.8	7.7841	1.1	6.3797	1.9	0.3602	1.5	0.79	1983.0	25.2	2029.5	16.4	2077.1	20.1	2077.1	20.1	95.5	4.5
SLAVEN RMA-135	55	44286	0.5	7.7833	0.9	6.7280	4.3	0.3798	4.1	0.97	2075.3	73.5	2076.3	37.6	2077.3	16.7	2077.3	16.7	99.9	0.1
SLAVEN RMA-157	88	64081	0.7	7.7802	0.4	6.6599	1.2	0.3758	1.2	0.93	2056.6	20.3	2067.3	10.9	2078.0	7.8	2078.0	7.8	99.0	1.0
SLAVEN RMA-141	70	70082	1.0	7.7762	0.6	6.7056	1.1	0.3782	0.9	0.85	2067.8	16.3	2073.4	9.6	2079.1	10.1	2079.1	10.1	99.5	0.5
SLAVEN RMA-196	45	83740	0.7	7.7750	0.7	6.7347	2.4	0.3798	2.3	0.96	2075.2	40.2	2077.2	20.8	2079.1	11.5	2079.1	11.5	99.8	0.2
SLAVEN RMA-24	72	108215	1.6	7.7602	0.6	6.8340	1.8	0.3846	1.7	0.95	2097.9	30.0	2090.1	15.7	2082.5	9.9	2082.5	9.9	100.7	-0.7
SLAVEN RMA-155	47	15484	1.4	7.7568	1.5	6.8005	1.8	0.3713	1.1	0.58	2035.6	18.6	2059.4	16.3	2083.3	26.6	2083.3	26.6	97.7	2.3
SLAVEN RMA-180	30	49020	0.5	7.7557	1.1	6.7514	2.4	0.3798	2.1	0.88	2075.2	37.1	2079.4	21.0	2083.5	20.0	2083.5	20.0	99.6	0.4
SLAVEN RMA-36	79	100177	0.8	7.7496	0.6	6.6402	2.4	0.3732	2.4	0.97	2044.5	41.2	2064.7	21.4	2084.9	10.3	2084.9	10.3	98.1	1.9
SLAVEN RMA-35	88	188835	2.2	7.7466	0.5	6.7425	1.3	0.3788	1.2	0.92	2070.8	20.5	2078.2	11.1	2085.6	8.6	2085.6	8.6	99.3	0.7
SLAVEN RMA-175	49	91812	1.4	7.7109	1.2	6.9796	2.5	0.3903	2.2	0.88	2124.4	39.9	2108.8	22.4	2093.7	21.4	2093.7	21.4	101.5	-1.5
SLAVEN RMA-166	64	84675	0.6	7.6457	0.5	6.9071	0.9	0.3830	0.7	0.84	2090.3	13.1	2099.6	7.8	2108.6	8.4	2108.6	8.4	99.1	0.9
SLAVEN RMA-61	154	57196	1.0	7.2775	0.5	7.3625	3.3	0.3886	3.3	0.99	2116.3	59.4	2156.4	29.8	2194.8	8.9	2194.8	8.9	96.4	3.6
SLAVEN RMA-69	60	23660	0.7	7.0508	1.3	7.9687	2.5	0.4075	2.2	0.86	2203.5	41.1	2227.5	23.0	2249.6	22.2	2249.6	22.2	97.9	2.1
SLAVEN RMA-58	120	176689	2.4	6.9650	0.3	8.1912	1.2	0.4138	1.1	0.96	2232.2	21.3	2252.4	10.6	2270.8	5.4	2270.8	5.4	98.3	1.7
SLAVEN RMA-189	39	42299	0.8	6.8352	1.1	8.7989	1.6	0.4362	1.2	0.74	2333.6	23.1	2317.4	14.5	2303.1	18.4	2303.1	18.4	101.3	-1.3
SLAVEN RMA-83	62	132694	0.7	6.8233	1.0	8.5611	1.2	0.4237	0.7	0.58	2277.1	13.9	2292.4	11.3	2306.1	17.3	2306.1	17.3	98.7	1.3
SLAVEN RMA-181	155	70268	1.0	6.6404	0.5	7.7200	1.7	0.3718	1.6	0.96	2037.9	28.4	2198.9	15.2	2352.7	7.8	2352.7	7.8	86.6	13.4
SLAVEN RMA-159	79	96153	0.9	6.5720	0.6	8.5122	3.9	0.4057	3.9	0.99	2195.4	72.0	2287.2	35.6	2370.3	9.8	2370.3	9.8	92.6	7.4
SLAVEN RMA-95	96	246129	2.6	6.5425	0.5	9.0011	2.1	0.4271	2.0	0.97	2292.7	38.5	2338.1	18.8	2378.0	8.8	2378.0	8.8	96.4	3.6
SLAVEN RMA-88	16	24833	0.8	6.4506	2.1	8.9540	6.7	0.4189	6.3	0.95	2255.5	120.4	2333.3	61.0	2402.1	36.1	2402.1	36.1	93.9	6.1
SLAVEN RMA-54	144	180847	1.2	6.2924	0.3	9.7455	1.5	0.4448	1.4	0.98	2371.9	28.6	2411.0	13.6	2444.2	5.5	2444.2	5.5	97.0	3.0
SLAVEN RMA-23	75	115115	2.2	6.2474	0.6	10.2028	3.2	0.4623	3.1	0.98	2449.7	63.7	2453.3	29.5	2456.0	10.7	2456.0	10.7	99.7	0.3
SLAVEN RMA-50	108	240510	1.6	5.9609	0.4	9.9722	2.2	0.4311	2.1	0.99	2310.8	41.5	2432.2	20.0	2536.4	6.0	2536.4	6.0	91.1	8.9
SLAVEN RMA-41	13	13259	0.4	5.8816	1.6	11.2177	2.8	0.4785	2.3	0.82	2520.8	47.9	2541.4	26.1	2557.8	26.8	2557.8	26.8	98.6	1.4
SLAVEN RMA-29	51	115551	1.4	5.8648	0.6	11.3669	2.1	0.4835	2.0	0.95	2542.5	42.1	2553.7	19.6	2562.6	10.4	2562.6	10.4	99.2	0.8
SLAVEN RMA-90	114	88337	1.6	5.8340	0.5	11.2707	1.0	0.4769	0.9	0.87	2513.7	18.0	2545.8	9.3	2571.4	8.2	2571.4	8.2	97.8	2.2
SLAVEN RMA-153	144	239479	2.4	5.8320	0.1	11.2377	1.4	0.4753	1.4	1.00	2506.9	29.1	2543.1	13.1	2572.0	1.9	2572.0	1.9	97.5	2.5
SLAVEN RMA-169	215	265095	2.9	5.7863	0.4	10.9722	2.5	0.4605	2.4	0.98	2441.6	49.4	2520.8	23.0	2585.2	7.4	2585.2	7.4	94.4	5.6
SLAVEN RMA-190	66	13637	0.5	5.7612	0.5	11.3294	1.0	0.4734	0.9	0.87	2498.4	17.8	2550.6	9.2	2592.4	8.1	2592.4	8.1	96.4	3.6
SLAVEN RMA-103	86	167372	1.0	5.7063	0.4	11.8294	1.6	0.4896	1.6	0.98	2568.8	33.3	2591.0	15.1	2608.4	5.9	2608.4	5.9	98.5	1.5
SLAVEN RMA-136	40	31484	0.9	5.6583	0.6	11.8790	1.7	0.4875	1.6	0.94	2559.8	33.3	2594.9	15.7	2622.4	9.5	2622.4	9.5	97.6	2.4
SLAVEN RMA-85	22	29092	0.6	5.6581	0.6	11.5086	2.0	0.4723	1.9	0.96	2493.5	39.9	2565.3	18.8	2622.5	9.2	2622.5	9.2	95.1	4.9
SLAVEN RMA-																				

## Notes:

1. Analyses with >10% uncertainty (1-sigma) in  $^{206}\text{Pb}/^{238}\text{U}$  age are not included.
2. Analyses with >10% uncertainty (1-sigma) in  $^{206}\text{Pb}/^{207}\text{Pb}$  age are not included, unless  $^{206}\text{Pb}/^{238}\text{U}$  age is <500 Ma.
3. Best age is determined from  $^{206}\text{Pb}/^{238}\text{U}$  age for analyses with  $^{206}\text{Pb}/^{238}\text{U}$  age <1000 Ma and from  $^{206}\text{Pb}/^{207}\text{Pb}$  age for analyses with  $^{206}\text{Pb}/^{238}\text{U}$  age > 1000 Ma.
4. Concordance is based on  $^{206}\text{Pb}/^{238}\text{U}$  age /  $^{206}\text{Pb}/^{207}\text{Pb}$  age. Value is not reported for  $^{206}\text{Pb}/^{238}\text{U}$  ages <500 Ma because of large uncertainty in  $^{206}\text{Pb}/^{207}\text{Pb}$  age.
5. Discordance is 100% - concordance.
6. Analyses with  $^{206}\text{Pb}/^{238}\text{U}$  age > 500 Ma and with >20% discordance (<80% concordance) are not included.
7. Analyses with  $^{206}\text{Pb}/^{238}\text{U}$  age > 500 Ma and with >5% reverse discordance (<105% concordance) are not included.
8. All uncertainties are reported at the 1-sigma level, and include only measurement errors.
9. External (systematic) errors are shown as  $^{206}\text{Pb}/^{238}\text{U}$  uncertainty,  $^{206}\text{Pb}/^{207}\text{Pb}$  uncertainty to the right of each sample (in %, at 2-sigma level).
10. Analyses conducted by LA-MC-ICPMS, as described by Gehrels et al. (2008).
11. U concentration and U/Th are calibrated relative to Sri Lanka zircon standard and are accurate to ~20%.
12. Common Pb correction is from measured  $^{204}\text{Pb}$  with common Pb composition interpreted from Stacey and Kramers (1975).
13. Uncertainties of 1.5 for  $^{206}\text{Pb}/^{204}\text{Pb}$ , 0.3 for  $^{207}\text{Pb}/^{204}\text{Pb}$ , and 2.0 for  $^{208}\text{Pb}/^{204}\text{Pb}$  are applied to common Pb composition.
14. U/Pb and  $^{206}\text{Pb}/^{207}\text{Pb}$  fractionation is calibrated relative to fragments of a large Sri Lanka zircon of  $563.5 \pm 3.2$  Ma (2-sigma).
15. U decay constants and composition as follows:  $^{235}\text{U} = 9.8485 \times 10^{-10}$ ,  $^{238}\text{U} = 1.55125 \times 10^{-10}$ ,  $^{238}\text{U}/^{235}\text{U} = 137.88$ .
16. Weighted mean and concordia plots determined with Isoplot (Ludwig, 2008).
17. Analytical methods as described by Gehrels and Pecha (2014).

## APPENDIX D

### Hafnium isotope data of selected Roberts Mountains allochthon strata

Table notes are at the end of the appendix.

	$(^{176}\text{Yb} + ^{176}\text{Lu}) / ^{177}\text{Hf}$ (%)	Volts Hf	$^{176}\text{Hf}/^{177}\text{Hf}$	$\pm (1\sigma)$	$^{176}\text{Lu}/^{177}\text{Hf}$	$^{176}\text{Hf}/^{177}\text{Hf}$ (T)	E-Hf (0)	E-Hf (0) $\pm (1\sigma)$	E-Hf (T)	Age (Ma)
<b>Sample: Snow Canyon Formation. Location: Snow Canyon, Independence Mountains; 0579760 4585698 (NAD 83 UTM 11T)</b>										
Snow-CYN-GG2000SPE-9	6.9	5.4	0.281448	0.000016	0.00041	0.281434	-47.3	0.6	-6.6	1826
Snow-CYAN-GG2000SPE-30	7.5	5.3	0.281335	0.000017	0.00044	0.281320	-51.3	0.6	-10.6	1827
Snow-CYAN-GG2000SPE-38	9.2	4.8	0.281420	0.000021	0.00055	0.281401	-48.3	0.7	-7.7	1829
Snow-CYAN-GG2000SPE-27	9.1	4.5	0.281408	0.000020	0.00055	0.281389	-48.7	0.7	-8.1	1832
Snow-CYN-GG2000SPE-14	7.5	5.1	0.281438	0.000021	0.00046	0.281422	-47.6	0.7	-6.9	1832
Snow-CYAN-GG2000SPE-36	14.4	4.0	0.281423	0.000022	0.00084	0.281394	-48.2	0.8	-7.9	1832
Snow-CYAN-GG2000SPE-58	6.8	5.5	0.281443	0.000020	0.00041	0.281429	-47.5	0.7	-6.4	1842
Snow-CYN-GG2000SPE-45	11.1	5.5	0.281453	0.000018	0.00063	0.281431	-47.1	0.6	-6.3	1844
Snow-CYN-GG2000SPE-48	9.0	4.8	0.281478	0.000022	0.00051	0.281460	-46.2	0.8	-5.2	1845
Snow-CYN-GG2000SPE-53L	17.5	4.7	0.281499	0.000022	0.00098	0.281465	-45.5	0.8	-5.1	1845
Snow-CYN-GG2000SPE-53R	7.7	5.0	0.281422	0.000021	0.00045	0.281406	-48.2	0.8	-7.2	1845
Snow-CYAN-GG2000SPE-195	10.9	5.5	0.281532	0.000019	0.00067	0.281509	-44.3	0.7	-3.4	1849
Snow-CYAN-GG2000SPE-181	15.9	4.5	0.281463	0.000020	0.00090	0.281431	-46.7	0.7	-6.1	1849
Snow-CYAN-GG2000SPE-186	6.9	4.5	0.281420	0.000019	0.00040	0.281406	-48.3	0.7	-7.0	1851
Snow-CYAN-GG2000SPE-57	7.0	5.2	0.281444	0.000016	0.00045	0.281428	-47.4	0.6	-6.2	1853
Snow-CYAN-GG2000SPE-190	12.2	4.2	0.281542	0.000025	0.00071	0.281517	-44.0	0.9	-2.9	1857
Snow-CYAN-GG2000SPE-129	2.7	5.1	0.281464	0.000023	0.00024	0.281455	-46.7	0.8	-5.1	1857
Snow-CYN-GG2000SPE-198	9.6	5.1	0.281831	0.000015	0.00060	0.281810	-33.7	0.5	7.8	1870
Snow-CYAN-GG2000SPE-184	6.4	5.2	0.281363	0.000021	0.00039	0.281349	-50.3	0.7	-8.3	1884
Snow-CYAN-GG2000SPE-193	7.0	5.5	0.281426	0.000023	0.00041	0.281411	-48.1	0.8	-6.0	1886
Snow-CYN-GG2000SPE-44	3.9	6.9	0.281816	0.000017	0.00034	0.281804	-41.3	0.6	1.2	1905
Snow-CYN-GG2000SPE-24	7.3	4.8	0.281465	0.000022	0.00043	0.281449	-46.7	0.8	-4.2	1906
Snow-CYN-GG2000SPE-12	3.4	5.7	0.281421	0.000022	0.00026	0.281411	-48.2	0.8	-5.4	1913
Snow-CYN-GG2000SPE-42	9.5	5.8	0.281502	0.000018	0.00055	0.281482	-45.4	0.6	-2.8	1917
Snow-CYAN-GG2000SPE-60	11.9	4.6	0.281362	0.000020	0.00065	0.281338	-50.3	0.7	-7.7	1924
Snow-CYN-GG2000SPE-199	6.7	4.7	0.281457	0.000019	0.00039	0.281443	-47.0	0.7	-4.0	1925
Snow-CYN-GG2000SPE-51	14.9	5.7	0.281261	0.000022	0.00084	0.281230	-53.9	0.8	-11.5	1926
Snow-CYN-GG2000SPE-16	5.8	5.4	0.281410	0.000016	0.00035	0.281397	-48.6	0.6	-5.6	1928
Snow-CYAN-GG2000SPE-137	10.7	4.0	0.281491	0.000020	0.00072	0.281463	-45.8	0.7	-1.2	2016
Snow-CYN-GG2000SPE-55	11.4	4.8	0.281497	0.000020	0.00067	0.281471	-45.5	0.7	0.1	2056
Snow-CYAN-GG2000SPE-180	8.2	5.0	0.281417	0.000021	0.00051	0.281397	-48.4	0.7	-2.6	2058
Snow-CYN-GG2000SPE-174	12.3	5.2	0.281450	0.000015	0.00077	0.281420	-47.2	0.5	-1.5	2070
Snow-CYAN-GG2000SPE-168	7.2	5.1	0.281438	0.000020	0.00045	0.281420	-47.6	0.7	-1.4	2072
Snow-CYN-GG2000SPE-46	8.9	6.4	0.281114	0.000015	0.00051	0.281089	-59.1	0.5	-2.6	2524
Snow-CYN-GG2000SPE-25	19.9	4.6	0.281254	0.000022	0.00120	0.281191	-54.2	0.8	5.0	2695
Snow-CYN-GG2000SPE-13	3.6	5.8	0.281007	0.000019	0.00022	0.280995	-62.9	0.7	-1.8	2704
Snow-CYN-GG2000SPE-41	7.6	5.3	0.281060	0.000015	0.00045	0.281036	-61.0	0.5	-0.3	2705
Snow-CYN-GG2000SPE-165	4.5	4.3	0.281052	0.000025	0.00027	0.281038	-61.3	0.9	0.2	2722
Snow-CYAN-GG2000SPE-34	3.0	5.3	0.280916	0.000022	0.00022	0.280905	-66.1	0.8	-4.4	2728
Snow-CYAN-GG2000SPE-192	8.6	3.5	0.281127	0.000023	0.00049	0.281102	-58.6	0.8	2.6	2731
Snow-CYN-GG2000SPE-26	12.0	5.3	0.281035	0.000019	0.00072	0.280997	-61.9	0.7	-0.9	2739
Snow-CYAN-GG2000SPE-123	5.9	4.7	0.281031	0.000028	0.00038	0.281010	-62.0	1.0	2.0	2843
<b>Sample: McAfee Quartzite. Location: McAfee Creek, Independence Mountains; 0590637 4599583 (NAD 83 UTM 11T)</b>										
McAFEE-97SF1-GG2000SPE-196	8.5	4.7	0.281278	0.000021	0.00054	0.281260	-53.3	0.7	-13.3	1803
McAFEE-97SF1-GG2000SPE-138	7.6	4.3	0.281454	0.000020	0.00045	0.281439	-47.1	0.7	-6.7	1815
McAFEE-97SF1-GG2000SPE-24	5.4	4.8	0.281485	0.000019	0.00034	0.281473	-46.0	0.7	-5.1	1830
McAFEE-97SF1-GG2000SPE-32	18.5	5.5	0.281386	0.000019	0.00110	0.281347	-49.5	0.7	-9.5	1833
McAFEE-97SF1-GG2000SPE-29	5.9	5.3	0.281438	0.000018	0.00037	0.281423	-47.7	0.6	-6.8	1834
McAFEE-97SF1-GG2000SPE-182	10.3	4.1	0.281427	0.000022	0.00062	0.281405	-48.0	0.8	-7.4	1835
McAFEE-97SF1-GG2000SPE-4	12.4	4.0	0.281818	0.000027	0.00074	0.281792	-34.2	1.0	6.3	1835
McAFEE-97SF1-GG2000SPE-199	10.6	4.7	0.281401	0.000017	0.00062	0.281380	-48.9	0.6	-8.3	1835
McAFEE-97SF1-GG2000SPE-1	11.2	4.0	0.281523	0.000021	0.00067	0.281499	-44.6	0.8	-4.0	1837
McAFEE-97SF1-GG2000SPE-171	7.1	5.1	0.281428	0.000019	0.00044	0.281413	-48.0	0.7	-7.1	1838
McAFEE-97SF1-GG2000SPE-10	6.2	5.1	0.281439	0.000022	0.00038	0.281426	-47.6	0.8	-6.6	1838
McAFEE-97SF1-GG2000SPE-2	11.4	4.4	0.281482	0.000027	0.00069	0.281458	-46.1	1.0	-5.3	1844
McAFEE-97SF1-GG2000SPE-181	5.7	5.2	0.281417	0.000016	0.00035	0.281405	-48.4	0.5	-7.2	1844
McAFEE-97SF1-GG2000SPE-9	8.2	3.8	0.281498	0.000024	0.00055	0.281479	-45.5	0.8	-4.5	1846
McAFEE-97SF1-GG2000SPE-7	7.0	4.6	0.281586	0.000024	0.00042	0.281571	-42.4	0.8	-1.3	1846
McAFEE-97SF1-GG2000SPE-21	9.2	4.6	0.281437	0.000021	0.00056	0.281418	-47.7	0.7	-6.7	1847
McAFEE-97SF1-GG2000SPE-15	15.7	4.6	0.281415	0.000021	0.00092	0.281383	-48.5	0.7	-7.9	1847
McAFEE-97SF1-GG2000SPE-179	11.1	4.2	0.281251	0.000020	0.00074	0.281225	-54.2	0.7	-13.5	1849
McAFEE-97SF1-GG2000SPE-174	7.0	4.4	0.281468	0.000022	0.00042	0.281453	-46.6	0.8	-5.2	1857
McAFEE-97SF1-GG2000SPE-186	19.0	3.4	0.281450	0.000030	0.00116	0.281409	-47.2	1.1	-6.7	1862
McAFEE-97SF1-GG2000SPE-188	10.4	5.7	0.281584	0.000019	0.00063	0.281562	-42.5	0.7	-1.2	1863
McAFEE-97SF1-GG2000SPE-198	6.8	3.8	0.281853	0.000021	0.00040	0.281839	-33.0	0.7	8.7	1866
McAFEE-97SF1-GG2000SPE-84	9.5	5.9	0.281276	0.000019	0.00056	0.281255	-53.4	0.7	-11.6	1884
McAFEE-97SF1-GG2000SPE-16	17.4	4.5	0.281507	0.000022	0.00102	0.281470	-45.2	0.8	-3.4	1910



### Hafnium isotope data of selected Roberts Mountains allochthon strata

	$^{176}\text{Yb} + ^{176}\text{Lu} / ^{177}\text{Hf}$ (%)	Volts Hf	$^{176}\text{Hf} / ^{177}\text{Hf}$	$\pm$ (1 $\sigma$ )	$^{176}\text{Lu} / ^{177}\text{Hf}$	$^{176}\text{Hf} / ^{177}\text{Hf}$ (T)	E-Hf (O)	E-Hf (O) $\pm$ (1 $\sigma$ )	E-Hf (T)	Age (Ma)
<b>Sample: McAfee Quartzite. Location: McAfee Creek, Independence Mountains; 0590637 4599583 (NAD 83 UTM 11T)</b>										
MCAFee-97SF1-GG2000SPE-197	5.2	5.5	0.281340	0.000020	0.00031	0.281328	-51.1	0.7	-8.3	1913
MCAFee-97SF1-GG2000SPE-175	5.9	3.7	0.281433	0.000026	0.00037	0.281419	-47.8	0.9	-5.0	1917
MCAFee-97SF1-GG2000SPE-12	2.0	4.9	0.281366	0.000018	0.00012	0.281362	-50.2	0.6	-7.0	1917
MCAFee-97SF1-GG2000SPE-176	13.4	4.8	0.281492	0.000022	0.00077	0.281464	-45.7	0.8	-3.4	1918
MCAFee-97SF1-GG2000SPE-166	6.7	5.6	0.281445	0.000020	0.00040	0.281430	-47.4	0.7	-4.5	1922
MCAFee-97SF1-GG2000SPE-187	8.0	5.5	0.281353	0.000019	0.00047	0.281335	-50.6	0.7	-7.7	1930
MCAFee-97SF1-GG2000SPE-20	8.0	4.5	0.281419	0.000024	0.00046	0.281402	-48.3	0.9	-5.3	1933
MCAFee-97SF1-GG2000SPE-172	3.9	4.9	0.281463	0.000019	0.00023	0.281455	-46.7	0.7	-3.3	1935
MCAFee-97SF1-GG2000SPE-169	6.4	4.9	0.281483	0.000022	0.00042	0.281467	-46.1	0.8	-2.8	1939
MCAFee-97SF1-GG2000SPE-17	9.0	5.1	0.281481	0.000026	0.00054	0.281460	-46.1	0.9	-0.1	2070
MCAFee-97SF1-GG2000SPE-159	7.3	5.8	0.281462	0.000016	0.00048	0.281443	-46.8	0.6	-0.4	2079
MCAFee-97SF1-GG2000SPE-157	10.1	5.1	0.281479	0.000018	0.00065	0.281453	-46.2	0.6	0.2	2091
MCAFee-97SF1-GG2000SPE-144	18.4	6.1	0.281458	0.000017	0.00111	0.281414	-46.9	0.6	-1.2	2091
MCAFee-97SF1-GG2000SPE-161	13.6	5.1	0.281399	0.000019	0.00083	0.281365	-49.0	0.7	-2.8	2095
MCAFee-97SF1-GG2000SPE-109	7.6	4.6	0.280890	0.000018	0.00046	0.280866	-67.0	0.6	-7.3	2665
MCAFee-97SF1-GG2000SPE-189	7.1	4.5	0.281061	0.000019	0.00047	0.281036	-61.0	0.7	-0.6	2694
MCAFee-97SF1-GG2000SPE-5	5.4	4.8	0.280917	0.000028	0.00035	0.280899	-66.1	1.0	-5.4	2697
<b>Sample: lower Vinini Formation. Location: Petes Summit, Toquima Range; 0518089 4337111 (NAD 83 UTM 11T)</b>										
Low-Vinini-97SF12-GG2000SPE-21	15.1	3.0	0.282370	0.000031	0.00092	0.282362	-14.7	1.1	-4.3	476
Low-Vinini-97SF12-GG2000SPE-19	19.8	3.1	0.282504	0.000027	0.00116	0.282493	-10.0	1.0	0.4	483
Low-Vinini-97SF12-GG2000SPE-165	47.6	3.7	0.282535	0.000030	0.00278	0.282510	-8.8	1.1	1.1	486
Low-Vinini-97SF12-GG2000SPE-22	11.9	5.6	0.282469	0.000021	0.00066	0.282463	-11.2	0.8	-0.4	493
Low-Vinini-97SF12-GG2000SPE-20	12.7	5.3	0.282504	0.000017	0.00070	0.282497	-10.0	0.6	0.8	494
Low-Vinini-97SF12-GG2000SPE-176	22.5	4.3	0.282514	0.000020	0.00127	0.282502	-9.6	0.7	1.0	495
Low-Vinini-97SF12-GG2000SPE-190	17.7	4.4	0.282508	0.000019	0.00100	0.282499	-9.8	0.7	0.9	495
Low-Vinini-97SF12-GG2000SPE-10	19.2	5.8	0.281800	0.000016	0.00114	0.281789	-34.8	0.6	-24.2	496
Low-Vinini-97SF12-GG2000SPE-35	28.9	6.1	0.281903	0.000022	0.00178	0.281886	-31.2	0.8	-20.7	497
Low-Vinini-97SF12-GG2000SPE-35	29.9	4.8	0.282567	0.000024	0.00163	0.282552	-7.7	0.9	2.9	497
Low-Vinini-97SF12-GG2000SPE-45	25.5	5.3	0.282253	0.000017	0.00133	0.282225	-18.8	0.6	5.2	1117
Low-Vinini-97SF12-GG2000SPE-2	36.3	4.6	0.282133	0.000023	0.00216	0.282074	-23.1	0.8	6.9	1422
Low-Vinini-97SF12-GG2000SPE-101	18.3	4.5	0.282014	0.000020	0.00105	0.281985	-27.3	0.7	4.0	1433
Low-Vinini-97SF12-GG2000SPE-121	14.0	4.8	0.281986	0.000018	0.00082	0.281963	-28.3	0.7	3.2	1434
Low-Vinini-97SF12-GG2000SPE-27	8.8	4.7	0.282086	0.000022	0.00053	0.282072	-24.7	0.8	7.1	1435
Low-Vinini-97SF12-GG2000SPE-87	15.0	4.8	0.281975	0.000021	0.00085	0.281952	-28.7	0.7	3.0	1444
Low-Vinini-97SF12-GG2000SPE-28	16.1	5.5	0.281936	0.000016	0.00098	0.281905	-30.0	0.6	7.1	1694
Low-Vinini-97SF12-GG2000SPE-172	25.6	4.3	0.281939	0.000026	0.00153	0.281890	-29.9	0.9	6.6	1694
Low-Vinini-97SF12-GG2000SPE-41	12.8	5.3	0.281905	0.000016	0.00083	0.281878	-31.1	0.6	6.5	1708
Low-Vinini-97SF12-GG2000SPE-43	19.6	5.9	0.281827	0.000018	0.00114	0.281790	-33.9	0.6	3.5	1713
Low-Vinini-97SF12-GG2000SPE-23	24.3	5.5	0.281764	0.000022	0.00139	0.281719	-36.1	0.8	1.0	1716
Low-Vinini-97SF12-GG2000SPE-8	12.1	4.6	0.281844	0.000019	0.00071	0.281820	-33.3	0.7	4.6	1717
Low-Vinini-97SF12-GG2000SPE-194	16.8	6.5	0.281822	0.000019	0.00100	0.281789	-34.0	0.7	3.7	1724
Low-Vinini-97SF12-GG2000SPE-3	4.9	4.5	0.281870	0.000020	0.00032	0.281859	-32.4	0.7	8.4	1733
Low-Vinini-97SF12-GG2000SPE-9	11.5	5.0	0.281915	0.000020	0.00067	0.281893	-30.8	0.7	7.5	1733
Low-Vinini-97SF12-GG2000SPE-164	13.3	4.0	0.282002	0.000020	0.00084	0.281975	-27.7	0.7	10.5	1736
Low-Vinini-97SF12-GG2000SPE-37	27.3	4.8	0.282009	0.000023	0.00162	0.281955	-27.4	0.8	10.0	1744
Low-Vinini-97SF12-GG2000SPE-26	6.5	5.3	0.281938	0.000021	0.00039	0.281925	-30.0	0.7	9.0	1744
Low-Vinini-97SF12-GG2000SPE-174	10.7	4.9	0.281843	0.000018	0.00067	0.281820	-33.3	0.6	5.8	1768
Low-Vinini-97SF12-GG2000SPE-39	15.9	5.6	0.281733	0.000016	0.00092	0.281702	-37.2	0.6	1.7	1774
Low-Vinini-97SF12-GG2000SPE-14	4.5	5.6	0.281841	0.000019	0.00027	0.281832	-33.4	0.7	6.6	1783
Low-Vinini-97SF12-GG2000SPE-49	14.3	5.4	0.281675	0.000019	0.00085	0.281646	-39.2	0.7	0.0	1785
Low-Vinini-97SF12-GG2000SPE-200	11.3	5.3	0.281202	0.000020	0.00067	0.281170	-56.0	0.7	-0.6	2489
<b>Sample: upper Vinini Formation. Location: Petes Summit, Toquima Range; 0518089 4337111 (NAD 83 UTM 11T)</b>										
upper VININI-97SF-11-GG2000SPE-138	11.4	5.5	0.281589	0.000022	0.00066	0.281566	-42.3	0.8	-2.7	1791
upper VININI-97SF-11-GG2000SPE-192	7.3	4.2	0.281493	0.000024	0.00044	0.281478	-45.7	0.9	-5.4	1809
upper VININI-97SF-11-GG2000SPE-25	6.4	5.3	0.281319	0.000024	0.00039	0.281306	-51.8	0.8	-11.4	1814
upper VININI-97SF-11-GG2000SPE-187	10.8	5.8	0.281500	0.000021	0.00068	0.281476	-45.4	0.7	-5.2	1823
upper VININI-97SF-11-GG2000SPE-27	9.7	4.6	0.281388	0.000024	0.00058	0.281368	-49.4	0.8	-8.9	1828
upper VININI-97SF-11-GG2000SPE-17	6.3	6.0	0.281504	0.000020	0.00038	0.281491	-45.3	0.7	-4.5	1828
upper VININI-97SF-11-GG2000SPE-4	11.5	4.5	0.281568	0.000021	0.00068	0.281544	-43.0	0.7	-2.6	1831
upper VININI-97SF-11-GG2000SPE-180	13.9	4.1	0.281649	0.000026	0.00090	0.281618	-40.2	0.9	0.1	1834
upper VININI-97SF-11-GG2000SPE-182	11.9	5.3	0.281476	0.000026	0.00069	0.281452	-46.3	0.9	-5.8	1834
upper VININI-97SF-11-GG2000SPE-1	13.2	4.4	0.281570	0.000022	0.00077	0.281544	-43.0	0.8	-2.5	1835
upper VININI-97SF-11-GG2000SPE-19	10.6	4.2	0.281416	0.000020	0.00065	0.281393	-48.4	0.7	-7.8	1836
upper VININI-97SF-11-GG2000SPE-145	7.6	5.9	0.281465	0.000019	0.00046	0.281449	-46.7	0.7	-5.7	1843
upper VININI-97SF-11-GG2000SPE-139	13.8	4.4	0.281767	0.000026	0.00092	0.281735	-36.0	0.9	4.5	1843
upper VININI-97SF-11-GG2000SPE-165	11.2	4.6	0.281518	0.000028	0.00067	0.281495	-44.8	1.0	-4.0	1844
upper VININI-97SF-11-GG2000SPE-169	10.9	3.9	0.281472	0.000022	0.00067	0.281449	-46.4	0.8	-5.6	1845
upper VININI-97SF-11-GG2000SPE-159	15.2	4.4	0.281510	0.000025	0.00090	0.281479	-45.1	0.9	-4.5	1846
upper VININI-97SF-11-GG2000SPE-6	8.7	5.0	0.281424	0.000021	0.00053	0.281405	-48.1	0.8	-7.1	1850
upper VININI-97SF-11-GG2000SPE-15	10.9	4.3	0.281576	0.000019	0.00067	0.281553	-42.7	0.7	-1.7	1854
upper VININI-97SF-11-GG2000SPE-198	15.4	3.2	0.281794	0.000029	0.00094	0.281761	-35.0	1.0	5.7	1855
upper VININI-97SF-11-GG2000SPE-8	4.3	5.6	0.281494	0.000020	0.00027	0.281485	-45.6	0.7	-3.9	1864
upper VININI-97SF-11-GG2000SPE-9	13.0	5.2	0.281844	0.000019	0.00081	0.281815	-33.3	0.7	8.0	1872
upper VININI-97SF-11-GG2000SPE-5	7.5	4.3	0.281495	0.000026	0.00044	0.281479	-45.6	0.9	-3.5	1891

### Hafnium isotope data of selected Roberts Mountains allochthon strata

	$(^{176}\text{Yb} + ^{176}\text{Lu}) / ^{177}\text{Hf}$ (%)	Volts Hf	$^{176}\text{Hf} / ^{177}\text{Hf}$	$\pm$ (1 $\sigma$ )	$^{176}\text{Lu} / ^{177}\text{Hf}$	$^{176}\text{Hf} / ^{177}\text{Hf}$ (T)	E-Hf (0)	E-Hf (0) $\pm$ (1 $\sigma$ )	E-Hf (T)	Age (Ma)
<b>Sample: upper Vinini Formation. Location: Petes Summit, Toquima Range; 0518089 4337111 (NAD 83 UTM 11T)</b>										
upper VININI-97SF-11-GG2000SPE-150	8.7	5.2	0.281498	0.000020	0.00054	0.281478	-45.5	0.7	-3.4	1895
upper VININI-97SF-11-GG2000SPE-149	11.7	4.5	0.281463	0.000023	0.00070	0.281438	-46.7	0.8	-4.8	1896
upper VININI-97SF-11-GG2000SPE-133	6.5	5.0	0.281477	0.000018	0.00040	0.281463	-46.2	0.6	-3.7	1906
upper VININI-97SF-11-GG2000SPE-2	8.2	5.0	0.281514	0.000024	0.00049	0.281496	-45.0	0.8	-2.5	1908
upper VININI-97SF-11-GG2000SPE-199	7.8	4.5	0.281562	0.000025	0.00049	0.281544	-43.2	0.9	-0.4	1926
upper VININI-97SF-11-GG2000SPE-197	8.6	4.9	0.281352	0.000023	0.00050	0.281334	-50.7	0.8	-7.8	1929
upper VININI-97SF-11-GG2000SPE-11	5.5	5.2	0.281162	0.000025	0.00034	0.281150	-57.4	0.9	-14.1	1937
upper VININI-97SF-11-GG2000SPE-7	9.0	5.3	0.281354	0.000022	0.00052	0.281335	-50.6	0.8	-7.6	1937
upper VININI-97SF-11-GG2000SPE-160	6.0	5.5	0.281464	0.000014	0.00037	0.281450	-46.7	0.5	-2.6	1974
upper VININI-97SF-11-GG2000SPE-12	10.2	4.8	0.281496	0.000018	0.00065	0.281471	-45.6	0.6	0.4	2074
upper VININI-97SF-11-GG2000SPE-179	7.8	5.1	0.281399	0.000023	0.00046	0.281381	-49.0	0.8	-2.7	2077
upper VININI-97SF-11-GG2000SPE-178	32.6	5.1	0.281508	0.000027	0.00191	0.281432	-45.2	0.9	-0.7	2086
upper VININI-97SF-11-GG2000SPE-189	9.3	5.1	0.281441	0.000020	0.00058	0.281418	-47.5	0.7	-1.2	2086
upper VININI-97SF-11-GG2000SPE-23	11.1	5.2	0.281410	0.000019	0.00067	0.281384	-48.6	0.7	-2.3	2089
upper VININI-97SF-11-GG2000SPE-135	12.1	4.8	0.281414	0.000023	0.00074	0.281384	-48.5	0.8	-2.0	2100
upper VININI-97SF-11-GG2000SPE-125	8.9	4.1	0.281371	0.000022	0.00055	0.281347	-50.0	0.8	1.7	2317
upper VININI-97SF-11-GG2000SPE-200	1.7	4.7	0.281044	0.000023	0.00011	0.281039	-61.6	0.8	-0.9	2677
upper VININI-97SF-11-GG2000SPE-194	9.5	3.9	0.280969	0.000025	0.00057	0.280939	-64.2	0.9	-3.6	2713
upper VININI-97SF-11-GG2000SPE-22	14.4	4.3	0.281085	0.000023	0.00081	0.281043	-60.1	0.8	0.1	2713
<b>Sample: Elder Sandstone. Location: Elder Creek, Shoshone Range; 0516196 4460270 (NAD 83 UTM 11T)</b>										
Elder-SS-GG2000SPE-59	13.0	3.7	0.281720	0.000022	0.00074	0.281695	-37.6	0.8	2.6	1823
Elder-SS-GG2000SPE-43	10.1	4.4	0.281560	0.000020	0.00060	0.281539	-43.3	0.7	-2.9	1824
Elder-SS-GG2000SPE-47	12.2	4.2	0.281403	0.000023	0.00067	0.281380	-48.9	0.8	-8.5	1826
Elder-SS-GG2000SPE-52	15.6	4.0	0.281483	0.000020	0.00087	0.281453	-46.0	0.7	-5.8	1832
Elder-SS-GG2000SPE-42	10.5	4.4	0.281431	0.000020	0.00063	0.281409	-47.9	0.7	-7.3	1835
Elder-SS-GG2000SPE-36	10.8	5.9	0.281460	0.000019	0.00069	0.281436	-46.9	0.7	-6.3	1835
Elder-SS-GG2000SPE-49	8.5	5.2	0.281431	0.000018	0.00049	0.281414	-47.9	0.6	-7.1	1836
Elder-SS-GG2000SPE-64	12.8	4.7	0.281571	0.000022	0.00075	0.281545	-42.9	0.8	-2.4	1836
Elder-SS-GG2000SPE-28	17.1	5.9	0.281385	0.000022	0.00096	0.281352	-49.5	0.8	-9.1	1845
Elder-SS-GG2000SPE-33	13.0	4.0	0.281749	0.000024	0.00077	0.281722	-36.6	0.8	4.1	1846
Elder-SS-GG2000SPE-72	8.5	4.8	0.281562	0.000016	0.00050	0.281544	-43.3	0.6	-2.2	1846
Elder-SS-GG2000SPE-16	7.4	4.5	0.281331	0.000027	0.00050	0.281314	-51.4	1.0	-10.4	1847
Elder-SS-GG2000SPE-101	12.0	4.2	0.281568	0.000020	0.00069	0.281544	-43.0	0.7	-2.2	1847
Elder-SS-GG2000SPE-102	17.8	3.0	0.281592	0.000038	0.00116	0.281551	-42.2	1.3	-1.9	1847
Elder-SS-GG2000SPE-30	37.8	3.9	0.281899	0.000035	0.00202	0.281828	-31.3	1.2	7.9	1849
Elder-SS-GG2000SPE-18	18.6	4.2	0.281581	0.000025	0.00105	0.281544	-42.6	0.9	-2.2	1849
Elder-SS-GG2000SPE-13	13.5	4.7	0.281545	0.000020	0.00077	0.281517	-43.9	0.7	-3.1	1850
Elder-SS-GG2000SPE-17	52.7	4.7	0.281635	0.000036	0.00270	0.281540	-40.7	1.3	-2.2	1853
Elder-SS-GG2000SPE-21	9.1	3.2	0.281590	0.000025	0.00053	0.281571	-42.3	0.9	-0.9	1860
Elder-SS-GG2000SPE-12	12.1	4.3	0.281539	0.000022	0.00076	0.281512	-44.1	0.8	-2.8	1871
Elder-SS-GG2000SPE-141	49.2	3.8	0.281879	0.000031	0.00311	0.281767	-32.0	1.1	6.6	1885
Elder-SS-GG2000SPE-148	10.8	4.0	0.281435	0.000027	0.00059	0.281414	-47.7	0.9	-5.7	1896
Elder-SS-GG2000SPE-148	15.9	3.7	0.281566	0.000025	0.00099	0.281530	-43.1	0.9	-1.6	1896
Elder-SS-GG2000SPE-41	7.2	4.9	0.281383	0.000017	0.00039	0.281368	-49.6	0.6	-6.9	1914
Elder-SS-GG2000SPE-92	10.3	4.5	0.281417	0.000017	0.00055	0.281398	-48.4	0.6	-5.8	1917
Elder-SS-GG2000SPE-66	11.4	5.5	0.281432	0.000020	0.00069	0.281404	-47.9	0.7	-2.0	2069
Elder-SS-GG2000SPE-57	10.7	4.6	0.281486	0.000021	0.00062	0.281461	-45.9	0.7	0.2	2078
Elder-SS-GG2000SPE-37	8.4	5.1	0.281490	0.000019	0.00054	0.281469	-45.8	0.7	0.6	2084
Elder-SS-GG2000SPE-11	7.7	5.4	0.281463	0.000021	0.00051	0.281443	-46.7	0.8	-0.3	2084
Elder-SS-GG2000SPE-27	14.9	5.4	0.281488	0.000020	0.00083	0.281455	-45.9	0.7	0.2	2088
Elder-SS-GG2000SPE-31	6.2	5.3	0.281858	0.000022	0.00036	0.281841	-32.8	0.8	26.0	2066
Elder-SS-GG2000SPE-62	7.1	4.3	0.281216	0.000017	0.00042	0.281195	-55.5	0.6	3.4	2621
Elder-SS-GG2000SPE-95	12.5	4.9	0.280973	0.000021	0.00066	0.280938	-64.1	0.7	-3.5	2717
<b>Sample: Slaven Chert. Location: Slaven Canyon, Shoshone Range; 0519428 4479302 (NAD 83 UTM 11T)</b>										
Slaven-RMA-GG2000SPE-172	44.8	4.3	0.281830	0.000027	0.00259	0.281742	-33.8	0.9	3.4	1786
Slaven-RMA-GG2000SPE-129	12.0	4.9	0.281468	0.000029	0.00069	0.281444	-46.6	1.0	-6.2	1825
Slaven-RMA-GG2000SPE-168	6.7	5.0	0.281442	0.000021	0.00040	0.281428	-47.5	0.8	-6.8	1825
Slaven-RMA-GG2000SPE-15	14.9	4.2	0.281670	0.000022	0.00079	0.281643	-39.4	0.8	0.8	1826
Slaven-RMA-GG2000SPE-100	5.5	3.9	0.281603	0.000029	0.00035	0.281591	-41.8	1.0	-1.0	1826
Slaven-RMA-GG2000SPE-167	9.0	4.0	0.281637	0.000023	0.00055	0.281618	-40.6	0.8	0.0	1829
Slaven-RMA-GG2000SPE-78	10.4	4.9	0.281640	0.000025	0.00065	0.281618	-40.5	0.9	0.0	1830
Slaven-RMA-GG2000SPE-180	7.5	5.2	0.281444	0.000016	0.00045	0.281428	-47.4	0.6	-6.7	1832
Slaven-RMA-GG2000SPE-152	9.2	5.4	0.281491	0.000026	0.00055	0.281472	-45.7	0.9	-5.1	1832
Slaven-RMA-GG2000SPE-142	11.2	5.6	0.281458	0.000027	0.00065	0.281435	-46.9	1.0	-6.4	1832
Slaven-RMA-GG2000SPE-186	10.2	4.8	0.281743	0.000028	0.00058	0.281723	-36.8	1.0	3.8	1832
Slaven-RMA-GG2000SPE-197	15.9	5.3	0.281339	0.000031	0.00089	0.281308	-51.1	1.1	-10.8	1838
Slaven-RMA-GG2000SPE-126	7.9	4.6	0.281476	0.000026	0.00048	0.281459	-46.3	0.9	-5.2	1846
Slaven-RMA-GG2000SPE-53	8.1	5.3	0.281454	0.000027	0.00049	0.281437	-47.1	0.9	-6.0	1846
Slaven-RMA-GG2000SPE-108	11.7	5.3	0.281385	0.000022	0.00068	0.281361	-49.5	0.8	-8.7	1846
Slaven-RMA-GG2000SPE-140	7.6	5.0	0.281458	0.000020	0.00046	0.281442	-46.9	0.7	-5.9	1846
Slaven-RMA-GG2000SPE-13	9.2	4.4	0.281474	0.000023	0.00056	0.281454	-46.4	0.8	-5.4	1847
Slaven-RMA-GG2000SPE-7	14.5	4.5	0.281591	0.000026	0.00086	0.281561	-42.2	0.9	-1.6	1847
Slaven-RMA-GG2000SPE-1	8.0	4.7	0.281530	0.000036	0.00048	0.281514	-44.4	1.3	-3.3	1848
Slaven-RMA-GG2000SPE-20	9.4	4.6	0.281455	0.000027	0.00054	0.281436	-47.0	1.0	-6.0	1849
Slaven-RMA-GG2000SPE-67	10.1	4.6	0.281498	0.000025	0.00061	0.281477	-45.5	0.9	-4.5	1849

### Hafnium isotope data of selected Roberts Mountains allochthon strata

	$(^{176}\text{Yb} + ^{176}\text{Lu}) / ^{176}\text{Hf}$ (%)	Volts Hf	$^{176}\text{Hf}/^{177}\text{Hf}$	$\pm$ (1 $\sigma$ )	$^{176}\text{Lu}/^{177}\text{Hf}$	$^{176}\text{Hf}/^{177}\text{Hf}$ (T)	E-Hf (0)	E-Hf (0) $\pm$ (1 $\sigma$ )	E-Hf (T)	Age (Ma)
<b>Sample: Slaven Chert. Location: Slaven Canyon, Shoshone Range; 0519428 4479302 (NAD 83 UTM 11T)</b>										
Slaven-RMA-GG2000SPE-137	4.9	5.1	0.281870	0.000017	0.00030	0.281859	-32.4	0.6	9.1	1852
Slaven-RMA-GG2000SPE-5	17.0	5.1	0.281519	0.000025	0.00102	0.281483	-44.8	0.9	-4.2	1852
Slaven-RMA-GG2000SPE-170	8.6	4.7	0.281522	0.000017	0.00054	0.281503	-44.7	0.6	-3.1	1872
Slaven-RMA-GG2000SPE-86	13.3	4.3	0.281479	0.000024	0.00080	0.281450	-46.2	0.9	-4.9	1874
Slaven-RMA-GG2000SPE-117	7.5	6.8	0.281536	0.000023	0.00047	0.281519	-44.2	0.8	-1.6	1913
Slaven-RMA-GG2000SPE-109	4.6	4.3	0.281358	0.000035	0.00028	0.281347	-50.5	1.2	-7.4	1925
Slaven-RMA-GG2000SPE-156	6.8	4.5	0.281482	0.000024	0.00042	0.281467	-46.1	0.9	-3.1	1927
Slaven-RMA-GG2000SPE-18	7.7	4.8	0.281656	0.000024	0.00047	0.281639	-39.9	0.9	3.0	1929
Slaven-RMA-GG2000SPE-107	13.1	5.3	0.281504	0.000039	0.00077	0.281474	-45.3	1.4	0.4	2066
Slaven-RMA-GG2000SPE-39	13.8	5.7	0.281482	0.000021	0.00083	0.281449	-46.1	0.7	-0.4	2070
Slaven-RMA-GG2000SPE-35	11.2	4.4	0.281547	0.000027	0.00086	0.281513	-43.8	0.9	2.2	2086
Slaven-RMA-GG2000SPE-175	18.2	5.1	0.281436	0.000017	0.00104	0.281395	-47.7	0.6	-1.8	2094
Slaven-RMA-GG2000SPE-90	11.5	5.4	0.281257	0.000021	0.00068	0.281224	-54.0	0.7	3.2	2571
Slaven-RMA-GG2000SPE-153	11.2	5.6	0.281245	0.000028	0.00064	0.281213	-54.5	1.0	2.9	2572
Slaven-RMA-GG2000SPE-48	15.1	5.1	0.280907	0.000033	0.00087	0.280863	-66.4	1.2	-7.3	2670
Slaven-RMA-GG2000SPE-82	8.3	4.2	0.281280	0.000033	0.00055	0.281251	-53.2	1.2	6.9	2686
Slaven-RMA-GG2000SPE-148	22.9	4.2	0.281257	0.000029	0.00135	0.281187	-54.0	1.0	4.8	2692

#### Notes:

1. Data reduction methodology is from Woodhead et al. (2004)
2. Analytical methods described in detail by Gehrels and Pecha (2014)
3.  $(^{176}\text{Yb} + ^{176}\text{Lu}) / ^{176}\text{Hf}$  (%) expresses the proportion of 176 due to  $^{176}\text{Yb} + ^{176}\text{Lu}$  versus the proportion due to  $^{176}\text{Hf}$ , in %.
4. Volts Hf is the sum of voltages of all Hf isotopes.
5.  $^{176}\text{Hf}/^{177}\text{Hf}$  is the measured  $^{176}\text{Hf}/^{177}\text{Hf}$ , corrected for fractionation and inferences. Shown with uncertainty expressed at 1-sigma.
6.  $^{176}\text{Lu}/^{177}\text{Hf}$  is the intensity of  $^{176}\text{Lu}$ , calculated from the measured intensity of  $^{175}\text{Lu}$  and  $^{176}\text{Lu}/^{175}\text{Lu}=0.02653$  (from Patchett, 1983), compared to the measured intensity of  $^{177}\text{Hf}$ . Fractionation of Lu isotopes is assumed to be the same as fractionation of Yb isotopes.
7.  $^{176}\text{Hf}/^{177}\text{Hf}$  (T) is the  $^{176}\text{Hf}/^{177}\text{Hf}$  corrected to the time of crystallization using a decay constant of  $1.867\text{e}-11$  (from Scherer et al., 2001 and Soderland et al., 2004)
8. E-Hf (0) is the present-day epsilon Hf value using  $^{176}\text{Hf}/^{177}\text{Hf}=0.282785$  and  $^{176}\text{Lu}/^{177}\text{Hf}=0.0336$  (from Bouvier et al., 2008). The uncertainty is expressed at 1-sigma.
9. E-Hf(T) is the epsilon Hf value at time of crystallization. Uncertainty is expressed at 1 sigma.
10. U-Pb ages are based on 206/238 for ages younger than  $\sim 1.0$  Ga, and on 206/207 for ages older than  $\sim 1.0$  Ga. This age cutoff may be slightly different for each sample.
11. Isotope ratios as follows:

180/177	1.8866600	Patchett (1983)
179/177	0.7325000	Patchett & Tatsumoto (1980)
178/177	1.4671800	Patchett (1983)
176/177	0.2821600	Patchett (1983)
174/177	0.0087100	Patchett (1983)
176/175	0.0265300	Patchett (1983)
176/171	0.9016910	Vervoort et al. (2004)
173/171	1.1323569	Vervoort et al. (2004)
172/171	1.5317360	Vervoort et al. (2004)

#### Notes for plots:

1. DM array is from Vervoort and Blichert-Toft (1999), using  $^{176}\text{Hf}/^{177}\text{Hf}=0.283225$  and  $^{176}\text{Lu}/^{177}\text{Hf}=0.0383$

2. CHUR is from Bouvier et al. (2008), using  $^{176}\text{Hf}/^{177}\text{Hf}=0.282785$  and  $^{176}\text{Lu}/^{177}\text{Hf}=0.0336$ .
3. Hf isotope evolution lines assume an average value of  $^{176}\text{Lu}/^{177}\text{Hf}=0.0115$  and a range of  $^{176}\text{Lu}/^{177}\text{Hf}=0.0036$  to  $^{176}\text{Lu}/^{177}\text{Hf}=0.0193$ . Values are from the average and 2-sigma range of values reported by Vervoort and Patchett (1996) and Vervoort et al. (1999).
4. Uncertainties shown at 2-sigma.
5. Uncertainty for EpsilonT is nearly identical for Epsilon 0 because of the very long half-life.



## APPENDIX E

### U-Pb Geochronologic analyses of selected Harmony Formation strata

Table notes are at the end of the appendix.

	Isotope ratios										Apparent ages (Ma)									
	U (ppm)	206Pb 204Pb	U/Th	206Pb* 207Pb*	± (%)	207Pb* 235U*	± (%)	206Pb* 238U*	± (%)	error corr.	206Pb* 238U*	± (Ma)	207Pb* 235U	± (Ma)	206Pb* 207Pb*	± (Ma)	Best age (Ma)	± (Ma)	Conc. (%)	Discor. (%)
<b>Sample: LCC #1. Location: Little Cottonwood Canyon, Galena Range; 0490600 4494792 (NAD 83 UTM 11T)</b>																				
LCC1-88-RM15-1	182	39487	2.0	13.2319	0.6	1.9196	1.5	0.1842	1.4	0.93	1090.0	14.4	1087.9	10.3	1083.8	11.6	1083.8	11.6	100.6	-0.6
LCC1-88-RM15-10	116	24982	1.4	12.7815	1.4	2.1130	3.6	0.1959	3.3	0.92	1153.1	34.7	1153.1	24.5	1152.9	27.1	1152.9	27.1	100.0	0.0
LCC1-88-RM15-100	64	17974	1.5	13.1335	1.8	1.9113	2.9	0.1821	2.2	0.77	1078.2	22.0	1085.0	19.3	1088.8	37.0	1088.8	37.0	98.1	1.9
LCC1-88-RM15-102	80	32781	0.7	9.8351	0.8	4.0861	2.6	0.2915	2.5	0.95	1648.8	35.7	1651.5	21.1	1654.9	15.3	1654.9	15.3	99.6	0.4
LCC1-88-RM15-103	58	24276	1.0	9.7848	1.6	4.2505	7.6	0.3010	7.4	0.97	1698.4	110.2	1683.8	62.6	1668.2	33.9	1668.2	33.9	101.7	-1.7
LCC1-88-RM15-104	117	40705	1.7	13.0278	1.1	2.0520	3.9	0.1939	3.8	0.96	1142.4	39.8	1133.0	27.0	1114.9	21.5	1114.9	21.5	102.5	-2.5
LCC1-88-RM15-106	21	5810	1.8	13.0514	5.5	1.9940	5.8	0.1887	2.0	0.34	1114.8	20.1	1113.5	39.2	1111.3	109.0	1111.3	109.0	100.3	-0.3
LCC1-88-RM15-107	261	25120	1.8	4.9085	0.3	13.7127	1.4	0.4882	1.4	0.98	2562.8	29.1	2730.1	13.2	2856.3	4.3	2856.3	4.3	89.7	10.3
LCC1-88-RM15-108	43	9139	0.9	11.6310	2.5	2.8138	3.7	0.2374	2.8	0.75	1373.0	34.3	1359.2	27.8	1337.7	47.4	1337.7	47.4	102.6	-2.6
LCC1-88-RM15-109	174	25420	2.0	13.3269	0.5	1.8943	1.5	0.1831	1.4	0.94	1083.9	13.8	1079.1	9.8	1069.5	10.0	1069.5	10.0	101.3	-1.3
LCC1-88-RM15-110	105	20397	2.0	13.3037	1.2	1.9242	2.2	0.1857	1.9	0.85	1097.8	19.2	1089.5	14.9	1073.0	23.3	1073.0	23.3	102.3	-2.3
LCC1-88-RM15-111	54	15414	1.6	13.3441	2.2	1.8598	3.2	0.1800	2.3	0.72	1067.0	23.1	1066.9	21.4	1066.9	45.2	1066.9	45.2	100.0	0.0
LCC1-88-RM15-112	63	17482	1.4	13.0287	1.6	2.0019	2.8	0.1892	2.2	0.81	1118.9	22.8	1116.2	18.7	1114.8	32.6	1114.8	32.6	100.2	-0.2
LCC1-88-RM15-113	136	55605	2.7	13.2656	1.1	1.8641	3.6	0.1793	3.5	0.96	1063.4	24.2	1068.4	24.1	1076.7	21.2	1076.7	21.2	98.6	1.4
LCC1-88-RM15-114	66	15720	3.0	12.5992	1.5	2.2009	2.5	0.2011	2.0	0.80	1181.3	21.8	1181.3	17.7	1181.4	30.3	1181.4	30.3	100.0	0.0
LCC1-88-RM15-115	282	82238	1.0	12.2627	0.5	2.3171	2.5	0.2061	2.4	0.98	1207.9	26.6	1217.5	17.5	1234.7	10.1	1234.7	10.1	97.8	2.2
LCC1-88-RM15-116	26	9370	1.5	13.5778	8.2	1.8388	8.6	0.1811	2.7	0.32	1072.8	27.0	1058.4	56.9	1031.9	165.9	1031.9	165.9	104.0	-4.0
LCC1-88-RM15-118	155	85081	2.3	9.1392	0.3	4.8430	1.1	0.3210	1.1	0.98	1794.7	16.6	1792.4	9.3	1789.7	5.8	1789.7	5.8	100.3	-0.3
LCC1-88-RM15-119	47	13087	2.1	12.8797	1.6	2.0915	2.7	0.1954	2.2	0.82	1150.4	23.4	1146.0	16.6	1137.7	30.9	1137.7	30.9	101.1	-1.1
LCC1-88-RM15-121	156	80163	1.7	13.2510	1.0	1.9284	1.7	0.1853	1.4	0.81	1096.0	13.9	1091.0	11.3	1080.9	19.6	1080.9	19.6	101.4	-1.4
LCC1-88-RM15-122	37	11074	0.9	13.3704	3.2	1.8838	4.1	0.1827	2.5	0.61	1081.6	24.8	1075.4	27.1	1062.9	65.1	1062.9	65.1	101.8	-1.8
LCC1-88-RM15-123	81	35495	1.2	10.9544	1.0	3.1932	2.1	0.2537	1.8	0.89	1457.5	24.0	1455.5	16.0	1452.6	18.3	1452.6	18.3	100.3	-0.3
LCC1-88-RM15-124	104	41683	1.9	13.0900	1.9	1.8333	3.2	0.1835	2.6	0.80	1086.3	25.8	1092.7	21.5	1105.4	38.3	1105.4	38.3	98.3	1.7
LCC1-88-RM15-125	187	47454	2.2	12.8428	0.6	2.0969	2.8	0.1953	2.7	0.98	1150.1	28.3	1147.8	18.9	1143.4	11.7	1143.4	11.7	100.6	-0.6
LCC1-88-RM15-126	52	10717	2.2	12.8719	2.4	2.0977	3.1	0.1958	1.8	0.61	1152.9	19.5	1148.0	21.0	1138.9	48.3	1138.9	48.3	101.2	-1.2
LCC1-88-RM15-127	57	7411	1.1	15.7953	4.1	1.0083	6.0	0.1155	4.3	0.72	704.7	28.8	708.0	30.5	718.6	88.1	704.7	28.8	98.1	1.9
LCC1-88-RM15-128	60	29628	1.6	8.7847	0.9	5.3318	5.3	0.3397	5.2	0.98	1885.2	85.6	1874.0	45.5	1861.5	18.6	1861.5	18.6	101.3	-1.3
LCC1-88-RM15-129	80	19756	1.4	12.3785	0.9	2.3201	2.1	0.2083	1.8	0.89	1219.7	20.4	1218.5	14.6	1216.2	18.4	1216.2	18.4	100.3	-0.3
LCC1-88-RM15-13	49	9226	0.8	12.5629	7.3	2.0939	8.9	0.1908	5.1	0.57	1125.6	52.3	1146.6	60.9	1187.1	143.8	1187.1	143.8	94.8	5.2
LCC1-88-RM15-130	180	54058	3.1	13.4363	0.7	1.8295	2.1	0.1783	2.0	0.94	1057.6	19.2	1056.1	13.8	1053.0	14.8	1053.0	14.8	100.4	-0.4
LCC1-88-RM15-131	151	89915	3.4	5.3569	0.1	13.4392	2.3	0.5221	2.3	1.00	2708.3	51.7	2711.0	22.1	2713.1	2.4	2713.1	2.4	99.8	0.2
LCC1-88-RM15-133	550	89612	2.6	13.1985	0.2	1.9031	2.4	0.1822	2.3	1.00	1078.8	23.3	1082.2	15.7	1088.9	4.4	1088.9	4.4	99.1	0.9
LCC1-88-RM15-134	78	37895	0.6	13.2022	1.8	1.9115	2.4	0.1830	1.6	0.67	1083.5	16.1	1085.1	16.1	1088.3	36.0	1088.3	36.0	99.6	0.4
LCC1-88-RM15-135	112	63794	1.3	9.2090	0.7	4.7209	1.7	0.3153	1.6	0.91	1766.8	24.3	1771.0	14.5	1775.9	13.3	1775.9	13.3	99.5	0.5
LCC1-88-RM15-136	66	28533	1.5	11.4427	1.4	2.9146	4.5	0.2419	4.3	0.95	1398.5	33.7	1385.7	34.0	1369.1	28.6	1369.1	28.6	102.0	-2.0
LCC1-88-RM15-137	126	20916	0.8	15.9911	2.2	0.9624	2.8	0.1116	1.7	0.60	882.1	10.8	884.5	13.9	892.5	47.8	892.1	10.8	98.5	1.5
LCC1-88-RM15-138	110	25720	1.9	13.2219	2.0	1.8807	2.8	0.1804	1.9	0.70	1068.9	19.2	1074.3	18.5	1085.3	40.0	1085.3	40.0	98.5	1.5
LCC1-88-RM15-139	191	56410	1.6	13.3174	1.2	1.8784	2.6	0.1814	2.3	0.89	1074.8	23.2	1073.5	17.4	1070.9	23.7	1070.9	23.7	100.4	-0.4
LCC1-88-RM15-14	74	16843	1.2	13.3982	0.9	1.8299	1.1	0.1778	0.6	0.58	1055.0	6.3	1056.2	7.2	1058.8	18.0	1058.8	18.0	99.6	0.4
LCC1-88-RM15-140	26	11190	1.0	10.8367	4.5	3.1929	5.2	0.2508	2.6	0.50	1443.3	34.0	1455.4	40.3	1473.1	85.3	1473.1	85.3	98.0	2.0
LCC1-88-RM15-141	69	19484	3.8	12.8033	3.3	2.0729	5.0	0.1925	3.7	0.75	1134.8	38.6	1139.9	34.0	1149.6	65.3	1149.6	65.3	98.7	1.3
LCC1-88-RM15-142	218	38398	2.1	13.5439	0.5	1.7628	1.1	0.1732	1.0	0.88	1029.5	9.2	1031.9	7.1	1036.9	10.7	1036.9	10.7	99.3	0.7
LCC1-88-RM15-143	281	11695	1.1	5.8923	0.2	9.0594	2.2	0.3872	2.2	0.99	2109.6	40.2	2344.0	20.6	2554.8	4.1	2554.8	4.1	82.6	17.4
LCC1-88-RM15-144	47	19484	1.8	12.0248	1.3	2.4874	4.0	0.2169	3.8	0.95	1265.6	43.3	1268.4	28.8	1273.0	24.9	1273.0	24.9	99.4	0.6
LCC1-88-RM15-145	88	68285	1.2	9.8781	0.8	4.0178	2.0	0.2878	1.9	0.92	1630.7	28.7	1637.8	18.4	1646.8	14.6	1646.8	14.6	99.0	1.0
LCC1-88-RM15-146	141	21183	1.1	12.9934	1.5	1.9878	2.5	0.1854	2.0	0.79	1096.5	19.9	1104.5	16.9	1120.2	30.9	1120.2	30.9	97.9	2.1
LCC1-88-RM15-147	283	85348	1.9	10.8200	0.3	3.2770	3.8	0.2572	3.8	1.00	1475.3	49.6	1475.6	29.4	1476.0	6.1	1476.0	6.1	98.9	0.1
LCC1-88-RM15-148	71	29738	2.1	9.7772	0.7	4.1894	2.8	0.2971	2.8	0.97	1676.8	40.7	1672.0	23.3	1665.9	13.1	1665.9	13.1	100.7	-0.7
LCC1-88-RM15-149	54	18147	2.8	13.1826	2.7	1.9368	3.5	0.1852	2.2	0.64	1095.2	22.5	1093.9	23.3	1091.3	53.5	1091.3	53.5	100.4	-0.4
LCC1-88-RM15-15	110	18505	1.0	13.2410	1.0	1.8939	2.3	0.1819	2.1	0.91	1077.2	21.1	1078.9	15.5	1082.5	19.7	1082.5	19.7	99.5	0.5
LCC1-88-RM15-150	178	62900	2.0	13.1520	0.8	1.9555	1.7	0.1865	1.5	0.87	1102.6	14.7	1100.3	11.3	1096.0	16.7	1096.0	16.7	100.6	-0.6
LCC1-88-RM15-151	80	33415	3.2	9.0739	0.4	4.7912	1.6	0.3153	1.6	0.97	1766.8	24.3	1783.4	13.6	1802.8	7.4	1802.8	7.4	98.0	2.0
LCC1-88-RM15-152	167	16190	2.0	13.5991	1.0	1.7027	1.9	0.1879	1.6	0.85	1000.8	14.7	1009.6	11.9	1028.7	19.8	1028.7	19.8	97.3	2.7
LCC1-88-RM15-153	261	144990	1.4	5.9124	0.1	10.9297	1.3	0.4687	1.3	1.00	2477.7	26.2	2517.2	11.9	2549.1	1.9	2549.1	1.9	97.2	2.8
LCC1-88-RM15-154	172	114327	1.2	9.8944	0.2	11.0631	1.2	0.4730	1.2	0.98	2498.5	24.8	2528.5	11.3	2554.2	3.7	2554.2	3.7	97.7	2.3
LCC1-88-RM15-155	201	62533	2.5	12.5935																





## U-Pb Geochronologic analyses of selected Harmony Formation strata

	Isotope ratios												Apparent ages (Ma)						Best age	Conc.	Discor.
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±					
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)					
<b>Sample: LCC #2. Location: Little Cottonwood Canyon, Galena Range; 0490529 4495008 (NAD 83 UTM 11T)</b>																					
LCC2-H971-190	86	33207	2.6	13.5968	2.2	1.7832	2.4	0.1759	0.40	1044.3	9.2	1039.4	15.3	1029.0	43.5	1029.0	43.5	101.5	-1.5		
LCC2-H971-171	178	41260	2.6	13.5919	0.7	1.7176	0.8	0.1693	0.4	0.55	1008.4	4.2	1015.1	5.2	1029.8	13.7	1029.8	13.7	97.9	-2.1	
LCC2-H971-112	145	30365	2.1	13.5873	0.9	1.8085	1.4	0.1782	1.1	0.77	1057.3	10.5	1048.6	9.1	1030.4	17.9	1030.4	17.9	102.6	-2.6	
LCC2-H971-164	266	52049	5.2	13.5422	0.8	1.8047	1.0	0.1773	0.6	0.60	1051.9	6.0	1047.2	6.8	1037.2	16.8	1037.2	16.8	101.4	-1.4	
LCC2-H971-53	186	33020	3.0	13.5339	1.0	1.7638	2.1	0.1751	1.9	0.89	1029.3	18.1	1032.2	13.8	1038.4	19.4	1038.4	19.4	99.1	-0.9	
LCC2-H971-3	46	14370	1.4	13.5032	1.7	1.8748	2.1	0.1836	1.2	0.58	1086.4	12.0	1072.2	13.6	1043.4	33.8	1043.4	33.8	104.1	-4.1	
LCC2-H971-170	131	38084	3.2	13.4693	1.1	1.7577	1.3	0.1717	0.6	0.45	1021.5	5.3	1030.0	8.1	1048.1	22.6	1048.1	22.6	97.5	-2.5	
LCC2-H971-110	87	14944	1.5	13.4665	2.3	1.8405	2.6	0.1798	1.1	0.44	1065.6	10.9	1060.0	16.8	1048.5	46.3	1048.5	46.3	101.6	-1.6	
LCC2-H971-30	52	10162	1.9	13.4524	3.6	1.8581	3.8	0.1813	1.0	0.26	1074.0	9.8	1066.3	24.9	1050.6	73.4	1050.6	73.4	102.2	-2.2	
LCC2-H971-47	44	10971	1.3	13.4359	3.0	1.8763	5.4	0.1828	4.5	0.83	1082.5	45.0	1072.8	35.8	1053.1	60.0	1053.1	60.0	102.8	-2.8	
LCC2-H971-36	77	47568	2.3	13.4221	1.0	1.7914	1.4	0.1744	1.0	0.72	1036.3	9.8	1042.3	9.3	1055.2	20.2	1055.2	20.2	98.2	-1.8	
LCC2-H971-142	40	6447	1.5	13.3801	4.4	1.9190	4.7	0.1862	1.6	0.34	1100.9	16.2	1087.7	31.5	1061.5	89.2	1061.5	89.2	103.7	-3.7	
LCC2-H971-29	47	18130	1.7	13.3653	3.5	1.9071	3.8	0.1849	1.4	0.38	1093.5	14.4	1083.6	25.4	1063.7	71.1	1063.7	71.1	102.8	-2.8	
LCC2-H971-104	67	24826	1.6	13.3640	1.1	1.8605	1.6	0.1803	1.2	0.74	1068.8	11.7	1067.2	10.7	1063.9	22.0	1063.9	22.0	100.5	-0.5	
LCC2-H971-54	60	10635	1.5	13.3268	2.6	1.9314	3.1	0.1867	1.6	0.52	1103.3	16.1	1092.0	20.5	1069.5	52.8	1069.5	52.8	103.2	-3.2	
LCC2-H971-126	87	15592	1.3	13.3173	1.1	1.9681	1.7	0.1901	1.3	0.78	1121.9	13.8	1104.7	11.7	1071.9	22.0	1071.9	22.0	104.8	-4.8	
LCC2-H971-192	123	48954	2.1	13.3161	1.4	1.9047	1.8	0.1839	1.0	0.58	1088.5	10.2	1082.7	17.7	1071.1	28.8	1071.1	28.8	101.6	-1.6	
LCC2-H971-194	58	19465	0.5	13.3105	2.3	1.9651	2.5	0.1897	1.0	0.38	1119.8	9.8	1103.6	16.7	1072.0	46.1	1072.0	46.1	104.5	-4.5	
LCC2-H971-49	42	10820	1.3	13.3055	3.6	1.9371	4.7	0.1869	3.0	0.64	1104.7	30.1	1094.0	31.2	1072.7	72.4	1072.7	72.4	103.0	-3.0	
LCC2-H971-179	27	9067	1.7	13.2869	4.2	1.9343	4.4	0.1864	1.3	0.29	1101.9	12.8	1093.0	29.3	1075.5	84.2	1075.5	84.2	102.4	-2.4	
LCC2-H971-27	78	51974	2.1	13.2842	1.8	1.8794	2.2	0.1811	1.2	0.55	1072.8	12.0	1073.9	14.6	1075.9	36.9	1075.9	36.9	99.7	-0.3	
LCC2-H971-187	40	20614	1.7	13.2732	2.5	1.8950	2.8	0.1824	1.1	0.41	1060.2	11.3	1079.3	18.5	1077.6	51.1	1077.6	51.1	100.2	-0.2	
LCC2-H971-198	105	35854	1.8	13.2695	1.2	1.9560	1.3	0.1882	0.6	0.44	1111.8	6.0	1100.5	8.9	1078.1	23.8	1078.1	23.8	103.1	-3.1	
LCC2-H971-136	128	63659	1.8	13.2618	0.7	1.9287	0.9	0.1855	0.5	0.58	1097.0	5.2	1091.1	5.9	1079.3	14.6	1079.3	14.6	101.6	-1.6	
LCC2-H971-5	76	25787	2.6	13.2558	1.4	1.9321	1.9	0.1858	1.2	0.64	1098.3	12.0	1092.3	12.5	1080.2	28.8	1080.2	28.8	101.7	-1.7	
LCC2-H971-123	305	34064	1.7	13.2444	0.4	1.9599	3.2	0.1883	3.1	0.99	1112.0	32.0	1101.8	21.2	1081.9	8.3	1081.9	8.3	102.8	-2.8	
LCC2-H971-84	208	58794	2.6	13.2117	1.2	1.9220	1.4	0.1832	0.9	0.59	1084.5	8.6	1085.3	9.7	1086.9	23.3	1086.9	23.3	99.9	-0.1	
LCC2-H971-174	120	65332	2.7	13.2086	1.2	1.9320	1.3	0.1851	0.6	0.43	1094.4	5.6	1092.2	6.8	1087.4	23.1	1087.4	23.1	100.7	-0.7	
LCC2-H971-63	62	12395	1.2	13.2056	2.6	1.8740	4.3	0.1795	3.4	0.80	1064.1	33.6	1071.9	28.3	1087.8	51.1	1087.8	51.1	97.8	-2.2	
LCC2-H971-78	92	27097	2.1	13.2015	1.5	1.9772	1.6	0.1893	0.6	0.35	1117.6	5.7	1107.8	10.8	1088.4	30.0	1088.4	30.0	107.2	-7.2	
LCC2-H971-118	399	63096	2.4	13.1933	0.4	1.9784	2.1	0.1893	2.0	0.98	1117.6	20.8	1108.2	14.0	1089.7	8.3	1089.7	8.3	102.6	-2.6	
LCC2-H971-173	149	57941	2.9	13.1846	0.6	1.9359	0.9	0.1851	0.6	0.69	1094.9	6.0	1093.6	5.7	1091.0	12.3	1091.0	12.3	100.4	-0.4	
LCC2-H971-152	138	32744	1.0	13.1785	1.0	1.9903	1.3	0.1902	0.9	0.66	1122.6	8.9	1112.2	8.9	1091.9	19.7	1091.9	19.7	102.8	-2.8	
LCC2-H971-11	81	21721	1.5	13.1705	2.2	1.9401	2.9	0.1853	1.9	0.64	1096.0	18.8	1095.0	19.5	1093.1	44.8	1093.1	44.8	100.3	-0.3	
LCC2-H971-188	80	69757	2.1	13.1690	0.9	1.9488	1.1	0.1861	0.7	0.62	1100.4	7.1	1098.0	7.6	1093.4	17.7	1093.4	17.7	100.6	-0.6	
LCC2-H971-28	46	15920	1.7	13.1685	2.9	1.9086	3.2	0.1823	1.4	0.44	1079.4	13.8	1084.1	21.1	1093.5	57.1	1093.5	57.1	98.7	-1.3	
LCC2-H971-39	76	34222	1.5	13.1679	2.5	1.9772	2.7	0.1888	0.9	0.32	1115.0	8.8	1107.8	18.0	1093.6	50.5	1093.6	50.5	102.0	-2.0	
LCC2-H971-109	177	71152	1.3	13.1660	1.3	1.9594	2.4	0.1871	2.0	0.84	1106.7	20.3	1101.7	16.0	1093.9	25.5	1093.9	25.5	101.1	-1.1	
LCC2-H971-50	99	20812	2.4	13.1587	1.7	1.9024	2.7	0.1816	2.1	0.77	1075.5	20.7	1081.9	18.0	1095.0	34.6	1095.0	34.6	98.2	-1.8	
LCC2-H971-115	126	53785	4.9	13.1439	1.0	2.0185	1.3	0.1924	0.8	0.64	1134.5	8.8	1121.8	9.0	1097.2	20.3	1097.2	20.3	103.4	-3.4	
LCC2-H971-189	41	15819	1.5	13.1414	3.8	2.0038	3.9	0.1910	0.8	0.19	1126.7	7.8	1116.8	26.2	1097.6	76.0	1097.6	76.0	102.7	-2.7	
LCC2-H971-34	126	24422	1.3	13.1064	0.7	1.9765	0.9	0.1879	0.6	0.66	1109.9	6.3	1107.5	6.3	1102.9	14.1	1102.9	14.1	100.6	-0.6	
LCC2-H971-169	263	74143	2.8	13.1002	1.0	1.9478	1.2	0.1851	0.7	0.55	1094.6	6.7	1097.7	8.1	1103.9	20.1	1103.9	20.1	99.2	-0.8	
LCC2-H971-193	109	27640	0.9	13.0878	1.5	2.0547	5.5	0.1952	5.3	0.96	1149.4	56.1	1133.9	37.8	1104.2	30.0	1104.2	30.0	104.1	-4.1	
LCC2-H971-96	99	33238	1.3	13.0878	2.3	1.9595	2.5	0.1861	0.9	0.37	1100.4	9.4	1101.7	16.8	1104.2	46.3	1104.2	46.3	99.7	-0.3	
LCC2-H971-145	81	26351	1.8	13.0829	2.0	1.9502	2.2	0.1852	0.9	0.39	1095.2	8.7	1098.5	14.7	1105.0	40.4	1105.0	40.4	99.1	-0.9	
LCC2-H971-125	52	32907	1.4	13.0912	2.8	1.9521	3.9	0.1853	2.8	0.72	1096.1	28.4	1099.2	26.5	1105.2	55.0	1105.2	55.0	99.2	-0.8	
LCC2-H971-88	124	46794	1.4	13.0885	1.7	2.0243	2.8	0.1919	2.2	0.79	1131.5	22.5	1123.7	18.7	1108.7	33.8	1108.7	33.8	102.1	-2.1	
LCC2-H971-133	30	23246	2.7	13.0566	5.7	1.9206	6.0	0.1819	1.8	0.31	1077.1	18.2	1083.3	40.0	1110.7	113.9	1110.7	113.9	97.9	-2.9	
LCC2-H971-150	472	21224	16.4	13.0259	0.5	1.9431	1.5	0.1836	1.4	0.94	1086.4	13.7	1096.1	9.7	1115.2	9.5	1115.2	9.5	97.4	-2.6	
LCC2-H971-51	65	19928	1.2	13.0237	2.2	1.9644	3.4	0.1856	2.6	0.77	1097.2	26.2	1103.4	22.8	1115.5	43.6	1115.5	43.6	98.4	-1.6	
LCC2-H971-60	111	39042	4.4	12.9972	1.1	2.0359	2.2	0.1919	1.9	0.87	1131.7	20.1	1127.6	15.2	1119.6	22.0	1119.6	22.0	101.1	-1.1	
LCC2-H971-9	85	54434	1.0	12.9936	0.9	2.0679	2.0	0.1949	1.8	0.90	1147.8	18.5	1138.3	13.4	1120.2	17.3	1120.2	17.3	102.5	-2.5	
LCC2-H971-166	35	13849	1.3	12.9875	4.6	1.9720	4.7	0.1858	1.1	0.23	1098.3	11.2	1106.0	31.8	1121.1	91.6	1121.1	91.6	98.0	-2.0	
LCC2-H971-131	348	62691	2.5	12.9813	0.4	1.9930	1.6	0.1876	1.6	0.97	1108.6	15.9	1113.1	10.9	1122.1	6.2	1122.1	6.2	98.8	-1.2	
LCC2-H971-32	48	17721	1.7	12.9746	2.9	2.0110	3.3	0.1892	1.6	0.50	1117.2	16.8	1119.2	22.5	1123.1	57.3	1123.1	57.3	99.5	-0.5	
LCC2-H971-16	33	11937	1.2	12.9634	5.2	1.9924	5.5	0.1873	1.8												



## U-Pb Geochronologic analyses of selected Harmony Formation strata

	Isotope ratios										Apparent ages (Ma)									
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc.	Discor.
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)
<b>Sample: LCC #2. Location: Little Cottonwood Canyon, Galena Range; 0490529 4495008 (NAD 83 UTM 11T)</b>																				
LCC2-H971-157	35	17210	1.6	11.6710	2.9	2.7728	7.2	0.2347	6.6	0.91	1359.1	80.8	1348.2	53.8	1331.0	56.3	1331.0	56.3	102.1	-2.1
LCC2-H971-92	205	72082	2.0	11.6600	0.7	2.7530	2.5	0.2328	2.4	0.96	1349.2	29.1	1342.9	18.5	1332.8	13.3	1332.8	13.3	101.2	-1.2
LCC2-H971-108	270	62133	0.6	11.8444	0.4	2.6780	1.4	0.2262	1.3	0.95	1314.3	15.3	1322.4	10.0	1335.4	8.1	1335.4	8.1	98.4	1.6
LCC2-H971-77	80	7702	1.5	11.6275	1.6	2.7401	1.9	0.2311	1.0	0.53	1340.1	12.2	1339.4	14.1	1338.2	31.0	1338.2	31.0	100.1	-0.1
LCC2-H971-168	45	9799	1.3	11.5702	3.5	2.7431	3.6	0.2302	0.9	2.42	1335.5	10.5	1340.2	26.7	1347.8	67.1	1347.8	67.1	99.1	0.9
LCC2-H971-45	93	23766	1.4	11.5105	1.6	2.7840	2.3	0.2324	1.7	0.73	1347.2	20.6	1351.3	17.3	1357.8	30.2	1357.8	30.2	99.2	0.8
LCC2-H971-44	52	12073	1.1	11.4511	2.2	2.8833	3.4	0.2395	2.6	0.75	1383.9	31.9	1377.6	25.7	1367.7	43.1	1367.7	43.1	101.2	-1.2
LCC2-H971-183	206	76217	2.0	11.4382	0.4	2.8478	0.6	0.2362	0.3	0.60	1367.1	4.1	1368.2	4.2	1369.9	8.6	1369.9	8.6	99.8	0.2
LCC2-H971-89	40	10623	1.6	11.4003	2.3	2.9177	3.1	0.2412	2.0	0.66	1393.2	25.6	1386.5	23.3	1376.3	44.3	1376.3	44.3	101.2	-1.2
LCC2-H971-67	64	14858	0.9	11.3897	2.2	2.8897	2.7	0.2387	1.5	0.55	1380.0	18.1	1379.2	20.1	1378.1	43.0	1378.1	43.0	100.1	-0.1
LCC2-H971-46	59	23260	1.1	11.3786	1.6	2.9337	2.5	0.2421	1.9	0.77	1397.6	24.3	1390.6	19.0	1379.9	30.9	1379.9	30.9	101.3	-1.3
LCC2-H971-180	184	109201	0.8	11.3774	0.4	2.9007	1.4	0.2394	1.4	0.97	1383.3	17.3	1382.1	10.8	1380.1	7.0	1380.1	7.0	100.2	-0.2
LCC2-H971-70	89	33843	1.5	11.3481	1.2	2.9509	2.2	0.2429	1.8	0.82	1401.6	22.4	1395.1	18.4	1385.1	23.5	1385.1	23.5	101.2	-1.2
LCC2-H971-84	39	9464	1.1	11.3361	1.8	3.0313	3.3	0.2492	1.8	0.54	1434.5	23.3	1415.5	25.4	1387.1	53.7	1387.1	53.7	103.4	-3.4
LCC2-H971-141	51	25758	1.6	11.3122	2.2	2.9579	3.1	0.2427	2.1	0.89	1400.6	26.9	1396.9	23.5	1391.2	42.9	1391.2	42.9	100.7	-0.7
LCC2-H971-76	65	12431	1.7	11.2425	1.8	2.9148	2.0	0.2377	0.8	0.42	1374.6	10.5	1385.8	15.2	1403.4	34.9	1403.4	34.9	98.0	2.0
LCC2-H971-116	168	9978	1.1	11.1777	0.9	2.4307	8.7	0.1971	8.7	0.99	1159.5	92.0	1251.7	62.8	1414.1	18.1	1414.1	18.1	82.0	18.0
LCC2-H971-153	87	41598	2.1	11.0639	1.2	3.1719	1.7	0.2545	1.2	0.72	1461.8	15.5	1450.4	12.8	1433.6	22.0	1433.6	22.0	102.0	-2.0
LCC2-H971-122	80	13652	1.5	11.0170	1.5	3.2803	1.9	0.2595	1.7	0.89	1487.1	22.3	1476.4	14.6	1461.0	16.0	1461.0	16.0	101.8	-1.8
LCC2-H971-195	97	44380	1.0	10.9837	0.8	3.2665	1.1	0.2602	0.7	0.64	1490.9	9.2	1473.1	8.3	1447.5	15.6	1447.5	15.6	103.0	-3.0
LCC2-H971-14	111	47690	1.2	10.9320	0.5	3.0728	1.1	0.2436	1.0	0.89	1405.6	12.4	1425.9	8.4	1456.5	9.5	1456.5	9.5	96.5	3.5
LCC2-H971-181	88	17070	1.4	10.9083	0.8	3.2213	1.0	0.2549	0.6	0.60	1463.5	7.9	1462.3	7.8	1460.6	15.2	1460.6	15.2	100.2	-0.2
LCC2-H971-23	96	48175	1.6	10.9063	0.8	3.2803	1.9	0.2595	1.7	0.89	1487.1	22.3	1476.4	14.6	1461.0	16.0	1461.0	16.0	101.8	-1.8
LCC2-H971-143	130	77961	1.7	10.9055	0.6	3.2859	1.2	0.2607	1.1	0.88	1493.4	14.2	1480.1	9.4	1461.1	10.7	1461.1	10.7	102.7	-2.2
LCC2-H971-113	57	12306	1.0	10.9009	2.0	3.2780	2.9	0.2592	2.2	0.74	1485.5	28.9	1475.8	22.8	1461.9	37.1	1461.9	37.1	101.6	-1.6
LCC2-H971-41	126	142999	1.2	10.8593	0.6	3.1741	1.0	0.2500	0.8	0.79	1438.4	10.4	1450.9	7.8	1469.2	11.8	1469.2	11.8	97.9	2.1
LCC2-H971-159	73	15906	1.6	10.8526	1.0	3.3635	2.0	0.2647	1.7	0.86	1514.1	23.3	1495.9	15.6	1470.3	19.2	1470.3	19.2	103.0	-3.0
LCC2-H971-167	81	28751	1.0	10.8242	1.4	3.3273	1.8	0.2612	1.1	0.60	1496.0	14.2	1487.5	13.8	1475.3	26.7	1475.3	26.7	101.4	-1.4
LCC2-H971-100	101	11881	0.8	10.7963	1.1	3.3178	1.2	0.2598	0.5	0.44	1488.8	7.2	1485.2	9.6	1480.2	10.0	1480.2	10.0	100.6	-0.6
LCC2-H971-118	89	106427	1.1	10.7915	0.8	3.3228	1.0	0.2601	0.6	0.58	1490.2	7.6	1486.4	7.7	1481.0	15.2	1481.0	15.2	100.6	-0.6
LCC2-H971-19	59	19979	4.5	10.7875	1.2	3.2623	1.5	0.2552	0.9	0.58	1465.4	11.6	1472.1	9.1	1481.8	23.6	1481.8	23.6	98.9	1.1
LCC2-H971-186	65	38538	1.4	10.7862	1.0	3.2802	1.2	0.2566	0.7	0.87	1477.5	6.9	1476.4	9.1	1482.0	16.2	1482.0	16.2	99.4	0.6
LCC2-H971-58	56	18219	0.7	10.0428	1.4	3.6315	1.8	0.2845	1.3	0.66	1512.9	17.8	1556.5	9.2	1616.1	25.9	1616.1	25.9	93.6	6.4
LCC2-H971-17	25	10457	1.9	10.0253	3.0	3.9859	3.2	0.2898	1.1	0.34	1640.6	16.0	1631.3	26.0	1619.4	56.0	1619.4	56.0	101.3	-1.3
LCC2-H971-21	132	63503	1.2	9.9285	0.5	4.0407	0.9	0.2910	0.7	0.84	1646.4	10.6	1642.4	7.1	1637.4	8.8	1637.4	8.8	100.5	-0.5
LCC2-H971-147	69	45640	1.2	9.8733	1.5	4.1311	1.7	0.2958	0.7	0.42	1670.5	10.2	1660.5	13.6	1647.8	28.0	1647.8	28.0	101.4	-1.4
LCC2-H971-154	72	83931	1.0	9.7673	1.1	4.2565	2.7	0.3015	2.4	0.91	1698.9	36.1	1685.0	21.9	1667.7	20.7	1667.7	20.7	101.9	-1.9
LCC2-H971-175	134	141660	1.1	9.7555	0.5	4.2859	0.7	0.3032	0.6	0.76	1707.4	8.5	1690.7	6.1	1670.0	8.9	1670.0	8.9	102.2	-2.2
LCC2-H971-158	96	20031	2.4	9.6870	1.1	3.6874	3.6	0.2591	3.4	0.95	1485.0	44.8	1568.6	28.5	1683.0	21.2	1683.0	21.2	88.2	11.8
LCC2-H971-97	187	99848	1.0	9.4931	0.6	4.5124	1.4	0.3107	1.3	0.91	1744.0	19.3	1733.3	11.6	1720.2	10.8	1720.2	10.8	101.4	-1.4
LCC2-H971-185	162	53185	1.8	9.4430	0.4	4.5220	0.9	0.3097	0.8	0.87	1739.2	11.7	1735.0	7.9	1729.9	7.9	1729.9	7.9	100.5	-0.5
LCC2-H971-200	85	40815	0.8	9.3222	0.8	4.7533	0.9	0.3214	0.5	0.53	1798.4	7.4	1776.7	7.5	1753.5	13.9	1753.5	13.9	102.4	-2.4
LCC2-H971-178	62	45541	1.3	9.2428	0.9	4.7159	1.1	0.3161	0.7	0.60	1770.8	10.7	1770.1	9.6	1769.2	16.8	1769.2	16.8	100.1	-0.1
LCC2-H971-198	99	39770	1.3	9.1855	0.9	4.8611	1.1	0.3238	0.7	0.58	1808.5	10.5	1795.5	9.6	1780.5	16.9	1780.5	16.9	101.6	-1.6
LCC2-H971-24	9	7316	3.2	9.1557	5.5	4.4735	5.9	0.2971	2.1	0.36	1676.7	31.7	1726.1	49.2	1786.4	100.6	1786.4	100.6	93.9	6.1
LCC2-H971-107	137	78563	1.9	9.1253	0.5	4.8691	1.4	0.3223	1.3	0.93	1800.7	19.9	1796.9	11.5	1792.5	9.3	1792.5	9.3	100.5	-0.5
LCC2-H971-69	299	257960	1.5	9.1183	0.3	4.9123	1.2	0.3249	1.1	0.97	1813.4	18.0	1804.4	9.9	1793.9	5.0	1793.9	5.0	101.1	-1.1
LCC2-H971-71	121	71484	1.5	9.1172	0.4	4.9654	1.3	0.3263	1.2	0.94	1830.3	19.6	1813.5	11.0	1794.1	8.0	1794.1	8.0	102.0	-2.0
LCC2-H971-90	98	41193	1.8	9.1076	0.5	4.8767	1.8	0.3221	1.8	0.96	1800.1	27.9	1796.2	15.5	1796.0	8.3	1796.0	8.3	100.2	-0.2
LCC2-H971-103	62	58780	3.3	9.0980	0.9	4.8045	1.3	0.3236	0.9	0.69	1807.4	14.4	1803.0	11.1	1798.0	17.9	1798.0	17.9	100.5	-0.5
LCC2-H971-68	63	22855	5.0	8.9311	0.9	4.9627	3.6	0.3215	3.5	0.97	1798.8	55.2	1810.0	30.8	1811.6	16.8	1811.6	16.8		

## U-Pb Geochronologic analyses of selected Harmony Formation strata

Sample	ID	Isotope ratios										Apparent ages (Ma)									
		U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc.	Discor.
		(ppm)	204Pb	207Pb*	(%)	235U*	(%)	238U*	(%)	corr.	238U*	(Ma)	235U*	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)	
Sample: LCC #3. Location: Little Cottonwood Canyon, Galena Range; 0490459 4494916 (NAD 83 UTM 11T)																					
LCC3-H973-34	50	27940	1.2	9.1245	1.2	4.8729	1.7	0.3225	1.1	0.68	1801.8	17.7	1797.6	14.0	1792.7	22.1	1792.7	22.1	100.5	-0.5	
LCC3-H973-189	169	23470	2.5	9.1240	0.4	4.8685	0.8	0.3220	0.7	0.90	1799.7	11.4	1796.5	6.8	1792.8	6.5	1792.8	6.5	100.4	-0.4	
LCC3-H973-88	213	12678	1.9	9.1215	0.5	4.8784	1.3	0.3227	1.2	0.91	1803.0	18.8	1798.5	11.1	1793.3	9.9	1793.3	9.9	100.5	-0.5	
LCC3-H973-8	78	8974	1.6	9.0824	1.6	4.8795	1.9	0.3214	1.1	0.58	1796.7	17.5	1798.7	16.2	1801.1	28.4	1801.1	28.4	99.8	0.2	
LCC3-H973-68	213	6033	0.8	9.0793	1.7	4.4766	1.9	0.2948	0.9	0.47	1665.4	13.3	1726.7	15.9	1801.7	30.7	1801.7	30.7	92.4	7.6	
LCC3-H973-169	168	4185	1.3	9.0274	1.7	4.6754	2.3	0.3061	1.5	0.65	1721.6	22.0	1762.9	18.8	1812.1	31.2	1812.1	31.2	95.0	5.0	
LCC3-H973-171	56	42844	1.5	9.0074	0.7	5.0487	1.3	0.3298	1.2	0.87	1837.5	18.8	1827.5	11.4	1816.2	12.0	1816.2	12.0	101.2	-1.2	
LCC3-H973-85	47	59631	0.8	8.9826	1.5	4.2240	2.0	0.2752	1.4	0.67	1567.1	18.8	1678.7	16.5	1821.2	27.0	1821.2	27.0	86.0	14.0	
LCC3-H973-25	160	16816	0.7	8.9540	0.2	3.9302	1.1	0.2552	1.1	0.98	1465.4	14.7	1619.9	9.3	1826.9	4.5	1826.9	4.5	80.2	19.8	
LCC3-H973-15	46	66150	1.2	8.9303	0.9	5.1461	1.5	0.3333	1.1	0.77	1854.4	18.2	1843.8	12.4	1831.8	16.8	1831.8	16.8	101.2	-1.2	
LCC3-H973-78	31	24454	1.0	8.9301	2.0	4.6449	2.4	0.3008	1.4	0.59	1823.5	21.2	1757.4	20.3	1831.8	35.7	1831.8	35.7	93.6	7.4	
LCC3-H973-98	108	193277	1.9	8.9186	0.6	5.1200	1.1	0.3313	1.0	0.64	1844.7	15.4	1838.8	9.7	1834.1	11.4	1834.1	11.4	100.6	-0.6	
LCC3-H973-23	50	50771	1.0	8.8953	0.8	5.0775	1.2	0.3278	0.9	0.75	1826.6	14.6	1832.4	10.4	1838.9	14.6	1838.9	14.6	99.3	0.7	
LCC3-H973-121	294	34843	3.3	8.8865	0.3	5.1481	1.2	0.3318	1.1	0.86	1847.1	18.0	1844.1	10.0	1840.7	6.3	1840.7	6.3	100.3	-0.3	
LCC3-H973-192	89	61697	2.0	8.8755	0.7	5.1890	1.5	0.3340	1.3	0.86	1857.8	20.3	1850.8	12.4	1842.9	13.4	1842.9	13.4	100.8	-0.8	
LCC3-H973-53	46	54628	0.7	8.8740	1.1	5.2185	1.6	0.3359	1.2	0.72	1866.7	18.6	1855.6	13.6	1843.2	20.0	1843.2	20.0	101.3	-1.3	
LCC3-H973-172	90	127628	1.0	8.8735	0.8	5.1778	1.0	0.3332	0.7	0.65	1854.0	10.8	1849.0	8.8	1843.3	14.1	1843.3	14.1	100.6	-0.6	
LCC3-H973-124	26	26506	1.1	8.8697	1.5	5.2267	2.0	0.3362	1.3	0.66	1868.5	21.7	1857.0	17.4	1844.1	27.9	1844.1	27.9	101.3	-1.3	
LCC3-H973-103	138	42283	0.8	8.8676	0.5	4.7522	0.8	0.3056	0.6	0.73	1719.2	8.5	1776.5	6.5	1844.5	9.6	1844.5	9.6	93.2	6.8	
LCC3-H973-167	151	103516	2.1	8.8527	0.4	4.9626	3.1	0.3186	3.1	0.99	1783.0	48.6	1813.0	26.6	1847.6	7.7	1847.6	7.7	96.5	3.5	
LCC3-H973-99	95	109548	2.2	8.8442	0.4	4.2848	0.8	0.2748	0.7	0.89	1585.4	10.2	1690.4	6.8	1849.3	6.7	1849.3	6.7	84.6	15.4	
LCC3-H973-118	65	70522	1.3	8.8379	1.3	4.6333	1.6	0.3270	2.6	0.89	1876.4	48.1	1755.3	24.2	1850.6	24.1	1850.6	24.1	90.6	9.8	
LCC3-H973-195	24	18635	0.4	8.8332	2.5	5.2903	2.9	0.3389	5.5	0.50	1881.5	24.0	1867.3	25.0	1851.5	45.7	1851.5	45.7	101.6	-1.6	
LCC3-H973-21	221	15230	1.4	8.8323	1.0	5.2246	2.0	0.3347	1.7	0.87	1861.0	27.7	1856.6	16.8	1851.7	17.6	1851.7	17.6	100.5	-0.5	
LCC3-H973-90	23	26557	0.7	8.8306	2.3	5.2527	2.7	0.3364	1.5	0.55	1869.2	24.1	1861.1	23.0	1852.1	40.7	1852.1	40.7	100.9	-0.9	
LCC3-H973-117	132	126267	1.9	8.8284	0.6	5.0822	2.4	0.3254	2.3	0.97	1816.1	36.5	1833.2	20.2	1852.0	10.8	1852.0	10.8	98.0	2.0	
LCC3-H973-196	74	47358	0.7	8.8258	0.9	5.2517	1.2	0.3362	0.7	0.59	1868.2	11.2	1861.0	9.9	1853.1	16.9	1853.1	16.9	100.8	-0.8	
LCC3-H973-59	21	24552	0.8	8.8149	2.0	5.2109	2.4	0.3331	1.4	0.57	1853.6	22.1	1854.4	20.5	1853.3	35.7	1853.3	35.7	99.9	0.1	
LCC3-H973-96	73	155037	0.6	8.8105	0.8	5.1881	1.2	0.3315	0.9	0.74	1845.7	13.8	1850.7	9.9	1856.2	14.3	1856.2	14.3	99.4	0.6	
LCC3-H973-29	122	6297	1.1	8.8037	1.1	4.3698	2.1	0.2790	1.8	0.84	1586.4	24.9	1706.6	17.4	1857.6	20.7	1857.6	20.7	85.4	14.6	
LCC3-H973-170	19	13201	0.4	8.7931	2.1	5.1740	2.5	0.3300	1.4	0.55	1838.2	22.0	1848.4	21.2	1859.8	37.4	1859.8	37.4	98.8	1.2	
LCC3-H973-109	98	92474	2.9	8.7888	0.5	5.2893	0.8	0.3317	0.7	0.65	1879.6	11.7	1867.1	7.2	1861.0	8.1	1861.0	8.1	100.6	-0.6	
LCC3-H973-44	102	89292	0.7	8.7812	0.4	5.2130	2.4	0.3320	2.3	0.98	1848.7	37.5	1854.7	20.7	1862.2	7.7	1862.2	7.7	99.2	0.8	
LCC3-H973-122	48	34121	0.9	8.7758	1.2	5.2352	1.9	0.3332	1.4	0.78	1853.9	23.3	1858.4	15.8	1863.3	21.0	1863.3	21.0	99.5	0.5	
LCC3-H973-11	64	55240	1.0	8.7742	0.6	5.3449	1.4	0.3401	1.2	0.89	1887.3	20.3	1876.1	11.9	1863.7	11.4	1863.7	11.4	101.3	-1.3	
LCC3-H973-1	111	150846	1.6	8.7396	0.3	5.3061	0.8	0.3363	0.7	0.93	1889.0	11.5	1869.8	6.5	1870.8	5.2	1870.8	5.2	99.9	0.1	
LCC3-H973-143	82	88944	1.0	8.7276	0.7	5.4370	1.3	0.3442	1.1	0.83	1906.6	17.9	1890.7	11.2	1873.3	13.0	1873.3	13.0	101.8	-1.8	
LCC3-H973-75	98	2315	0.7	8.7122	1.3	4.6252	3.6	0.2923	3.4	0.93	1652.8	49.2	1753.8	30.3	1876.4	23.9	1876.4	23.9	88.1	11.9	
LCC3-H973-43	90	101565	0.7	8.6983	1.0	5.2695	1.1	0.3324	0.5	0.45	1850.1	8.0	1863.9	9.3	1879.3	17.5	1879.3	17.5	98.4	1.6	
LCC3-H973-95	60	64431	0.8	8.6233	0.8	5.4544	1.1	0.3411	0.8	0.73	1892.1	13.8	1893.4	9.7	1894.9	13.9	1894.9	13.9	99.9	0.1	
LCC3-H973-135	82	96248	1.1	8.5990	0.8	5.4866	1.4	0.3423	1.1	0.79	1897.7	18.1	1898.8	11.9	1900.0	15.2	1900.0	15.2	99.9	0.1	
LCC3-H973-94	77	21290	1.5	8.5484	3.9	4.9428	5.9	0.3064	4.4	0.75	1723.2	68.7	1809.6	49.8	1910.6	69.4	1910.6	69.4	90.2	9.8	
LCC3-H973-22	31	31965	1.4	8.3769	4.3	5.3621	5.1	0.3258	7.7	0.54	1817.9	45.3	1878.8	43.5	1946.9	78.7	1946.9	78.7	93.4	6.6	
LCC3-H973-130	31	42475	2.2	8.2877	1.6	5.8864	2.1	0.3538	1.4	0.65	1952.8	22.9	1959.2	18.0	1968.8	28.0	1968.8	28.0	99.3	0.7	
LCC3-H973-100	51	86313	0.8	8.8635	0.9	6.6044	1.2	0.3767	0.9	0.71	2060.7	15.2	2059.9	10.8	2059.2	15.3	2059.2	15.3	100.1	-0.1	
LCC3-H973-93	14	21093	0.9	7.7030	3.8	7.0521	4.1	0.3940	1.5	0.38	2141.3	27.8	2118.0	36.1	2095.5	66.0	2095.5	66.0	102.2	-2.2	
LCC3-H973-110	163	46433	1.2	7.5945	0.8	6.6702	5.2	0.3674	5.2	0.99	2017.1	89.6	2068.7	46.3	2124.0	14.8	2124.0	14.8	95.1	4.9	
LCC3-H973-177	65	90579	3.1	7.0633	0.6	7.8378	1.4	0.4015	1.3	0.91	2176.0	24.0	2212.6	12.9	2246.6	10.3	2246.6	10.3	96.9	3.1	
LCC3-H973-64	114	117601	1.3	6.8296	0.6	8.3231	4.7	0.4123	4.7	0.99	2225.3	88.2	2266.8	42.8	2304.5	9.5	2304.5	9.5	96.6	3.4	
LCC3-H973-193	267	18480	0.8	6.8289	0.5	8.7417	1.7	0.4330	1.7	0.96	2319.0	32.2	2311.4	15.6	2304.7	7.8	2304.7	7.8	100.6	-0.6	
LCC3-H973-146	177	271278	1.2	6.7574	0.2	8.7842	1.1	0.4305	1.1	0.98	2308.0	21.1	2315.9	10.1	2322.8	4.0	2322.8	4.0	99.4	0.6	
LCC3-H973-1																					



## U-Pb Geochronologic analyses of selected Harmony Formation strata

	Isotope ratios											Apparent ages (Ma)											Conc.	Discor.
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	Best age	±	Conc.	Discor.				
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)				
<b>Sample: LCC #3. Location: Little Cottonwood Canyon, Galena Range: 0490459 4494916 (NAD 83 UTM 11T)</b>																								
LCC3-H973-108	110	357936	1.6	5.4627	0.2	12.6260	0.5	0.5002	0.5	0.90	2614.8	10.3	2652.2	5.0	2680.8	3.8	2680.8	3.8	97.5	2.5				
LCC3-H973-148	209	23237	1.0	5.4598	0.2	13.2400	1.8	0.5243	1.7	0.99	2717.3	38.7	2696.9	16.6	2681.6	3.0	2681.6	3.0	101.3	-1.3				
LCC3-H973-127	16	10317	0.7	5.4501	1.4	13.0248	1.7	0.5148	1.0	0.58	2677.3	21.5	2681.5	16.0	2684.6	22.8	2684.6	22.8	99.7	0.3				
LCC3-H973-9	24	69676	1.2	5.4384	0.5	12.9251	1.6	0.5098	1.5	0.95	2655.8	32.4	2674.2	14.7	2688.2	7.8	2688.2	7.8	98.8	1.2				
LCC3-H973-77	29	78596	0.5	5.4341	0.9	13.1620	1.4	0.5187	1.1	0.78	2693.8	23.9	2691.3	13.2	2689.5	14.6	2689.5	14.6	100.2	-0.2				
LCC3-H973-57	24	30729	1.1	5.4259	0.8	13.2228	1.6	0.5203	1.4	0.87	2700.6	30.7	2695.7	15.2	2692.0	13.3	2692.0	13.3	100.3	-0.3				
LCC3-H973-27	46	72568	1.5	5.4210	0.6	13.0786	0.9	0.5142	0.7	0.77	2674.6	15.3	2685.3	8.6	2693.4	9.7	2693.4	9.7	99.3	0.7				
LCC3-H973-166	79	91251	1.4	5.3881	0.2	12.5344	1.0	0.4898	1.0	0.98	2589.9	20.5	2645.3	9.3	2703.5	3.4	2703.5	3.4	95.1	4.9				
LCC3-H973-80	15	13488	0.3	5.3831	1.0	13.2351	3.7	0.5167	3.6	0.96	2685.3	79.0	2696.6	35.3	2705.0	16.7	2705.0	16.7	99.3	0.7				
LCC3-H973-37	97	220310	1.4	5.3598	0.3	13.3717	1.0	0.5198	1.0	0.97	2698.3	21.0	2706.3	9.3	2712.2	4.1	2712.2	4.1	99.5	0.5				
LCC3-H973-163	177	13382	1.9	5.3594	0.3	13.5407	0.8	0.5263	0.7	0.94	2726.0	16.4	2718.1	7.4	2712.3	4.4	2712.3	4.4	100.5	-0.5				
LCC3-H973-157	175	113946	1.0	5.3410	1.0	14.1216	3.0	0.5470	3.0	1.00	2812.8	69.4	2757.9	28.9	2718.0	2.3	2718.0	2.3	100.5	-3.5				
LCC3-H973-70	92	142962	0.8	5.3395	0.2	13.5657	0.9	0.5253	0.8	0.97	2721.8	18.4	2719.9	8.1	2718.4	3.7	2718.4	3.7	100.1	-0.1				
LCC3-H973-101	35	40920	1.6	5.3386	0.7	14.2757	2.7	0.5527	2.7	0.97	2838.6	60.9	2768.2	26.0	2718.7	11.2	2718.7	11.2	104.3	-4.3				
LCC3-H973-83	47	118041	1.2	5.3154	0.5	12.5486	2.4	0.4838	2.3	0.97	2543.6	48.4	2646.4	22.7	2725.9	9.0	2725.9	9.0	93.3	6.7				
LCC3-H973-40	81	136102	1.6	5.2968	0.3	13.7024	0.7	0.5264	0.6	0.89	2726.1	13.4	2729.4	6.4	2731.7	5.0	2731.7	5.0	99.8	0.2				
LCC3-H973-4	28	92323	1.5	5.2861	0.4	13.6849	1.2	0.5247	1.2	0.95	2718.9	26.1	2728.1	11.7	2735.0	6.2	2735.0	6.2	99.4	0.6				
LCC3-H973-87	50	71229	1.6	5.2709	0.6	13.9752	0.8	0.5343	0.5	0.65	2759.3	12.3	2748.0	8.0	2739.7	10.5	2739.7	10.5	100.7	-0.7				
LCC3-H973-91	157	147286	3.9	5.2707	0.7	13.7061	4.4	0.5239	4.3	0.99	2715.8	96.2	2729.6	41.6	2739.8	10.8	2739.8	10.8	99.1	0.9				
LCC3-H973-187	28	57696	1.1	4.9625	0.5	15.1794	1.2	0.5463	1.1	0.91	2809.9	24.4	2826.6	11.2	2838.5	7.8	2838.5	7.8	99.0	1.0				
LCC3-H973-155	55	83128	0.8	4.9602	0.6	15.8261	1.2	0.5693	1.0	0.87	2905.1	23.6	2866.4	11.0	2839.2	9.2	2839.2	9.2	102.3	-2.3				
LCC3-H973-140	53	131907	1.6	4.6767	0.3	16.6499	1.4	0.5647	1.4	0.97	2886.2	31.7	2814.9	13.4	2934.8	5.1	2934.8	5.1	98.9	1.7				
LCC3-H973-67	8	12793	0.3	4.5888	1.8	17.3004	2.9	0.5756	2.2	0.78	2931.5	52.2	2951.6	27.4	2965.4	29.1	2965.4	29.1	99.3	1.1				
LCC3-H973-111	150	129946	0.5	4.0842	0.3	16.8515	1.8	0.4967	1.8	0.88	2599.7	37.5	2826.4	17.1	3159.5	5.5	3159.5	5.5	82.3	17.7				
LCC3-H973-82	110	254422	1.2	3.5012	0.4	26.5414	2.1	0.7441	2.1	0.98	3321.2	64.4	3365.7	33.9	3393.9	6.7	3393.9	6.7	97.9	2.1				
LCC3-H973-17	108	208359	1.1	2.9199	0.2	35.7474	1.3	0.7570	1.3	0.99	3633.4	35.5	3659.5	12.8	3673.8	3.4	3673.8	3.4	98.9	1.1				
<b>Sample: LCC #4. Location: Little Cottonwood Canyon, Galena Range: 0490623 4494977 (NAD 83 UTM 11T)</b>																								
LCC4-88RM13-122	164	63542	1.7	12.7286	1.1	2.1680	3.0	0.2001	2.8	0.93	1176.1	29.6	1170.9	20.5	1161.1	21.2	1161.1	21.2	101.3	-1.3				
LCC4-88RM13-157	105	49618	1.1	9.5690	0.8	0.6960	3.3	0.2824	3.2	0.97	1603.4	45.3	1648.1	26.8	1705.6	14.2	1705.6	14.2	94.0	6.0				
LCC4-88RM13-80	18	7740	0.8	9.4151	3.2	4.6997	4.7	0.3209	3.4	0.73	1794.2	53.3	1767.2	39.3	1735.4	59.2	1735.4	59.2	103.4	-3.4				
LCC4-88RM13-52	182	24329	1.6	9.3259	0.5	1.5554	5.5	0.2811	5.5	0.99	1596.7	77.1	1665.3	44.9	1752.8	10.1	1752.8	10.1	91.1	8.9				
LCC4-88RM13-38	407	14111	1.5	9.3087	0.2	3.8655	2.5	0.2610	2.5	1.00	1494.8	33.2	1606.5	20.1	1756.2	4.4	1756.2	4.4	85.1	14.9				
LCC4-88RM13-46	52	35011	1.1	9.3028	1.1	4.8549	1.6	0.3276	1.1	0.71	1826.6	17.5	1794.5	13.1	1757.4	20.0	1757.4	20.0	103.9	-3.9				
LCC4-88RM13-61	482	123777	3.9	9.2839	0.3	3.8328	3.4	0.2581	3.4	1.00	1480.0	45.1	1599.7	27.5	1761.1	4.6	1761.1	4.6	84.0	16.0				
LCC4-88RM13-13	375	10400	1.0	9.2818	0.4	4.0430	6.9	0.2722	6.9	1.00	1551.8	95.4	1642.9	56.5	1761.5	7.4	1761.5	7.4	88.1	11.9				
LCC4-88RM13-42	92	35125	1.2	9.2576	0.4	4.4184	0.7	0.2967	0.7	0.87	1674.9	9.5	1715.8	6.1	1766.3	6.6	1766.3	6.6	94.8	5.3				
LCC4-88RM13-94	355	15842	2.7	9.2495	0.4	4.4413	3.3	0.2979	3.3	0.99	1681.1	48.3	1720.1	27.2	1767.9	7.0	1767.9	7.0	95.1	4.9				
LCC4-88RM13-57	205	36400	1.0	9.2397	0.4	4.8706	0.8	0.3264	0.7	0.84	1820.9	10.8	1797.2	6.8	1769.8	8.0	1769.8	8.0	102.9	-2.9				
LCC4-88RM13-24	35	23766	1.0	9.2260	1.2	4.7864	1.6	0.3189	1.1	0.70	1784.5	17.8	1779.0	13.8	1772.5	21.5	1772.5	21.5	100.7	-0.7				
LCC4-88RM13-54	115	55259	1.1	9.2235	0.6	4.8130	1.1	0.3220	0.9	0.83	1799.3	14.0	1787.2	9.0	1773.0	10.8	1773.0	10.8	101.5	-1.5				
LCC4-88RM13-82	211	29672	2.3	9.2191	0.6	4.4530	0.8	0.2977	0.5	0.60	1680.1	6.8	1722.3	6.4	1773.9	11.4	1773.9	11.4	94.7	5.3				
LCC4-88RM13-108	38	35099	0.6	9.2177	1.0	4.9242	2.2	0.3292	1.9	0.89	1834.5	30.7	1806.4	18.2	1774.1	17.9	1774.1	17.9	103.4	-3.4				
LCC4-88RM13-48	193	14078	0.6	9.2168	0.4	4.3541	1.5	0.2911	1.4	0.96	1646.8	20.4	1703.7	12.1	1774.3	7.5	1774.3	7.5	92.8	7.2				
LCC4-88RM13-39	87	31614	1.5	9.2108	0.4	4.8638	1.3	0.3249	1.2	0.95	1813.7	19.5	1796.0	10.9	1775.5	7.4	1775.5	7.4	102.1	-2.1				
LCC4-88RM13-86	245	22636	1.3	9.2054	0.4	4.4574	1.1	0.2976	1.0	0.94	1679.3	15.3	1723.1	9.1	1776.6	6.6	1776.6	6.6	94.5	5.5				
LCC4-88RM13-177	113	52669	0.7	9.2026	0.6	4.7746	0.8	0.3187	0.8	0.74	1783.2	13.8	1792.4	6.9	1777.1	10.2	1777.1	10.2	100.3	-0.3				
LCC4-88RM13-166	76	54241	1.0	9.1977	0.8	4.7655	1.1	0.3179	0.7	0.63	1779.4	10.5	1778.8	8.9	1778.1	15.0	1778.1	15.0	100.1	-0.1				
LCC4-88RM13-130	74	30558	1.4	9.1975	0.9	4.9571	1.7	0.3307	1.5	0.86	1841.6	23.6	1812.0	14.4	1776.2	15.7	1776.2	15.7	103.6	-3.6				
LCC4-88RM13-71	176	21688	3.2	9.1930	0.3	4.8430	1.6	0.3229	1.6	0.98	1803.9	24.4	1792.4	13.3	1779.0	5.1	1779.0	5.1	101.4	-1.4				
LCC4-88RM13-87	85	8108	1.6	9.1924	0.9	4.7628	3.0	0.3175	2.8	0.95	1777.7	44.1	1778.4	25.1	1779.2	16.8	1779.2	16.8	99.9	0.1				
LCC4-88RM13-191	301	38477	2.5	9.1912	0.3	4.4010	0.8	0.2934	0.8	0.94	1658.4	11.1	1712.5	6.7	1779.4	4.9	1779.4	4.9	93.2	6.8				
LCC4-88RM13-183	197	38098	1.4	9.1882	0.3	4.7635	1.1	0																

## U-Pb Geochronologic analyses of selected Harmony Formation strata

	Isotope ratios														Apparent ages (Ma)									
	U (ppm)	206Pb 204Pb	U/Th	206Pb* 207Pb*	± (%)	207Pb* 235U*	± (%)	206Pb* 238U*	± (%)	error corr.	206Pb* 238U*	± (Ma)	207Pb* 235U	± (Ma)	208Pb* 207Pb*	± (Ma)	Best age (Ma)	± (Ma)	Conc. (%)	Discor. (%)				
																					206Pb* 207Pb*	± (%)	206Pb* 238U*	± (%)
<b>Sample: LCC #4. Location: Little Cottonwood Canyon, Galena Range: 0490623 4494977 (NAD 83 UTM 11T)</b>																								
LCC4-88RM13-74	71	49204	3.7	8.6461	0.5	5.6543	1.0	0.3546	0.8	0.83	1956.4	13.6	1924.4	8.4	1890.2	9.9	1890.2	9.9	103.5	-3.5				
LCC4-88RM13-14	18	26282	0.8	8.5246	2.2	5.5974	2.4	0.3461	0.9	0.38	1915.8	14.8	1915.7	20.4	1915.6	39.3	1915.6	39.3	100.0	0.0				
LCC4-88RM13-141	42	47070	4.0	8.4410	1.0	5.9362	1.1	0.3450	0.4	0.38	1910.9	6.8	1821.7	9.3	1833.2	17.8	1833.2	17.8	98.8	1.2				
LCC4-88RM13-75	101	30628	0.9	7.8383	0.5	6.7180	1.7	0.3819	1.7	0.95	2085.2	29.4	2075.0	15.4	2064.9	9.6	2064.9	9.6	101.0	-1.0				
LCC4-88RM13-179	254	23372	1.9	7.8005	2.1	5.5538	2.7	0.3142	1.7	0.64	1761.4	26.5	1909.0	23.3	2073.4	36.8	2073.4	36.8	85.0	15.0				
LCC4-88RM13-122	122	71177	1.8	7.7097	0.4	6.9111	0.8	0.3864	0.6	0.84	2106.3	11.6	2100.1	6.8	2094.0	7.4	2094.0	7.4	100.0	-0.6				
LCC4-88RM13-178	91	89611	1.6	6.5373	0.8	8.9396	1.8	0.4239	1.6	0.90	2277.9	31.2	2331.9	16.4	2329.4	13.1	2379.4	13.1	95.7	4.3				
LCC4-88RM13-13	359	32405	2.5	6.3819	2.4	9.5544	7.4	0.4422	7.0	0.95	2360.6	138.1	2392.8	67.9	2420.3	40.0	2420.3	40.0	97.5	2.5				
LCC4-88RM13-120	174	24333	2.1	6.3179	0.8	9.0122	2.4	0.4130	2.2	0.95	2228.4	42.0	2339.2	21.5	2437.4	12.8	2437.4	12.8	91.4	8.6				
LCC4-88RM13-43	195	14744	0.7	6.2962	0.7	7.9881	3.2	0.3648	3.1	0.97	2004.8	53.7	2229.7	28.9	2443.2	12.0	2443.2	12.0	82.1	17.9				
LCC4-88RM13-133	308	16722	0.8	6.2622	0.3	8.8874	3.0	0.4036	3.0	1.00	2185.8	54.8	2326.5	27.1	2452.3	4.9	2452.3	4.9	89.1	10.9				
LCC4-88RM13-9	182	53176	1.4	6.0618	0.2	10.8600	0.7	0.4775	0.6	0.93	2516.2	12.8	2511.2	6.2	2507.2	4.2	2507.2	4.2	100.0	-0.4				
LCC4-88RM13-38	21	14830	1.8	6.0487	2.0	10.2875	2.7	0.4513	1.9	0.69	2401.0	37.7	2461.0	25.1	2510.8	32.9	2510.8	32.9	95.6	4.4				
LCC4-88RM13-19	191	32279	0.9	6.9507	0.8	10.4833	1.5	0.4529	1.3	0.95	2408.0	25.6	2479.3	6.9	2538.3	13.3	2538.3	13.3	94.9	5.1				
LCC4-88RM13-151	64	50122	1.8	5.8883	0.6	11.0957	1.3	0.4739	1.1	0.88	2500.8	23.8	2531.2	12.2	2555.7	10.4	2555.7	10.4	97.9	2.1				
LCC4-88RM13-115	205	27165	1.1	5.8676	1.0	10.7859	1.8	0.4590	1.7	0.95	2435.1	35.2	2504.9	17.0	2561.8	9.8	2561.8	9.8	95.1	4.9				
LCC4-88RM13-83	55	64497	0.9	5.8622	1.3	11.1763	3.5	0.4752	3.2	0.93	2506.3	66.6	2537.9	32.2	2563.4	21.7	2563.4	21.7	97.8	2.2				
LCC4-88RM13-126	235	24306	1.0	5.8168	0.4	10.7707	4.6	0.4544	4.6	1.00	2414.7	92.8	2503.5	43.0	2576.4	6.3	2576.4	6.3	93.7	6.3				
LCC4-88RM13-192	132	34846	1.5	5.8127	0.5	10.6774	3.1	0.4501	3.1	0.99	2395.9	61.6	2495.5	29.0	2577.5	8.6	2577.5	8.6	93.0	7.0				
LCC4-88RM13-32	264	18331	1.2	5.8076	0.1	11.0938	0.7	0.4673	0.7	0.98	2471.6	13.4	2531.0	6.2	2579.0	2.2	2579.0	2.2	95.8	4.2				
LCC4-88RM13-185	126	27904	1.1	5.7988	0.3	9.6784	1.2	0.4070	1.2	0.96	2201.4	21.5	2404.7	11.0	2581.6	5.5	2581.6	5.5	85.3	14.7				
LCC4-88RM13-154	61	31979	0.9	5.7877	0.6	11.7012	1.2	0.4920	1.0	0.86	2579.4	22.1	2580.8	11.3	2581.9	10.4	2581.9	10.4	99.0	0.1				
LCC4-88RM13-11	23	25181	1.1	5.7835	1.3	11.3457	2.7	0.4759	2.4	0.87	2509.4	48.9	2552.0	25.1	2586.0	21.9	2586.0	21.9	97.9	3.0				
LCC4-88RM13-44	186	19921	1.0	5.6951	0.2	10.4846	1.2	0.4647	1.1	0.97	2495.5	23.0	2530.1	10.9	2586.4	4.3	2586.4	4.3	95.1	4.9				
LCC4-88RM13-22	120	39399	1.5	5.7811	0.4	11.3616	0.6	0.4764	0.4	0.71	2511.5	8.3	2553.3	5.2	2586.6	6.6	2586.6	6.6	97.1	2.9				
LCC4-88RM13-81	146	77202	1.0	5.7584	0.2	11.9238	0.8	0.4980	0.8	0.98	2605.1	16.4	2598.4	7.3	2593.2	2.6	2593.2	2.6	100.0	-0.5				
LCC4-88RM13-8	244	27697	1.2	5.7505	0.2	10.7217	1.2	0.4472	1.2	0.99	2382.6	24.2	2499.3	11.5	2595.5	3.4	2595.5	3.4	91.8	8.2				
LCC4-88RM13-163	75	101593	2.1	5.7474	0.7	11.9667	1.1	0.4988	0.8	0.76	2608.7	18.1	2601.8	10.4	2596.4	12.0	2596.4	12.0	100.0	-0.5				
LCC4-88RM13-73	123	15791	0.8	5.7464	0.3	9.5333	1.2	0.3973	1.1	0.96	2156.7	20.5	2390.8	10.7	2596.7	5.5	2596.7	5.5	83.1	16.9				
LCC4-88RM13-199	163	63955	1.0	5.7437	0.2	12.2107	4.8	0.5087	4.7	1.00	2650.9	103.2	2620.7	44.6	2597.5	3.9	2597.5	3.9	102.1	-2.1				
LCC4-88RM13-45	112	184197	1.1	5.7055	0.3	12.4651	1.0	0.5158	1.0	0.96	2681.4	21.0	2640.1	9.4	2608.6	4.9	2608.6	4.9	102.8	-2.8				
LCC4-88RM13-79	124	68769	1.9	5.6980	0.3	12.2531	1.1	0.5064	1.1	0.97	2641.1	23.9	2624.0	10.6	2610.8	4.5	2610.8	4.5	101.2	-1.2				
LCC4-88RM13-57	106	186879	1.9	5.6636	3.8	11.7190	5.4	0.4814	3.8	0.71	2533.2	80.4	2592.2	50.6	2620.9	63.2	2620.9	63.2	96.7	3.3				
LCC4-88RM13-149	159	34833	0.9	5.6561	0.2	10.4846	1.2	0.4647	1.1	0.97	2495.5	23.0	2530.1	10.9	2586.4	4.3	2586.4	4.3	95.1	4.9				
LCC4-88RM13-7	157	283348	1.4	5.6432	0.2	12.2830	0.5	0.5027	0.5	0.95	2625.5	8.7	2626.3	4.9	2626.9	3.0	2626.9	3.0	99.9	0.1				
LCC4-88RM13-90	81	149965	5.2	5.6222	0.3	13.0026	1.2	0.5245	1.2	0.96	2718.4	26.5	2679.8	11.7	2650.9	5.8	2650.9	5.8	102.5	-2.5				
LCC4-88RM13-78	80	55216	1.2	5.4866	0.6	13.0765	6.9	0.5204	6.9	1.00	2700.7	152.8	2685.2	65.6	2673.5	9.9	2673.5	9.9	101.0	-1.0				
LCC4-88RM13-84	130	12290	1.8	5.4749	0.6	10.9983	5.5	0.4367	5.4	0.99	2335.9	106.8	2523.0	51.1	2677.1	9.8	2677.1	9.8	87.3	12.7				
LCC4-88RM13-118	86	120604	1.0	5.4444	0.4	13.7106	1.4	0.5414	1.4	0.97	2789.2	31.0	2729.9	13.4	2686.3	5.8	2686.3	5.8	103.8	-3.8				
LCC4-88RM13-172	155	15888	0.8	5.4353	0.3	10.1149	2.1	0.3987	2.0	0.99	2163.2	37.3	2445.3	19.0	2689.1	5.2	2689.1	5.2	80.4	19.6				
LCC4-88RM13-16	146	18934	1.0	5.4215	0.3	12.1920	1.6	0.4794	1.6	0.99	2524.6	33.0	2619.3	15.0	2693.3	4.5	2693.3	4.5	93.7	6.3				
LCC4-88RM13-60	36	37590	0.4	5.4213	1.2	13.6816	2.8	0.5379	2.6	0.91	2774.8	57.9	2727.9	26.7	2693.4	19.0	2693.4	19.0	103.0	-3.0				
LCC4-88RM13-169	103	18603	0.6	5.4158	0.3	10.1913	2.3	0.4003	2.3	0.99	2170.5	42.4	2452.3	21.4	2695.0	4.2	2695.0	4.2	80.5	19.5				
LCC4-88RM13-194	72	48567	0.6	5.3970	0.3	13.0560	0.8	0.5275	0.8	0.93	2730.7	17.1	2713.5	7.8	2700.8	5.1	2700.8	5.1	101.1	-1.1				
LCC4-88RM13-134	51	51777	2.4	5.2944	0.6	14.1070	1.1	0.5417	1.0	0.87	2790.5	22.6	2756.9	10.9	2732.3	9.4	2732.4	9.4	102.1	-2.1				
LCC4-88RM13-136	52	29452	1.5	5.2534	0.4	13.8998	1.0	0.5296	0.9	0.92	2739.7	19.8	2742.9	9.0	2742.6	6.0	2742.6	6.0	99.8	0.2				
LCC4-88RM13-127	119	103494	1.8	5.1717	0.3	14.7895	0.9	0.5547	0.9	0.95	2844.8	20.6	2801.8	8.9	2770.9	4.6	2770.9	4.6	102.7	-2.7				
LCC4-88RM13-180	29	41080	1.1	5.0667	0.9	14.7251	1.0	0.5411	0.6	0.56	2788.1	13.3	2797.6	10.0	2804.5	14.2	2804.5	14.2	99.4	0.6				
LCC4-88RM13-12	88	37914	0.7	4.9035	0.4	13.3720	1.4	0.4756	1.4	0.95	2507.9	26.6	2706.3	13.6	2857.9	7.0	2857.9	7.0	87.8	12.2				
LCC4-88RM13-150	114	202774	1.7	4.8374	0.2	15.6955	1.3	0.5507	1.3	0.98	2827.9	29.2	2858.4	12.4	2880.0	4.1	2880.0	4.1	98.2	1.8				
<b>Sample: LCC #9. Location: Little Cottonwood Canyon, Galena Range: 0491232 4495602 (NAD 83 UTM 11T)</b>																								
LCC9-09HA-76	148	34636	1.6	9.5251	1.1	4.4450	3.5	0.3071	3.3	0.95	1726.3	49.8	1720.8	28.7	1714.0	19.9	1714.0	19.9	100.7	-0.7				
LCC9-09HA-9	253	12741	1.4	9.4099	0.5	3.5777	4.1	0.2442	4.1	0.99	1408.3	51.9	1544.6	32.9	1736.4	10.0	1736.4	10.0	81.1	18.9				
LCC9-09HA-51	420	7629	5.4	9.3216	0.5	3.9124	4.2	0.2645	4.2	0.99	1512.8	56.7	1616.3	34.3	1753.7	9.6	1753.7	9.6	86.3	13.7				
LCC9-09HA-65	434	24921	7.2	9.3056	0.4	3.9646	2.8	0.2765	2.7	0.99	1528.6	37.4	1627.0	22.6	1763.2	7.9	1763.2	7.9	97.9	12.0				
LCC9-09HA-83	333	20660	0.9	9.2859	0.8	3.2367	1.9	0.2180	1.7	0.91	1712.1	19.7	1746.0	14.6	1760.7	14.3	1760.7	14.3	72.2	27.8				
LCC9-																								



## U-Pb Geochronologic analyses of selected Harmony Formation strata

	Isotope ratios										Apparent ages (Ma)										Best age (Ma)	Conc. (%)	Discor. (%)
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±	207Pb*	±	206Pb*	±			
	(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	235U	(Ma)	207Pb*	(Ma)			
<b>Sample: LCC #9. Location: Little Cottonwood Canyon, Galena Range; 0491232 4495602 (NAD 83 UTM 11T)</b>																							
LCC-09HA-98	101	29345	1.1	8.6433	0.7	5.1826	4.0	0.3249	3.9	0.99	1813.5	61.8	1849.8	33.8	1890.7	11.8	1890.7	11.8	95.9	4.1			
LCC-09HA-41	119	27555	2.5	8.5544	1.0	5.7972	4.1	0.3597	3.9	0.97	1980.6	67.1	1940.6	35.1	1909.3	17.5	1909.3	17.5	103.7	-3.7			
LCC-09HA-40	90	10988	2.2	8.5298	1.2	4.9018	1.4	0.3032	0.8	0.56	1707.4	11.9	1802.6	11.8	1914.5	20.8	1914.5	20.8	89.2	10.8			
LCC-09HA-15	257	52961	1.1	8.4791	0.4	5.6289	2.3	0.3462	2.3	0.98	1916.2	37.6	1920.5	19.9	1925.2	7.9	1925.2	7.9	99.5	0.5			
LCC-09HA-80	148	44166	1.3	8.4778	0.9	5.9687	4.3	0.3870	4.2	0.98	2015.2	73.1	1971.3	37.5	1925.5	15.6	1925.5	15.6	104.7	-4.7			
LCC-09HA-36	134	52604	2.1	7.0730	0.9	7.6484	2.4	0.3923	2.2	0.93	2133.7	39.9	2190.6	21.3	2244.2	15.4	2244.2	15.4	95.1	4.9			
LCC-09HA-82	80	26866	1.2	6.1156	0.9	10.4224	2.4	0.4623	2.2	0.93	2449.6	44.4	2473.0	21.8	2492.3	14.9	2492.3	14.9	98.3	1.7			
LCC-09HA-75	221	48289	2.3	5.9916	0.5	10.4072	6.1	0.4522	6.1	1.00	2405.2	122.0	2471.7	56.5	2526.8	8.1	2526.8	8.1	95.2	4.8			
LCC-09HA-52	120	43744	3.2	5.7198	0.4	11.7241	3.2	0.4864	3.1	0.99	2554.9	66.0	2582.6	29.6	2604.4	7.3	2604.4	7.3	98.1	1.9			
LCC-09HA-54	223	22607	1.4	5.7195	0.4	10.1058	1.4	0.4192	1.3	0.95	2256.9	24.5	2444.5	12.6	2604.5	7.3	2604.5	7.3	86.7	13.3			
LCC-09HA-27	195	67476	2.0	5.7069	0.5	11.5631	0.9	0.4786	0.8	0.85	2521.2	16.6	2569.7	8.8	2608.2	8.2	2608.2	8.2	96.7	3.3			
LCC-09HA-71	171	65244	3.6	5.6653	0.3	11.8508	3.8	0.4869	3.8	1.00	2557.4	80.0	2592.7	35.6	2620.4	5.1	2620.4	5.1	97.6	2.4			
LCC-09HA-19	170	54182	1.3	5.6453	0.3	11.4288	1.9	0.4679	1.9	0.99	2474.5	39.2	2558.8	18.0	2626.2	5.2	2626.2	5.2	94.2	5.8			
LCC-09HA-7	128	19736	1.1	5.4432	1.2	8.8708	4.3	0.3502	4.1	0.96	1935.6	69.2	2324.8	39.4	2686.7	20.3	2686.7	20.3	72.0	28.0			
LCC-09HA-1	86	40571	1.4	5.3353	0.4	13.8045	1.2	0.5342	1.1	0.94	2759.0	24.4	2736.4	11.0	2719.7	6.8	2719.7	6.8	101.1	-1.4			
LCC-09HA-29	105	42749	1.6	2.7331	0.2	34.2676	0.5	0.6793	0.4	0.88	3341.5	11.6	3617.8	5.0	3774.5	3.7	3774.5	3.7	88.5	11.5			
<b>Sample: LCC #10. Location: Little Cottonwood Canyon, Galena Range; 0490510 4495682 (NAD 83 UTM 11T)</b>																							
LCC-10-HQ-86	533	33864	4.1	17.0454	1.9	0.7362	3.3	0.0910	2.7	0.82	561.5	14.7	560.2	14.3	554.8	40.9	561.5	14.7	101.2	-1.2			
LCC-10-HQ-97	115	14748	2.4	14.2236	3.0	1.6021	4.3	0.1653	3.2	0.73	986.0	29.0	971.0	27.2	937.3	60.9	937.3	60.9	105.2	-5.2			
LCC-10-HQ-50	103	11260	2.3	13.9139	3.1	1.6653	3.6	0.1681	1.8	0.49	1001.4	16.3	995.4	22.9	982.3	64.0	982.3	64.0	101.9	-1.9			
LCC-10-HQ-16	71	6330	2.5	13.8173	4.0	1.6603	4.4	0.1664	1.9	0.43	992.1	17.6	993.5	27.8	996.5	60.3	996.5	60.3	99.6	0.4			
LCC-10-HQ-18	98	11792	1.7	13.7737	3.4	1.6939	4.0	0.1692	2.0	0.51	1007.8	18.8	1006.2	25.4	1002.9	69.5	1002.9	69.5	100.5	-0.5			
LCC-10-HQ-23	85	10008	2.3	13.6900	4.5	1.8518	4.7	0.1839	1.2	0.27	1088.0	12.5	1064.1	30.8	1015.2	91.1	1015.2	91.1	107.2	-7.2			
LCC-10-HQ-89	229	27578	2.6	13.6207	1.0	1.7500	2.0	0.1729	1.7	0.86	1028.0	16.2	1027.2	12.8	1025.5	20.6	1025.5	20.6	100.2	-0.2			
LCC-10-HQ-41	267	29787	2.5	13.6083	1.6	1.7548	2.4	0.1732	1.8	0.74	1029.7	16.8	1028.9	15.6	1027.3	33.0	1027.3	33.0	100.2	-0.2			
LCC-10-HQ-33	266	35180	5.8	13.6082	1.0	1.7099	2.7	0.1688	2.5	0.92	1005.3	23.4	1012.2	17.4	1027.3	21.0	1027.3	21.0	97.9	2.1			
LCC-10-HQ-91	213	24710	2.1	13.5556	1.3	1.7255	3.2	0.1697	2.9	0.91	1010.7	26.8	1018.4	20.3	1036.6	26.8	1036.6	26.8	97.6	2.4			
LCC-10-HQ-78	142	17208	2.5	13.5256	1.5	1.9233	1.8	0.1886	0.9	0.52	1113.6	9.7	1088.9	12.1	1039.6	31.1	1039.6	31.1	107.1	-7.1			
LCC-10-HQ-54	159	19089	2.2	13.5182	1.5	1.7352	3.2	0.1701	2.8	0.89	1012.8	26.6	1021.7	20.5	1040.8	29.3	1040.8	29.3	97.3	2.7			
LCC-10-HQ-40	54	7937	2.1	13.4720	4.8	1.9284	5.3	0.1884	2.2	0.42	1112.8	22.6	1091.0	35.4	1047.7	97.1	1047.7	97.1	106.2	-6.2			
LCC-10-HQ-70	268	32089	4.6	13.4414	1.2	1.8429	2.6	0.1797	2.3	0.88	1065.1	22.4	1060.9	17.1	1052.3	24.8	1052.3	24.8	101.2	-1.2			
LCC-10-HQ-36	540	56405	5.8	13.4356	0.5	1.8504	1.9	0.1803	1.8	0.97	1068.7	17.8	1063.6	12.3	1053.1	9.4	1053.1	9.4	101.5	-1.5			
LCC-10-HQ-22	136	17385	2.0	13.4188	2.1	1.8696	2.3	0.1820	1.0	0.42	1077.6	9.6	1070.4	15.2	1055.7	41.8	1055.7	41.8	102.1	-2.1			
LCC-10-HQ-09	147	17896	1.0	13.4116	1.2	1.8614	2.5	0.1811	2.2	0.88	1072.8	21.5	1067.5	16.4	1056.8	23.9	1056.8	23.9	101.5	-1.5			
LCC-10-HQ-07	112	11850	2.6	13.3655	2.3	1.8051	3.6	0.1750	2.7	0.76	1039.5	26.3	1047.3	23.5	1063.7	46.9	1063.7	46.9	97.7	2.3			
LCC-10-HQ-17	142	8968	2.4	13.3099	2.2	1.8075	3.7	0.1745	2.9	0.79	1036.8	28.1	1048.2	24.2	1072.1	45.2	1072.1	45.2	96.7	3.3			
LCC-10-HQ-59	236	28521	2.6	13.2598	1.2	1.8339	4.5	0.1762	4.4	0.97	1046.4	42.2	1057.7	29.8	1081.0	23.7	1081.0	23.7	96.8	3.2			
LCC-10-HQ-62	737	33135	2.2	13.1787	0.6	1.9371	2.3	0.1851	2.2	0.97	1095.0	22.6	1094.0	15.5	1091.9	12.1	1091.9	12.1	100.3	-0.3			
LCC-10-HQ-48	102	17271	3.8	13.1621	2.6	1.8960	4.7	0.1810	3.9	0.83	1072.4	38.2	1079.7	31.0	1094.5	52.1	1094.5	52.1	98.0	2.0			
LCC-10-HQ-100	227	35625	2.6	13.1518	0.7	1.9323	3.6	0.1843	3.5	0.98	1090.5	35.5	1092.3	24.2	1096.0	14.2	1096.0	14.2	99.5	0.5			
LCC-10-HQ-69	372	55727	3.1	13.1445	0.8	1.9592	1.9	0.1868	1.7	0.90	1103.9	17.6	1101.6	12.9	1097.1	16.3	1097.1	16.3	100.6	-0.6			
LCC-10-HQ-72	57	6928	1.8	13.0709	3.3	1.9685	4.5	0.1866	3.1	0.68	1103.0	31.1	1104.8	30.5	1108.3	66.9	1108.3	66.9	99.5	0.5			
LCC-10-HQ-85	53	7430	2.2	13.0444	3.1	1.9572	3.3	0.1852	1.1	0.33	1095.1	11.1	1100.9	22.4	1112.4	62.8	1112.4	62.8	98.4	1.6			
LCC-10-HQ-92	152	17089	2.9	12.9702	2.0	1.9394	8.4	0.1824	8.2	0.97	1080.3	81.3	1094.8	56.5	1123.8	40.7	1123.8	40.7	96.1	3.9			
LCC-10-HQ-01	502	63910	4.2	12.9025	0.5	2.0630	1.2	0.1930	1.1	0.91	1137.9	11.0	1136.6	7.9	1134.2	9.7	1134.2	9.7	100.3	-0.3			
LCC-10-HQ-47	86	9291	1.7	12.9009	3.6	1.9402	4.7	0.1815	3.0	0.84	1075.4	29.8	1085.1	31.6	1134.4	72.2	1134.4	72.2	94.8	5.2			
LCC-10-HQ-52	135	18179	2.0	12.8889	1.5	1.9653	3.2	0.1838	2.9	0.89	1087.8	28.5	1104.0	20.3	1136.3	29.2	1136.3	29.2	97.6	4.3			
LCC-10-HQ-60	197	21019	2.1	12.8495	1.1	2.0409	5.0	0.1902	4.9	0.98	1122.5	50.5	1129.3	34.2	1142.4	21.9	1142.4	21.9	98.3	1.7			
LCC-10-HQ-19	77	10624	2.3	12.8266	3.5	2.1548	3.9	0.2005	1.7	0.43	1177.8	18.3	1166.6	27.3	1145.9	70.4	1145.9	70.4	102.8	-2.8			
LCC-10-HQ-27	408	52605	2.9	12.7986	1.0	2.0535	2.7	0.1906	2.5	0.93	1124.7	25.5	1133.5	18.2	1150.3	19.9	1150.3	19.9	97.8	2.2			
LCC-10-HQ-95	202	19119	3.2	12.7513	1.4	2.1574	4.6	0.1995	4.4	0.95	1172.7	47.2	1167.4	32.1	1157.6	28.2	1157.6	28.2	101.3	-1.3			
LCC-10-HQ-99	156	12799	2.4	12.7107	1.8	2.0534	2.5	0.1893	1.8	0													

## U-Pb Geochronologic analyses of selected Harmony Formation strata

		Isotope ratios								Apparent ages (Ma)										
U	<sup>206</sup> Pb	U/Th	<sup>206</sup> Pb*	±	<sup>207</sup> Pb*	±	<sup>206</sup> Pb*	±	error	<sup>206</sup> Pb*	±	<sup>207</sup> Pb*	±	<sup>206</sup> Pb*	±	Best age	±	Conc.	Discor.	
(ppm)	204Pb		207Pb*	(%)	235U*	(%)	238U	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)	
<b>Sample: LCC #10. Location: Little Cottonwood Canyon, Galena Range; 0490510 4495682 (NAD 83 UTM 11T)</b>																				
LCC-10-HQ-39	462	104227	6.3	9.0721	0.2	4.8298	2.1	0.3178	2.1	0.99	1778.9	32.4	1790.1	17.7	1803.2	4.5	1803.2	4.5	98.7	1.3
LCC-10-HQ-31	201	41402	2.5	9.0634	0.5	4.8497	2.8	0.3188	2.7	0.98	1783.8	42.1	1793.6	23.2	1804.9	9.6	1804.9	9.6	98.8	1.2
LCC-10-HQ-33	463	70528	5.7	9.0388	0.4	4.9429	2.2	0.3240	2.1	0.99	1809.4	33.7	1809.8	18.3	1809.8	6.4	1809.8	6.4	100.0	0.0
LCC-10-HQ-73	246	41028	2.5	8.2894	0.4	5.7968	1.7	0.3498	1.6	0.97	1827.4	27.0	1845.9	14.4	1865.6	8.9	1865.6	8.9	98.1	1.9
LCC-10-HQ-06	82	16380	2.0	7.9671	1.3	9.2991	3.1	0.3640	2.9	0.92	2001.0	49.3	2018.3	27.4	2036.1	22.2	2036.1	22.2	98.3	1.7
LCC-10-HQ-29	72	12111	1.4	6.0141	1.0	10.9695	2.5	0.4785	1.3	0.97	2520.6	47.2	2520.5	23.0	2520.5	16.7	2520.5	16.7	100.0	0.0
LCC-10-HQ-28	82	21363	0.5	5.5824	0.7	11.0392	4.4	0.4470	4.3	0.99	2381.7	86.5	2526.4	41.0	2644.8	12.0	2644.8	12.0	90.1	9.9
LCC-10-HQ-02	29	10551	1.0	5.5316	1.0	12.8852	1.3	0.5169	0.8	0.95	2686.2	18.4	2671.3	12.2	2660.0	16.3	2660.0	16.3	101.0	-1.0
LCC-10-HQ-32	82	25072	2.0	5.4337	0.5	13.0495	7.1	0.5143	7.0	1.00	2674.8	154.4	2683.2	66.7	2689.6	7.9	2689.6	7.9	99.5	0.5
LCC-10-HQ-61	80	28074	2.2	5.4181	0.5	13.6582	1.5	0.5367	1.4	0.94	2769.7	30.9	2726.3	13.8	2694.3	8.2	2694.3	8.2	102.8	-2.8
LCC-10-HQ-93	274	52535	2.3	5.4003	0.2	12.5629	2.6	0.4928	2.6	1.00	2582.9	54.3	2648.9	24.1	2699.8	3.3	2699.8	3.3	95.7	4.3
LCC-10-HQ-82	56	21460	1.8	5.3148	0.5	13.1857	2.3	0.5083	2.2	0.97	2649.2	47.6	2693.0	21.3	2726.1	8.8	2726.1	8.8	97.2	2.8
LCC-10-HQ-83	53	30076	1.3	4.3131	0.6	18.0614	4.0	0.5650	3.9	0.99	2887.2	91.8	2993.0	38.4	3064.9	9.0	3064.9	9.0	94.2	5.8
LCC-10-HQ-67	178	56861	2.8	3.4889	0.6	27.4114	4.5	0.6396	4.4	0.99	3396.4	117.5	3398.3	44.0	3399.4	8.8	3399.4	8.8	99.9	0.1
<b>Sample: Gough's Canyon. Location: Gough's Canyon, Osgood Mountains; 0471723 4556266 (NAD 83 UTM 11T)</b>																				
GC02HA-42	180	22024	3.7	9.6916	1.0	4.2465	3.5	0.2985	3.3	0.96	1683.8	49.5	1683.1	28.6	1682.1	17.6	1682.1	17.6	100.1	-0.1
GC02HA-21	111	4729	1.9	9.3243	0.6	4.7832	2.0	0.3235	1.9	0.95	1806.6	30.5	1791.9	17.2	1753.1	11.9	1753.1	11.9	103.1	-3.1
GC02HA-22	191	5928	1.6	9.2697	0.5	4.6201	1.6	0.3106	1.5	0.96	1743.7	23.6	1752.9	13.5	1763.9	8.5	1763.9	8.5	98.9	1.1
GC02HA-83	149	9943	1.8	9.2339	0.8	4.9524	3.2	0.3317	3.1	0.97	1846.4	49.7	1811.2	27.1	1770.9	15.0	1770.9	15.0	104.3	-4.3
GC02HA-41	273	28599	3.1	9.2125	0.5	4.7496	1.9	0.3173	1.8	0.96	1776.2	28.6	1776.0	16.1	1775.2	9.8	100.1	9.8	100.1	-0.1
GC02HA-04	186	6410	1.0	9.2104	1.3	4.7135	2.1	0.3149	1.7	0.79	1764.6	25.6	1769.6	17.5	1775.6	23.1	1775.6	23.1	99.4	0.6
GC02HA-50	280	11142	2.6	9.1971	1.0	4.6602	2.3	0.3109	2.0	0.99	1744.9	31.1	1760.1	19.0	1778.2	18.5	1778.2	18.5	98.1	1.9
GC02HA-59	222	25328	1.9	9.1962	0.8	4.5972	1.1	0.3066	0.8	0.70	1724.1	11.8	1748.8	9.4	1778.4	14.7	1778.4	14.7	96.9	3.1
GC02HA-38	380	14526	3.4	9.1924	0.5	4.8739	2.1	0.3249	2.0	0.97	1813.8	31.5	1797.7	17.3	1779.2	8.6	100.1	8.6	100.1	-1.9
GC02HA-36	395	16307	1.8	9.1893	0.4	4.7370	1.5	0.3157	1.5	0.96	1768.7	22.8	1773.8	12.8	1779.8	7.4	1779.8	7.4	99.4	0.6
GC02HA-87	84	3583	1.3	9.1889	3.1	4.8530	3.3	0.3234	3.0	0.32	1806.4	16.2	1794.1	27.5	1779.8	56.4	1779.8	56.4	101.5	-1.5
GC02HA-03	215	8443	3.2	9.1868	0.6	4.6384	1.8	0.3091	1.7	0.94	1736.0	25.2	1756.2	14.7	1780.3	10.9	1780.3	10.9	97.5	2.5
GC02HA-40	209	9131	3.0	9.1882	0.6	4.6854	1.6	0.3122	1.5	0.92	1751.6	22.9	1766.4	13.6	1784.0	11.5	1784.0	11.5	98.2	1.8
GC02HA-74	110	7832	1.6	9.1620	1.2	5.0049	1.6	0.3326	0.9	0.59	1850.8	14.8	1820.1	13.1	1785.2	22.7	1785.2	22.7	103.7	-3.7
GC02HA-31	232	9141	3.6	9.1571	0.8	4.9557	1.8	0.3291	1.6	0.91	1834.1	25.8	1811.8	15.1	1786.2	13.8	1786.2	13.8	102.7	-2.7
GC02HA-86	116	2875	1.6	9.1444	2.8	4.7421	3.0	0.3145	1.1	0.36	1762.8	16.9	1774.7	25.4	1788.7	51.4	1788.7	51.4	98.6	1.4
GC02HA-80	214	9146	1.3	9.1423	1.1	4.9644	1.7	0.3292	1.3	0.75	1834.3	20.3	1813.3	14.3	1789.1	20.2	1789.1	20.2	102.5	-2.5
GC02HA-86	411	21883	7.5	9.1357	1.2	4.7592	1.7	0.3153	1.3	0.74	1766.9	20.0	1777.7	14.6	1790.4	21.3	1790.4	21.3	98.7	1.3
GC02HA-05	295	10089	4.3	9.0887	1.0	4.9368	2.9	0.3254	2.7	0.94	1816.2	43.3	1808.6	24.5	1799.8	17.4	1799.8	17.4	100.9	-0.9
GC02HA-78	126	9078	1.6	9.0579	1.9	5.0199	2.1	0.3298	0.8	0.39	1837.3	12.9	1822.7	17.5	1806.0	34.7	1806.0	34.7	101.7	-1.7
GC02HA-24	41	2700	0.6	9.0514	2.9	5.0190	3.8	0.3295	2.4	0.63	1835.9	36.3	1822.5	32.0	1807.3	53.1	1807.3	53.1	101.6	-1.6
GC02HA-29	207	9232	5.2	9.0227	0.9	4.4142	3.8	0.2869	3.7	0.97	1635.8	53.2	1715.0	31.5	1813.1	16.7	1813.1	16.7	90.2	9.8
GC02HA-49	127	2593	2.5	9.0259	2.0	5.1294	2.6	0.3321	1.6	0.82	1848.4	25.6	1841.0	21.7	1832.6	36.2	1832.6	36.2	100.9	-0.9
GC02HA-30	84	4090	1.3	8.8818	1.1	5.5139	1.7	0.3552	1.2	0.74	1959.3	20.9	1902.8	14.4	1841.6	20.6	1841.6	20.6	106.4	-8.4
GC02HA-34	67	3320	1.1	8.8714	4.0	5.2225	4.6	0.3360	2.3	0.49	1867.5	36.7	1856.3	39.2	1843.7	72.5	1843.7	72.5	101.3	-1.3
GC02HA-32	85	3424	4.0	8.8688	1.8	5.2005	2.0	0.3345	0.8	0.42	1860.2	13.2	1852.7	16.7	1844.3	32.2	1844.3	32.2	100.9	-0.9
GC02HA-25	252	17538	0.7	8.8677	0.4	5.2661	2.3	0.3387	2.3	0.99	1880.3	37.6	1863.4	19.9	1844.5	7.1	1844.5	7.1	101.9	-1.9
GC02HA-37	296	11658	2.2	8.7462	0.3	5.4612	1.1	0.3464	1.1	0.96	1917.5	18.1	1894.5	9.8	1869.4	5.8	1869.4	5.8	102.6	-2.6
GC02HA-54	50	2700	1.5	8.7433	3.9	5.5708	4.8	0.3532	2.8	0.58	1950.1	47.0	1911.6	41.7	1870.0	71.3	1870.0	71.3	104.3	-4.3
GC02HA-44	268	26708	1.0	8.4886	0.4	5.5009	2.7	0.3387	2.6	0.99	1880.2	43.1	1900.7	22.9	1923.2	6.5	1923.2	6.5	97.8	2.2
GC02HA-50	243	16192	2.5	8.4334	0.7	5.5138	1.9	0.3373	1.7	0.93	1873.4	28.0	1902.8	15.9	1934.9	12.2	1934.9	12.2	96.8	3.2
GC02HA-79	164	9446	1.1	8.4078	1.3	5.9164	1.8	0.3608	1.2	0.70	1985.9	21.4	1993.7	15.4	1940.3	22.6	1940.3	22.6	102.3	-2.3
GC02HA-96	66	3300	1.5	8.3992	2.7	6.6281	3.0	0.3428	1.3	0.43	1900.4	21.4	1920.4	26.1	1942.1	48.9	1942.1	48.9	97.8	2.2
GC02HA-93	309	8170	6.8	8.3227	0.6	5.9235	0.9	0.3576	0.7	0.72	1970.6	11.2	1964.7	8.0	1958.5	11.4	1958.5	11.4	100.6	-0.6
GC02HA-57	275	16916	2.0	8.1691	0.3	5.8976	1.8	0.3494	1.7	0.98	1931.8	29.1	1960.9	15.4	1991.7	6.1	1991.7	6.1	97.0	3.0
GC02HA-58	279	49585	3.4	7.9427	0.5	6.2291	2.3	0.3588	2.2	0.98	1976.6	37.8	2008.5	19.9	2041.5	8.4	2041.5	8.4	96.8	3.2
GC02HA-09	98	3933	2.5	7.8541	3.7	6.5310	3.8	0.3720	1.1	0.30	2038.9	19.9	2050.1	33.7	2061.3	64.5	2061.3	64.5	98.9	1.1
GC02HA-75	44	2756	1.2	7.7866	2.2	6.8430	2.4	0.3865	1.0	0.43	2106.3	18.5	2091.3	21.2	2076.5	38.1	2076.5	38.1	101.4	-1.4
GC02HA-56	166	20815	3.8	7.6487	0.5	6.9399	2.0	0.3850	1.9	0.97	2099.5	34.1	2103.8	17.4	2107.9	8.3	2107.9	8.3	99.6	0.4
GC02HA-71	234	11876	1.8	6.9047	0.6	7.9182	2.7	0.3965	2.6	0.97	2153.0	48.4	2221.8	24.6	2285.7	11.1	2285.7	11.1	94.2	5.8
GC02HA-23	77	3873	1.4	6.7989	0.8	8.6660	2.0	0.4372	1.8	0.91	2338.0	35.7	2324.3	18.3	2312.3	14.5	2312.3	14.5	101.1	-1.1
GC02HA-65	319	13147	1.9	6.7455	0.5	9.2991	2.1	0.4667	1.9	0.90	2694.7	41.0	2351.1	19.5	2325.8	9.3	2325.8	9.3	102.3	-2.3
GC02HA-10	83	11919	2.4	6.2438	0.7	10.3625	1.8	0.4693	1.7	0.93	2480.3	35.2	2467.7	17.1	2457.3	11				





## U-Pb Geochronologic analyses of selected Harmony Formation strata

	Isotope ratios										Apparent ages (Ma)										
	U		U/Th		206Pb*		207Pb*		206Pb*/207Pb*		206Pb*		207Pb*		206Pb*/207Pb*		Best age		Conc.		Discor.
	(ppm)	204Pb			(%)	235U	(%)	238U	(%)	error corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)	
<b>Sample: Harmony Canyon. Location: Harmony Canyon, Sonoma Range: 0446225 4533064 (NAD 83 UTM 11T)</b>																					
HARMCYN-173	81	64547	1.4	9.0451	1.0	4.9068	2.0	0.3219	1.7	0.85	1799.0	26.5	1803.4	16.7	1808.8	18.7	1808.8	18.7	99.5	0.5	
HARMCYN-118	72	59144	0.6	8.9847	1.3	4.4771	1.6	0.2917	0.9	0.56	1650.2	12.9	1726.7	13.2	1820.7	23.9	1820.7	23.9	90.6	9.4	
HARMCYN-187	493	23339	1.1	8.9561	1.6	4.6868	10.1	0.3044	10.0	0.99	1713.3	150.2	1764.9	84.9	1826.5	29.5	1826.5	29.5	93.8	6.2	
HARMCYN-176	82	133891	2.3	8.9548	0.9	5.1573	1.5	0.3349	1.3	0.83	1862.3	20.9	1845.6	13.2	1826.8	15.5	1826.8	15.5	101.9	-1.9	
HARMCYN-182	47	45093	1.6	8.9446	1.5	5.1745	1.7	0.3357	0.7	0.45	1865.9	12.0	1848.4	14.1	1828.9	26.8	1828.9	26.8	102.0	-2.0	
HARMCYN-117	287	20420	1.0	8.9305	0.5	4.5776	2.5	0.2965	2.5	0.98	1673.9	36.1	1745.2	20.8	1831.7	8.2	1831.7	8.2	91.4	6.6	
HARMCYN-115	14	13377	7.5	8.9232	6.3	5.2074	6.7	0.3370	2.3	0.35	1872.3	38.1	1853.8	57.3	1833.2	114.3	1833.2	114.3	102.1	-2.1	
HARMCYN-158	90	60528	0.7	8.8897	0.8	5.0439	1.3	0.3252	1.0	0.78	1815.1	15.6	1826.7	10.7	1840.0	14.4	1840.0	14.4	98.6	1.4	
HARMCYN-121	203	211768	1.0	8.8792	0.4	5.1684	1.0	0.3328	0.9	0.91	1852.1	14.7	1847.4	8.6	1842.1	7.7	1842.1	7.7	100.5	-0.5	
HARMCYN-130	37	27557	1.0	8.8550	2.0	5.1723	2.6	0.3322	1.6	0.62	1848.9	26.2	1848.1	22.2	1847.1	37.0	1847.1	37.0	100.1	-0.1	
HARMCYN-137	51	31839	0.7	8.8484	1.9	5.1543	2.4	0.3307	1.6	0.64	1841.8	25.0	1845.1	20.8	1848.8	34.0	1848.8	34.0	99.6	0.4	
HARMCYN-111	86	56745	1.7	8.8402	0.6	4.9564	1.2	0.3178	1.0	0.86	1778.9	15.8	1811.9	10.0	1850.1	11.0	1850.1	11.0	96.1	3.9	
HARMCYN-164	52	58492	0.8	8.8219	0.7	5.2391	1.5	0.3352	1.3	0.88	1863.6	21.1	1859.0	12.6	1853.9	12.5	1853.9	12.5	100.5	-0.5	
HARMCYN-112	78	98840	1.2	8.7919	0.7	5.3315	1.2	0.3400	1.0	0.82	1886.5	16.5	1873.9	10.6	1860.0	12.9	1860.0	12.9	101.4	-1.4	
HARMCYN-193	43	41359	2.9	8.7909	2.2	5.3392	2.4	0.3404	1.1	0.44	1888.7	17.7	1875.2	20.9	1860.2	39.6	1860.2	39.6	101.5	-1.5	
HARMCYN-120	41	51447	4.8	8.7878	1.0	5.2794	1.3	0.3365	0.9	0.69	1869.7	15.2	1865.5	11.5	1860.8	17.5	1860.8	17.5	100.5	-0.5	
HARMCYN-128	87	115310	2.3	8.7655	0.8	5.3966	1.5	0.3431	1.3	0.85	1901.5	21.4	1884.3	13.2	1865.4	14.8	1865.4	14.8	101.9	-1.9	
HARMCYN-195	33	26072	0.6	8.7549	1.6	5.2798	2.0	0.3353	1.2	0.60	1863.8	19.3	1865.6	17.1	1867.6	29.0	1867.6	29.0	99.8	0.2	
HARMCYN-107	62	6663	0.7	8.7279	1.0	5.3070	1.9	0.3359	1.6	0.85	1867.1	25.5	1870.0	15.9	1873.2	17.9	1873.2	17.9	99.7	0.3	
HARMCYN-105	80	49141	3.0	8.7169	1.1	5.2934	1.4	0.3347	0.9	0.66	1860.9	19.3	1867.8	12.2	1875.5	19.3	1875.5	19.3	99.2	0.8	
HARMCYN-196	100	69057	2.3	8.7002	0.8	5.4636	3.3	0.3451	3.2	0.97	1911.3	32.5	1958.8	28.0	1878.9	13.6	1878.9	13.6	101.7	-1.7	
HARMCYN-187	297	33372	1.0	8.6609	0.2	4.8954	4.9	0.3075	4.9	1.00	1728.4	74.7	1801.5	41.6	1887.1	4.2	1887.1	4.2	91.6	8.4	
HARMCYN-136	75	95186	1.5	8.6508	0.7	5.3715	1.1	0.3370	0.9	0.77	1872.3	14.2	1880.3	9.7	1889.2	12.9	1889.2	12.9	99.1	0.9	
HARMCYN-181	180	115527	1.4	8.6480	0.7	5.5423	1.9	0.3476	1.7	0.92	1923.2	28.6	1907.2	16.1	1889.8	13.0	1889.8	13.0	101.8	-1.8	
HARMCYN-139	99	147553	1.4	8.3920	0.9	5.8193	1.3	0.3542	0.9	0.68	1954.6	14.6	1949.3	11.0	1943.7	16.8	1943.7	16.8	100.6	-0.6	
HARMCYN-198	85	76588	0.5	8.0926	0.6	5.1398	1.8	0.3017	1.8	0.95	1699.6	26.2	1842.7	15.6	2008.4	9.9	2008.4	9.9	84.6	15.4	
HARMCYN-165	46	49055	0.5	7.9194	1.3	6.2079	3.6	0.3566	3.3	0.94	1965.8	56.6	2005.6	31.2	2046.7	22.2	2046.7	22.2	96.0	4.0	
HARMCYN-156	201	81591	1.4	7.8572	0.4	6.5688	2.8	0.3743	2.8	0.99	2049.7	49.4	2055.2	25.1	2060.6	7.5	2060.6	7.5	99.5	0.5	
HARMCYN-186	318	36857	1.4	6.7597	2.1	8.5031	6.8	0.4169	6.4	0.95	2246.3	122.0	2296.3	61.7	2322.2	36.8	2322.2	36.8	96.7	3.3	
HARMCYN-135	113	120030	4.4	6.2586	0.4	10.0898	1.2	0.4480	1.1	0.93	2430.6	22.9	2443.0	11.2	2453.3	7.3	2453.3	7.3	99.1	0.9	
HARMCYN-140	208	125235	1.2	6.9303	0.6	10.8912	1.7	0.4684	1.7	0.92	2476.7	32.0	2513.9	15.7	2544.0	10.9	2544.0	10.9	97.1	2.6	
HARMCYN-116	228	167271	5.4	6.9151	0.2	10.3768	1.7	0.4452	1.7	0.99	2373.7	34.2	2469.0	16.1	2548.3	4.0	2548.3	4.0	93.1	6.9	
HARMCYN-170	177	239606	1.4	6.9088	0.7	11.7548	6.9	0.5037	6.8	0.99	2629.9	147.9	2585.1	64.5	2550.1	12.0	2550.1	12.0	103.1	-3.1	
HARMCYN-123	173	220842	1.2	6.8039	0.4	11.4463	1.0	0.4818	1.0	0.93	2535.2	20.0	2560.2	9.7	2580.1	6.3	2580.1	6.3	98.3	1.7	
HARMCYN-174	158	35611	1.2	6.7229	0.3	11.1290	1.5	0.4619	1.5	0.98	2448.0	29.9	2534.0	13.9	2603.5	4.8	2603.5	4.8	94.0	6.0	
HARMCYN-171	39	76755	2.4	6.7197	0.7	12.1651	1.4	0.5046	1.2	0.87	2633.7	26.9	2617.2	13.5	2604.5	11.9	2604.5	11.9	101.1	-1.1	
HARMCYN-162	142	117727	1.2	6.6844	0.6	12.1023	3.1	0.4998	3.0	0.98	2613.0	64.5	2612.4	28.6	2611.8	9.2	2611.8	9.2	100.0	0.0	
HARMCYN-157	129	50997	1.2	6.5851	0.3	11.5037	1.3	0.4671	1.2	0.97	2470.7	25.2	2567.1	11.8	2644.0	4.7	2644.0	4.7	93.4	6.6	
HARMCYN-152	136	167553	1.1	6.5473	0.3	12.3552	2.2	0.4971	2.2	0.99	2601.2	47.6	2631.8	21.1	2655.3	5.6	2655.3	5.6	98.0	2.0	
HARMCYN-119	157	17538	0.7	6.5072	0.2	9.8822	3.3	0.3947	3.3	1.03	2144.7	60.9	2438.8	30.9	2667.3	4.0	2667.3	4.0	80.4	19.6	
HARMCYN-145	91	149808	1.1	6.5068	0.4	12.6627	1.0	0.5057	0.9	0.93	2638.4	19.8	2654.9	6.3	2667.5	6.0	2667.5	6.0	98.9	1.1	
HARMCYN-101	130	110380	0.9	6.4106	0.6	12.8200	4.4	0.5031	4.4	0.99	2627.0	95.3	2666.5	41.9	2696.6	9.2	2696.6	9.2	97.4	2.6	
HARMCYN-120	91	116123	1.6	6.4033	0.6	13.2790	1.1	0.5204	0.9	0.86	2700.8	20.9	2699.7	10.4	2698.9	9.2	2698.9	9.2	100.1	-0.1	
HARMCYN-104	146	186503	1.7	6.3696	0.2	13.6173	1.5	0.5303	1.5	0.99	2742.8	32.9	2723.5	14.1	2709.2	3.4	2709.2	3.4	101.2	-1.2	
HARMCYN-124	115	227017	2.2	6.2858	0.3	14.1244	1.9	0.5415	1.9	0.99	2789.6	42.7	2758.1	18.1	2735.1	4.5	2735.1	4.5	102.0	-2.0	
<b>Sample: Elbow Canyon. Location: Elbow Canyon, Sonoma Range: 0441026 4514450 (NAD 83 UTM 11T)</b>																					
ELBOW CYN-148	89	22814	1.2	14.3469	2.6	1.3210	2.8	0.1375	1.1	0.41	830.2	8.9	854.9	16.3	919.6	53.1	830.2	8.9	90.3	9.7	
ELBOW CYN-69	157	208587	2.6	9.5321	0.3	4.4292	0.8	0.3062	0.8	0.92	1722.0	11.7	1717.8	7.0	1712.7	6.1	1712.7	6.1	100.5	-0.5	
ELBOW CYN-119	43	23689	3.1	9.4623	1.8	4.6104	2.0	0.3164	0.9	0.46	1722.1	14.3	1751.2	16.9	1726.2	33.0	1726.2	33.0	102.7	-2.7	
ELBOW CYN-139	20	16811	1.0	9.3961	2.9	4.7135	3.2	0.3212	1.2	0.39	1795.6	19.5	1769.6	26.7	1739.1	53.8	1739.1	53.8	103.3	-3.3	
ELBOW CYN-174	27	9332	1.1	9.3256	2.3	4.6852	2.4	0.3168	0.7	0.29	1744.5	10.8	1794.8	20.4	1752.9	42.8	1752.9	42.8	101.2	-1.2	
ELBOW CYN-196	30	12192	0.7	9.2960	1.7	4.7607	1.7	0.3210	1.5	0.65	1796.5	17.9	1778.0	19.4	1758.7	31.9	1758.7	31.9	102.0	-2.0	
ELBOW CYN-178	181	120477	0.8	9.2807	0.3	4.7015	0.9	0.3158	0.9	0.95	1769.1	13.2	1767.5	7.5	1765.6	5.1	1765.6	5.1	100.2	-0.2	
ELBOW CYN-112	43	31712	1.2	9.2590	1.0	4.7849	1.4	0.3213	1.0	0.72	1796.2	15.4	1782.3	11.5	1766.0	17.5	1766.0	17.5	101.7	-1.7	
ELBOW CYN-188	39	17222	1.2	9.2390	1.1	4.7613	1.5	0.3190	0.9	0.63	1785.1	14.6	1778.1	12.4	1769.9	20.9	1769.9	20.9	100.9	-0.9	
ELBOW CYN-175	22	11965	1.1	9.2332	3.4	4.7465	3.6	0.3179	1.3	0.37	1779.2	21.0	1775.5	30.4	1771.1	61.5	1771.1	61.5	100.5	-0.5	
ELBOW CYN-102	57	19865	1.2	9.2213	0.7	4.7645	1.1	0.3186	0.9	0.79	1783.1	13.6	1778.7	9.2	1773.4	12.1	1773.4	12.1	100.5	-0.5	
ELBOW CYN-137	79	35461	1.1	9.2186	0.5	4.8185	0.7	0.3222	0.5	0.72	1800.3	7.9	1788.1	5.9	1774.0	8.8	1774.0	8.8	101.5	-1.5	



## U-Pb Geochronologic analyses of selected Harmony Formation strata

	Isotope ratios										Apparent ages (Ma)						Conc.	Discor.		
	U	206Pb	U/Th	206Pb*	±	207Pb*	±	206Pb*	±	error	206Pb*	±	207Pb*	±	206Pb*	±			Best age	±
(ppm)	204Pb			(%)	207Pb*	(%)	235U*	(%)	corr.	238U*	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)	
<b>Sample: Elbow Canyon. Location: Elbow Canyon, Sonoma Range; 0441026 4514450 (NAD 83 UTM 11T)</b>																				
ELBOW CYN-56	66	77524	1.5	9.1581	0.6	4.8612	1.2	0.3229	1.1	0.86	1803.8	16.7	1795.5	10.4	1786.0	11.7	1786.0	11.7	100.1	-1.0
ELBOW CYN-173	176	183336	2.4	9.1568	0.3	4.8159	1.2	0.3198	1.1	0.97	1788.9	17.7	1787.7	9.8	1786.2	4.9	1786.2	4.9	101.0	-0.1
ELBOW CYN-160	175	128010	2.3	9.1561	0.2	4.8401	1.1	0.3214	1.1	0.98	1796.6	16.6	1791.9	9.1	1786.4	3.7	1786.4	3.7	100.6	-0.6
ELBOW CYN-140	153	53615	1.5	9.1558	0.3	4.7972	0.9	0.3186	0.8	0.92	1782.7	12.6	1784.4	7.4	1786.4	6.1	1786.4	6.1	99.8	0.2
ELBOW CYN-171	59	19384	1.7	9.1534	1.0	4.8175	1.3	0.3198	0.8	0.64	1788.8	12.8	1788.0	10.7	1786.9	17.7	1786.9	17.7	100.1	-0.1
ELBOW CYN-127	115	118431	1.4	9.1527	0.5	4.8313	1.1	0.3207	0.9	0.86	1793.2	14.2	1790.4	8.9	1787.1	9.9	1787.1	9.9	100.3	-0.3
ELBOW CYN-145	179	96716	2.3	9.1521	0.4	4.8789	1.5	0.3238	1.4	0.96	1809.5	22.7	1799.6	12.7	1787.2	8.0	1787.2	8.0	101.2	-1.2
ELBOW CYN-138	225	89705	4.9	9.1516	0.2	4.8132	0.8	0.3195	0.7	0.98	1787.1	11.5	1787.2	6.3	1787.3	3.0	1787.3	3.0	100.0	0.0
ELBOW CYN-92	27	28066	1.2	9.1510	2.2	4.8249	2.4	0.3202	1.0	0.41	1790.8	15.7	1789.3	20.6	1787.4	40.7	1787.4	40.7	100.2	-0.2
ELBOW CYN-115	122	128374	2.1	9.1478	0.4	4.9398	1.2	0.3277	1.2	0.94	1827.4	18.5	1809.1	10.4	1788.0	7.7	1788.0	7.7	102.2	-2.2
ELBOW CYN-159	177	262841	2.3	9.1464	0.3	4.8343	0.7	0.3207	0.6	0.90	1793.1	9.8	1790.9	5.9	1788.3	5.5	1788.3	5.5	100.3	-0.3
ELBOW CYN-11	109	98857	1.9	9.1462	0.7	4.8397	1.1	0.3210	0.9	0.78	1794.8	13.4	1791.8	9.2	1788.3	12.5	1788.3	12.5	100.4	-0.4
ELBOW CYN-187	34	13360	1.5	9.1409	1.7	4.7836	2.2	0.3171	1.4	0.63	1775.7	21.9	1782.0	18.8	1789.4	31.7	1789.4	31.7	99.2	0.8
ELBOW CYN-7	79	34518	1.4	9.1383	0.8	4.7960	1.3	0.3179	1.0	0.78	1779.3	15.9	1784.2	11.0	1789.9	14.8	1789.9	14.8	99.4	0.6
ELBOW CYN-9	216	98059	3.1	9.1371	0.2	4.8668	0.6	0.3225	0.6	0.94	1802.0	9.2	1796.5	5.2	1790.2	3.7	1790.2	3.7	100.7	-0.7
ELBOW CYN-36	42	35071	1.1	9.1356	0.9	4.9205	1.3	0.3260	0.9	0.72	1819.1	14.8	1805.8	10.9	1790.5	16.3	1790.5	16.3	101.6	-1.6
ELBOW CYN-129	83	89980	2.1	9.1307	0.5	4.8906	0.9	0.3239	0.7	0.81	1808.6	11.7	1800.6	7.7	1791.4	9.9	1791.4	9.9	101.0	-1.0
ELBOW CYN-118	235	46984	1.7	9.1294	0.3	4.7036	1.1	0.3116	1.0	0.96	1748.7	15.9	1769.4	9.1	1791.7	5.7	1791.7	5.7	97.6	2.4
ELBOW CYN-57	47	42256	1.5	9.1276	1.1	4.8895	1.2	0.3237	0.5	0.61	1807.7	8.2	1804.3	10.5	1792.0	20.6	1792.0	20.6	100.9	-0.9
ELBOW CYN-89	146	96978	1.6	9.1247	0.4	4.9209	1.3	0.3257	1.2	0.94	1817.3	18.9	1805.9	10.7	1792.6	7.7	1792.6	7.7	101.4	-1.4
ELBOW CYN-18	75	46576	1.3	9.1243	0.6	4.9373	2.0	0.3267	1.9	0.96	1822.5	29.6	1808.7	16.5	1792.7	10.5	1792.7	10.5	101.7	-1.7
ELBOW CYN-149	152	50766	2.0	9.1235	0.4	4.8226	0.7	0.3191	0.6	0.85	1785.4	8.9	1788.8	5.7	1792.9	6.5	1792.9	6.5	99.6	0.4
ELBOW CYN-26	130	64192	1.7	9.1195	0.4	4.7901	0.8	0.3168	0.7	0.83	1774.2	10.2	1783.2	6.7	1793.7	8.0	1793.7	8.0	99.9	1.1
ELBOW CYN-116	258	88002	2.1	9.1115	0.3	4.7728	3.1	0.3154	3.1	1.00	1767.2	47.5	1780.1	25.9	1795.3	5.4	1795.3	5.4	98.4	1.6
ELBOW CYN-35	36	24854	1.0	9.1092	2.7	4.7612	3.4	0.3146	2.1	0.62	1763.1	32.7	1778.1	28.7	1795.7	48.9	1795.7	48.9	99.2	1.8
ELBOW CYN-21	91	45798	1.3	9.1076	0.6	4.9416	1.2	0.3264	1.0	0.85	1821.0	16.3	1809.4	10.2	1796.1	11.7	1796.1	11.7	101.4	-1.4
ELBOW CYN-60	99	45879	1.3	9.0920	0.7	4.9341	2.0	0.3254	1.9	0.93	1815.9	29.4	1808.1	16.8	1799.2	13.2	1799.2	13.2	100.9	-0.9
ELBOW CYN-117	48	20522	1.4	9.0902	1.2	4.9399	1.8	0.3256	1.3	0.72	1817.1	20.0	1805.9	14.9	1799.5	22.4	1799.5	22.4	101.0	-1.0
ELBOW CYN-114	32	20996	1.2	9.0789	2.7	4.9217	2.8	0.3247	0.8	0.53	1813.8	19.2	1807.7	23.5	1803.8	49.3	1803.8	49.3	100.6	-0.6
ELBOW CYN-8	35	19068	1.3	9.0764	1.3	4.8781	2.4	0.3210	2.0	0.83	1794.5	30.8	1796.1	19.9	1802.3	23.8	1802.3	23.8	99.8	0.2
ELBOW CYN-156	35	35442	0.8	9.0642	0.8	5.1288	1.7	0.3314	1.5	0.88	1855.1	24.7	1840.9	14.9	1824.9	25.2	1824.9	25.2	101.7	-1.7
ELBOW CYN-70	39	58387	1.2	9.0633	1.0	5.0427	1.2	0.3278	0.7	0.56	1827.8	11.0	1826.5	10.4	1825.1	18.3	1825.1	18.3	100.1	-1.1
ELBOW CYN-161	39	17851	1.9	8.9998	0.8	5.1006	1.3	0.3314	1.0	0.78	1845.4	16.7	1836.2	11.3	1825.8	15.0	1825.8	15.0	101.1	-0.1
ELBOW CYN-131	78	17379	0.7	8.9595	0.7	5.0244	1.0	0.3265	0.7	0.72	1821.4	11.4	1823.4	8.4	1825.8	12.5	1825.8	12.5	99.8	-0.2
ELBOW CYN-91	20	11265	0.5	8.9469	2.6	5.1775	3.4	0.3360	2.2	0.65	1867.2	35.8	1848.9	29.1	1828.4	47.3	1828.4	47.3	102.1	-2.1
ELBOW CYN-146	28	15319	0.5	8.9435	1.5	5.1336	2.1	0.3330	1.5	0.70	1852.8	23.4	1841.7	17.6	1829.1	26.7	1829.1	26.7	101.3	-1.3
ELBOW CYN-189	132	171375	1.4	8.9344	0.6	5.1645	0.9	0.3347	0.7	0.77	1860.9	11.5	1846.8	7.9	1830.9	10.7	1830.9	10.7	101.6	-1.6
ELBOW CYN-133	150	101588	1.0	8.9102	0.4	5.1409	0.7	0.3322	0.6	0.83	1849.1	9.9	1842.9	6.4	1835.8	7.6	1835.8	7.6	100.7	-0.7
ELBOW CYN-55	80	49558	1.2	8.9028	0.7	4.6322	2.2	0.2991	2.1	0.95	1698.8	30.5	1755.1	18.2	1837.3	12.8	1837.3	12.8	91.8	8.2
ELBOW CYN-118	252	76731	2.1	8.9257	0.2	5.1558	0.8	0.3312	0.5	0.41	1847.7	8.2	1845.3	7.1	1846.8	11.9	1846.8	11.9	99.8	0.2
ELBOW CYN-199	24	28099	1.8	8.8400	1.4	5.1422	1.8	0.3297	1.1	0.61	1836.8	17.3	1843.1	15.1	1850.2	25.4	1850.2	25.4	99.3	0.7
ELBOW CYN-155	224	164060	5.8	8.8303	0.5	5.1976	0.7	0.3329	0.4	0.68	1852.3	7.2	1852.2	5.6	1852.1	8.6	1852.1	8.6	100.0	0.0
ELBOW CYN-29	127	87734	1.6	8.8257	0.4	5.2005	0.6	0.3329	0.4	0.70	1852.4	6.5	1852.7	4.9	1853.1	7.4	1853.1	7.4	100.0	0.0
ELBOW CYN-32	70	60139	2.2	8.8243	0.6	5.1749	1.0	0.3312	0.9	0.84	1844.2	13.9	1845.8	8.8	1853.4	10.2	1853.4	10.2	99.5	0.5
ELBOW CYN-176	155	156674	1.5	8.8178	0.3	5.2335	0.9	0.3347	0.8	0.93	1861.1	13.2	1858.1	7.5	1854.7	5.9	1854.7	5.9	100.3	-0.3
ELBOW CYN-167	74	56348	2.7	8.8114	0.7	5.2063	1.1	0.3327	0.9	0.79	1851.5	14.5	1853.6	9.7	1856.0	12.7	1856.0	12.7	99.8	0.2
ELBOW CYN-23	50	30383	1.3	8.8044	1.1	5.2331	1.4	0.3342	0.9	0.65	1858.5	15.0	1858.0	12.2	1857.4	19.6	1857.4	19.6	100.1	-0.1
ELBOW CYN-184	95	119541	0.9	8.8031	0.5	5.2208	1.3	0.3333	1.3	0.93	1854.5	20.2	1856.0	11.5	1857.7	8.9	1857.7	8.9	99.8	0.2
ELBOW CYN-165	202	311502	8.9	8.7980	0.3	5.2149	0.6	0.3328	0.5	0.86	1851.7	8.8	1855.0	5.4	1858.8	5.9	1858.8	5.9	99.6	0.4
ELBOW CYN-118	252	76731	2.1	8.7957	0.2	5.2117	0.8	0.3312	0.5	0.53	1853.8	19.2	1857.7	23.5	1853.8	49.3	1853.8	49.3	100.6	-0.6
ELBOW CYN-65	95	100687	3.8	8.7886	0.4	5.3255	0.8	0.3395	0.7	0.86	1884.0	11.3	1873.0	6.9	1860.7	7.5	1860.7	7.5	101.3	-1.3
ELBOW CYN-158	50	80414	1.3	8.7808	1.3	5.2337	1.6													

## U-Pb Geochronologic analyses of selected Harmony Formation strata

	Isotope ratios										Apparent ages (Ma)									
	U (ppm)	<sup>206</sup> Pb 204Pb	U/Th	<sup>206</sup> Pb* 207Pb*	± 235U*	<sup>207</sup> Pb* (%)	± 238U*	<sup>206</sup> Pb* (%)	± 238U*	error (%)	<sup>206</sup> Pb* (Ma)	± 235U	<sup>207</sup> Pb* (Ma)	± 235U	<sup>206</sup> Pb* (Ma)	± 207Pb*	Best age (Ma)	± (Ma)	Conc. (%)	Discor. (%)
Sample: Elbow Canyon. Location: Elbow Canyon, Sonoma Range: 0441026 4514450 (NAD 83 UTM 11T)																				
ELBOW CYN-56	66	77524	1.5	9.1581	0.6	4.8612	1.2	0.3229	1.1	0.96	1803.8	16.7	1795.5	10.4	1786.0	11.7	1786.0	11.7	101.0	-1.0
ELBOW CYN-173	176	163336	2.4	9.1568	0.3	4.8159	1.2	0.3198	1.1	0.97	1788.9	17.7	1787.7	9.8	1786.2	4.9	1786.2	4.9	100.1	-0.1
ELBOW CYN-160	175	128010	2.3	9.1561	0.2	4.8401	1.1	0.3214	1.1	0.98	1796.6	16.6	1791.9	9.1	1786.4	3.7	1786.4	3.7	100.6	-0.6
ELBOW CYN-140	153	53615	1.5	9.1558	0.3	4.7972	0.9	0.3186	0.8	0.92	1782.7	12.6	1784.4	7.4	1786.4	6.1	1786.4	6.1	99.8	0.2
ELBOW CYN-171	59	19384	1.7	9.1534	1.0	4.8175	1.3	0.3198	0.8	0.64	1788.8	12.8	1788.0	10.7	1786.9	17.7	1786.9	17.7	100.1	-0.1
ELBOW CYN-127	115	118431	1.4	9.1527	0.5	4.8313	1.1	0.3207	0.9	0.86	1793.2	14.2	1790.4	8.9	1787.1	9.9	1787.1	9.9	100.3	-0.3
ELBOW CYN-145	179	96716	2.3	9.1521	0.4	4.8789	1.5	0.3238	1.4	0.96	1808.5	22.7	1798.6	12.7	1787.2	8.0	1787.2	8.0	101.2	-1.2
ELBOW CYN-138	225	89705	4.9	9.1516	0.2	4.8132	0.8	0.3195	0.7	0.98	1787.1	11.5	1787.2	6.3	1787.3	3.0	1787.3	3.0	100.0	0.0
ELBOW CYN-92	27	28066	1.2	9.1510	2.2	4.8249	2.4	0.3202	1.0	0.41	1790.8	15.7	1789.3	20.6	1787.4	40.7	1787.4	40.7	100.2	-0.2
ELBOW CYN-115	122	128374	2.1	9.1478	0.4	4.9398	1.2	0.3277	1.2	0.94	1827.4	18.5	1809.1	10.4	1788.0	7.7	1788.0	7.7	102.2	-2.2
ELBOW CYN-159	177	262841	2.3	9.1464	0.3	4.8343	0.7	0.3207	0.6	0.90	1793.1	9.8	1790.9	5.9	1788.3	5.3	1788.3	5.3	100.3	-0.3
ELBOW CYN-111	109	98857	1.9	9.1462	0.7	4.8397	1.1	0.3210	0.9	0.78	1794.8	13.4	1791.8	9.2	1789.3	12.5	1788.3	12.5	100.4	-0.4
ELBOW CYN-187	34	13360	1.5	9.1409	1.7	4.7836	2.2	0.3171	1.4	0.63	1775.7	21.9	1782.0	18.8	1789.4	31.7	1789.4	31.7	99.2	0.8
ELBOW CYN-7	79	34518	1.4	9.1383	0.8	4.7960	1.3	0.3179	1.0	0.78	1779.3	15.9	1784.2	11.0	1789.9	14.8	1789.9	14.8	99.4	0.6
ELBOW CYN-9	216	98059	3.1	9.1371	0.2	4.8668	0.6	0.3225	0.6	0.94	1802.0	9.2	1796.5	5.2	1790.2	3.7	1790.2	3.7	100.7	-0.7
ELBOW CYN-36	42	35071	1.1	9.1356	0.9	4.9205	1.3	0.3260	0.9	0.72	1819.1	14.8	1805.8	10.9	1790.5	16.3	1790.5	16.3	101.6	-1.6
ELBOW CYN-129	83	89980	2.1	9.1307	0.5	4.8906	0.9	0.3239	0.7	0.81	1808.6	11.7	1800.6	7.7	1791.4	9.9	1791.4	9.9	101.0	-1.0
ELBOW CYN-4	235	26864	1.7	9.1294	0.3	4.7066	1.1	0.3116	1.0	0.96	1748.7	15.9	1768.4	9.1	1791.7	5.7	1791.7	5.7	97.6	2.4
ELBOW CYN-57	47	42256	1.5	9.1276	1.1	4.8895	1.2	0.3237	0.5	0.41	1807.7	8.1	1800.4	10.5	1792.0	20.6	1792.0	20.6	100.9	-0.9
ELBOW CYN-89	146	96978	1.6	9.1247	0.4	4.9209	1.3	0.3257	1.2	0.94	1817.3	18.9	1805.9	10.7	1792.6	7.7	1792.6	7.7	101.4	-1.4
ELBOW CYN-18	75	46576	1.3	9.1243	0.6	4.9373	2.0	0.3267	1.9	0.96	1822.5	29.5	1808.7	16.5	1792.7	10.5	1792.7	10.5	101.7	-1.7
ELBOW CYN-149	152	50766	2.0	9.1235	0.4	4.8226	0.7	0.3191	0.6	0.85	1785.4	8.9	1789.8	4.7	1793.9	8.5	1792.9	8.5	99.6	0.4
ELBOW CYN-26	130	64192	1.7	9.1195	0.4	4.7901	0.8	0.3168	0.7	0.83	1774.2	10.2	1783.2	6.7	1793.0	8.0	1793.0	8.0	98.9	1.1
ELBOW CYN-116	258	88002	2.1	9.1115	0.3	4.7728	3.1	0.3154	3.1	1.00	1767.2	47.5	1780.1	25.9	1795.3	5.4	1795.3	5.4	98.4	1.6
ELBOW CYN-35	36	24854	1.0	9.1092	2.7	4.7612	3.4	0.3146	2.1	0.62	1763.1	32.7	1778.1	28.7	1795.7	48.9	1795.7	48.9	98.2	1.8
ELBOW CYN-21	91	45798	1.3	9.1076	0.6	4.9416	1.2	0.3264	1.0	0.85	1821.0	16.3	1809.4	10.2	1796.1	11.7	1796.1	11.7	101.4	-1.4
ELBOW CYN-60	99	45879	1.3	9.0920	0.7	4.9341	2.0	0.3254	1.9	0.93	1815.9	29.4	1808.1	16.8	1799.2	13.2	1799.2	13.2	100.9	-0.9
ELBOW CYN-117	48	20522	1.4	9.0902	1.2	4.9389	1.8	0.3256	1.3	0.72	1817.1	20.0	1808.9	14.9	1799.5	22.4	1799.5	22.4	101.0	-1.0
ELBOW CYN-114	32	20996	0.7	9.0789	2.7	4.9317	2.8	0.3247	0.6	0.23	1812.8	10.2	1807.7	23.5	1801.8	49.3	1801.8	49.3	100.6	-0.6
ELBOW CYN-8	35	19066	1.3	9.0764	1.3	4.8761	2.4	0.3210	2.0	0.83	1794.5	30.8	1798.1	19.9	1802.3	23.8	1802.3	23.8	99.6	0.4
ELBOW CYN-156	35	35442	0.8	9.0642	0.8	5.1268	1.7	0.3334	1.5	0.88	1855.1	24.7	1840.9	14.9	1824.9	15.2	1824.9	15.2	101.7	-1.7
ELBOW CYN-70	39	17851	1.3	9.0633	1.0	5.0427	1.2	0.3278	0.7	0.86	1837.8	11.0	1828.5	13.1	1828.5	13.1	100.0	-1.1		
ELBOW CYN-161	39	17851	1.3	9.0598	0.8	5.1006	1.3	0.3314	1.0	0.78	1845.6	16.7	1836.2	11.1	1825.8	15.0	1825.8	15.0	101.1	-1.1
ELBOW CYN-131	78	17379	0.7	9.0595	0.7	5.0244	1.0	0.3265	0.7	0.72	1821.4	11.4	1823.4	8.4	1825.8	12.5	1825.8	12.5	99.8	0.2
ELBOW CYN-91	20	11265	0.5	9.0469	2.6	5.1775	3.4	0.3360	2.2	0.65	1867.2	35.8	1848.9	29.1	1828.4	47.3	1828.4	47.3	102.1	-2.1
ELBOW CYN-146	28	15319	0.5	9.0435	1.5	5.1336	2.1	0.3330	1.5	0.70	1852.8	23.4	1841.7	17.6	1829.1	26.7	1829.1	26.7	101.3	-1.3
ELBOW CYN-189	132	171375	1.4	9.0344	0.6	5.1645	0.9	0.3347	0.7	0.77	1860.9	11.5	1846.8	7.9	1830.0	10.7	1830.0	10.7	101.6	-1.6
ELBOW CYN-133	150	101588	1.0	9.0102	0.4	5.1409	0.7	0.3322	0.6	0.83	1849.1	9.9	1842.9	6.4	1835.8	7.6	1835.8	7.6	100.7	-0.7
ELBOW CYN-55	80	49058	1.2	9.0028	0.7	4.6322	2.2	0.2991	2.1	0.95	1686.8	30.5	1755.1	18.2	1837.3	12.8	1837.3	12.8	91.8	8.2
ELBOW CYN-53	94	96856	1.0	8.9562	0.7	5.1558	0.8	0.3312	0.5	0.61	1844.0	8.2	1845.3	7.1	1846.8	11.9	1846.8	11.9	99.8	0.2
ELBOW CYN-199	24	28099	1.8	8.8400	1.4	5.1422	1.8	0.3297	1.1	0.61	1836.8	17.3	1843.1	15.1	1850.2	25.4	1850.2	25.4	99.3	0.7
ELBOW CYN-155	224	164080	5.8	8.8303	0.5	5.1976	0.7	0.3329	0.4	0.68	1852.3	7.2	1852.2	5.6	1852.1	8.6	1852.1	8.6	100.0	0.0
ELBOW CYN-29	127	87734	1.6	8.8257	0.4	5.2005	0.6	0.3329	0.4	0.70	1854.4	8.5	1852.7	4.9	1853.1	7.4	1853.1	7.4	100.0	0.0
ELBOW CYN-32	70	60139	2.2	8.8243	0.6	5.1749	1.0	0.3312	0.9	0.84	1844.2	13.9	1848.5	8.8	1853.4	10.2	1853.4	10.2	99.5	0.5
ELBOW CYN-176	155	155674	1.5	8.8178	0.3	5.2335	0.9	0.3347	0.8	0.93	1861.1	13.2	1858.1	7.5	1854.7	5.9	1854.7	5.9	100.3	-0.3
ELBOW CYN-167	74	56348	2.7	8.8114	0.7	5.2063	1.1	0.3327	0.9	0.79	1851.5	14.5	1853.6	9.7	1856.2	12.7	1856.2	12.7	99.8	0.2
ELBOW CYN-23	50	30383	1.3	8.8044	1.1	5.2331	1.4	0.3342	0.9	0.65	1858.5	15.0	1858.0	12.2	1857.4	19.6	1857.4	19.6	100.1	-0.1
ELBOW CYN-184	95	119541	0.9	8.8031	0.5	5.2208	1.3	0.3333	1.3	0.93	1854.5	20.2	1856.0	11.5	1857.7	8.9	1857.7	8.9	99.8	0.2
ELBOW CYN-165	202	311502	8.9	8.7980	0.3	5.2149	0.6	0.3328	0.5	0.86	1851.7	8.8	1855.0	5.4	1858.6	5.9	1858.6	5.9	99.6	0.4
ELBOW CYN-118	253	76721	2.4	8.7957	0.2	5.2218	0.5	0.3331	0.5	0.95	1853.5	7.9	1856.2	4.4	1859.2	3.0	1859.2	3.0	99.7	0.3
ELBOW CYN-65	95	100687	3.8	8.7886	0.4	5.2355	0.8	0.3395	0.7	0.86	1884.0	11.3	1873.0	6.9	1860.7	7.5	1860.7	7.5	101.3	-1.3
ELBOW CYN-158	50	80414	1.3	8.7808	1.3	5.2337	1.6	0.3333	0.9	0.58	1854.4	14.7	1858.1	13.5	1862.3	23.3	1862.3	23.3	99.6	0.4
ELBOW CYN-132	138	176986	2.4	8.7678	0.2	5.3257	1.0	0.3377	1.2	0.98	1890.3	18.7	1876.0	8.9	1864.9	3.8	1864.9	3.8	100.8	-0.8
ELBOW CYN-164	43	8567	0.6	8.7674	1.3	5.2413	1.6	0.3333	0.9	0.58	1854.4	14.7	1858.1	13.5	1862.3	23.3	1862.3	23.3	99.4	0.6
ELBOW CYN-123	18	15282	2.5	8.7647	1.6	5.3473	2.4	0.3399	1.8	0.74	1886.3	29.3	1876.5	20.7	1865.6	29.4	1865.6	29.4	101.1	-1.1
ELBOW CYN-67	93	42284	1.9	8.7592	0.4	5.1911	4.4	0.3298	4.3	0.99	1837.3	69.4	1851.2	37.2	1866.7	8.1	1866.7	8.1	98.4	1.6
ELBOW CYN-194	224	252747	2.2	8.7580	0.3	5.4068	0.8	0.3434	0.7	0.93	1903.2	12.1	1885.9	6.8	1867.0	5.3	1867.0	5.3	101.9	-1.9
ELBOW CYN-85	47	86571	0.7	8.7502	1.1	5.3033	1.3													



## U-Pb Geochronologic analyses of selected Harmony Formation strata

		Isotope ratios								Apparent ages (Ma)										
	U	<sup>206</sup> Pb	U/Th	<sup>206</sup> Pb*	±	<sup>207</sup> Pb*	±	<sup>206</sup> Pb*	±	error	<sup>206</sup> Pb*	±	<sup>207</sup> Pb*	±	<sup>206</sup> Pb*	±	Best age	±	Conc.	Discor.
	(ppm)	<sup>204</sup> Pb		<sup>207</sup> Pb*	(%)	<sup>235</sup> U*	(%)	<sup>238</sup> U	(%)	corr.	<sup>238</sup> U*	(Ma)	<sup>235</sup> U	(Ma)	<sup>207</sup> Pb*	(Ma)	(Ma)	(Ma)	(%)	(%)
<b>Sample: Elbow Canyon. Location: Elbow Canyon, Sonoma Range: 0441026 4514450 (NAD 83 UTM 11T)</b>																				
ELBOW CYN-72	79	100089	0.9	5.4353	0.3	13.2192	0.5	0.5211	0.4	0.83	2703.9	9.7	2695.4	4.9	2689.1	4.8	2689.1	4.8	100.6	-0.6
ELBOW CYN-96	20	39801	0.8	5.4294	0.7	13.2444	1.3	0.5215	1.1	0.85	2705.7	25.2	2697.2	12.7	2690.9	11.8	2690.9	11.8	100.6	-0.6
ELBOW CYN-134	15	40812	0.8	5.4250	1.1	13.3400	1.5	0.5249	1.0	0.68	2719.8	22.1	2704.0	13.9	2692.2	17.9	2692.2	17.9	101.0	-1.0
ELBOW CYN-97	21	26060	1.7	5.3978	1.1	13.4605	1.4	0.5270	0.8	0.62	2728.6	18.6	2712.5	12.8	2700.5	17.5	2700.5	17.5	101.0	-1.0
ELBOW CYN-152	29	27819	2.8	5.3989	2.7	14.0342	3.1	0.5493	1.6	0.52	2622.4	36.9	2752.0	29.7	2700.8	44.3	2700.8	44.3	104.5	-4.5
ELBOW CYN-38	13	34636	0.7	5.3923	1.9	13.2394	2.2	0.5178	1.2	0.52	2689.7	25.5	2696.9	21.1	2702.2	31.6	2702.2	31.6	99.5	0.5
ELBOW CYN-30	132	322942	0.8	5.3881	0.2	13.2659	0.8	0.5184	0.7	0.98	2692.4	16.2	2698.8	7.1	2703.5	2.6	2703.5	2.6	99.6	0.4
ELBOW CYN-13	21	43739	0.9	5.3684	1.6	12.8093	2.3	0.4987	1.7	0.72	2606.4	35.6	2665.7	21.7	2709.5	26.2	2709.5	26.2	96.3	3.7
ELBOW CYN-28	22	42263	1.2	5.3488	1.4	13.8035	2.3	0.5355	1.9	0.81	2764.5	42.4	2736.3	22.1	2715.6	22.8	2715.6	22.8	101.8	-1.8
ELBOW CYN-191	19	33098	1.0	5.3417	0.6	13.5470	1.3	0.5248	1.2	0.91	2719.7	26.9	2718.6	12.7	2717.8	9.3	2717.8	9.3	100.1	-0.1
ELBOW CYN-150	108	127157	0.8	5.3248	0.1	13.5966	0.8	0.5251	0.8	0.99	2720.7	18.4	2722.0	7.9	2723.0	2.0	2723.0	2.0	99.9	0.1
ELBOW CYN-82	152	20635	0.6	5.1931	0.2	13.9427	0.9	0.5251	0.8	0.96	2720.9	18.5	2745.8	8.2	2764.2	4.0	2764.2	4.0	98.4	1.6
ELBOW CYN-3	130	313414	0.8	4.9648	0.2	15.5384	0.5	0.5595	0.5	0.95	2664.6	11.9	2648.8	5.2	2637.7	2.7	2637.7	2.7	100.9	-0.9
ELBOW CYN-153	12	19658	1.3	4.9147	1.2	15.4133	1.3	0.5494	0.6	0.49	2622.7	14.8	284.1	12.6	2854.2	19.7	2854.2	19.7	98.9	1.1
ELBOW CYN-126	72	26451	1.2	4.8616	0.5	14.9380	2.0	0.5267	1.9	0.97	2727.6	42.8	2811.3	18.8	2871.9	7.4	2871.9	7.4	95.0	5.0
ELBOW CYN-80	85	97530	1.0	4.7964	0.2	16.2022	3.6	0.5636	3.6	1.00	2681.6	84.2	2888.8	34.7	2893.8	3.6	2893.8	3.6	99.6	0.4
ELBOW CYN-75	100	119046	0.8	4.6187	0.2	17.1784	0.7	0.5754	0.6	0.95	2930.1	14.9	2944.8	6.4	2954.9	3.5	2954.9	3.5	99.2	0.8
ELBOW CYN-157	27	30847	1.5	3.7066	0.4	24.7610	1.0	0.6657	0.9	0.91	3289.1	24.1	3298.8	10.0	3304.8	6.7	3304.8	6.7	99.5	0.5

### Notes:

- Analyses with >10% uncertainty (1-sigma) in <sup>206</sup>Pb/<sup>238</sup>U age are not included.
- Analyses with >10% uncertainty (1-sigma) in <sup>206</sup>Pb/<sup>207</sup>Pb age are not included, unless <sup>206</sup>Pb/<sup>238</sup>U age is <500 Ma.
- Best age is determined from <sup>206</sup>Pb/<sup>238</sup>U age for analyses with <sup>206</sup>Pb/<sup>238</sup>U age <1000 Ma and from <sup>206</sup>Pb/<sup>207</sup>Pb age for analyses with <sup>206</sup>Pb/<sup>238</sup>U age > 1000 Ma.
- Concordance is based on <sup>206</sup>Pb/<sup>238</sup>U age / <sup>206</sup>Pb/<sup>207</sup>Pb age. Value is not reported for <sup>206</sup>Pb/<sup>238</sup>U ages <500 Ma because of large uncertainty in <sup>206</sup>Pb/<sup>207</sup>Pb age.
- Discordance is 100% - concordance.
- Analyses with <sup>206</sup>Pb/<sup>238</sup>U age > 500 Ma and with >20% discordance (<80% concordance) are not included.
- Analyses with <sup>206</sup>Pb/<sup>238</sup>U age > 500 Ma and with >5% reverse discordance (<105% concordance) are not included.
- All uncertainties are reported at the 1-sigma level, and include only measurement errors.
- External (systematic) errors are shown as <sup>206</sup>Pb/<sup>238</sup>U uncertainty, <sup>206</sup>Pb/<sup>207</sup>Pb uncertainty to the right of each sample (in %, at 2-sigma level).
- Analyses conducted by LA-MC-ICPMS, as described by Gehrels et al. (2008).
- U concentration and U/Th are calibrated relative to Sri Lanka zircon standard and are accurate to ~20%.
- Common Pb correction is from measured <sup>204</sup>Pb with common Pb composition interpreted from Stacey and Kramers (1975).
- Uncertainties of 1.5 for <sup>206</sup>Pb/<sup>204</sup>Pb, 0.3 for <sup>207</sup>Pb/<sup>204</sup>Pb, and 2.0 for <sup>208</sup>Pb/<sup>204</sup>Pb are applied to common Pb composition.
- U/Pb and <sup>206</sup>Pb/<sup>207</sup>Pb fractionation is calibrated relative to fragments of a large Sri Lanka zircon of 563.5 ± 3.2 Ma (2-sigma).
- U decay constants and composition as follows: <sup>235</sup>U = 9.8485 x 10<sup>-10</sup>, <sup>238</sup>U = 1.55125 x 10<sup>-10</sup>, <sup>238</sup>U/<sup>235</sup>U = 137.88.
- Weighted mean and concordia plots determined with Isoplot (Ludwig, 2008).
- Analytical methods as described by Gehrels and Pecha (2014).

## APPENDIX F

### Hafnium isotope data of selected Harmony Formation strata

Table notes are at the end of the appendix.

Sample: LCC #1. Location: Little Cottonwood Canyon, Galena Range; 0490600 4494792 (NAD 83 UTM 11T)	$(^{176}\text{Yb} + ^{176}\text{Lu}) / ^{176}\text{Hf}$ (%)	Volts Hf	$^{176}\text{Hf}/^{177}\text{Hf}$	$\pm$ (1s)	$^{176}\text{Lu}/^{177}\text{Hf}$	$^{176}\text{Hf}/^{177}\text{Hf}$ (T)	E-Hf (0)	E-Hf (0) $\pm$ (1s)	E-Hf (T)	Age (Ma)
LCC1-88-RM15-81	15.3	4.1	0.282381	0.000037	0.00081	0.282371	-14.3	1.3	0.4	673
LCC1-88-RM15-137	13.1	3.9	0.282213	0.000037	0.00074	0.282204	-20.2	1.3	-5.3	682
LCC1-88-RM15-127	7.6	4.2	0.282300	0.000026	0.00043	0.282294	-17.2	0.9	-1.6	705
LCC1-88-RM15-93	7.8	4.3	0.282392	0.000031	0.00044	0.282386	-13.9	1.1	1.9	717
LCC1-88-RM15-152	12.2	5.7	0.282285	0.000029	0.00072	0.282271	-17.7	1.0	4.9	1029
LCC1-88-RM15-190	6.7	5.3	0.282195	0.000025	0.00040	0.282187	-20.9	0.9	1.9	1031
LCC1-88-RM15-130	9.0	6.1	0.282255	0.000027	0.00053	0.282245	-18.7	0.9	4.5	1053
LCC1-88-RM15-157	17.4	5.3	0.282280	0.000026	0.00104	0.282259	-17.9	0.9	5.4	1070
LCC1-88-RM15-139	13.8	6.1	0.282308	0.000030	0.00079	0.282292	-16.9	1.1	6.6	1071
LCC1-88-RM15-110	9.0	5.4	0.282308	0.000020	0.00051	0.282298	-16.9	0.7	6.8	1073
LCC1-88-RM15-62	11.3	4.8	0.282303	0.000027	0.00069	0.282289	-17.0	0.9	6.6	1075
LCC1-88-RM15-133	42.4	7.2	0.282361	0.000029	0.00265	0.282306	-15.0	1.0	7.5	1089
LCC1-88-RM15-192	16.2	5.2	0.282387	0.000032	0.00093	0.282368	-14.1	1.1	9.7	1091
LCC1-88-RM15-32	9.1	5.2	0.282311	0.000039	0.00053	0.282300	-16.7	1.4	7.3	1091
LCC1-88-RM15-149	16.0	4.7	0.282378	0.000036	0.00084	0.282361	-14.4	1.3	9.5	1091
LCC1-88-RM15-7	17.1	5.7	0.282326	0.000028	0.00102	0.282305	-16.2	1.0	7.5	1092
LCC1-88-RM15-95	23.8	4.8	0.282291	0.000031	0.00134	0.282263	-17.5	1.1	6.3	1106
LCC1-88-RM15-36	8.0	5.8	0.282358	0.000033	0.00049	0.282347	-15.1	1.2	9.4	1108
LCC1-88-RM15-3	20.8	4.2	0.282321	0.000034	0.00119	0.282296	-16.4	1.2	7.6	1108
LCC1-88-RM15-162	14.4	6.7	0.282299	0.000022	0.00081	0.282282	-17.2	0.8	7.1	1109
LCC1-88-RM15-86	5.9	4.8	0.282367	0.000030	0.00045	0.282357	-14.8	1.1	10.2	1129
LCC1-88-RM15-79	19.2	4.9	0.282194	0.000024	0.00108	0.282171	-20.9	0.9	3.6	1129
LCC1-88-RM15-10	14.1	5.0	0.282225	0.000026	0.00086	0.282206	-19.8	0.9	5.4	1153
LCC1-88-RM15-80	9.6	2.8	0.282166	0.000032	0.00060	0.282153	-21.9	1.1	3.6	1154
LCC1-88-RM15-13	18.4	4.7	0.282283	0.000028	0.00109	0.282258	-17.8	1.0	8.0	1187
LCC1-88-RM15-30	8.9	5.9	0.282159	0.000022	0.00056	0.282146	-22.1	0.8	4.1	1190
LCC1-88-RM15-57	11.5	5.0	0.282272	0.000032	0.00072	0.282256	-18.1	1.1	8.5	1213
LCC1-88-RM15-67	40.1	3.1	0.282289	0.000040	0.00196	0.282240	-17.6	1.4	10.3	1313
LCC1-88-RM15-75	10.2	5.8	0.282142	0.000020	0.00058	0.282128	-22.7	0.7	7.1	1350
LCC1-88-RM15-45	17.9	5.1	0.282211	0.000028	0.00105	0.282184	-20.3	1.0	9.5	1369
LCC1-88-RM15-174	14.5	4.3	0.282177	0.000026	0.00090	0.282154	-21.5	0.9	8.5	1370
LCC1-88-RM15-188	14.7	6.2	0.282015	0.000023	0.00087	0.281990	-27.2	0.8	4.9	1468
LCC1-88-RM15-165	11.7	5.6	0.282036	0.000024	0.00073	0.282015	-26.5	0.8	5.9	1471
LCC1-88-RM15-103	12.1	5.0	0.281911	0.000019	0.00078	0.281887	-30.9	0.7	5.8	1668
LCC1-88-RM15-175	9.5	3.2	0.282344	0.000027	0.00057	0.282325	-15.6	0.9	22.7	1724
LCC1-88-RM15-94	19.2	4.5	0.282006	0.000029	0.00112	0.281969	-27.5	1.0	10.3	1736
LCC1-88-RM15-153	11.7	5.4	0.280939	0.000027	0.00076	0.280902	-65.3	1.0	-8.7	2549
<b>Sample: LCC #2. Location: Little Cottonwood Canyon, Galena Range; 0490529 4495008 (NAD 83 UTM 11T)</b>										
LCC2-H971-121	13.7	4.7	0.282352	0.000021	0.00076	0.282337	-15.3	0.7	6.8	1009
LCC2-H971-91	12.4	5.6	0.282246	0.000022	0.00072	0.282232	-19.0	0.8	3.5	1027
LCC2-H971-39	22.1	3.9	0.282344	0.000023	0.00128	0.282319	-15.6	0.8	6.8	1040
LCC2-H971-3	8.5	4.3	0.282266	0.000021	0.00050	0.282257	-18.3	0.7	4.7	1044
LCC2-H971-36	7.9	4.4	0.282338	0.000023	0.00046	0.282329	-15.8	0.8	7.5	1055
LCC2-H971-29	13.3	4.1	0.282296	0.000025	0.00077	0.282280	-17.3	0.9	6.0	1064
LCC2-H971-49	23.8	3.7	0.282353	0.000026	0.00138	0.282325	-15.3	0.9	7.8	1073
LCC2-H971-27	12.4	4.1	0.282343	0.000028	0.00071	0.282329	-15.6	1.0	8.0	1076
LCC2-H971-63	22.5	4.0	0.282260	0.000027	0.00128	0.282234	-18.6	0.9	4.9	1088
LCC2-H971-28	15.9	4.2	0.282291	0.000023	0.00089	0.282273	-17.5	0.8	6.4	1093
LCC2-H971-34	10.9	4.7	0.282352	0.000022	0.00067	0.282338	-15.3	0.8	8.9	1103
LCC2-H971-20	14.6	5.2	0.282201	0.000018	0.00085	0.282183	-20.7	0.6	3.5	1105
LCC2-H971-51	17.4	4.6	0.282296	0.000024	0.00095	0.282276	-17.3	0.8	7.0	1116
LCC2-H971-131	17.3	6.4	0.282340	0.000016	0.00097	0.282320	-15.7	0.6	8.7	1122
LCC2-H971-134	19.4	4.7	0.282351	0.000022	0.00113	0.282327	-15.3	0.8	9.2	1133
LCC2-H971-8	17.7	4.7	0.282138	0.000022	0.00106	0.282115	-22.9	0.8	2.0	1146
LCC2-H971-72	19.6	5.4	0.282292	0.000017	0.00122	0.282265	-17.4	0.6	7.5	1152
LCC2-H971-33	26.2	4.1	0.282186	0.000026	0.00151	0.282152	-21.2	0.9	3.6	1157
LCC2-H971-6	26.8	5.5	0.282248	0.000019	0.00164	0.282213	-19.0	0.7	5.7	1158
LCC2-H971-42	31.5	4.0	0.282262	0.000026	0.00181	0.282222	-18.5	0.9	6.3	1167
LCC2-H971-37	40.1	3.0	0.282266	0.000032	0.00199	0.282222	-18.4	1.1	6.3	1169
LCC2-H971-68	10.0	5.8	0.282121	0.000017	0.00057	0.282109	-23.5	0.6	2.6	1184
LCC2-H971-65	15.3	5.2	0.282277	0.000021	0.00094	0.282255	-18.0	0.7	8.6	1216
LCC2-H971-98	62.7	3.3	0.282377	0.000033	0.00333	0.282300	-14.4	1.2	10.5	1233
LCC2-H971-79	5.3	4.5	0.282241	0.000019	0.00033	0.282233	-19.2	0.7	8.6	1253
LCC2-H971-26	5.2	4.5	0.282116	0.000018	0.00032	0.282108	-23.7	0.6	4.5	1265
LCC2-H971-92	29.3	4.9	0.282251	0.000025	0.00154	0.282213	-18.9	0.9	9.7	1333
LCC2-H971-77	18.5	4.3	0.282150	0.000022	0.00104	0.282124	-22.4	0.8	6.7	1338
LCC2-H971-44	23.3	4.3	0.282094	0.000024	0.00150	0.282055	-24.5	0.8	4.9	1368
LCC2-H971-67	28.6	3.9	0.282176	0.000025	0.00169	0.282132	-21.5	0.9	7.9	1378
LCC2-H971-70	23.4	4.3	0.282215	0.000023	0.00129	0.282181	-20.2	0.8	9.8	1385
LCC2-H971-84	14.3	5.1	0.282194	0.000019	0.00090	0.282171	-20.9	0.7	9.5	1387
LCC2-H971-14	18.1	4.9	0.282107	0.000025	0.00110	0.282076	-24.0	0.9	7.7	1456
LCC2-H971-23	10.7	4.9	0.282121	0.000022	0.00072	0.282101	-23.5	0.8	8.7	1461
LCC2-H971-41	16.4	4.7	0.282127	0.000020	0.00097	0.282100	-23.3	0.7	8.8	1469
LCC2-H971-19	43.4	5.6	0.282109	0.000019	0.00256	0.282037	-23.9	0.7	6.9	1482
LCC2-H971-24	13.2	4.8	0.281328	0.000020	0.00093	0.281296	-51.5	0.7	-12.4	1786
LCC2-H971-69	19.4	4.9	0.281902	0.000019	0.00119	0.281861	-31.2	0.7	7.8	1794
LCC2-H971-48	7.9	4.8	0.281207	0.000021	0.00051	0.281180	-55.8	0.7	4.4	2687

### Hafnium isotope data of selected Harmony Formation strata

	$(^{176}\text{Yb} + ^{176}\text{Lu}) / ^{176}\text{Hf}$ (%)	Volts Hf	$^{176}\text{Hf}/^{177}\text{Hf}$	$\pm$ (1s)	$^{176}\text{Lu}/^{177}\text{Hf}$	$^{176}\text{Hf}/^{177}\text{Hf}$ (T)	E-Hf (0)	E-Hf (0) $\pm$ (1s)	E-Hf (T)	Age (Ma)
<b>Sample: LCC #3. Location: Little Cottonwood Canyon, Galena Range; 0490459 4494916 (NAD 83 UTM 11T)</b>										
LCC3-H973-1	8.3	5.4	0.281539	0.000017	0.00055	0.281519	-44.1	0.6	-2.5	1871
LCC3-H973-2	38.9	4.2	0.281556	0.000029	0.00242	0.281474	-43.5	1.0	-6.6	1765
LCC3-H973-4	8.4	5.0	0.281215	0.000018	0.00055	0.281186	-55.5	0.6	5.7	2735
LCC3-H973-49	11.3	5.2	0.281586	0.000018	0.00062	0.281565	-42.4	0.6	-2.9	1786
LCC3-H973-58	18.2	4.3	0.281943	0.000022	0.00107	0.281907	-29.8	0.8	9.0	1775
LCC3-H973-90	10.4	4.4	0.281500	0.000024	0.00062	0.281478	-45.4	0.8	-4.4	1852
LCC3-H973-86	6.7	5.0	0.281209	0.000020	0.00040	0.281188	-55.7	0.7	3.0	2614
LCC3-H973-88	18.1	5.2	0.281230	0.000019	0.00111	0.281192	-55.0	0.7	-15.9	1793
LCC3-H973-115	15.8	3.6	0.281227	0.000022	0.00097	0.281177	-55.1	0.8	4.0	2673
LCC3-H973-37	8.0	5.3	0.280842	0.000016	0.00048	0.280817	-68.7	0.6	-7.9	2712
LCC3-H973-33	10.1	5.1	0.281244	0.000017	0.00060	0.281214	-54.5	0.6	3.9	2612
LCC3-H973-34	9.6	5.1	0.281281	0.000018	0.00058	0.281261	-53.2	0.6	-13.5	1793
LCC3-H973-27	11.0	4.2	0.281158	0.000021	0.00068	0.281123	-57.5	0.7	2.5	2693
LCC3-H973-26	17.6	2.3	0.281784	0.000047	0.00116	0.281745	-35.4	1.7	3.5	1787
LCC3-H973-30	29.2	4.4	0.281078	0.000025	0.00164	0.280995	-60.4	0.9	-3.0	2652
LCC3-H973-9	2.8	5.0	0.281089	0.000019	0.00020	0.281079	-60.0	0.7	0.8	2688
LCC3-H973-8	8.9	5.0	0.281568	0.000017	0.00057	0.281549	-43.0	0.6	-3.1	1801
LCC3-H973-7	9.6	6.1	0.281182	0.000016	0.00053	0.281155	-56.7	0.6	1.3	2593
LCC3-H973-10	11.3	5.9	0.281623	0.000018	0.00074	0.281598	-41.1	0.7	-1.7	1787
LCC3-H973-11	8.3	4.8	0.281506	0.000019	0.00051	0.281488	-45.2	0.7	-3.8	1864
LCC3-H973-53	12.3	4.0	0.281520	0.000022	0.00089	0.281489	-44.7	0.8	-4.2	1843
LCC3-H973-15	7.7	4.8	0.281526	0.000016	0.00050	0.281509	-44.5	0.6	-3.8	1832
LCC3-H973-21	11.8	5.2	0.281775	0.000016	0.00069	0.281750	-35.7	0.6	5.2	1852
LCC3-H973-20	10.8	4.9	0.281251	0.000016	0.00064	0.281219	-54.3	0.6	3.5	2590
LCC3-H973-25	9.7	5.7	0.281082	0.000014	0.00056	0.281063	-60.2	0.5	-19.8	1827
LCC3-H973-70	11.6	5.1	0.281149	0.000017	0.00070	0.281112	-57.9	0.6	2.7	2718
LCC3-H973-71	15.1	4.8	0.281630	0.000020	0.00085	0.281601	-40.8	0.7	-1.5	1791
LCC3-H973-73	13.1	3.8	0.281554	0.000021	0.00079	0.281527	-43.5	0.7	-4.5	1771
LCC3-H973-74	13.9	4.5	0.281227	0.000021	0.00084	0.281186	-55.1	0.8	2.7	2608
LCC3-H973-96	24.4	4.4	0.281501	0.000021	0.00142	0.281451	-45.4	0.7	-5.3	1856
<b>Sample: LCC #4. Location: Little Cottonwood Canyon, Galena Range; 0490623 4494977 (NAD 83 UTM 11T)</b>										
LCC4-88RMA13-1	11.1	5.0	0.281525	0.000022	0.00063	0.281503	-44.6	0.8	-5.0	1787
LCC4-88RMA13-2	18.4	4.2	0.281891	0.000021	0.00112	0.281851	-31.6	0.7	9.2	1867
LCC4-88RMA13-22	13.5	4.6	0.281213	0.000020	0.00079	0.281174	-55.6	0.7	1.8	2587
LCC4-88RMA13-23	7.6	4.8	0.281680	0.000023	0.00043	0.281666	-39.1	0.8	0.7	1783
LCC4-88RMA13-24	8.0	4.3	0.281378	0.000019	0.00046	0.281362	-49.8	0.7	-10.4	1773
LCC4-88RMA13-152	7.5	4.2	0.281844	0.000021	0.00045	0.281828	-33.3	0.7	8.4	1870
LCC4-88RMA13-11	4.2	4.5	0.281024	0.000021	0.00023	0.281013	-62.3	0.7	-3.9	2586
LCC4-88RMA13-16	9.3	4.4	0.281166	0.000020	0.00060	0.281135	-57.2	0.7	3.0	2693
LCC4-88RMA13-17	8.5	5.1	0.281788	0.000020	0.00049	0.281771	-35.3	0.7	4.4	1784
LCC4-88RMA13-189	13.3	4.4	0.281448	0.000022	0.00068	0.281424	-47.3	0.8	-6.3	1856
LCC4-88RMA13-48	19.5	4.8	0.281793	0.000020	0.00107	0.281757	-35.1	0.7	3.7	1774
LCC4-88RMA13-44	13.4	4.5	0.281247	0.000018	0.00069	0.281213	-54.4	0.6	3.2	2586
LCC4-88RMA13-42	8.7	2.9	0.281210	0.000035	0.00056	0.281192	-55.7	1.2	-16.6	1766
LCC4-88RMA13-126	14.9	0.4	0.281312	0.000062	0.00102	0.281262	-52.1	2.2	4.7	2576
LCC4-88RMA13-35	1.8	4.6	0.281824	0.000022	0.00013	0.281819	-34.0	0.8	7.8	1857
LCC4-88RMA13-37	13.8	5.5	0.281269	0.000018	0.00080	0.281242	-53.6	0.6	-14.7	1770
LCC4-88RMA13-38	28.5	5.1	0.281778	0.000020	0.00149	0.281728	-35.6	0.7	2.2	1756
LCC4-88RMA13-39	9.2	4.7	0.281510	0.000020	0.00052	0.281493	-45.1	0.7	-5.7	1776
LCC4-88RMA13-181	18.1	5.5	0.281676	0.000020	0.00102	0.281641	-39.2	0.7	-0.2	1785
LCC4-88RMA13-177	13.6	4.1	0.281541	0.000020	0.00077	0.281515	-44.0	0.7	-4.9	1777
LCC4-88RMA13-109	9.5	5.0	0.281609	0.000018	0.00065	0.281587	-41.6	0.6	-1.2	1825
LCC4-88RMA13-170	11.2	4.4	0.281659	0.000023	0.00066	0.281636	-39.8	0.8	-0.4	1783
LCC4-88RMA13-192	10.1	4.2	0.281207	0.000022	0.00067	0.281174	-55.8	0.8	1.6	2578
LCC4-88RMA13-111	18.8	4.4	0.281560	0.000022	0.00113	0.281521	-43.3	0.8	-3.4	1831
LCC4-88RMA13-118	11.2	5.2	0.281221	0.000017	0.00069	0.281186	-55.3	0.6	4.6	2686
LCC4-88RMA13-126	20.2	4.3	0.281189	0.000022	0.00115	0.281133	-56.4	0.8	0.1	2576
LCC4-88RMA13-197	9.9	4.6	0.281677	0.000020	0.00056	0.281658	-39.2	0.7	0.4	1784
<b>Sample: LCC #9. Location: Little Cottonwood Canyon, Galena Range; 0491232 4495602 (NAD 83 UTM 11T)</b>										
1312-LCC-09HA-9	14.3	5.2	0.281550	0.000024	0.00078	0.281525	-43.7	0.8	-5.4	1736
1312-LCC-09HA-51	19.6	5.6	0.281600	0.000018	0.00113	0.281562	-41.9	0.6	-3.7	1754
1312-LCC-09HA-12	10.1	5.1	0.281544	0.000020	0.00060	0.281524	-43.9	0.7	-4.7	1768
1312-LCC-09HA-37	9.7	4.6	0.281399	0.000022	0.00057	0.281380	-49.0	0.8	-9.8	1771
1312-LCC-09HA-2	12.6	5.1	0.281527	0.000024	0.00074	0.281502	-44.5	0.8	-5.3	1776
1312-LCC-09HA-14	9.3	5.3	0.281552	0.000018	0.00055	0.281533	-43.8	0.6	-4.0	1784
1312-LCC-09HA-20	16.7	6.1	0.281604	0.000016	0.00098	0.281571	-41.8	0.6	-2.5	1791
1312-LCC-09HA-8	10.9	5.0	0.281567	0.000016	0.00065	0.281545	-43.1	0.6	-3.4	1793
1312-LCC-09HA-38	8.7	5.5	0.281600	0.000021	0.00052	0.281582	-41.9	0.8	-2.0	1798
1312-LCC-09HA-50	13.2	5.2	0.281576	0.000019	0.00076	0.281550	-42.7	0.7	-3.1	1798
1312-LCC-09HA-31	10.1	5.0	0.281553	0.000018	0.00064	0.281531	-43.6	0.6	-3.7	1802
1312-LCC-09HA-34	10.5	3.3	0.281579	0.000028	0.00065	0.281556	-42.7	1.0	-2.7	1805
1312-LCC-09HA-55	19.4	4.2	0.281828	0.000029	0.00107	0.281791	-33.9	1.0	6.2	1833
1312-LCC-09HA-44	13.5	4.3	0.281438	0.000020	0.00081	0.281410	-47.6	0.7	-7.1	1841
1312-LCC-09HA-3	16.4	5.1	0.281828	0.000022	0.00097	0.281793	-33.9	0.8	6.7	1851
1312-LCC-09HA-42	14.3	4.6	0.281942	0.000019	0.00096	0.281908	-29.8	0.7	11.1	1862
1312-LCC-09HA-4	14.8	4.2	0.281451	0.000025	0.00091	0.281418	-47.2	0.9	-6.3	1863
1312-LCC-09HA-28	5.4	4.7	0.281834	0.000021	0.00034	0.281822	-33.6	0.7	8.2	1870
1312-LCC-09HA-84	11.0	5.1	0.281862	0.000022	0.00072	0.281836	-32.6	0.8	9.0	1885
1312-LCC-09HA-89	5.9	5.1	0.281875	0.000023	0.00036	0.281862	-32.2	0.8	10.3	1898
1312-LCC-09HA-82	3.8	2.7	0.281187	0.000027	0.00025	0.281175	-56.5	1.0	-0.3	2492
1312-LCC-09HA-71	8.1	5.4	0.281232	0.000016	0.00049	0.281207	-54.9	0.6	3.8	2620
1312-LCC-09HA-19	28.4	3.6	0.281313	0.000030	0.00164	0.281230	-52.1	1.1	4.8	2626
1312-LCC-09HA-1	11.1	4.9	0.281039	0.000020	0.00067	0.281004	-61.8	0.7	-1.1	2720



## Hafnium isotope data of selected Harmony Formation strata

	$(^{176}\text{Yb} + ^{176}\text{Lu}) / ^{176}\text{Hf}$ (%)	Volts Hf	$^{176}\text{Hf}/^{177}\text{Hf}$	$\pm$ (1s)	$^{176}\text{Lu}/^{177}\text{Hf}$	$^{176}\text{Hf}/^{177}\text{Hf}$ (T)	E-Hf (0)	E-Hf (0) $\pm$ (1s)	E-Hf (T)	Age (Ma)
<b>Sample: LCC #10. Location: Little Cottonwood Canyon, Galena Range; 0490510 4495682 (NAD 83 UTM 11T)</b>										
LCC-10HQ-7	13.1	5.7	0.282309	0.000020	0.00080	0.282293	-16.8	0.7	6.4	1064
LCC-10HQ-8	29.9	3.3	0.282055	0.000026	0.00185	0.282004	-25.8	0.9	5.2	1460
LCC-10HQ-9	32.6	2.6	0.282360	0.000031	0.00202	0.282320	-15.0	1.1	7.2	1057
LCC-10HQ-10	16.6	2.7	0.282160	0.000038	0.00107	0.282132	-22.1	1.3	8.3	1396
LCC-10HQ-12	14.3	2.6	0.282325	0.000030	0.00085	0.282306	-16.3	1.0	9.6	1183
LCC-10HQ-11	35.5	2.6	0.282188	0.000027	0.00211	0.282135	-21.1	1.0	6.8	1325
LCC-10HQ-90	9.5	4.1	0.282171	0.000021	0.00072	0.282152	-21.7	0.7	9.1	1399
LCC-10HQ-86	12.7	0.4	0.282394	0.000083	0.00115	0.282361	-13.8	2.9	18.7	1493
LCC-10HQ-45	19.9	3.8	0.282013	0.000027	0.00120	0.281986	-27.3	1.0	-0.9	1221
LCC-10HQ-99	15.3	2.6	0.282252	0.000029	0.00087	0.282232	-18.9	1.0	6.6	1164
LCC-10HQ-100	18.2	3.2	0.282278	0.000031	0.00108	0.282256	-17.9	1.1	5.9	1097
LCC-10HQ-96	10.2	2.9	0.281761	0.000023	0.00063	0.281741	-36.2	0.8	2.1	1729
LCC-10HQ-95	21.2	2.8	0.282253	0.000032	0.00130	0.282224	-18.8	1.1	6.1	1158
LCC-10HQ-93	11.3	3.2	0.281040	0.000024	0.00069	0.281005	-61.7	0.8	-1.5	2700
LCC-10HQ-16	16.4	1.2	0.282849	0.000098	0.00400	0.282774	2.3	3.5	22.0	996
LCC-10HQ-17	20.4	4.4	0.281904	0.000025	0.00128	0.281878	-31.2	0.9	-8.1	1072
LCC-10HQ-18	13.5	5.3	0.282292	0.000021	0.00080	0.282277	-17.4	0.8	4.5	1003
LCC-10HQ-19	16.3	3.2	0.282387	0.000027	0.00093	0.282367	-14.1	1.0	10.9	1146
LCC-10HQ-20	6.7	3.0	0.281528	0.000023	0.00044	0.281513	-44.4	0.8	-4.5	1797
LCC-10HQ-79	20.5	5.7	0.282277	0.000018	0.00116	0.282250	-18.0	0.6	8.2	1209
LCC-10HQ-77	23.7	3.4	0.281867	0.000024	0.00136	0.281823	-32.5	0.9	4.3	1700
LCC-10HQ-39	9.9	6.4	0.281794	0.000019	0.00063	0.281772	-35.1	0.7	4.9	1803
LCC-10HQ-1	16.6	3.6	0.282336	0.000024	0.00098	0.282315	-15.9	0.8	8.8	1134
LCC-10HQ-2	9.9	2.2	0.281206	0.000032	0.00065	0.281173	-55.8	1.1	3.5	2660
LCC-10HQ-3	6.2	4.6	0.282332	0.000022	0.00044	0.282324	-16.0	0.8	6.7	1027
LCC-10HQ-4	12.2	5.3	0.281949	0.000019	0.00078	0.281928	-29.6	0.7	2.5	1460
LCC-10HQ-5	22.0	0.6	0.282171	0.000091	0.00157	0.282127	-21.7	3.2	10.1	1483
<b>Sample: Harmony Canyon. Location: Harmony Canyon, Sonoma Range; 0446225 4533064 (NAD 83 UTM 11T)</b>										
HARM-CYN-143	16.4	5.4	0.281662	0.000027	0.00089	0.281633	-39.7	1.0	-1.3	1750
HARM-CYN-126	38.8	5.4	0.281693	0.000028	0.00210	0.281623	-38.6	1.0	-1.4	1762
HARM-CYN-17	9.1	3.7	0.281588	0.000037	0.00059	0.281568	-42.3	1.3	-3.0	1773
HARM-CYN-19	16.6	3.0	0.281565	0.000039	0.00098	0.281532	-43.2	1.4	-4.3	1775
HARM-CYN-146	11.4	5.2	0.281580	0.000033	0.00069	0.281557	-42.6	1.2	-3.4	1776
HARM-CYN-134	7.3	5.0	0.281592	0.000026	0.00044	0.281577	-42.2	0.9	-2.6	1777
HARM-CYN-94	14.1	4.7	0.281449	0.000028	0.00087	0.281420	-47.2	1.0	-8.2	1778
HARM-CYN-129	14.4	5.0	0.281737	0.000034	0.00093	0.281706	-37.1	1.2	2.1	1784
HARM-CYN-59	8.4	5.3	0.281591	0.000026	0.00050	0.281574	-42.2	0.9	-2.6	1785
HARM-CYN-32	13.1	5.2	0.281607	0.000030	0.00077	0.281581	-41.6	1.1	-2.3	1785
HARM-CYN-46	8.5	5.9	0.281554	0.000025	0.00049	0.281538	-43.5	0.9	-3.8	1788
HARM-CYN-58	12.4	4.8	0.281522	0.000030	0.00072	0.281498	-44.6	1.1	-5.2	1788
HARM-CYN-89	9.7	4.6	0.281580	0.000034	0.00064	0.281559	-42.6	1.2	-3.0	1788
HARM-CYN-169	11.9	4.5	0.281500	0.000029	0.00072	0.281475	-45.5	1.0	-6.0	1788
HARM-CYN-78	10.7	5.8	0.281764	0.000026	0.00067	0.281741	-36.1	0.9	4.1	1818
HARM-CYN-187	20.0	5.9	0.281518	0.000034	0.00124	0.281475	-44.8	1.2	-5.1	1827
HARM-CYN-117	49.9	4.2	0.281923	0.000037	0.00275	0.281827	-30.5	1.3	7.5	1832
HARM-CYN-115	7.1	4.7	0.281872	0.000020	0.00048	0.281856	-32.3	0.7	8.5	1833
HARM-CYN-121	3.6	4.4	0.281070	0.000028	0.00026	0.281061	-60.6	1.0	-19.5	1842
HARM-CYN-130	5.3	2.8	0.281553	0.000043	0.00037	0.281540	-43.6	1.5	-2.3	1847
HARM-CYN-35	11.5	2.9	0.281824	0.000035	0.00069	0.281800	-34.0	1.2	6.9	1849
HARM-CYN-111	9.0	4.5	0.281855	0.000027	0.00051	0.281837	-32.9	1.0	8.3	1850
HARM-CYN-68	17.1	4.6	0.281512	0.000029	0.00102	0.281476	-45.0	1.0	-4.5	1850
HARM-CYN-64	12.8	5.1	0.281460	0.000025	0.00078	0.281432	-46.9	0.9	-5.7	1867
HARM-CYN-195	11.0	3.9	0.281456	0.000034	0.00069	0.281431	-47.0	1.2	-5.7	1868
HARM-CYN-62	15.6	4.4	0.281815	0.000027	0.00104	0.281778	-34.3	0.9	6.6	1868
HARM-CYN-107	20.9	4.8	0.281574	0.000024	0.00126	0.281529	-42.8	0.8	-2.1	1873
HARM-CYN-105	14.6	2.5	0.281891	0.000034	0.00094	0.281858	-31.6	1.2	9.6	1875
HARM-CYN-135	5.8	6.0	0.281182	0.000026	0.00033	0.281167	-56.7	0.9	-1.5	2453
HARM-CYN-11	13.1	2.8	0.281184	0.000039	0.00077	0.281146	-56.6	1.4	0.9	2591
HARM-CYN-162	22.1	5.0	0.281243	0.000039	0.00143	0.281171	-54.5	1.4	2.3	2612
HARM-CYN-37	14.2	3.0	0.281162	0.000037	0.00087	0.281117	-57.4	1.3	0.9	2634
HARM-CYN-72	10.2	4.1	0.281259	0.000029	0.00076	0.281221	-54.0	1.0	4.6	2634
HARM-CYN-157	14.2	5.7	0.281703	0.000033	0.00085	0.281660	-38.3	1.2	20.5	2644
<b>Sample: Elbow Canyon. Location: Elbow Canyon, Sonoma Range; 0441026 4514450 (NAD 83 UTM 11T)</b>										
Elbow-CYN-74	9.2	5.0	0.281512	0.000021	0.00053	0.281494	-45.0	0.7	-5.6	1776
Elbow-CYN-66	15.1	4.9	0.281664	0.000019	0.00085	0.281636	-39.6	0.7	-0.6	1777
Elbow-CYN-42	8.6	5.3	0.281542	0.000024	0.00049	0.281525	-44.0	0.9	-4.5	1777
Elbow-CYN-195	6.4	3.7	0.281438	0.000027	0.00035	0.281426	-47.6	1.0	-8.0	1778
Elbow-CYN-45	10.0	5.5	0.281386	0.000026	0.00057	0.281366	-49.5	0.9	-10.1	1780
Elbow-CYN-163	19.1	4.4	0.281522	0.000029	0.00102	0.281487	-44.7	1.0	-5.7	1782
Elbow-CYN-170	13.4	4.8	0.281538	0.000036	0.00076	0.281512	-44.1	1.3	-4.8	1782
Elbow-CYN-76	39.5	6.0	0.281676	0.000040	0.00161	0.281621	-39.2	1.4	-0.9	1783
Elbow-CYN-162	8.5	4.7	0.281408	0.000025	0.00049	0.281392	-48.7	0.9	-9.1	1784
Elbow-CYN-182	16.2	5.2	0.281651	0.000018	0.00092	0.281620	-40.1	0.6	-0.9	1786
Elbow-CYN-160	8.7	5.1	0.281563	0.000021	0.00048	0.281547	-43.2	0.7	-3.5	1786
Elbow-CYN-171	9.5	4.3	0.281595	0.000026	0.00061	0.281574	-42.1	0.9	-2.5	1787
Elbow-CYN-127	11.4	4.8	0.281570	0.000018	0.00063	0.281549	-43.0	0.6	-3.4	1787
Elbow-CYN-92	7.7	4.4	0.281379	0.000019	0.00046	0.281364	-49.7	0.7	-10.0	1787
Elbow-CYN-115	10.9	5.0	0.281633	0.000019	0.00063	0.281612	-40.7	0.7	-1.2	1788
Elbow-CYN-159	9.8	5.4	0.281585	0.000017	0.00057	0.281565	-42.5	0.6	-2.8	1788
Elbow-CYN-187	6.3	5.4	0.281372	0.000018	0.00038	0.281360	-50.0	0.6	-10.1	1789
Elbow-CYN-129	7.4	5.4	0.281525	0.000022	0.00043	0.281510	-44.6	0.8	-4.7	1791
Elbow-CYN-116	10.6	5.2	0.281749	0.000020	0.00060	0.281729	-36.6	0.7	3.2	1795
Elbow-CYN-70	6.0	4.0	0.281370	0.000030	0.00036	0.281358	-50.0	1.1	-9.3	1825
Elbow-CYN-133	11.5	5.0	0.281774	0.000025	0.00069	0.281750	-35.7	0.9	4.9	1836

### Hafnium isotope data of selected Harmony Formation strata

	$(^{176}\text{Yb} + ^{176}\text{Lu}) / ^{176}\text{Hf}$ (%)	Volts Hf	$^{176}\text{Hf}/^{177}\text{Hf}$	$\pm$ (1s)	$^{176}\text{Lu}/^{177}\text{Hf}$	$^{176}\text{Hf}/^{177}\text{Hf}$ (T)	E-Hf (0)	E-Hf (0) $\pm$ (1s)	E-Hf (T)	Age (Ma)
<b>Sample: Elbow Canyon. Location: Elbow Canyon, Sonoma Range; 0441026 4514450 (NAD 83 UTM 11T)</b>										
Elbow-CYN-53	12.7	4.5	0.281584	0.000016	0.00075	0.281557	-42.5	0.6	-1.7	1847
Elbow-CYN-23	4.7	3.9	0.281357	0.000030	0.00030	0.281347	-50.5	1.1	-9.0	1857
Elbow-CYN-132	7.4	4.8	0.281823	0.000023	0.00045	0.281807	-34.0	0.8	7.6	1865
Elbow-CYN-67	3.6	4.0	0.281577	0.000027	0.00021	0.281569	-42.7	1.0	-0.9	1867
Elbow-CYN-93	10.9	5.0	0.281504	0.000019	0.00067	0.281480	-45.3	0.7	-3.9	1871
Elbow-CYN-112	7.5	4.7	0.281389	0.000017	0.00044	0.281369	-49.4	0.6	5.2	2436
Elbow-CYN-141	11.0	5.3	0.281135	0.000023	0.00065	0.281103	-58.3	0.8	-0.7	2586
Elbow-CYN-87	24.4	4.6	0.281237	0.000020	0.00118	0.281178	-54.7	0.7	2.0	2589
Elbow-CYN-190	24.4	4.5	0.281249	0.000020	0.00130	0.281185	-54.3	0.7	2.3	2592
Elbow-CYN-19	7.2	5.2	0.281195	0.000033	0.00042	0.281174	-56.2	1.2	2.4	2612
Elbow-CYN-58	11.6	3.9	0.281272	0.000020	0.00068	0.281238	-53.5	0.7	4.7	2614
Elbow-CYN-37	10.7	4.2	0.281237	0.000025	0.00064	0.281205	-54.7	0.9	4.0	2634
Elbow-CYN-79	13.5	4.0	0.281163	0.000016	0.00080	0.281121	-57.4	0.6	2.1	2678
Elbow-CYN-96	3.1	5.2	0.280906	0.000025	0.00017	0.280897	-66.4	0.9	-5.6	2691
Elbow-CYN-134	6.0	4.1	0.281033	0.000022	0.00037	0.281013	-62.0	0.8	-1.4	2692
Elbow-CYN-28	8.2	4.0	0.281082	0.000021	0.00053	0.281055	-60.2	0.7	0.6	2716
Elbow-CYN-191	8.1	4.5	0.281178	0.000022	0.00047	0.281154	-56.8	0.8	4.2	2718

#### Notes:

1. Data reduction methodology is from Woodhead et al. (2004)
2. Analytical methods described in detail by Gehrels and Pecha (2014)
3.  $(^{176}\text{Yb} + ^{176}\text{Lu}) / ^{176}\text{Hf}$  (%) expresses the proportion of  $^{176}\text{Lu}$  due to  $^{176}\text{Yb} + ^{176}\text{Lu}$  versus the proportion due to  $^{176}\text{Hf}$ , in %.
4. Volts Hf is the sum of voltages of all Hf isotopes.
5.  $^{176}\text{Hf}/^{177}\text{Hf}$  is the measured  $^{176}\text{Hf}/^{177}\text{Hf}$ , corrected for fractionation and inferences. Shown with uncertainty expressed at 1-sigma.
6.  $^{176}\text{Lu}/^{177}\text{Hf}$  is the intensity of  $^{176}\text{Lu}$ , calculated from the measured intensity of  $^{175}\text{Lu}$  and  $^{176}\text{Lu}/^{175}\text{Lu}=0.02653$  (from Patchett, 1983), compared to the measured intensity of  $^{177}\text{Hf}$ . Fractionation of Lu isotopes is assumed to be the same as fractionation of Yb isotopes.
7.  $^{176}\text{Hf}/^{177}\text{Hf}$  (T) is the  $^{176}\text{Hf}/^{177}\text{Hf}$  corrected to the time of crystallization using a decay constant of  $1.867\text{e-}11$  (from Scherer et al., 2001 and Soderland et al., 2004)
8. E-Hf (0) is the present-day epsilon Hf value using  $^{176}\text{Hf}/^{177}\text{Hf}=0.282785$  and  $^{176}\text{Lu}/^{177}\text{Hf}=0.0336$  (from Bouvier et al., 2008). The uncertainty is expressed at 1-sigma.
9. E-Hf(T) is the epsilon Hf value at time of crystallization. Uncertainty is expressed at 1 sigma.
10. U-Pb ages are based on 206/238 for ages younger than  $\sim 1.0$  Ga, and on 206/207 for ages older than  $\sim 1.0$  Ga. This age cutoff may be slightly different for each sample.
11. Isotope ratios as follows:

180/177	1.8866600	Patchett (1983)
179/177	0.7325000	Patchett & Tatsumoto (1980)
178/177	1.4671800	Patchett (1983)
176/177	0.2821600	Patchett (1983)
174/177	0.0087100	Patchett (1983)
176/175	0.0265300	Patchett (1983)
176/171	0.9016910	Vervoort et al. (2004)
173/171	1.1323569	Vervoort et al. (2004)
172/171	1.5317360	Vervoort et al. (2004)

#### Notes for plots:

1. DM array is from Vervoort and Blichert-Toft (1999), using  $^{176}\text{Hf}/^{177}\text{Hf}=0.283225$  and  $^{176}\text{Lu}/^{177}\text{Hf}=0.0383$

2. CHUR is from Bouvier et al. (2008), using  $^{176}\text{Hf}/^{177}\text{Hf}=0.282785$  and  $^{176}\text{Lu}/^{177}\text{Hf}=0.0336$ .
3. Hf isotope evolution lines assume an average value of  $^{176}\text{Lu}/^{177}\text{Hf}=0.0115$  and a range of  $^{176}\text{Lu}/^{177}\text{Hf}=0.0036$  to  $^{176}\text{Lu}/^{177}\text{Hf}=0.0193$ . Values are from the average and 2-sigma range of values reported by Vervoort and Patchett (1996) and Vervoort et al. (1999).
4. Uncertainties shown at 2-sigma.
5. Uncertainty for EpsilonT is nearly identical for Epsilon 0 because of the very long half-life.

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